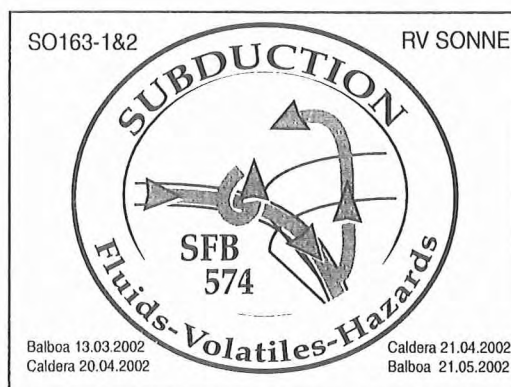




**FS/RV SONNE
FAHRTBERICHT SO163
CRUISE REPORT SO163
SUBDUCTION I**

**MULTI-SYSTEM ANALYSIS OF FLUID RECYCLING AND GEODYNAMICS
AT THE CONTINENTAL MARGIN OFF COSTA RICA**

**SO163-1
BALBOA - CALDERA, MARCH 13 - APRIL 20, 2002
SO163-2
CALDERA - BALBOA, APRIL 21 - MAY 21, 2002**



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**Edited by
Wilhelm Weinrebe and Ernst Flüh
with contributions of cruise participants**

**Investigations in the frame of SFB 574
"Volatiles and Fluids in Subduction Zones:
Climate Feedback and Trigger Mechanisms for Natural Disasters"**

Contribution No. 14 of SFB 574

GEOMAR
Forschungszentrum
für marine Geowissenschaften
der Christian-Albrechts-Universität
zu Kiel

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GEOMAR
Forschungszentrum
für marine Geowissenschaften
Wischhofstr. 1-3
D - 24148 Kiel
Tel. (0431) 600-2555, 600-2505

Editors of this issue:
Wilhelm Weinrebe and Ernst Flüh

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GEOMAR
Research Center
for Marine Geosciences
Wischhofstr. 1-3
D - 24148 Kiel
Tel. (49) 431 / 600-2555, 600-2505

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1. Summary

During cruise SO163 (13. March to 20. Mai 2002, Balboa to Balboa) the Pacific continental margin off Costa Rica was surveyed using a variety of tools. These investigations follow several previous cruises and are part of the new research initiative (Sonderforschungsbereich) SFB574: "Volatiles and Fluids in Subduction Zones: Climate Feedback and Trigger mechanisms for Natural Desasters". Several other cruises will follow in 2002 and 2003.

The cruise was split into two legs with an intermediate port call on April 20 in Caldera., Costa Rica.

Leg 1 was set about to investigate the material input, mobilisation, and return flow of volatiles in the subduction zone off Costa Rica. During leg 1 a comprehensive study of the area of investigation was carried out, applying a whole set of different mapping tools with different resolution. The results comprise an important database for all forthcoming expeditions and will be an invaluable help in finding appropriate target areas for future sampling work. The investigations during leg 1 also aimed in quantifying the processes in the material cycle which directly affect the seafloor and its morphology, such as sediment deposition, erosion, submarine landslides as well as manifestations of fluid venting. In particular deep-towed systems were used. The TOBI system of the Southampton Oceanography Centre was used for a reconnaissance survey to detect places of active mud diapirism and mass wasting. In total, this system was used for 315 hours to map an area of about 8000 km² with a resolution on the order of around 10 metres. Together with the results of the TOBI survey during the SO-144 cruise in 1999, nearly a complete coverage of the whole continental margin from Santa Elena Peninsula in the northwest to the Cocos Ridge in the southeast was achieved. In order to image parts of the surveyed area with greater resolution, the newly acquired GEOMAR DTS-1 deep-towed sidescan sonar system was used to map five key areas. In total an area of about 650 km² was mapped in 175 hours with a resolution of less than one metre. As a further step to even greater resolution and ground truthing, direct observations of the seafloor by TV-sled OFOS was carried out at 24 deployment sites. The survey was completed with sampling the seafloor by the TV-guided grab at nine positions.

During leg 2 numerous seismic and seismological experiments and a water sampling programme dominated the work carried out. The active seismic work focussed on three areas. The Megalens structure, being interpreted as a key factor for erosional processes, the decollement reflection and the BSR occurrences off Nicoya. In all three areas numerous wide angle data sets were collected using ocean bottom hydrophones (OBH) and seismometers (OBS). The Megalens reflections are not easily identified in the records processed onboard. The profiles collected on the decollement reflection show a pronounced lateral variability of the interface, and surprisingly several deep penetrating faults that cut the subducting oceanic plate have been identified on preliminarily processed sections. Off Nicoya several profiles were shot to collect three component data in areas where a strong BSR reflection had been observed before. These three component data should allow to identify shear waves and be used for receiver transfer function methods to investigate the shear wave velocity of gas hydrate bearing sediment. Numerous records were collected where p-s conversions observed at about the time expected for a conversion at the BSR.

Two seismological networks were operated. The first one comprising ten instruments was placed on and around an active vent site and was operational for about 40 days. Preliminary analysis of the data showed a special kind of events, termed hongo events, which could be localised at the vent site. These events consist of an approximate 1 sec long monochromatic (36 Hz) burst with increasing amplitudes, followed by a two second long monochromatic (20 Hz) arrival of lower frequency. These events are most probably related to the hydrologic system in the vent site. A second network was installed around Jaco Scar. This network comprises 28 stations (5 of them are actually centered on the near by Quepos mound) and shall be operational until October 2002. It is

further augmented by an 15 element land array. In addition, two tiltmeters were deployed in this network to monitor long term slope changes.

In total, the seismic programme comprised 212 OBH/S deployments, with 184 successful recoveries and 28 instruments left on the seafloor for monitoring. More than 800 nm of seismic profiles were collected, with a total of about 25000 airgun shots fired.

The water column sampling program comprised a total of 24 stations and had two main objectives. (a) The investigation of the general oxygen and methane distribution off Costa Rica. Two sections with 6 CTD/rosette stations spread over 100 nm were sampled in the northern and southern part of the working area. The results will provide information about the production of methane within the upper 600m of the water column due to the geostrophical upwelling along the western coast of Middle America (Costa Rica Dome), and will define the methane background for the detection of venting-related methane-rich plumes in bottom-near waters. (b) Detailed surveys of the methane distribution at potential vent sites in order to quantify the release of CH₄ to the water column and to define targets for a detailed geochemical program during expeditions later this year. Multiple CTD casts were deployed in the Jaco Scarp area as well as at Quepos Mound and Mound Culebra. The results revealed evidence for active venting of methane rich fluids in all 3 areas. The water column program was accompanied by a small sediment sampling program, which lead to first data on the geochemical composition of the non-vent-influenced upper sediment in the SFB target area.

The Simrad system of Sonne was operated continuously, and magnetic data were collected along 4 profiles extending over more than 200nm.

2. Introduction

2.1 Objectives of the cruise SO-163/1&2

SO-163/1 continues the series of cruises with RV SONNE to the Pacific convergent continental margin off Costa Rica and Nicaragua, which started with SO-76 in 1991. That cruise was followed by SO-81 in 1993, SO-107 in 1996 and SO-144 with 6 legs in 1999. Moreover, this area was also investigated by several cruises of other institutions with different research vessels (RV EWING, ROGER REVELLE), and in addition, probed by drilling with JOIDES RESOLUTION. Furthermore, more cruises with RV SONNE and RV METEOR are already scheduled for 2002 and 2003. Why is this section of the margins of the Pacific plates so in the focus of current geoscientific research?

The segment of the Central American continental margin off Costa Rica and Nicaragua provide along a relatively short distance a spectrum of quite different subduction styles and scenarios. The differences are: Age and source of subducted oceanic crusts, subduction angle, depth of melting zone and earthquake focal depths. Characteristic features are also the seamount province parallel to the Cocos Ridge, which is subducted under Costa Rica, or the heavily fragmented plate segment cut by numerous normal faults, subducted off Nicaragua. As such, the segment off Costa Rica and Nicaragua is a natural laboratory to investigate causes, consequences, and relationships between these different parameters.

Volatiles and fluids play a major role in the global geological material budget. At subduction zones, material from the oceanic plate (the oceanic crust and its overlying sediments) is brought to great depth. Sediments from the continental margin transported downslope by mass wasting, add to the subducted material. These processes extract considerable amounts of carbon, sulphur, water, and halogenes out of the exosphere. Inside the subduction zone, the downgoing material is remobilized, fractionated, and forms new reservoirs with different volatile compositions. A portion of the subducted and remobilized material is partly fed back into the exosphere through different transport pathways and mechanisms, partly subducted into the lower mantle. The return flow is accomplished through three distinct pathways:

- fluid venting at the deformation front
- mud diapirism and gas hydrate dynamics at the margin
- magmatism and exhalations at the volcanic arc

The long-term development of the Earth's climate, the geochemical evolution of the hydrosphere and atmosphere, and the causes of natural disasters (earthquakes, volcanic eruptions, submarine landslides, tsunamis) are all connected with the return flow and impact of volatiles and fluids from subduction zones. To understand and quantify these processes and relationships is the main objective of the new research initiative (Sonderforschungsbereich) SFB 574: "Volatiles and Fluids in Subduction Zones: Climate Feedback and Trigger Mechanisms for Natural Disasters". SO-163 is the first cruise carried out in the framework of this SFB.

The SFB 574 addresses its objectives in three overarching themes:

- A. Material input and tectonic behaviour during plate subduction
- B. Formation and role of gas hydrates and mud diapirs as exotic volatile reservoirs
- C. Devolatilization through magma and metamorphic processes

This first cruise in the framework of the SFB 574 focuses on the processes, which directly affect the seafloor by changing its topography, nature, and compositional characteristics, such as:

- deposition of sediments
- erosion
- downslope transport by mass wasting
- gas and fluid expulsion due to tectonic activity and dewatering of the lower slope wedge.

Because of convergent movements of several centimetres per year and resulting erosional and accretional processes, subduction zones are dynamic areas that are subject to rapid changes. Among other parameters, the kind of mechanic coupling between upper and lower plate, trench fill volume, geological history and the formation and decomposition of gas hydrates all play a major role. Convergent plate boundaries are areas of the earth's greatest topographic relief and are characterised by their morphological variation.

Due to rapid accumulation of sediments high fluid pressure build up within the tectonic compression setting of a subduction zone. The mixture of materials, e.g. pelagic mud with high water content and silt/sand containing less water but being more permeable, allows an overpressurization of the sediment water mixture. For the pressure to be released, fluid pathways will open towards the surface and fluids and volatiles move upwards. On the seafloor, manifestations of fluid venting should appear.

At convergent margins, gravitational downslope mass transport takes place in varying dimensions. Submarine landslides occur as large megaslumps of several hundreds to thousands of km³, but also as a sequence of individual slumps, generated by oversteepened submarine slopes. Major slides are often attributed to high sedimentation rates leading to overpressuring. Earthquakes are considered the likely immediate triggers for failure along the continental margins, yet compelling evidence for these causal factors is inconclusive. In the Costa Rica area, special downslope transport is associated to seamount subduction (Ranero and von Huene, 2000; von Huene et al., 2000).

The formation of gas hydrates may increase the shear strength of sediments immediately below the seafloor, and thus may play an important role in stabilizing the continental margins (Mienert et al., 1998). Decomposition of gas hydrates consequently results in weak zones, where failure could be triggered by gravitational loading or seismic disturbances and hence submarine landslides would result. The slides may generate large tsunamis, possibly larger than those generated by earthquakes alone. Numerous large landslides occur in areas where gas hydrates have been documented by geophysical investigations. Furthermore, there is abundant evidence for the presence of fluid and gas escape structures in these areas. Therefore, a causal relationship between gas hydrates and slides has been proposed, but direct proof is missing.

All these processes shape and change the seafloor. Cruise SO-163/1 is set about to map in detail the surface of the continental margin off Costa Rica in a multiscale approach, applying a whole set of different mapping tools with different resolution. High resolution surveys will aim to identify local features such as tectonic structures, slides, and fluid-venting related manifestations on the seafloor. To assist interpretation, direct observations and detailed optical imaging of the seafloor by TV-sled are necessary. Simultaneously, the results of this investigations constitute the most important basis for sample site selections for most other subprojects of the SFB 574.

In the following leg, SO-163/2, a variety of geophysical, mainly seismic experiments are planned, in addition to the installation of a temporary (6 month) local seismic network, with is further augmented by an onshore array. These experiments focus on target reflections that can help to better understand the erosion process along the margin, the nature of vent sites and gashydrates.

A second topic of high priority during this leg are water column and upper sediment core sampling. Fluid flow associated with subduction leads to the upward transport of reduced, energy-rich compounds, which can be either consumed within the upper sediment column and at the sediment water interface, or be injected into the water column. Both processes are complimentary and need to be assessed to close the carbon budget at seep sites. In addition, methane plumes within the lower water column are an unquestionable sign for active venting, and the results will help to define targets for the extensive geochemical field work planned for this summer.

2.2 Tectonic Setting and Previous Studies

(César R. Ranero)

The Pacific convergent margin of Middle America, particularly along Costa Rica and Nicaragua has become a classical example of forearc tectonics driven by the variable character of the subducting ocean plate [Corrigan and Mann, 1990; Fisher et al., 1998; Ranero and von Huene, 2000; von Huene et al., 1995; von Huene et al., 2000]. The subducting ocean plate has a variable character because it was formed in two different spreading centers (East Pacific Rise and Cocos-Nazca spreading center) and has also been modified by the volcanism of the Galapagos hot spot (Barckhausen et al., 2001). The Galapagos hot spot has created the Cocos Ridge and flanking seamounts (Werner et al., 1999) and has thickened the preexisting crust in some areas. The different types of ocean plate are well displayed in the bathymetry. Opposite SE Costa Rica the Cocos Ridge stands 2.5 km shallower than the surrounding seafloor due to a ~ 20 km thick crust. Offshore Central Costa Rica subducts a seamount province with large conical seamounts and elongated ridges (e.g. Quepos Ridge). Facing Nicoya Peninsula there is a smooth ocean plate. A small linear ridge that subducts offshore middle Nicoya Peninsula separates the previous 3 areas formed at the Cocos-Nazca spreading center from the lithosphere located to the NW formed at the East Pacific Rise. The plate formed at the East Pacific Rise exhibits a rather different morphology, characterised by pervasive normal faulting that forms a rough terrain approaching the trench.

The differences in character of the ocean plate are probably caused by a combination of variations in crustal thickness and age of the subducting plate and lead to dramatic differences along the trench in the behaviour of the incoming plate entering the subduction zone. Other subduction parameters like convergence rate or sediment input vary little along the research area and thus the effects on forearc tectonics of subducting segmented plate can be isolated.

The subduction of the buoyant Cocos Ridge shows little bending into the subduction zone. The underthrust portion of the Cocos Ridge has uplifted Osa Peninsula at the front of the margin and also the volcanic arc forming the Talamanca Cordillera (deBoer et al., 1995). In that area, subduction related volcanism ceased at ~7 Ma, perhaps associated to the arrival of Cocos Ridge type of crust, but the age of the impact of Cocos Ridge in to the margin of Costa Rica is a matter of debate, ranging from ~ 1 Ma (Lonsdale & Klitgord, 1978) to ~ 5 Ma (Graeffe, 1998). Offshore central Costa Rica, the subduction of the large conical seamounts and ridges has produced large embayments in the continental slope (von Huene et al., 2000). The subduction of the smooth ocean plate offshore Nicoya Peninsula is paralleled by the most stable morphology of the continental slope. The subduction of the lithosphere formed at the East Pacific Rise, located to the NW, and pervasively faulted is related to the deepest trench and a steep continental slope.

The nature of the rocks units forming the overriding upper plate in the area has been a matter of considerable recent research and the concepts of the origin and structure of onshore and offshore rock bodies have dramatically changed in the last decade.

Onshore, the basic igneous complexes cropping out along much of Pacific Costa Rica were previously interpreted as ophiolites resulting from accretion of several oceanic crust nappes. However, recent studies show that they are part of the large Caribbean Flood Basalt province. This province formed possibly at the Galapagos Hot Spot, between 74 and 94 Ma, intruding older lithosphere (Hauff et al., 1997; Hauff, 1998; Sinton et al., 1997; Sinton et al., 1998). Also, those studies find that some of the smaller igneous rocks bodies of oceanic origin cropping out along Middle America are younger accreted seamounts and aseismic ridges.

Offshore, near-vertical and wide-angle seismic data (Ye et al., 1996; Ranero and von Huene, 2000; Walther et al., 2000) indicate a basic structural configuration of the frontal ~60 km of the continental margin from the Costa Rica to Nicaragua. It consists of a wedge shaped body of relatively high velocity rock comprising the bulk of the continental margin (the so-called margin wedge) covered by 1-2 km of slope sediment, underthrust by trench sediment, and fronted by a

small (3-10 km wide) sediment prism. The origin of the margin wedge has been long debated. Some data sets were interpreted as representing a large, active accretionary prism that grows through frontal accretion and underplating of the sediment overlying the ocean plate (e.g. Silver et al., 1985; Moore and Shipley, 1988; Shipley et al., 1992). Alternatively, it was also interpreted as formed by the offshore extension of the igneous complexes observed onshore (von Huene and Flueh., 1994; Ye et al., 1996). Drilling offshore Nicoya Peninsula, during ODP leg 170 demonstrated that basically all sediment of the ocean plate is underthrust past the frontal sediment prism and that accretion is not an important process at the margin (Kimura et al., 1997). The frontal sediment prism is formed by reworked slope sediment at the location of the ODP 170, although may locally involved small amount of accretion elsewhere along the margin (von Huene et al., 2000). In addition, drilling during leg 170 also supported that the bulk of the margin is formed by the seaward extension of the igneous complexes outcropping along Pacific Costa Rica.

More recently geophysical data sets along the Nicaraguan (Ranero et al., 2000a) and Costa Rican (Ranero and von Huene, 2000; von Huene et al., 2000) segments of the margin have shown evidence for widespread extension and long-term subsidence of the margin that can best be explained as produced by subduction erosional processes. Further independent evidence has come from the study of samples from the unconformity separating the margin wedge from the slope sediment collected during ODP 170 drilling. The evidence for erosion comes from the identification of beach rock overlying the unconformity (Vannucchi et al, 2001). The unconformity, cut by wave surf, is currently ~ 4 km deep at the lower slope and has subsided since early-middle Miocene time implying a continental slope retreat of ~50 km.

Submersible and deep towed camera investigations have revealed a widespread cold venting system along the margin. The main evidence for vents is the presence of chemoautotrophic communities and carbonate mounds observed at several localities (Bohrmann et al, 2002, Kahn et al, 1996) and mapped along much of the margin with swath bathymetry, sidescan data and seismic reflection images (Shipley et al., 1990; Shipley et al., 1992; Bialas et al., 1999; Ranero et al., 2000b). The areas where venting might be the most vigorous are areas of rapid tectonic activity related to seamount subduction (headwalls of slides and uplifted regions) and at the carbonate mounds (Bohrmann et al, 2002). Submersible investigations on the lower continental slope indicate non-localized diffuse flow to be the dominant process there (McAdoo et al, 1996).

Samples from Leg 170 from the decollement yielded freshened water containing thermogenic methane. Samples from the upper plate sediment revealed fresh water from dissociated gas hydrate in addition to freshened water with thermogenic methane (Silver et al., 2000). A different hydrological system seems to exist beneath the decollement. There, water has a near seawater salinity. Measured temperatures are insufficient generate mineral dehydration and thermogenic methane, implying a fluid flow from ~ 15 km depth in the subduction zone (Kimura et al, 1997).

On the incoming ocean plate heat flux measurements near the trench axis offshore Nicoya Peninsula yielded very low values that were interpreted as indication of fluid flow into the ocean basement (Langseth and Silver, 1996). More recently, new measurements at several locations of the ocean plate opposite Nicoya Peninsula, have confirmed previous results near the trench (Fisher et al., 2001). Also, the measurements collected in areas away from the trench show a complex pattern that suggests active fluid flow into the ocean igneous crust. The robust evidence for fluid flow is very significant because those measurements have been obtained in the area where the ocean plate displays the least deformation associated to the bending approaching the trench. The segments of the ocean plate located to the NW and to the SE, where faulting is more pervasively developed and fault offsets are larger, leading to the exposure of ocean basement to the seawater, awaits as an exciting area for further investigations.

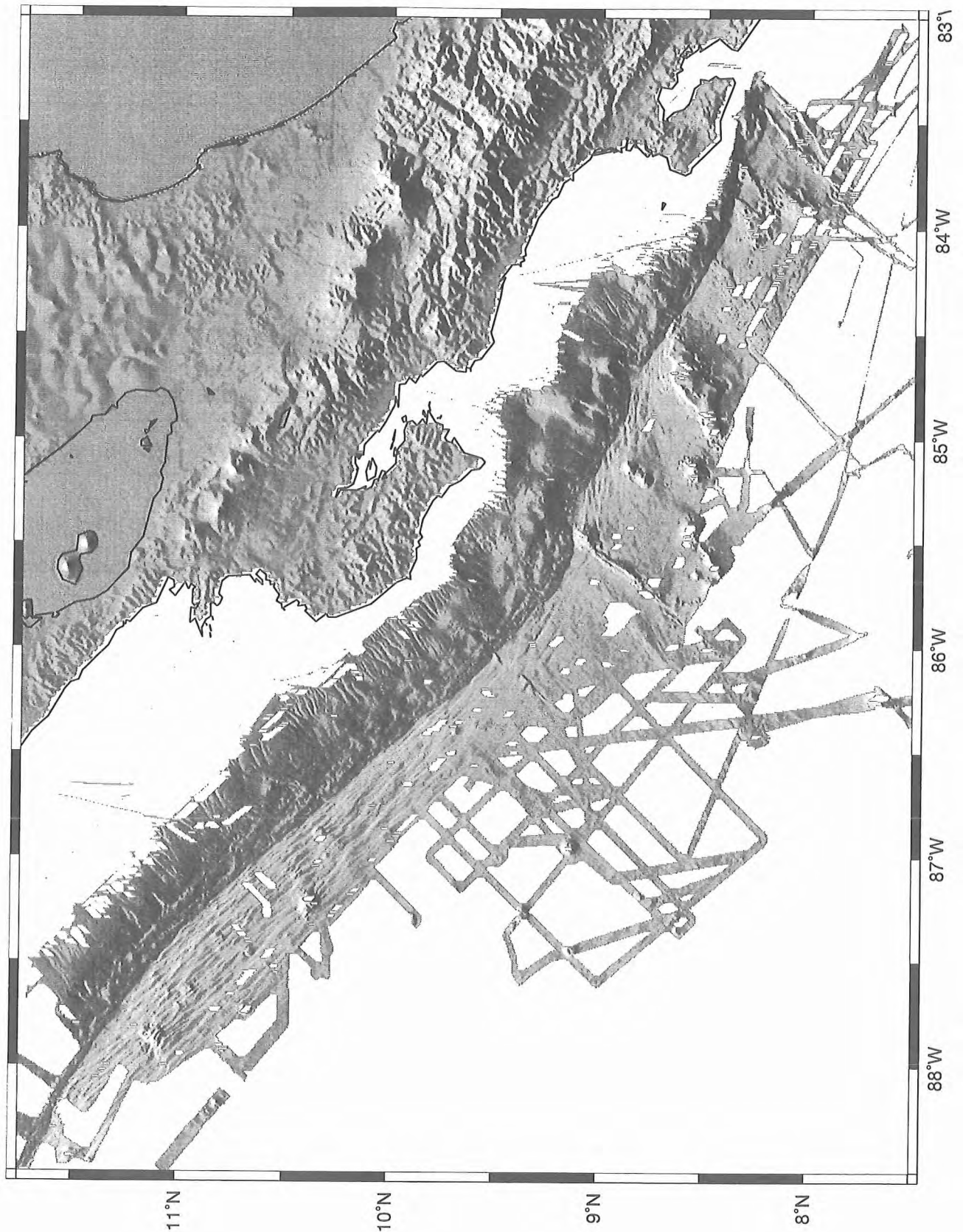


Figure 2.2.1 Bathymetric compilation along the Middle American trench showing the different morphological provinces of the ocean plate and the continental slope (SONNE 76, 81, 107, 144, 150, 163-1 and EWING 0005, 0104 cruises).

3. Participants

3.1 Scientists

3.1.1 Scientists - SO 163-Leg 1

Dr. Wilhelm Weinrebe	GEOMAR, chief scientist
Frank Eichelhardt	SFB 574
Rolf Hamann	SFB 574
Christian Hartmann	SFB 574
Dr. Caroline Huguen	SFB 574
Veit Hühnerbach	SOC
Dr. Ingo Klaucke	GEOMAR
Dr. Douglas Masson	SOC
Duncan Matthew	SOC
Thomas Nadler	SFB 574
Dr. Cesar Ranero	GEOMAR
Ian Rouse	SOC
Dr. Heiko Sahling	SFB 574
Thorsten Schott	Octopus
Dirk Schwerdtfeger	SFB 574
Peter Oliver Thierer	GEOMAR

3.1.2 Scientists - SO 163-Leg 2

Prof. Dr. Ernst R. Flüh	GEOMAR, chief scientist
Ivonne Arroyo	ICE
Arnim Berhorst	GEOMAR
Ileana Boschini	ICE
Guido Calvo	ICE
Denise De Nil	SFB 574
Noemi Fekete	SFB 574
Dr. Jürgen Goßler	SFB 574
Dr. Sibylle Grandel	SFB 574
Rolf Hamann	SFB 574
Dr. Heidrun Kopp	GEOMAR
Susan Mau	SFB 574
Virginie Pierre	GEOMAR
Lars Planert	GEOMAR
Dr. Gregor Rehder	GEOMAR
Michael Schnabel	SFB 574
Karen Stange	SFB 574
Klaus-Peter Steffen	KUM
Peter Oliver Thierer	GEOMAR
Tanja Woronowicz	GEOMAR

3.2 Crew

3.2.1 Crew - SO 163-Leg1

Henning Papenhagen	Master
Jörn Löffler	Chief Mate
Walter Bascheck	1st Mate
Frank Göldner	Radio Officer
Anke Walther	Surgeon
Volker Hartig	Chief Engineer
Norman Lindhorst	2nd Engineer
Andreas Rex	2nd Engineer

Uwe Rieper
 Hilmar Hoffmann
 Jörg Leppin
 Paul Wintersteller
 Matthias Großmann
 Herrmann Rademacher
 Christian Kunze
 Rene Brendel
 Heinrich Riedler
 Rudolf Prammer
 Volkhard Falk
 Andreas Wege
 Anja Baumgärtel
 Justine Hasler
 Karl-Heinz Lohmüller
 Ulrich-Bruno Hampel
 Andreas Woltmann
 Peter Rosin
 Norbert Bosselmann
 Dirk Schachel
 Karsten Bosselmann

Electrician
 Chief Electronic Engineer
 Electronic Engineer
 System Operator
 System Operator
 Fitter
 Motorman
 Motorman
 Motorman
 Chief Cook
 2nd Cook
 Chief Steward
 2nd Steward
 2nd Steward
 Boatswain
 A. B.
 A. B.
 A. B.
 A. B.
 A. B.
 A. B.

3.2.2 Crew - SO 163-Leg 2

Hartmut Andresen
 Jörn Löffler
 Nils-Arne Aden
 Frank Göldner
 Anke Walther
 Volker Hartig
 Normann Lindhorst
 Andreas Rex
 Thorsten Dammann
 Jörg Leppin
 Kurt Stammer
 Paul Wintersteller
 Matthias Großmann
 Jochen Stenzler
 Rene Brendel
 Christian Kunze
 Klaus Parschow
 Heinrich Riedler
 Rudolf Prammer
 Volkhard Falk
 Andreas Wege
 Justine Hasler
 Anja Baumgärtel
 Norbert Bosselmann
 Dirk Schachel
 Torsten Bierstedt
 Andreas Woltmann
 Karl-Michael Braasch
 Jörn Dettmer
 Peter Rosin

Master
 Chief Mate
 1st Mate
 Radio Officer
 Surgeon
 Chief Engineer
 2nd Engineer
 2nd Engineer
 Electrician
 Chief Electronic Engineer
 Electronic Engineer
 System Operator
 System Operator
 Fitter
 Motorman
 Motorman
 Motorman
 Motorman
 Chief Cook
 2nd Cook
 Chief Steward
 2nd Steward
 2nd Steward
 Boatswain
 A.B.
 A.B.
 A.B.
 A.B.
 A.B.
 A.B.

3.3 Addresses of Participating Institutions

- GEOMAR:** GEOMAR Forschungszentrum für marine Geowissenschaften
der Christian-Albrechts-Universität zu Kiel
Wischhofstr. 1-3
24148 Kiel
Germany
Tel.: +49 - 431 - 600 - 2972
Fax: +49 - 431 - 600 - 2922
eMail: Iname@geomar.de
Internet: www.geomar.de
- SFB 574:** Sonderforschungsbereich 574
Christian-Albrechts-Universität zu Kiel
c/o GEOMAR Forschungszentrum
Wischhofstr. 1-3
24148 Kiel
Germany
Tel.: +49 – 431 – 600 - 2972
Fax: +49 – 431 – 600 – 2922
eMail: Iname@geomar.de
Internet: www.geomar.de
- SOC:** Southampton Oceanography Centre
Empress Dock
Dock Gate 4
Southampton, SO14 3ZH
United Kingdom
Tel.: +43-1703-596568
Fax.: +43-1703-596554
eMail: d.masson@soc.soton.ac.uk
- Oktopus:** OKTOPUS
Gesellschaft für angewandte Wissenschaft
Innovative Technologien und Service in der Meerestechnik mbH
Wischhofstr. 1-3, Geb. D5
24148 Kiel
Germany
Tel.: +49 - 431 - 7209 - 350
Fax: +49 - 431 - 7209 - 356
- ICE:** Insituto Costaricense de Electricidad
C.S. Exploratió Subterránea, Sabana Norte
Apartado 10032 San José, Costa Rica
Tel.: 00506-220-6394
00506-695-6522
Fax: 00506-220-8212
e-mail: igarroyo@cariari.ucr.ac.cr
- KUM:** K.U.M. Umwelt- und Meerestechnik Kiel GmbH
Wischhofstraße 1-3, Geb. D5
24148 Kiel, Germany
Tel.: 0049-431-7209-220
Fax: 0049-431-7209-244
e-mail: KUM.Umweltsmeersestechnik@t-online.de



Figure 3.1 Cruise participants of SO-163/1

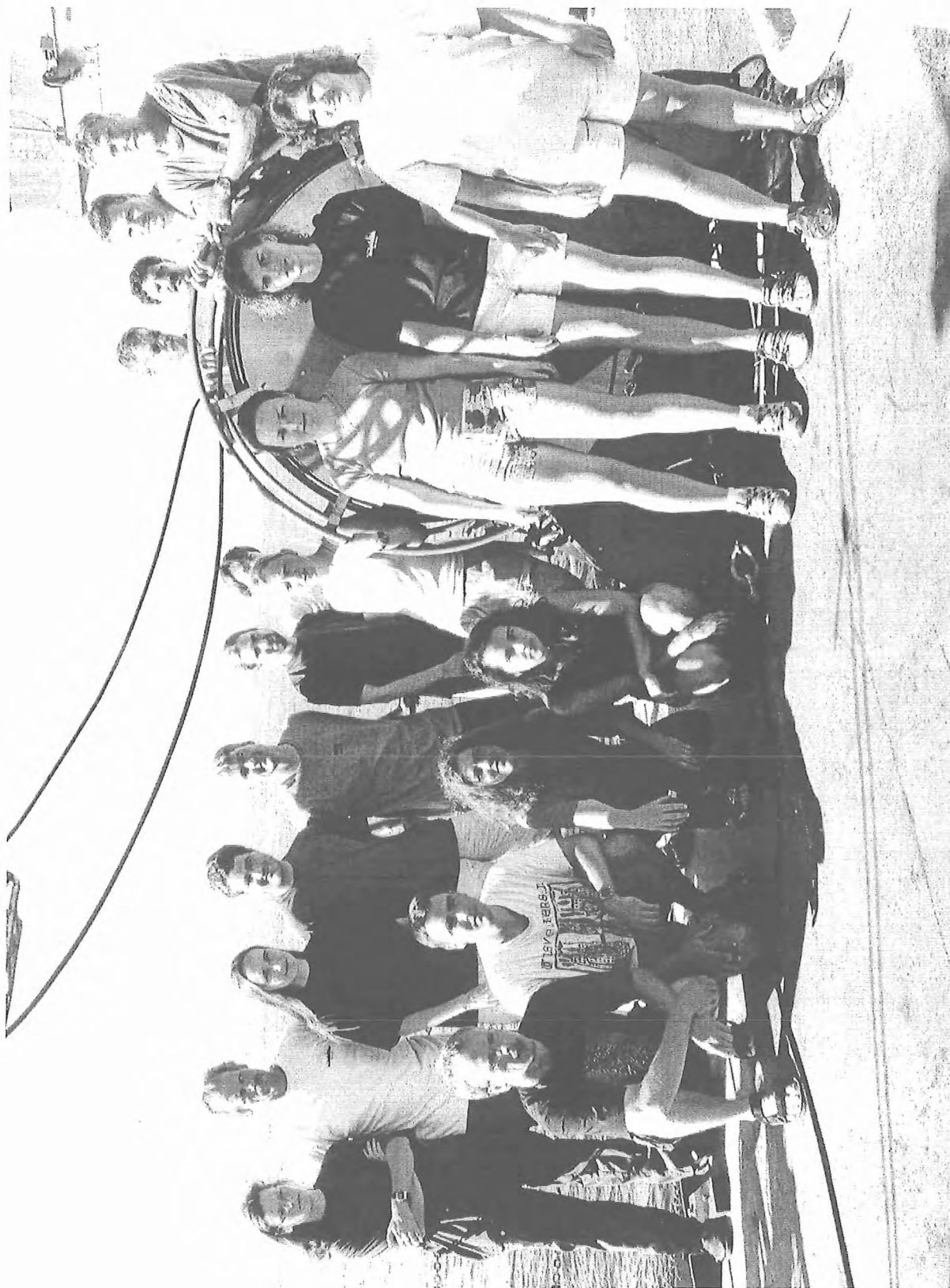


Figure 3.1.2.: Scientists of SO 163-2

4.1 Agenda of the cruise SO163-1

(W. Weinrebe)

Cruise SO 163 „SUBDUCTION I“ started on March 13, 2002, in Balboa, Panama. 13 scientists embarked on RV SONNE in Balboa, whereas 3 participants of cruise SO 162 stayed onboard and completed the international group of scientists from Great Britain, France, Spain, and Germany. Although most of the scientific equipment was already set up onboard since the previous cruise, additional work was necessary to load and install the TOBI system of our colleagues from Southampton.

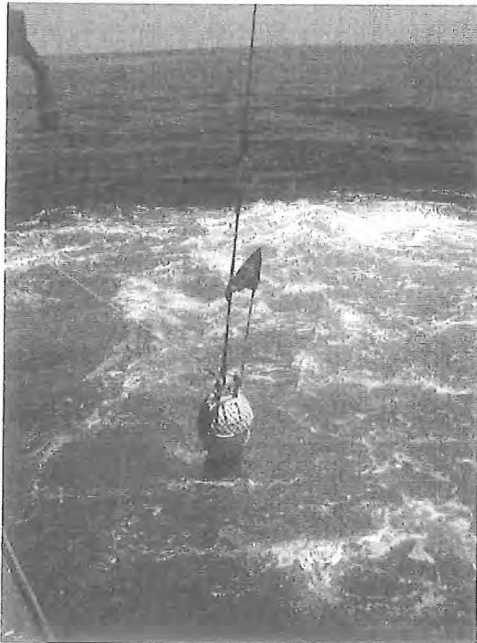


Figure 4.1.1 Deployment of OBH

At 11:00 in the morning of March 14 SONNE left the port of Balboa and steamed towards the research area offshore Pacific Costa Rica. The first target was an area west of Nicoya Peninsula, where several mud volcanoes had been found on previous cruises. It was planned to deploy a seismological network comprising 10 ocean bottom hydrophones (OBH) on top and around a mud volcano. These OBH will be recovered during the subsequent leg. During the transit to the working area the Simrad multibeam system was used to map several small areas in order to fill gaps in the existing bathymetric coverage.

Before starting to deploy the OBHs, a releaser test was carried out successfully in a water depth of 1,750 m. The deployment of the 10 OBHs was accomplished in 4 hours in the morning of March 16. After running a CTD down to a depth of 1,900 m to determine the water sound velocity profile, the calibration procedure for the new POSIDONIA USBL-system was started in the afternoon. Unfortunately this calibration failed due to a system error. As the first launch of TOBI was determined to take place in daylight, an additional mapping with Simrad EM120 multibeam, Parasound, and a magnetometer was carried out during the night.

TOBI with a weight of 2 tons as well as the depressor of around 650 kg were launched smoothly in 45 minutes in the morning of March 17. For this first deployment 8 profiles each of 40 nm in length and spaced 5 km apart were planned, adding up to a total time of around 6 days. First data from the upper continental margin showed many incised, probably active canyons and gullies.

Unfortunately, due to a faulty connection between umbilical- and deep sea cable, data transfer was interrupted. The first profile was abandoned incomplete and TOBI was recovered in daylight in the morning of March 18. The time span for repairs was used for additional bathymetric and magnetic mapping, before TOBI could be relaunched again successfully in the afternoon of the same day.

TOBI started surveying at the beginning of the second profile and operated well for more than three and a half days. Many steeply incised canyons and gullies could be seen. The upper slope is dominantly characterised by mass wasting processes. Very few circular features, interpreted as mud diapirs, were imaged, though many more of these structures show up in the bathymetry data. Obviously, just these few still seem to be active, as derived by their higher backscattering strength.



Figure 4.1.2 Deployment of TOBI

As on the lower slope no more manifestations of fluid venting were detected, TOBI was recovered on the morning of March 22 before finishing profile 6. It was decided to investigate the most prominent mud volcano seen so far at 10°17.85' N, 86°18.36' W with the TV-sled OFOS. Due to strong winds and currents OFOS runs could only be carried out in a distinct direction. Two missions with OFOS in the afternoon of the same day revealed considerable amount of carbonate crusts and vent fauna (vesicomylid clams and pogonophora tube worms). Thus it seemed worth to sample this location with a TV-grab. Two deployments were carried out. Both grabs felt over during the closing of the jaws, however, we retrieved some pieces of rocks, sediment with sulfide odour and shell fragments. The calcareous cemented sediments are probably formed in the course of methane venting.

Mapping work was continued southwards with an additional survey with TOBI from March 23 to March 25. This survey area included the characteristic mud volcano field seaward of Central Nicoya Peninsula locally studied during cruise SO-144 in 1999. The work during the following days was concentrated in that area.

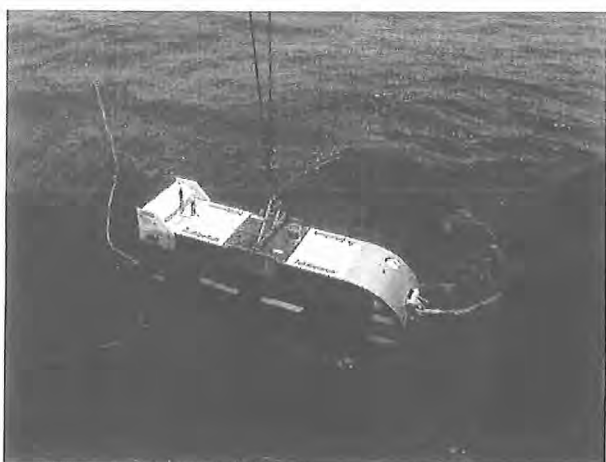


Figure 4.1.3 Deployment of DTS-1

A deployment of the DTS-1 deep-towed side-scan sonar system followed the TOBI survey. This system operates with higher frequencies (75 kHz, 410 kHz) than TOBI (30 kHz), and is therefore capable of mapping the seafloor with a much higher resolution but also with a considerably smaller range. The two systems are complementary.

After the new USBL underwater positioning system POSIDONIA had been calibrated, DTS-1 was launched in the evening of March 25 for a 36 hours survey around the mud volcano area at 9°36' N, 85°50' W. Later on, this survey was completed by five missions with the TV-sled OFOS. in order to acquire more detailed information.

So far, target areas for OFOS deployments concentrated on features interpreted as mud-volcanoes. These structures are small circular features, several hundred meters in diameter, some showing a central crater, which can be well detected in the bathymetry, and display high backscatter in TOBI images. During the OFOS surveys it became evident that these structures are actually related to fluid venting. At the top or the flanks of the mounds we found massive chemoherm-like carbonates, and at some of these structures, vesicomylid clams and pogonophoran tubeworms. However, we found no geological evidence for active fluid expulsion.

The next working area was located on the oceanic plate, whereas all other targets so far were situated on the continental plate. A recent cruise with RV EWING found an area with local anomalously high heat flux. This area occurs at the inception of the trench parallel flexural faults created by the bending of the oceanic lithosphere into the subduction zone. The anomalous heat flux and the faulting point towards fluid flow in the area, and we looked for venting activity and specific fauna. Due to the limited length of the deep-sea cable on RV SONNE (around 7,000 m), the success of a survey of this area with the DTS-1 is at least highly questionable. An operation of the deep-towed vehicle at a depth of 3,300 m requires a longer cable. Based on experience, a cable length 2.5 to 3.5 times the towing depth is necessary. Therefore we used the Simrad EM120 multibeam system for a detailed survey switching the beams focused in the center, and the ship's speed adjusted in order to obtain a dense coverage along-track. We processed and displayed the „sidescan“ information of the system to obtain the backscattering properties. The area of high heat-flux is characterised by a gentle high and marked by higher backscatter. However, two missions with OFOS failed to find typical vent fauna.

Following this survey towards the end of the third week, an area just downslope of the headwall of the prominent Nicoya Slide was surveyed with DTS-1. The TOBI images of this region, acquired during SO-144 cruise in 1999, show spots of high backscatter not justified by seafloor topography. This area was mapped with DTS-1 75 kHz side-scan and subbottom profiler on five parallel lines.

Work continued on April 2 offshore Central Costa Rica in the area of the upper plate opposite the seamount province of the oceanic plate. The margin here is sculptured by the furrows of seamounts subducting with the lower plate. First target was the Christmas Ridge. OFOS observations revealed massive outcropping carbonates at the whole top area of this local high (1 mile in diameter). Neither at the top nor at the western flanks we found sufficient recent venting activity that would have justified further investigations.

The next target was the uplift landward of Jaco Scarp, caused by a subducting seamount. The top of the structure, the area of highest uplift, is characterised by numerous faults. We ran two OFOS missions across these areas characterised by high backscatter in TOBI images. OFOS 10 revealed a seafloor either paved by carbonates, or covered by fragmented and donut-shaped carbonates. In the whole area we observed scattered bivalve shells. OFOS 11 verified that most of the lineated high backscatter areas are caused by faults indicated by small depressions with depths on the order of 5 to 15 m. Only at the rim of the depressions talus boulders were found. The top of the uplifted area is structured by massive chemoherm-like carbonates. A dense field of vesicomylid clams were observed at a place, where the massive chemoherm complex were probably cut by fractures.

Generally, all OFOS deployments had to cope with relatively strong currents. Planning the tracks often had to follow operational rather than geological requirements.

Starting on April 3 in the morning, we continued to survey the uplifted and faulted area upslope Jaco Scarp with the side-scan sonar and the subbottom profiler of the DTS-1. Gaining more and more experience after several launch and recovery operations, the deployment was carried out smoothly in less than 15 minutes. The whole uplifted area with the pronounced fault pattern was imaged with 8 parallel profiles each around 10 miles in length. The northeastern flank of the uplifted area, showing a steady slope in the bathymetric data, turned out to be a series of rather steep, staircase like precipices. As the DTS-1 vehicle is towed at a height of just 100 m above the seafloor, steaming uphill required total attention of the DTS-1 watchkeepers. The survey across the uphill area of Jaco Scarp was completed with a track operating the DTS-1 system in the high-frequency mode. The faults and cracks on top of the structure, and the headwall of Jaco Scarp were imaged in great detail.



Figure 4.1.4 Deployment of OFOS

Work continued in the morning of April 5 with a TV-grab in order to sample donut-shaped carbonates observed by OFOS 10 two days ago. The grab felt over again, yet it yielded eroded carbonate chimneys and calcitic cemented sediments forming partly rounded boulders. In addition, more than 90 specimen of the seep-typical bivalve *Acharax* (*Solemyidae*) were extracted from the sediment indicating a weak venting activity at this site.

The mosaic of the processed TOBI data recorded during SO-144 in 1999 show a small area with striking high backscatter between Jaco und Parrita Scarp. This structure was investigated with an OFOS mission on

the afternoon of the same day. Beside very infrequent clamshells in soft sediment along the whole track we observed no changes in the sediment texture or lithology which could explain the high backscatter patterns.

In the evening of the same day, following the OFOS survey, TOBI was again launched for a more than 4 days mapping of the lower part of the margin and the deformation front. Three parallel profiles, each about 80 miles in length, extend the investigated area southward until Osa Peninsula. The structures of the middle slope and the frontal prism were mapped in great detail. In the area between Quepos Plateau and Cocos Ridge, we observed several small indentations of the deformation front, scars of seamounts currently impacting the margin. Close to the western crest of the Cocos Ridge a circular cone showing extremely high backscatter, surrounded by several concentric rings, was imaged.

The TOBI survey stopped on wednesday, April 10, near Cocos Ridge in order to run several OFOS missions in between wednesday morning and thursday afternoon, aimed at ground truthing the structures with high backscatter in the TOBI images.

In the afternoon of April 11, TOBI was relaunched again for a mission to map the area landward of Quepos Plateau up to the upper slope. This area, characterised by the large Quepos Slide, was surveyed during cruise SO-144 with an OFOS (OFOS-4). The observations then showed considerable indications for fluid venting. The 40 hours TOBI survey mapped this part completely. TOBI data showed several mounds with high backscatter as well as small bands of slightly increased backscatter, downwards of these mounds. TOBI was recovered on saturday, April 14 in the late afternoon.

During the last week of the cruise, there were some highlights as a reward for the efforts made in the course of the preceding weeks. The results obtained from OFOS, used twice in the area of the Quepos-slide, confirmed that bacteria mats cover a large surface at the basis of the landslide edge, as had been observed on cruise SO-144. Using the TV-guided grab, samples were taken from this area. Although no bacteria mat was found in the TV-G, the sediment, containing large amounts of sulphide, allows the conclusion that anaerobic methane oxidation is the predominant geochemical process.

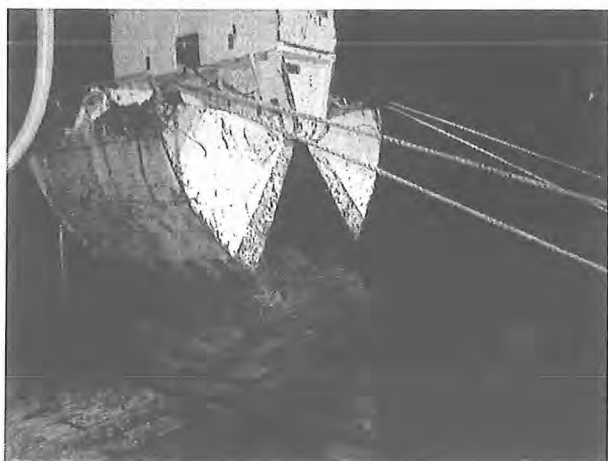


Figure 4.1.4 TV-grab on deck

Work continued on April 15 and 16 with a deployment of the DTS-1 sidescan sonar above Parrita Scarp. The faults caused by the uplift could be mapped well. Then OFOS was again deployed several times north of the Parrita Scarp. A small mound of about 30 m height showed clear signs of active fluid-venting. Here, the same species occur as had been found previously, however, throughout a considerably larger surface. In an area of about 150 m carbonates occur massively. Gaps between carbonate surfaces are covered with shells and pogonophoran tubeworms. This gigantic seep area is excellently suitable for future investigations. On Wednesday, we spontaneously decided to operate DTS-1 for high-resolution mapping in 410 kHz mode to record the aerial extent of this

structure. Due to the limited time only one profile could be recorded since a ship's turn to obtain another profile would have taken too long. Owing to the joint efforts of the work groups and the Captain's bridge the mound could be measured successfully.

Work continued during April 17 and 18 with several OFOS and TV-grab operations at Jaco Scarp. Bathymetric measurements with the Simrad EM-120 multibeam system on the passage to Caldera terminated the scientific programme at 2.00 h a.m. of 19 April. In the morning of 19 April at 09:15 h RV SONNE moored at pier 3 in port of Caldera, Costa Rica. Thus, the first leg of cruise SO-163 ended after 37 days at sea.

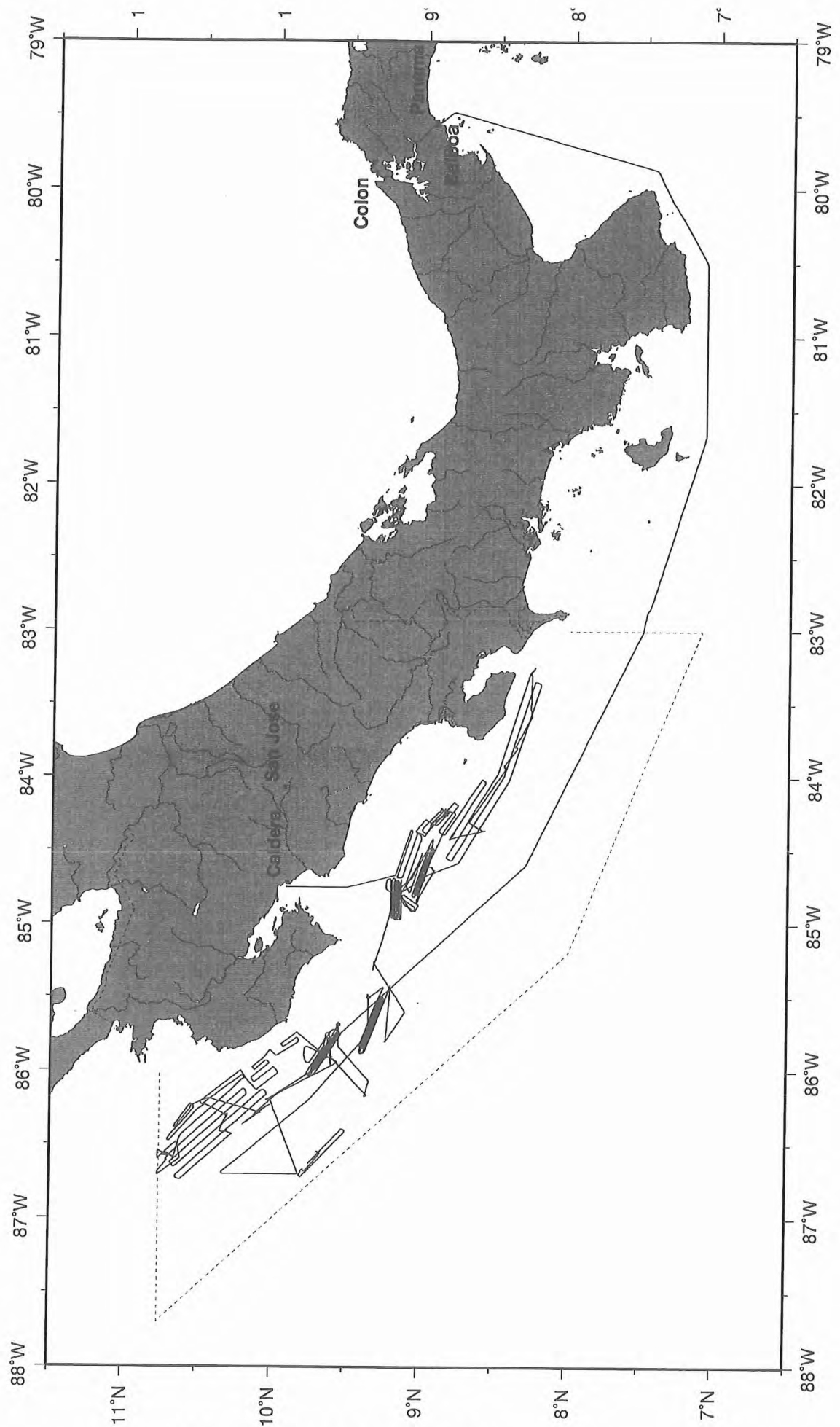


Figure 4.1.6 Total track of the cruise SO163 Leg 1

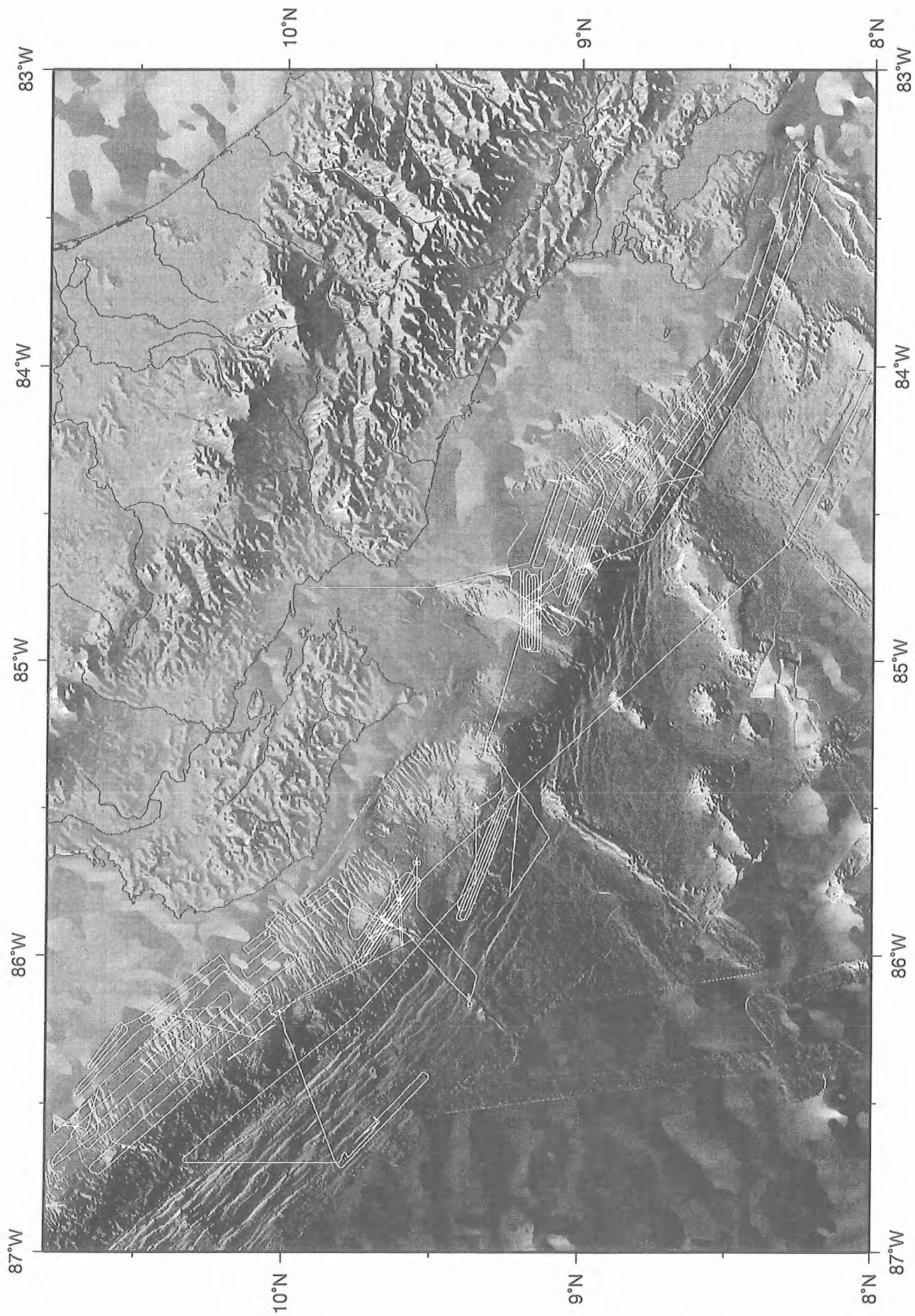


Figure 4.1.7 Cruise track of the cruise SO163 Leg 1 in the area off Costa Rica

4.2 Agenda of the cruise SO163-2

Cruise SO 163 "SUBDUCTION I-Leg 2" started when leaving the pier of Puerto Caldera on 22. April at 10:40. Four containers with scientific equipment had been loaded and installed before, and the computer network needed for the seismic data processing remained onboard from the two previous cruises. After a short transit of only 3 hours we reached the Jaco Scar area, where a seismological network was to be installed for several month (to be picked up during cruise M54/3B in October 2002). Three instruments were ready for immediate deployment, before the remaining acoustic releases were tested using W6 winch at a waterdepth of 3000 m. Then four more instruments and a new tiltmeter were also installed.

We then headed northwest to an active vent site, where during the previous leg a local seismic array comprising 10 instruments had been installed. We first added two more instruments and shot six short airgun profiles across the array, thereby varying the gun depth from 3.5 to 14 m to monitor the far field signal. Only two of the instruments were recovered, before we started a regional transect of CTD/water sampling/multicorer investigations. In between stations CDT02 and 03 the magnetometer was deployed during the transit.

On 25. April two more profiles, now using the g-gun clusters, were shot across the array before all instruments were safely recovered at 09:00 25.04.

Next we deployed 12 OBS/H on the Nicoya slide and shot into them along two long lines (profiles P9 and P10). While shooting on the oceanic plate, two more magnetic profiles were collected.

After the recovery of three instruments two more were deployed (amongst them a new OBS design) and several short cross lines were shot using different airgun configurations. All instruments were recovered by 17:00 27.04.

CDT05 and 06 completed the northern regional transect and Sonne returned to the Jaco Scar area. Here the tiltmeter was recovered, 5 instruments were deployed for the seismic network and a CTD was run in the center of Jaco Scar.

On 29. April 17 instruments (OBH43 to 59) were deployed on the margin above a previously identified reflection, that is interpreted to indicate upper plate erosion through the detachment of megalense type structures. Shooting across this ca 40 nm long linear array was done twice, again using different airgun configurations (P15 and 16). Amongst the 17 instruments was a new OBS design (OBS2002), which failed to return to the surface upon interrogation. The release command was acknowledged, but the instrument did not pop up. Since the waterdepth was only 200 m, an attempt to dredge the instrument was made but failed due to problems with the cable.

Following two more CTD stations (CTD08 and 09) another attempt to dredge the instrument was finally successful.

In the evening of 01. May two tiltmeter stations (OBT60 and 61) were deployed on Quepos plateau, and after CDT10 a 30 nm long cross line on the megalense structure was covered with 14 OBS/H (OBH62 to 75), shooting was done twice (P17, 18) across the array. All instruments were safely recovered 03. May in the morning.

Two more CTD stations (CDT11 and 12) were run on Quepos mound, using the Sonne USBL positioning system attached to the instrument, so samples could be taken along profiles across the mound.

A second strike line across the megalense structure was made after the deployment of a 15 instrument linear Array (OBH76 to 90). This profile was shot four times using different airgun configurations (P18, 20, 21, and 22), and a short reflection profile (P19) was included across Quepos mound, using a high frequency airgun array and a shot interval of only 10 sec. All instruments were recovered 05. May,

followed by two CTD-stations on Quepos mound (CTD13 and 14). The last profile on the megalense was collected parallel to Profiles 15 and 16, with 15 instruments (OBH92 to 106) deployed. Again, this line was shot using two different airgun configurations (P24 and 25). This work was completed in the morning of 07. May.

A regional CTD-transect off southern Costa Rica comprising 6 stations was started with the first two measurements (CTD15 and 16) in the afternoon of 7 May.

The next seismic work focussed on the detachment reflection south of Quepos Plateau. In Total 15 instruments (OBH107 to 121) were deployed in the evening of 07 May, the first and the last

instruments were deployed with a 1000 m long cable such that the far field signature of the guns could be monitored without interference from the sea floor reflection. Shooting along this 8.5 nm long profile was made at a minimum speed of 2 kn with different airguns and trigger rates (profiles P 26,27,28), and was followed by a series of testshots using various combinations throughout the G-Gun cluster (P29). All instruments were successfully recovered 08. May in the evening, before two more CTD stations were made on the southern regional transect (CTD17 and 18).

On May 09 the two tiltmeters and the 7 instruments from the first deployment (OBS14 to 21) were all recovered, followed by a transit to the outermost two CDT stations of the southern transect (CTD 19 and 20).

Two parallel strike lines on the detachment reflection were collected May 10 and 11 (profiles P30 and 31), and three more instruments (OBH 139 to 141) were deployed in the Jaco Scar seismological network, where another CTD (21) was deployed on Jaco Scar.

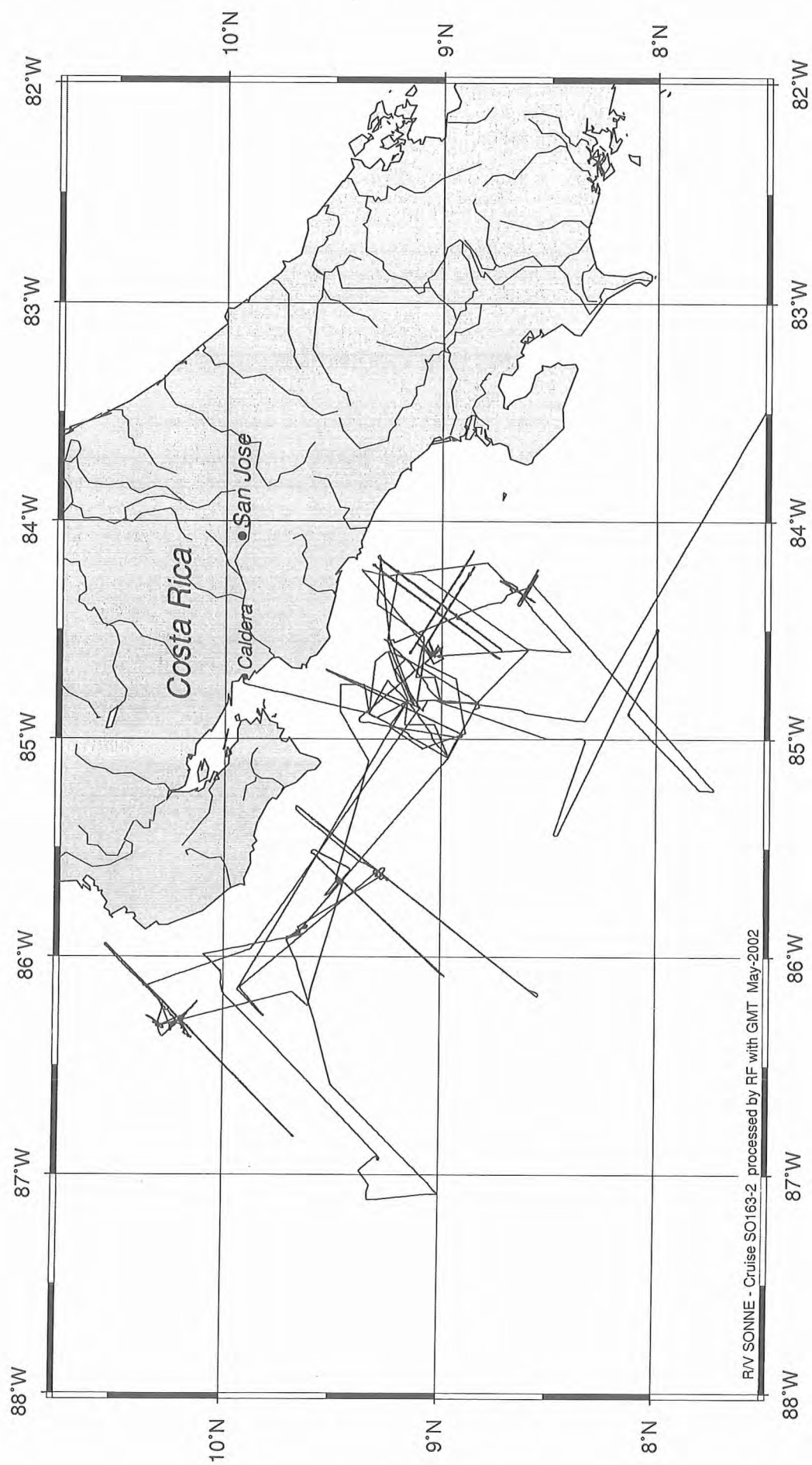
On May 12 20 instruments were deployed on a regional profile across the Culebra Mound, including a short strike line and a densely sample line (Profile 33) across Culebra Mound. Here also two CTD deployment were made, and subsequently two attempts to dredge rocks from this mound unfortunately failed.

On the transit back to Nicoya slide a high resolution bathymetric profile including Parasound echosounding was collected around the planned ODP-leg 205 drilling.

On the Nicoya slide another 20 instruments were deployed, and a regional line plus two short strike lines were collected May 15 and 16.

Finally, 20 additional instruments (OBH192 to 212) were all deployed in the Jaco Scar seismological network, and three profiles (P38 to P40) were shot into the then 28 instruments deployed in the network. Shooting terminated at 06:00 on May 18, and following a last CDT deployment at Jaco Scar (CTD 24) sonne set course to Balboa, Panama. The transit through Costa Rican waters was chosen such that an hithero unsurveyed part of the oceanic plate could be mapped with the Simrad system, and we also deployed the magnetometer (magnetic profile M04). Upon reaching Panamenian waters at 83°W the Simrad system was shut down and the magnetometer was recovered at 04:00 on 19. May, which terminated the scientific work.

Sonne finally reached the pilot station in Balboa at 06:00 May 20 after cruising a total of 4000 nm within 28 days at sea. A trackplot of leg SO163-2 is shown inFigure 4.2.1.



SUBDUCTION Fluid-Volatiles-Hazards

Caldera 21.04.2002 - Balboa 21.05.2002

GEOMAR Kiel / SFB 574

Figure 4.21: Cruise Track of 163-2



Scale 1:3.000.000

Mercator Projection (WGS 84)

SO-163-1/2: Seismic Profiles and Ocean-Bottom Stations

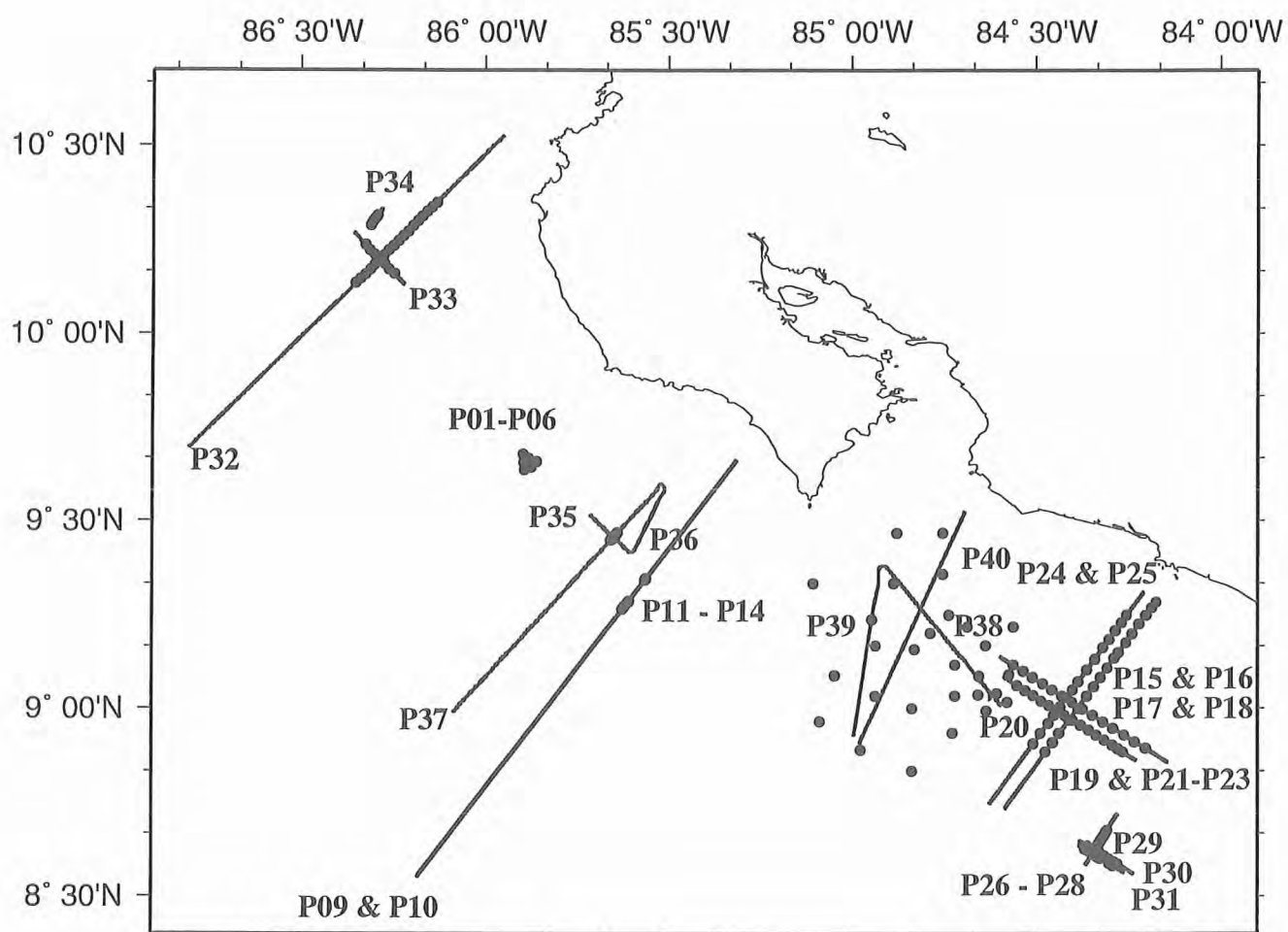


Figure 4.2.2: Location map of profiles P01-P40 and stations OBH1-OBH212.

5.1.1 Navigation

(W. Weinrebe)

A crucial prerequisite for all kinds of marine surveys is the precise knowledge of position information (latitude, longitude, altitude above/below a reference level). Since 1993 the global positioning system (GPS) is commercially available and widely used for marine surveys. It operates 24 satellites in synchronous orbits, thus at least 3 satellites are visible anywhere at any moment (Seeber, 1996). The full precision of this originally military service yields positioning accuracies of a few meters. In the past this was restricted to military forces and inaccessible to commercial users (Blondel and Murton, 1997). Since about two years the full resolution is generally available.

The resolution of GPS can be enhanced with the Differential GPS (D-GPS) scheme (Blondel and Murton, 1997, Knickmeyer, 1996). Using several reference stations the determination of the ship's position can be corrected in real time and enhanced to a 1 m to 5 m accuracy. Since the cruise SO-109 (1996) D-GPS service is available onboard R/V SONNE. The ships ASHTEC system provides a validated accuracy better than 5 - 10 m in the area off Costa Rica.

D-GPS-values as well as most other cruise parameters are continuously stored in the navigation database, and are distributed via the DVS- („data distribution system“) on the ship's network.

Unfortunately, the precision of the position information does not correspond to the accuracy of the time base in the navigation database, as the navigation processing unit Atlas ANP 2000 does not copy the precise GPS-time values, but adds time stamps of its internal unsynchronized clock.

5.1.2 Simrad EM120 Multibeam System

(C. Huguen)

The EM120 system is a multibeam echosounder (with 191 beams) providing accurate bathymetric mapping up to depths higher than 11000 m. This system is composed of two transducer arrays fixed on the hull of the ship, which send successive frequency coded acoustic signals (11.25 to 12.6 kHz). Data acquisition is based on successive emission-reception cycles of this signal. The emission beam is 150° wide across track, and 2° along track direction (Fig. 5.1.2.1). The reception is obtained from 191 overlapping beams, with widths of 2° across track and 20° along it (Fig. 1). The beam spacing can be defined as equidistant or equiangular, and the maximum seafloor coverage fixed or not. The echoes from the intersection area (2°*2°) between transmission and reception patterns (Fig. 5.1.2.1), produce a signal from which depth and reflectivity are extracted.

For depth measurements, 191 isolated depth values are obtained perpendicular to the track for each signal. Using the 2-way-travel-time and the beam angle known for each beam, and taking into account the ray bending due to refraction in the water column by sound speed variations, depth is estimated for each beam. A combination of phase (for the central beams) and amplitude (lateral beams) is used to provide a measurement accuracy practically independant of the beam pointing angle. The raw depth datas need then to be processed to obtain depth-contour maps. In the first step, the data are merged with navigation files to compute their geographic position, and the depth values are plotted on a regular grid to obtain a digital terrain model (DTM). In the last stage, the grid is interpolated, and finally smoothed to obtain a better graphic representation. Together with depth measurements, the acoustic signal is sampled each 3.2ms and processed to obtain a cartographic representation, commonly named mosaic, where grey levels are representative of backscatter amplitudes. These data provide thus informations on the sea-floor nature and texture; it can be simply said that a smooth and soft seabed will backscatter little energy, whereas a rough and hard relief will return a stronger echo.

During the SO 163 cruise, the Simrad EM 120 Multibeam echosounder, available on R/V SONNE since June 2001, was used continuously. Bathymetric data were processed routinely onboard during the survey, using the NEPTUNE software from Simrad, available on board and the academic software MB-System from Lamont-Doherty Earth Observatory.

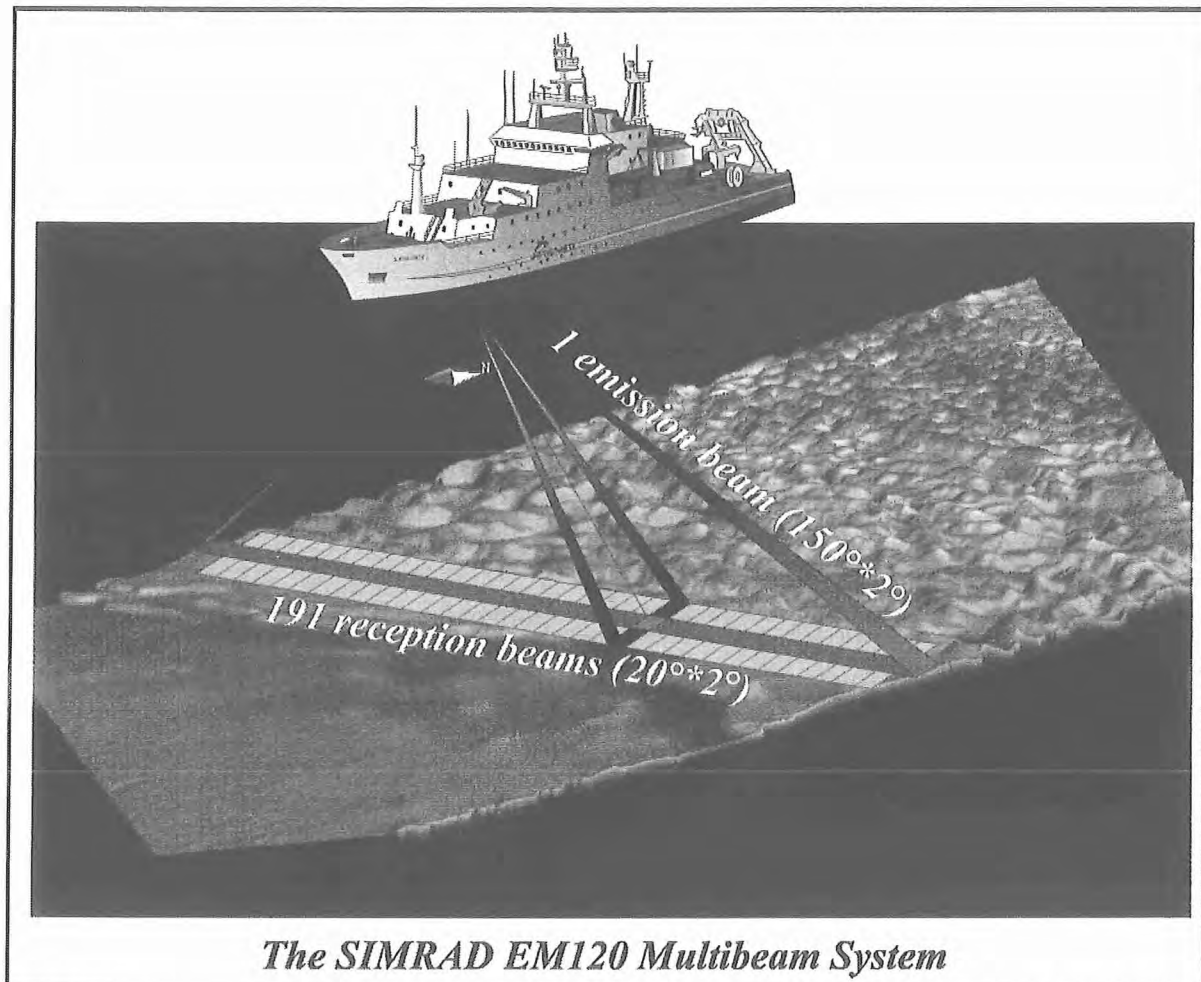


Figure 5.1.1 Acquisition method for bathymetric and backscatter data from the Simrad EM120 system (crossed beams technique)

5.1.3 PARASOUND

(T. Nadler)

The PARASOUND system works both as a low-frequency sediment echosounder and as a high-frequency narrow beam sounder to determine the water depth. It utilizes the parametric effect, which produces additional frequencies through nonlinear acoustic interaction of finite amplitude waves. If two sound waves of similar frequencies (here 18 kHz and e.g. 22 kHz) are emitted simultaneously, a signal of the difference frequency (e.g. 4 kHz) is generated for sufficiently high primary amplitudes. The new component travels within the emission cone of the original high frequency waves, which are limited to an angle of only 4° for the equipment used. Therefore, the footprint size of 7% of the water depth is much smaller than for conventional systems and both vertical and lateral resolution are significantly improved.

The PARASOUND system is permanently installed on the ship. The hull-mounted transducer array has 128 elements within an area of 1 m². It requires up to 70 kW of electric power due to the low degree of efficiency of the parametric effect. In 2 electronic cabinets, beam formation, signal generation and the separation of the primary (18, 22 kHz) and secondary frequencies (4 kHz) is carried out. Using the third electronic cabinet located in the echosounder control room, the system is operated on a 24 hour watch schedule.

Since the two-way travel time in the deep sea is long compared to the length of the reception window of up to 266 ms, the PARASOUND System sends out a burst of pulses at 400 ms intervals, until the first echo returns. The coverage in this discontinuous mode is dependent on the water depth and also produces non-equidistant shot distances between bursts.

The main tasks of the operators are system and quality control and to adjust the start of the reception window. Because of the limited penetration of the echosounding signal into the sediment, only a short time window close to the sea floor is recorded.

In addition to the analog recording features with the b/w DESO 25 device, the PARASOUND System is equipped with the digital data acquisition system ParaDigMA, developed at the University of Bremen. The data is stored on removable hard disks using the standard, industry-compatible SEG-Y-format. The 486-processor based PC allows for buffering, transfer and storage of the digital seismograms at very high repetition rates. Of the emitted series of pulses, usually only every second pulse can be digitized and stored, resulting in recording intervals of 800 ms for a given pulse sequence. The seismograms were sampled at a frequency of 40 kHz, with a typical registration length of 266 ms for a depth window of ~200 m. The source signal was a band limited, 2-6 kHz sinusoidal wavelet with a dominant frequency of 4 kHz and duration of 1 period (250 μ s total length). Data was stored on DAT-tapes using Windows NT backup software.

5.1.4 Ocean Floor Observation System (OFOS) and TV-guided grab (TVG)

(H. Sahling, T. Nadler)

OFOS

The OFOS system on board RV Sonne is a most sophisticated camera sled, which has been specifically adapted to completely satisfy all scientific requirements. OFOS is equipped with following instruments: two video cameras (colour camera: Deep Sea Power and Light, black and white camera: Photosea), a stereo still camera system (Photosea), various xenon and halogen lights (Deep Sea Power and Light), CTD (Seabird), a compass, pitch and roll sensor, and a Simrad SSBL responder.

The sled is towed behind the ship at a speed of about 0.8 knots (kn). The distance of about 1.5 m to the seafloor is manually adjusted by the winch operator. For that, a weight is hanging on a rope of 2 m length below the sled. Two laser pointer can be used to scale the video image and the still camera images. Two laser pointer are parallel and 20 cm apart, while a third one points in an oblique angel. The still camera is loaded with a slide film (Kodak Ektachrom 200) allowing a maximum of 800 shots. After using an aperture of 5.6 (camera A) and 8 (camera B) during OFOS 1 & 2, which resulted in too-dark seafloor images an aperture of 4 resulted in a satisfactory illumination. The images were taken manually or at regular time intervals. On the video signal and on the images the date and time (UTC) is overlaid. The ships data, the SSBL position and some OFOS data were recorded in the database of the ship. These data were downloaded and stored in the database Access referenced by the UTC time. Additionally, the system operators downloaded the raw data of the CTD mounted on OFOS (sampling rate 0.5 sec.). These data were recalculated to the time of deployment and integrated also into the database.

During SO 163 cruise we were confronted with peculiarities of the investigation area that influenced the deployments of OFOS considerably. Strong currents and, at some times winds made navigation difficult. Thus, most tracks were planed in arrangement with the bridge. Furthermore, OFOS should be preferential towed downslope. As a result the directions of survey lines were strongly controlled by these factors rather than solely on scientific targets, which we wanted to image.

There are a few improvements concerning OFOS we would like to mention. The data of the CTD, which is mounted on OFOS, are not stored in the ships database. That would be most important in the future. The front end for OFOS Bottom travel should include the SSBL position, too.

TVG

The TV-guided grab (TVG) is a heavy tool to collect sediment, rock or biological samples from the seafloor. Its lits can be closed hydraulically. It is powered by deep-sea batteries allowing closing and opening it about three times. The opening of the grab is about 1.8 m². For the deployments the grab was equipped with the black-and-white camera and the colour camera used in OFOS. Accurate position can be acquired by a SSBL responder, which is tied on the fiberoptic cable 50 m above the tool.

5.1.5 CTD

(H. Sahling)

A CTD was deployed during SO 163-1 cruise in order to acquire a sound velocity profile through the water column. This profile was then used to calibrate the echosounding systems (Simrad, Posidonia).

The CTD/rosette aboard Sonne consists of a deck unit (SBE 11 plus), which supports the underwater unit of the CTD (SBE 9 plus) and the SBE 32 carousel water sampler equipped with 24 * 10 l Niskin bottles. The used CTD system supplies a modular temperature sensor (SBE 3 plus), a conductivity sensor (SBE 4), a Digiquartz pressure sensor, and an oxygen sensor (SBE 13). An external pump (SBE 5) maintains an optimum and constant water flushing speed for temperature, conductivity and oxygen sensors via a TC Duct, thus guaranteeing that all sensors measure the same water.

The deck unit provides DC power to the sea cable, decodes the serial data stream arriving from the underwater unit and passes the data to a companion computer. Binary data from the sensors are transmitted serially 24 times a second. The deck unit has a separate communication channel for controlling the Seabird carousel and an audible bottom contact alarm. Acquired data are placed on the computer where postprocessing is achieved with the software provided by Seabird. The data were averaged and bottle files were produced.

5.2 Computer facilities for bathymetry, magnetic, and seismic data processing (W. Weinrebe)

The experiments and investigations during SO163 required special computing facilities in addition to the existing shipboard systems. For programming of ocean bottom stations, processing and interpretation of seismic data and analysis of magnetics several workstations and a dedicated PC-laptop were installed by the wide angle and seismology groups of GEOMAR. Due to the large amount of data transfer GEOMAR installed a workstation cluster onboard comprising the following systems:

1	"moho"	SUN Sparc20 2 CPU	256 MB memory	20 GB disks DAT, CD	Sun Solaris 5.8
2	"devonia"	SUN Ultra 60 2 CPU	1 GB 1 GB	150 GB disks 1x DLT8000 2x DAT 2x Exabyte	Sun Solaris 5.6
3	"hotblack"	SUN Ultra 1 1 CPU	128 MB memory	36 GB disks CD	Sun Solaris 5.8
4	"galicia"	SUN Sparc 10 1 CPU	96 MB memory	12 GB disks DAT	Sun O.S.4.1.4
5	"crimea"	AMD DURON 700 MHz	128 MB memory	68 GB disks 6x PCMCIA	Windows2000
6	„pinta“	AMD DURON 700 MHz	128 MB memory	68 GB disks 6x PCMCIA	Windows2000

In addition to these computers, several laptops were used. For plotting and printing two HP Postscript Laserprinters (papersize A3 and A4) as well as the shipboard color plotters were available.

The workstation cluster was placed in the Magnetiklabor and the Reinlabor where it was set up according to a "client-server" model, with "moho" being the server. All important file systems from the main server at GEOMAR were duplicated onto the "moho"-disks. Using NFS-, NIS-, and automounter services the computing environment was nearly identical to that at GEOMAR, so every user found his/her familiar user interface. The convenience of network mounted file systems has to be paid for with a heavy network load, particularly during playback of OBH-data (c.f. SO123 cruise report, Flueh et al., 1997). This required a high-performance network, which was accomplished by a switched twisted-pair ethernet. A 12-port ethernet switching-hub (3COM-SuperstackII 1000) with an uplink connection of 100 Mbps to the server "moho" and dedicated 10 Mbps ports for the client workstations maintained the necessary network performance. In order to keep the shipboard network undisturbed by the workstation cluster, but to allow for communication between them, the server "moho" was equipped with two network interfaces and served as a router. This provided the additional benefit of a simplified network configuration. Considerable setup work was dedicated to "moho", while the other workstations used the same IP-addresses and network configuration as at GEOMAR. In addition, „hotblack“ was set up as an redundant server, so in case „moho“ would have failed, „hotblack“ could easily switched to replace „moho“ as a server.

This network setup showed a reliable and stable performance, and no breakdowns were observed.

5.3 Seismic Instrumentation

The Ocean Bottom Hydrophone

The first GEOMAR Ocean Bottom Hydrophone was built in 1991 and tested at sea in January 1992. This type of instrument has proved to have a high reliability; in fact during this cruise we celebrated the 2000 successful deployment. A total of 20 OBH and 9 OBS instruments were available for SO163. Altogether 212 sites were occupied during the SO163 cruise.

The principle design of the instrument is shown in Figure 5.3.1, and a photograph showing the instrument upon deployment can be seen in Figure 5.3.2. The design is described in detail by Flueh and Bialas (1996).

The system components are mounted on a steel pipe which holds the buoyancy body on its top. The buoyancy is made of syntactic foam and is rated, as are all other components of the system, for a water depth of 6000 m, except for the pressure cylinders holding the recording electronics. Here, various models are available for variable depths (2500 m, 3000 m, and 6000 m). Attached to the buoyant body are a radio beacon, a flash light, a flag and a swimming line for retrieving from aboard the vessel. The hydrophone for the acoustic release is also mounted here. The release transponder is a model *RT661CE* made by *MORS Technology*. Communication with the instrument is possible through the ship's transducer system, and even at maximum speed and ranges of 4 to 5 miles release and range commands are successful. For anchors, we use pieces of railway tracks weighing about 40 kg each. The anchors are suspended 2 to 3 m below the instrument. The sensor is an *E-2PD* hydrophone from *OAS Inc.*, or the *HTI-01-PCA* hydrophone from *HIGH TECH INC.* and the recording device is a *MBS* recorder of *SEND GmbH*, which is contained in its own pressure tube and mounted below the buoyant body opposite the release transponder (see Figures 5.3.1 and 5.3.2).

The Ocean Bottom Seismometer

The Ocean Bottom Seismometer (OBS) construction (Bialas and Flueh, 1999; Fig. 5.3.3) is based on the experiences with the GEOMAR OBH. For system compatibility the acoustic release, pressure tubes, and the hydrophone are identical to those used for the OBH. Syntactic foam was used as floatation again but of larger diameter due to the increased payload. In contrast to the OBH the OBS has three legs around a center post to which the anchor weight is attached (Fig. 5.3.4). While the OBH is floating about 1 m above the sea bottom, the OBS is positioned on the sea bottom to avoid collisions between the the seismometer cable and the anchor. The sensible seismometer is deployed about 1 m to the side of the system once the OBS lands on the sea floor. The three component seismometer is housed in a titanium tube, modified from a package build by Tim Owen (Cambridge) earlier. Geophones of 4.5 or 30 Hz natural frequency are available. During descent to the ocean bottom, the footplate of the seismometer release lever is about one meter below the base of the anchor and therefore hits the seafloor first. At touch down the baseplate forces an upward movement of the lever which lays out the seismometer hook until the seismometer anchor is about 0.5 m above the seafloor. At about 45 degrees to the vertical the seismometer is released from its hook and falls to the sea floor from about 1 m height, ensuring coupling between the seismometer and the sea floor. At this time the only connection from the seismometer to the instrument is a cable and an attached wire which retracts the seismometer during ascent to the sea surface. An oscillation of the instrument caused by possible currents is therefore not transmitted mechanically to the seismometer. All three channels are preamplified within the seismometer housing and recorded by the standard Methusalem recorder as used in the OBH units. Parallel to these three channels the standard hydrophone is recorded on the fourth channel. A self-levelling seismometer with 4.5 Hz geophones was also successfully used during SO163. The levelling of the three component geophone (termed LG01) is done at a preset time by unlocking springs that clamp the geophone package against the titanium housing.

Beside the standard OBS type used for active seismic recording broadband seismometer were available and routinely deployed within the profiles. Two different types of sensors can be attached to the

to 4 storage cards may be used. Currently, up to 640 MB per card are available. Data compression allows to increase this capacity. Recently technical specifications of flashdisks (disk drives of PCMCIA technology) have been modified to operate below 10 °C, therefore 2 GB drives are now available for data storage. After recording the flashcards need to be copied to a PC workstation. During this transcription the data are decompressed and data files from a maximum of four flash memory are combined into one data set and formatted according to the PASSCAL data scheme used by the *Methusalem* system. This enables full compatibility with the established processing system. While the *Methusalem* system did provide 16 bit integer data, the 18 bit data resolution of the *MBS* can be fully utilized using a 32 bit data format.

The Marine Longtime Seismograph

For the purpose of low frequent recordings such as seismological observations of earthquakes during long term deployments of about one year time a new data logger, the Marine Longtime Seismograph (MLS) was developed by *SEND GmbH* with support from GEOMAR.

The MLS is again a four channel data logger whose input channels have been optimized for 3-component seismometers and one hydrophone channel. The modular design of the analogue front end allows to adopt for different seismometers and hydrophones or pressure sensors. Currently front ends for the Spahr Webb, PMD and Guralp seismometers as well as for a differential pressure gauge (DPG) and the OAS hydrophone are available. With these sensors we are able to record events between 50 Hz and 120 s. The very low power consumption of 250 mW during recording together with a high precision internal clock (0.05 ppm drift) allows data acquisition for one year. Data storage is done on up to 12 PCMCIA type II flashcards. The instrument can be parameterized and programmed via a RS232 interface. After low pass filtering the signals of the input channels are digitized using Sigma-Delta A/D converters. A final decimating sharp digital low-pass filter is realized in software by a Digital Signal Processor. The effective signal resolution depends on the sample rate and varies between 18.5 bit at 20 ms and 22 bits at 1 s. Playback of the data is done under the same scheme as described for the *MBS* above. After playback and decompression the data is provided in PASSCAL format from where it could be easily transformed into standard seismological data formats.

Towed Streamer

In addition to the ocean bottom seismic recorders also a mini-streamer was used during seismic profiling. This streamer was manufactured by *SIG (Service et Instruments de Geophysique, France)* The system comprises several parts: four 50 m long active sections with 20 hydrophones spaced at 2.50 m, two 2.50 m long lead in sections separating the depth transducer (Philips P30) in the tail and the depth transducer and preamplifier in the head from the active sections. The lead-in cable is 150 m, and a 50 m long deck cable can be laid out to connect the winch to the lab. The individual hydrophones are omnidirectional and have a flat frequency response from 10 to 1000 Hz. The sensitivity is -90db, re 1V/ μ bar, \pm 1 db. The hydrophones are mounted in an oilfilled polyurethane pipe of 34 mm diameter, with a nominal density of 1.12 gr/cm³. The lead-in cable can be trimmed to the required depth using air and seawater. A control unit provides power to the preamplifier, displays the depths of the head and tail depth transducers and provides the analogue signals of the four channels. The depth readings are also available on RS232 interfaces for storage on a PC.

The signals recorded by the streamer were stored on a four-channel *MBS* recorder, identical to those used in the ocean bottom seismic recorders. The streamer winch was placed midships about 8 m away from the aft of the vessel.

instrument. The "Spahr Webb" type seismometer is based on *Mark-L4* sensors which are operated with a feedback loop to enable recordings of frequencies as low as about 60 sec. As the sensors are sensitive to horizontal or vertical adjustment the complete construction is fully gimbaled. Tilt is measured at selected intervals and two electric motors are used to adjust and fix for a proper positioning. The second sensor is a *PMD-113* which has a flat frequency response curve from 95 sec. up to 30 Hz. This sensor type operates on the base of measuring levels within electrolytical tubes. This principle is less sensitive to horizontal adjustment. The sensor is fixed like a pendulum while its lower third is surrounded by a viscous oil filling that gives freedom to very slow movements (within a circle of 18°) and could be assumed to be solid within the measured frequency range. Both systems are mounted within a 17" glass sphere with an additional weight at the bottom (20 kg weight in water) which should ensure a good coupling to the ground. The above described lever system was not able to handle this size of sensor and therefore a slide system was designed (Fig. 5.3.5 and 5.3.6) which allows to deploy the sensor about three meters to the side of the instrument carrier. A clock controlled burning wire is used first as release of the slide. Then the sensor is pulled by an elastic rope along the slide until it falls off at the end of the boom. Secondly the clock releases the sledge itself to enable deployment of the sensor even if it does not drop off the slide for any reason. Both sensors are recorded by use of the *Marine Longtime Recorder (MLS)* which is manufactured by *SEND GmbH* and specially designed for longtime recordings of low frequency bands. Together with the broadband sensors the fourth channel records a Differential Pressure Gauge (DPG) as described by Cox et al (1984).

During cruise SO163 a prototype of a new instrument design was tested. This design, termed OBS2002, is characterized by a very shallow design, which should potentially decrease current induced noise. Also, the mechanical design is such that only a few pieces of hardware are needed, which potentially will reduce the price of manufacturing. In addition, the instrument can be disassembled rapidly and can be packed in a rather dense fashion, such that up to 50 instruments could be stored in a single container. A picture of the instrument is shown in Figure 5.3.7. All sensors and operational devices remain identical to those described above.

Marine Broadband Seismic Recorder (MBS)

The so-called *Marine Broadband Seismic recorder (MBS)* (Bialas and Flueh, 1999), manufactured by *SEND GmbH*, was developed based upon experience with the DAT based recording unit *Methusalem* (Flueh and Bialas, 1996) over the last few years. This new recorder avoids a mechanically driven recording media, and the PCMCIA technology enables static flash memory cards to be used as unpowered storage media. Read/write errors due to failure in tape handling operations should not occur with this system. In addition, a data compression algorithm is implemented to increase data capacity. Redesign of the electronic layout enables a decreased power consumption (1.5 W) of about 25% compared to the *Methusalem* system. Depending on the sampling rate, data output could be in 16 to 18 bit signed data. Based on digital decimation filtering, the system was developed to serve a variety of seismic recording requirements. Therefore, the bandwidth reaches from 0.1 Hz for seismological observations to the 50 Hz range for refraction seismic experiments and up to 10 kHz for high resolution seismic surveys. The basic system is adapted to the required frequency range by setting up the appropriate analog front module. Alternatively, 1, 2, 3 or 4 analogue input channels may be processed. Operational handling of the recording unit is similar to the *Methusalem* system or by loading a file via command or automatically after power-on. The time base is based on a DTCXO with a 0.05 ppm accuracy over temperature. Setting and synchronizing the time as well as monitoring the drift is carried out automatically by synchronization signals (DCF77 format) from a GPS-based coded time signal generator. Clock synchronization and drift are checked after recovery and compared with the original GPS units. After software preamplification the signals are low-pass filtered using a 5-pole Bessel filter with a -3 dB corner frequency of 10 kHz. Then each channel is digitised using a sigma-delta A/D converter at a resolution of 22 bits producing 32-bit signed digital data. After delta modulation and Huffman coding the samples are saved on PCMCIA storage cards together with timing information. Up

Tiltmeter

Research and surveying in risk areas such as the vicinity of volcanoes and earthquake-prone areas have been supported on land by tiltmeters for years (e.g. Decker and Decker, 1981). Small-scale as well as integral, large-scale surveys are carried out - i.e. gauges with short (ca. 2 m) as well as long (several tens to hundreds of m) base lines are being used. These systems have been successfully deployed especially in areas with tunnels and mining. The combination of tiltmeters with seismic surveys yielded good results in the field of short-term forecasts of eruptions (Dzurisin et al., 1982). Tiltmeters for marine usage have been successfully used only recently (e.g. Dünnebier and Harris, 1996; Anderson et al., 1997). Many of the systems employed up to now have been developed and constructed in academia (Simamura and Kazanawa, 1988; Wyatt et al., 1984 and 1996; Anderson et al., 1997). Short baseline systems constituted the majority among these. They are less expensive to develop, their deployment is less complicated and they can be mounted onto a variety of different platforms.

As the electronics of the referenced marine tiltmeters are all in house developments that have not been designed to be used and operated by a wide community we decided to complete a new easy-to-use instrument for the SFB574 making use of existing components as far as possible. Based on the experiences with marine long baseline instruments published by Anderson et al. (1997) we favored a short baseline solution. Future operation should then result into a larger number of instruments distributed over the area of investigation to allow minimization of station effects by an overall integration of (parts of) the deployed systems. Restricting to short baseline enables the use of the GEOMAR OBS (Bialas and Flueh, 1999) as a system carrier further using the available periphery of support (e.g. release system, relocation tools, etc.). Internet research and personal discussion with other long term users of tiltmeters led to the decision of selecting the type 501 and 900-45 tiltmeters from Applied Geomechanics. The company had already been involved in marine tiltmeter developments with other customers. In general tilt values should be expected to be in the range of a few 2.5 urad to 200 urad (Wyatt et al, 1996; Tolstoy et al., 1998; a.o.) which could be easily resolved by the type 510 (+/- 1400 urad, < 10 nrad resolution). From earlier developments of a marine tiltmeter Applied Geomechanics could offer a gimbaled system based on freezing Jojoba oil. Repetitive leveling would not be possible with this solution and therefore we decided to use a motor driven leveling. The basic type 510 tiltmeter was developed for drill hole deployments and could therefore compensate for a small amount of tilt only. To adopt for sea floor applications and possible larger angles of tilt caused by the slope of the sea floor the leveling system was extended to +/- 30 degrees for the GEOMAR sensor. The type 900-45 clinometer can be operated over a range of +/- 50 degrees with a resolution of 0.02 degrees. It is used to give an overall impression about the large scale tilt variations throughout the deployment. As the complete system is designed to free fall to the seafloor a C-100 compass from KVH Industries is implemented within the tiltmeter housing to allow later orientation of the tilt axis. It is well known from earlier studies that such high precision instruments are sensitive against temperature changes which requires continuous recording of this value. For data recording the easiest solution was to modify two of the long time operated MLS type seismic data loggers (Bialas et al., 2002) from SEND. X- and Y-axis of the 510 tiltmeter and temperature are permanently recorded with 1 s sample interval while the 900-45 and the magnetic compass are read after each leveling operation.

Seismic Sources

32 I BOLT Airgun

A Model 800 CT *BOLT* airguns was used during the cruise; a photo of the guns is shown in Figure 5.3.8. The gun has a volume of 32 liters (2000 inch³), and generates a signal with a main frequency centered around 6 to 8 Hz and including higher harmonics. The gun was deployed and towed from the assistant winch and beam normally used for the piston corer recovery. Trigger cables and airhoses were deployed manually. The gun was suspended on two floats with an additional float attached to the supply lines to prevent contact between the gun and the towing wire. The gun was towed 60 m behind the vessel and operated at 145 bar in 7 to 8 m depth. Due to good weather conditions the handling of the gun was smooth all the time.

During cruise SO162 the gun was used twice. The total operation time was about 10 hours, with more than 1500 shots being fired, either at a 20 s or a 15 s shot interval. The ship's compressor system worked smoothly and caused no delays or interruptions.

G-Gun-Cluster

During this cruise two G-Gun clusters were available as a loan from Sodera, France. Each G-gun cluster comprises two identical guns mounted in a fixed frame. This assembly guarantees an interaction between the bubbles, and thus increases peak to bubble (P/B) ratio compared to a single gun. One cluster consisted of G-Gun 520 cu. in. each, the other one G-Gun 380 cu. in.. The P/B ratio should be between 16 and 17 according to the manufacturers specifications, being slightly higher for the smaller gun.

The clusters were towed over the A-frame of Sonne using the anchor winches and worked without any problem. In Figure 5.3.9 one of the clusters is shown upon deployment.

The G-Guns can and should be operated at a pressure of 3000 psi, unfortunately the Sonne compressor only provides a maximum pressure of 2000 psi.

Prakla –Array

In addition to the high and low frequency guns, a small airgun array was set up using the airgun slide at port aft of R/V Sonne. The array comprised of 7 Prakla VLA and VLF type guns, with volumes of 2.0 l, 1.2 l, 2 x 0.65 l, and 3 x 0.33l (Fig. 5.3.10). Parts of the array were provided on loan from the Institut für Geowissenschaften, Sektion Geophysik, Universität Kiel. Compared to trials undertaken during earlier cruises with R/V Sonne, some modifications in terms of single chamber volume, gun spacing and towing depth were done. As a consequence the frequency content of the array signal could be enhanced and the bubble period could be further suppressed. Single gun volumes and separations are shown on Fig. 5.3.11. Towing depth of the array was 2.5 m when three fenders were used as floatation buoys to carry the array. Shooting of the array was synchronized using a LongShot source controller with two FourShot power supplies. Two ports of the power supplies were reconfigured to operate the GI-gun, while the remaining 6 ports were used for the Prakla array. As a consequence, one of the 7 guns was always deployed as a spare. The source controller read the sensor signals submitted from each gun at shot release to synchronize the individual shot delays to the same aiming point with an accuracy of 1 ms. The array was operated at a pressure of 140 bar and a shot interval of 10 s. Despite problems with irregular sensor signals, all guns operated without major failures.

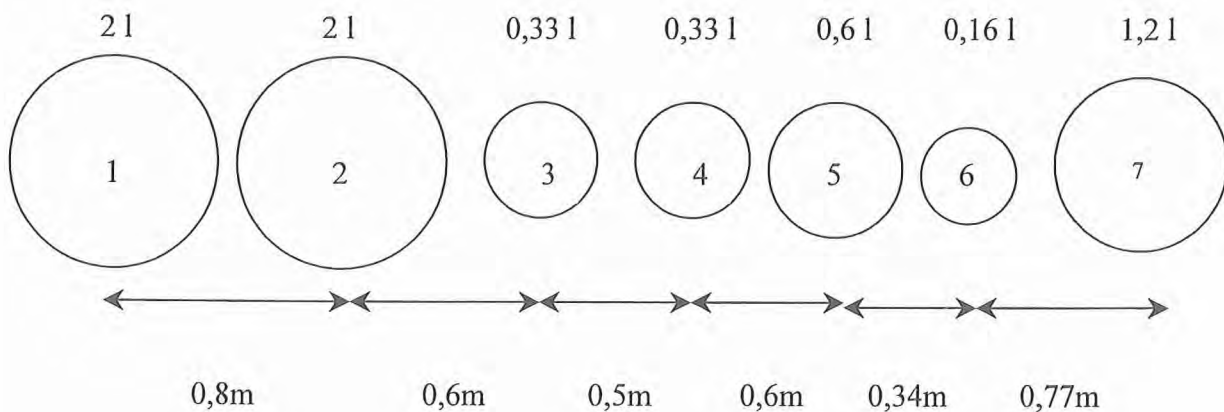


Fig. 5.3.10: Configuration for Prakla Type Airgun Array
Guns were towed at 2,5m depth

External Trigger

The trigger signal was supplied from the ships *Ashtech* GG24 GPS/Glonass receiver, and was available in the Geology Lab and the Seismic Lab. The receiver can provide a one millisecond long 5 V-TTL pulse at intervals between 0.2 and 999 s. The impulse should be stable to within the accuracy of the GPS Time, which is 70 nanoseconds. The impulse was delivered to the *Longshot* trigger box, which can handle different guns (G-gun, GI-Gun, Bolt abd Prakla guns). The shotbreaks, necessary for subsequent data processing and instrument location, were stored on a MBS recorder and displayed in real time to double check. For this process the same time basis was used that is used for the OBH (see chapter 4.4) and the trigger signal was converted into a 5 V TTL pulse of 250 ms length by a circuit provided from the ships technical support staff (WTD). Exact position calculation for the shot time should be done by later post-processing using shot time and UTC time values stored with DGPS coordinates in the ship's data base. As kwon from earlier cruises the coordinates stored within the data base were provided by the *Atlas ANP 2000* system, which does not copy the exact GPS time values but adds time stamps of its internal uncontrolled clock to the high precision coordinates of the DGPS system. Accuracy of the time values mainly depends on the operators skills by manually setting the ANP clock to GPS time. This is clearly a somewhat conservative method compared to the efforts of precise positioning. To enable the most accurate GPS related time stamps within the ANP system prior to each seismic survey the system operators were informed to reset the ANP.

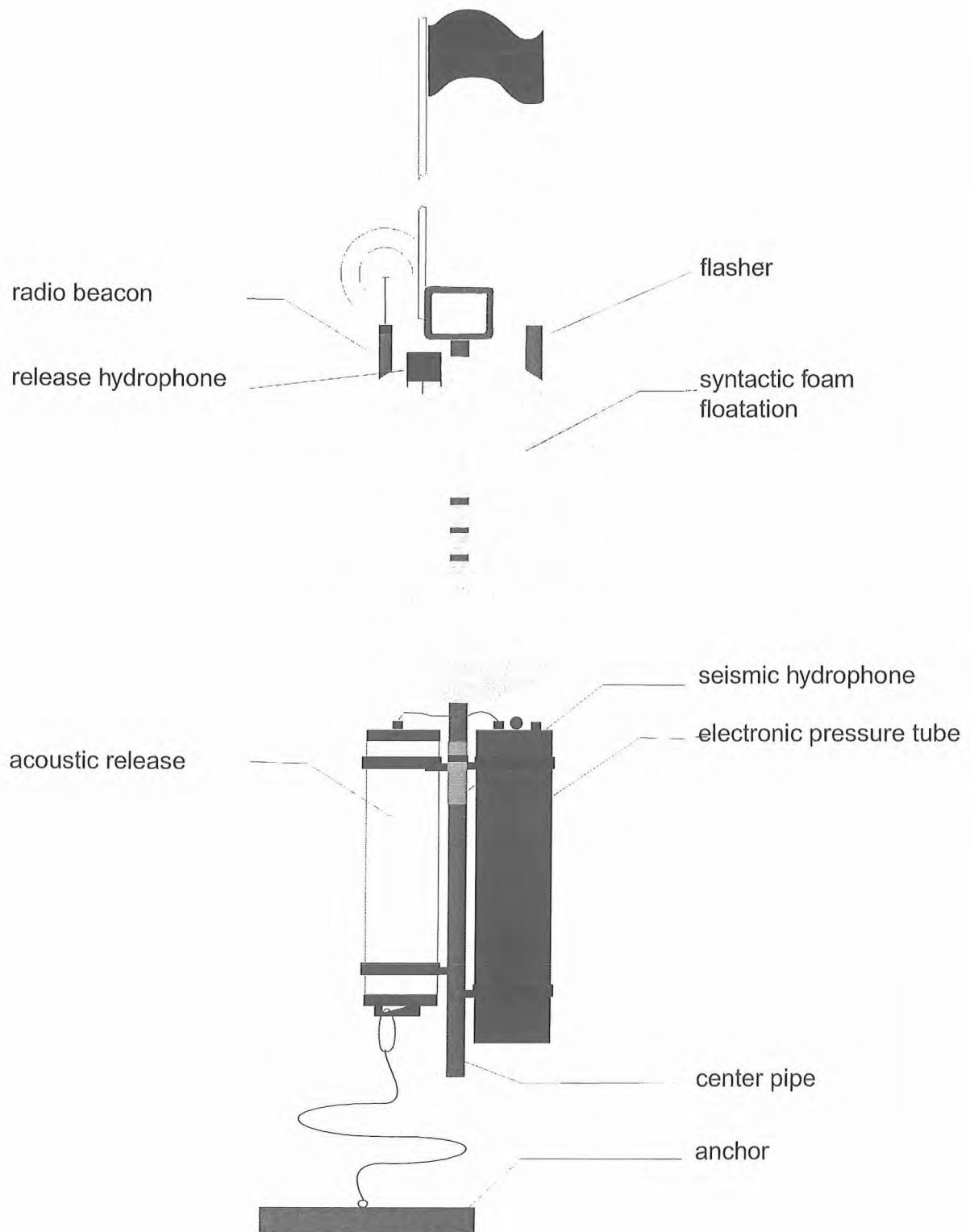


Figure 5.3.1: Principle design of the GEOMAR OBH (after Flueh and Bialas, 1996)

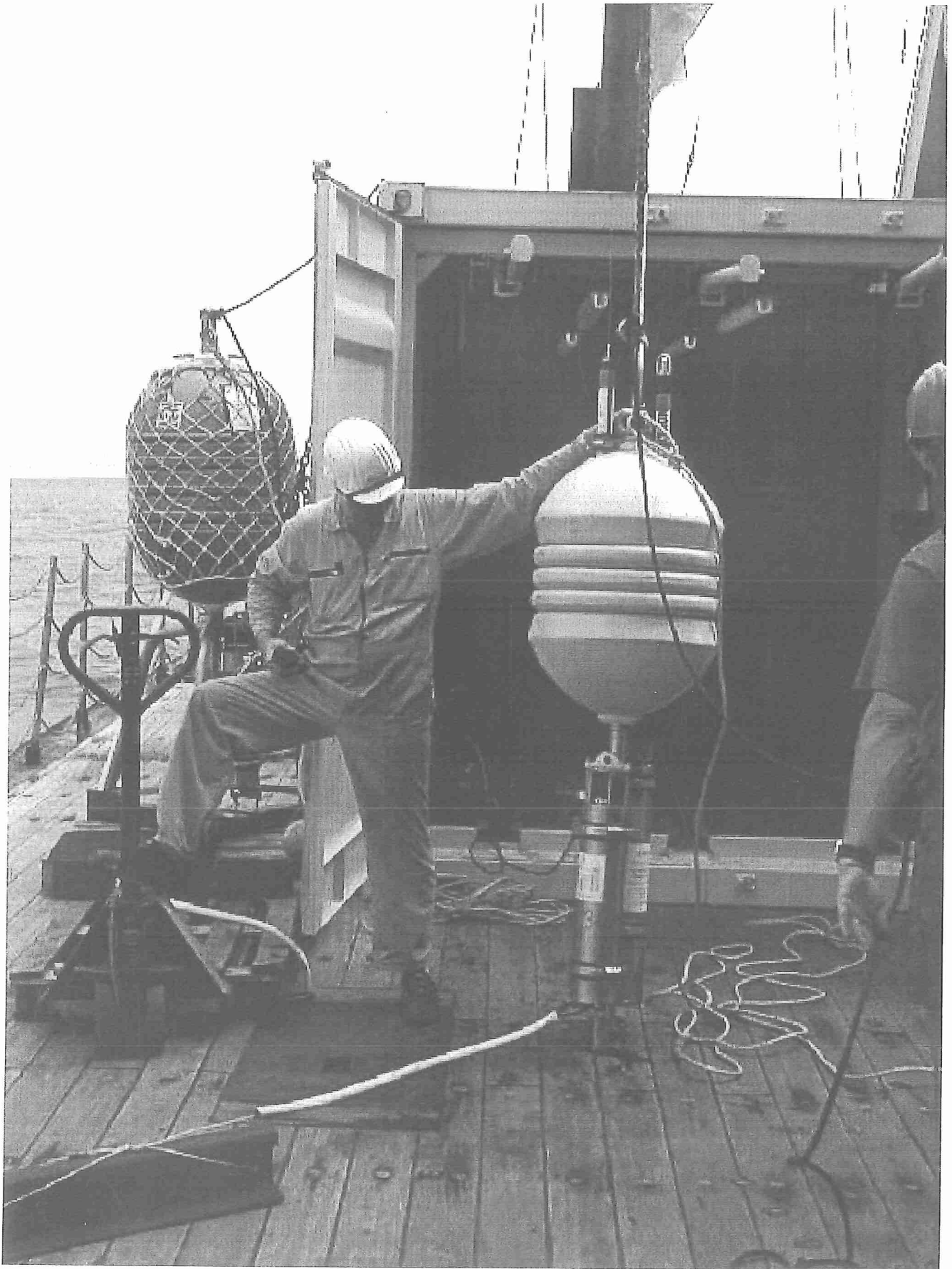


Figure 5.3.2 : The GEOMAR OBH ready for deployment
Background left: GEOMAR OBS
Background right: OBH container

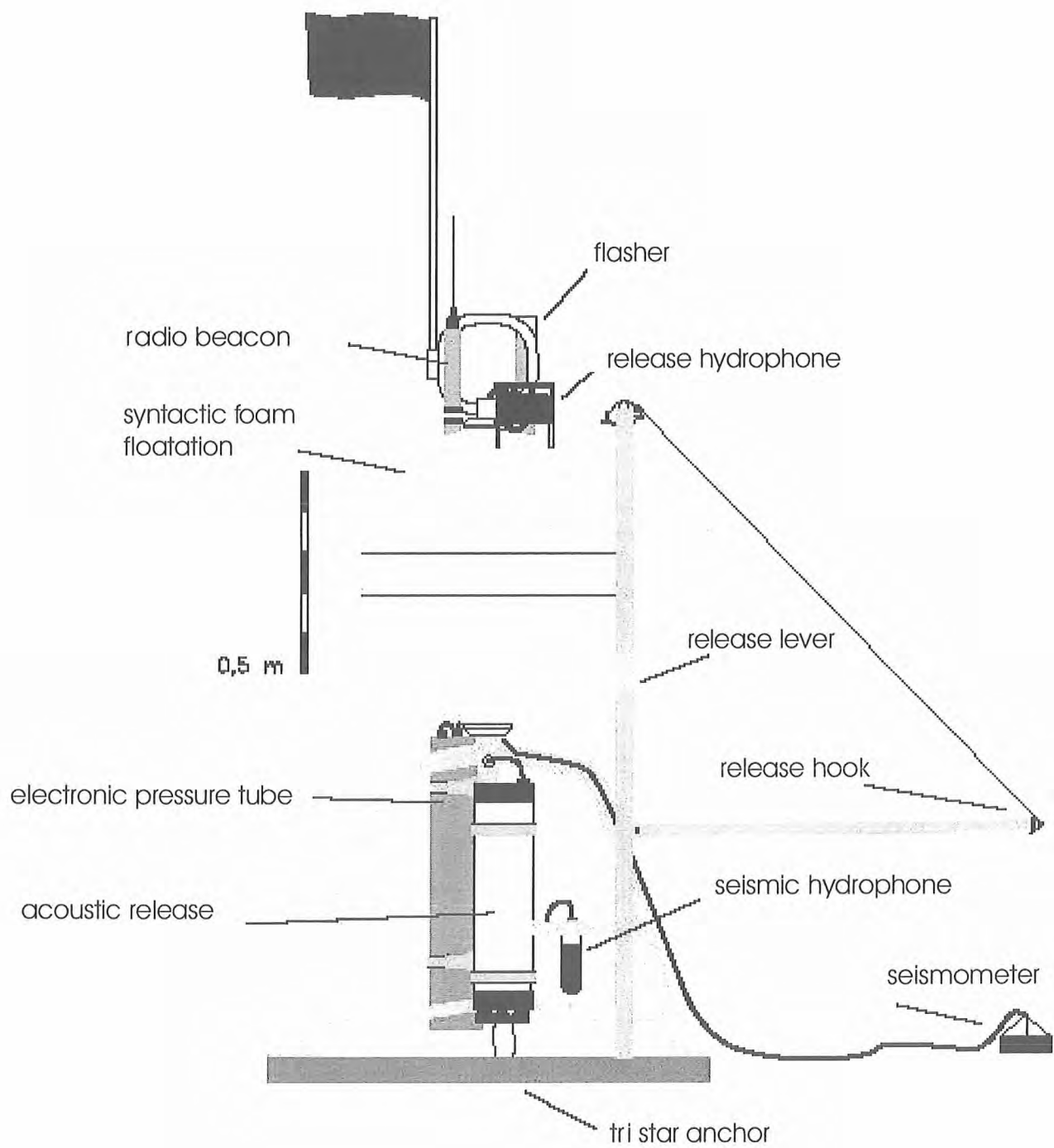


Figure 5.3.3: Principle of the GEOMAR OBS (after Bialas and Flueh, 1999)

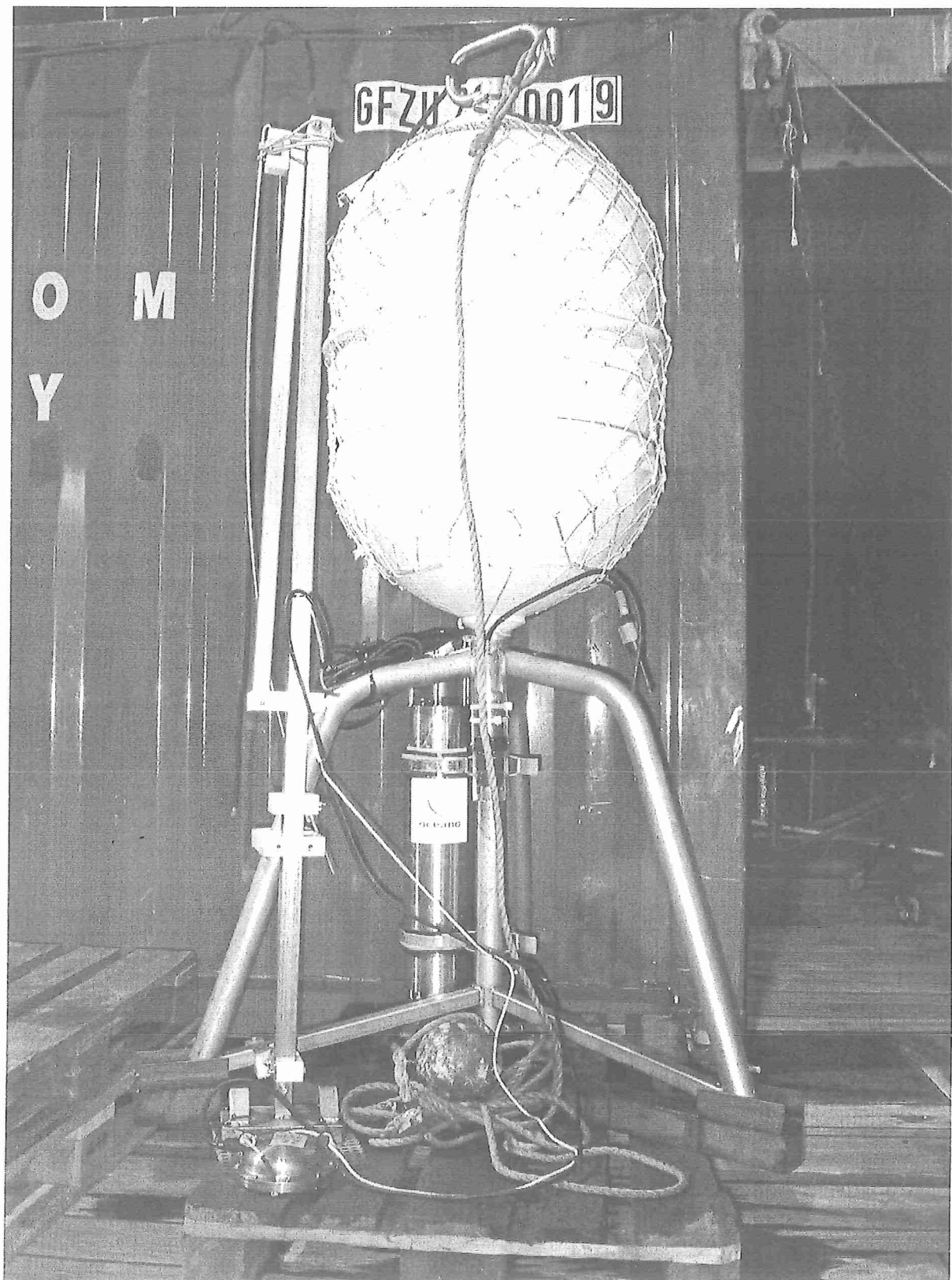


Figure 5.3.4: The GEOMAR OBS

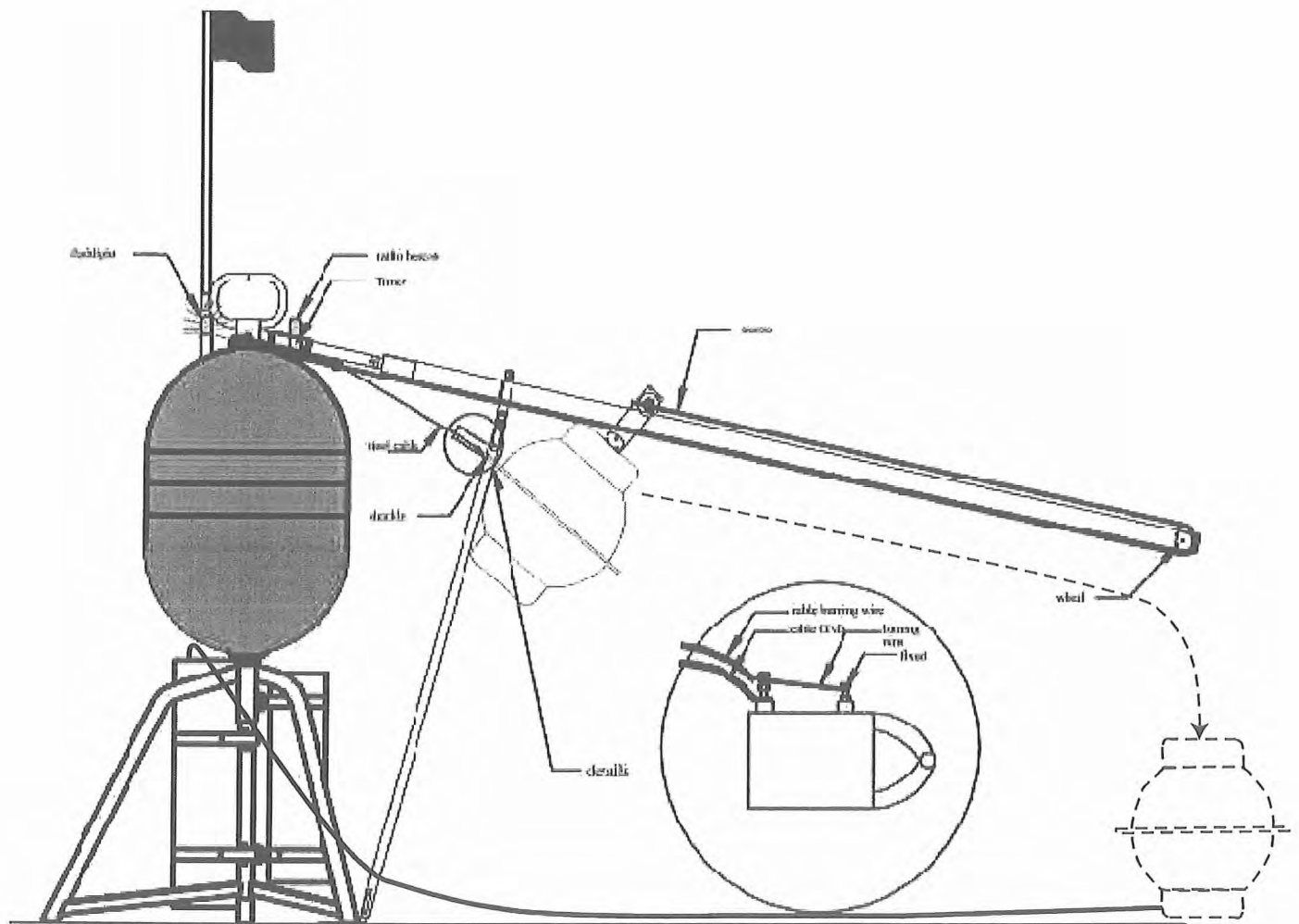


Figure 5.3.5: Sketch of GEOMAR broadband seismometer
dashed lines indicate deployment of the external broadband sensor at the seafloor

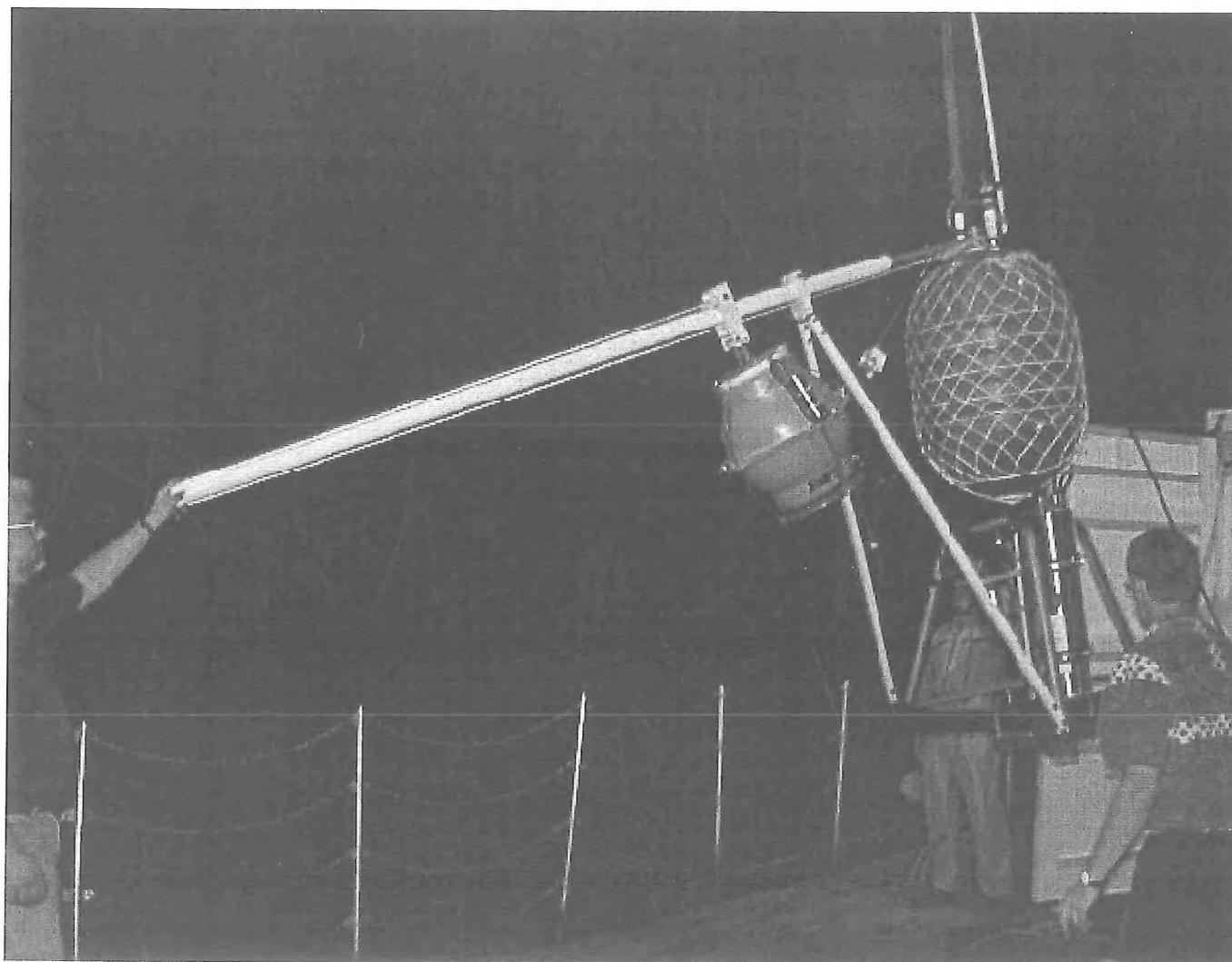


Figure 5.3.6: The GEOMAR broadband seismometer equipped with Webb seismometer at deployment

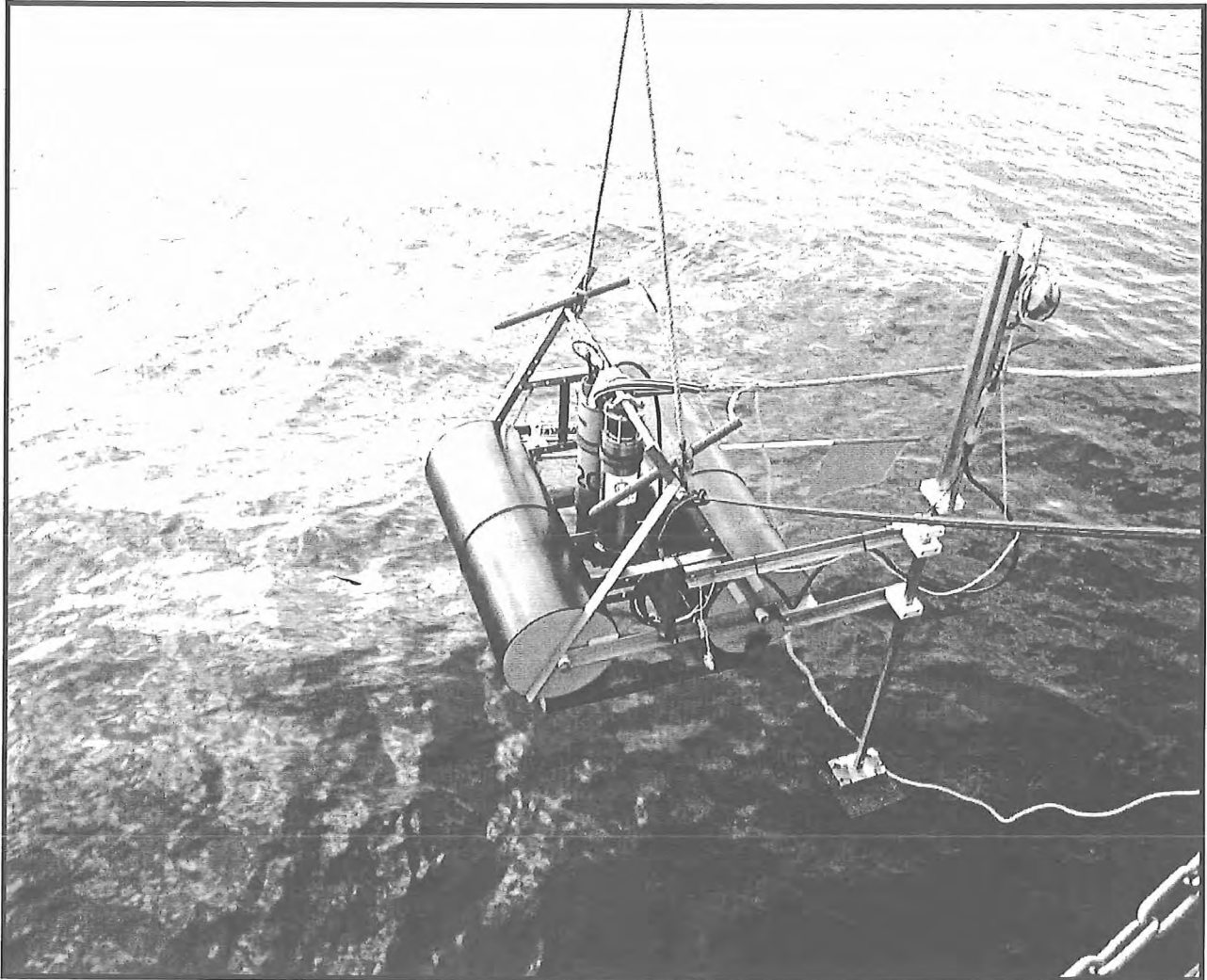


Figure 5.3.7: OBS 2002 upon deployment

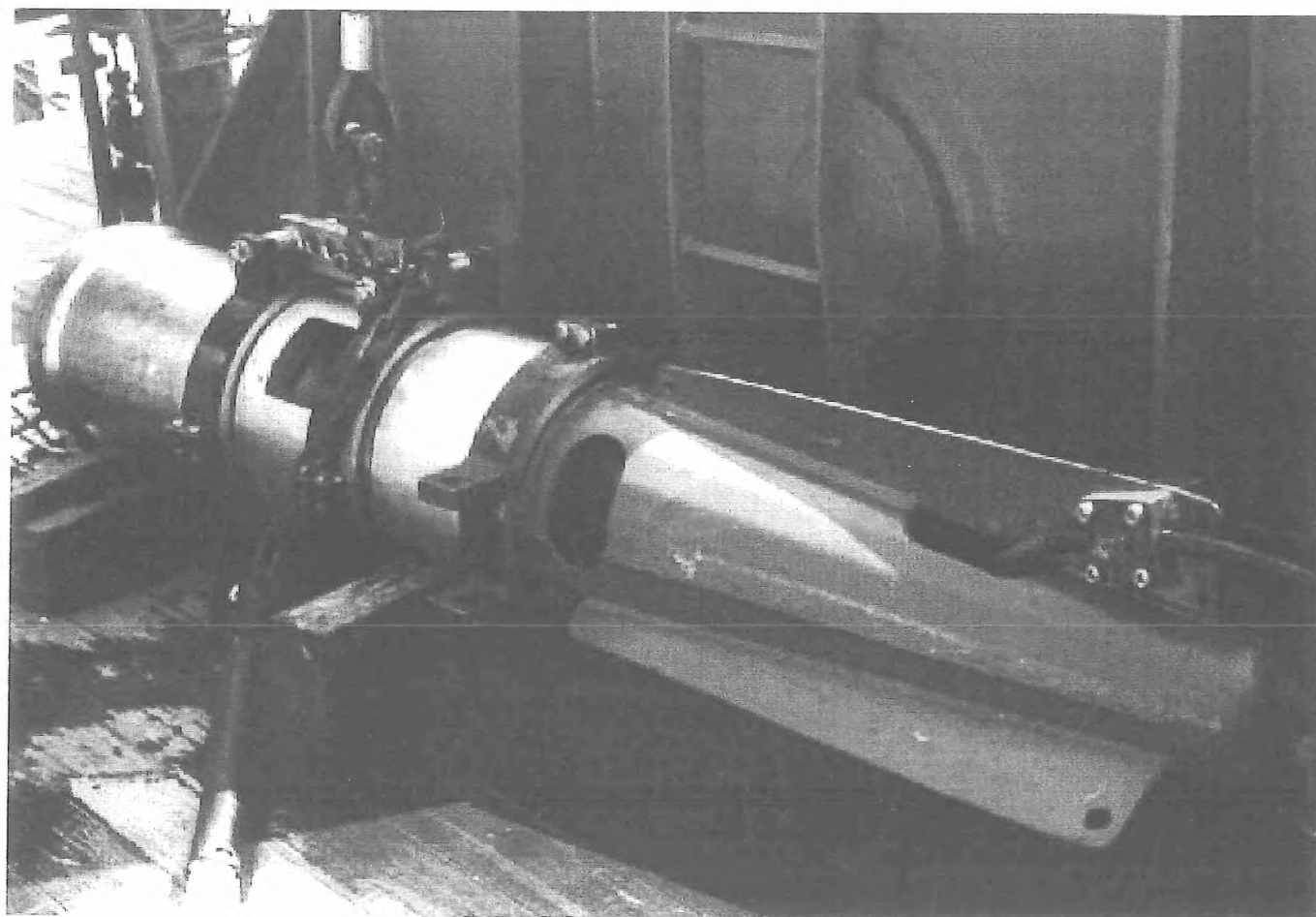


Figure 5.3.8: Photograph of a Bolt PAR 800 CT airgun with 32 l chamber

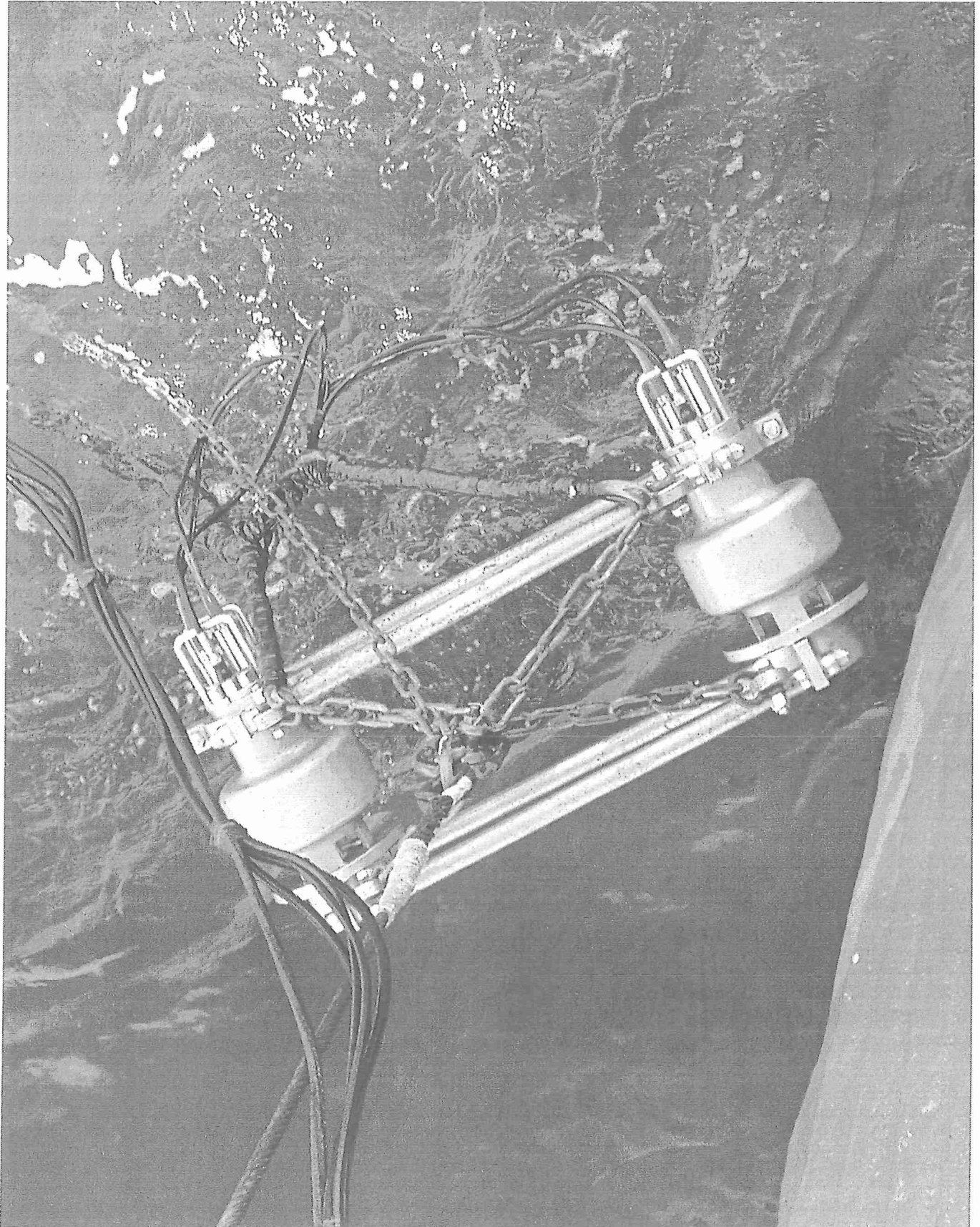


Figure 5.3.9: The G-Gun Cluster.

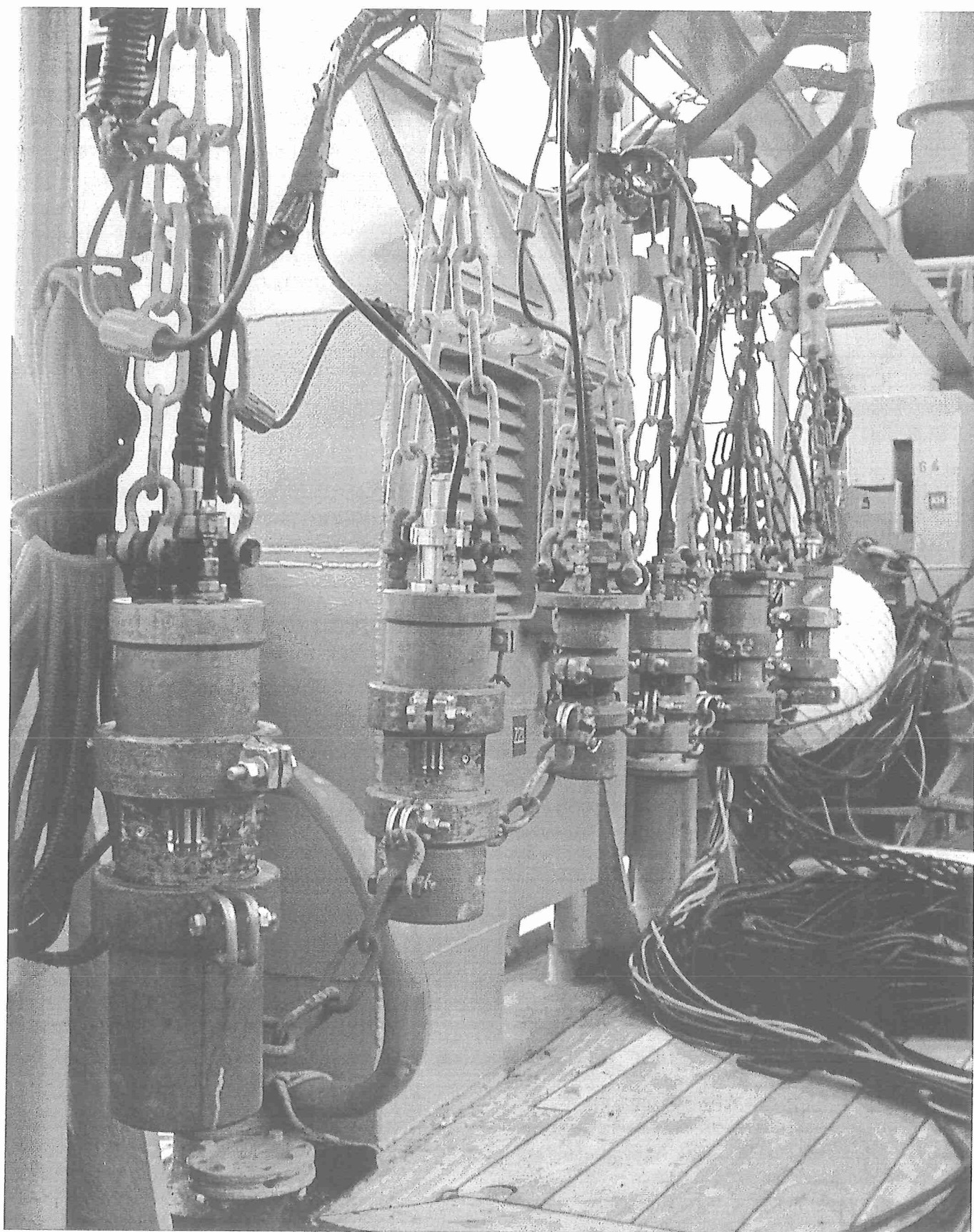


Figure 5.3.11: The Prakla Array, consisting of seven air guns of different volume.

5.4 The Magnetometer

(P. O. Thierer)

During cruise SO163-1 we used a GeoMetrics G801/3 Marine Proton Magnetometer. This unit uses a gasoline-filled sensor with 350 m marine cable and a control unit. During a polarisation cycle an electric current generates a strong magnetic field in the coil and forces the magnetic moments of the protons to be aligned for a short time parallel to the existing field. During the following measuring cycle, i.e. when the electric current is turned off, the previously excited field is removed and the protons “try” to realign themselves with the Earth’s magnetic field. According to the moment preservation law, this happens by precession of the protons with a certain frequency, which is directly proportional to the intensity of the Earth’s magnetic field. Basically, this frequency is measured as AC electric current created by magnetic induction in the coil, amplified, counted and transformed to the magnetic field intensity values (measuring unit: 10^{-9} Tesla = 1 nT), which are recorded.

In order to minimize the influence of the ship’s hull, the sensor of the Magnetometer is towed 180 m behind the ship. This distance of the sensor from the ship was long enough, so that the magnetic influence of the vessel could not produce any disturbances in the recorded magnetic field, so we achieved a resolution of about 5 nT.

On board of RV SONNE, the winch was placed on the port back deck and the sensor was towed to the port side of the vessel. A boom leads the cable about 7 m to the side of the ship in order to prevent it from being tangled with the ship. After having some minor problems at the beginning, the system worked well throughout the trip.

The measured values of the total intensity magnetic field were displayed on a console and written as digital output coded in BCD values. The system was set to deliver one data value every 3 seconds via digital multiport interface to a PC, where a special software was used to store the data together with UTC-time in ASCII tables.

After data backup the files were transferred to a SUN workstation. GPS coordinates and time were taken from the ship’s navigation system and assigned to each magnetic stamp on the basis of the recorded time. The magnetic and the navigation data were resampled to 10 s interval. After optional median filtering they were displayed using GMT plot routines (Wessel and Smith, 1995).

5.5 USBL Underwater positioning

(I. Klaucke, C. Huguen)

During cruise SO163/1 the ship was equipped with a POSIDONIA deployable acoustic array for underwater navigation from OCEANO Technologies (figure 5.5.1). This system is based on a bi-directional exchange of submarine acoustic signals between one or several acoustic transponders and the acoustic array consisting of 1 transmission transducer and two pairs of hydrophones. The acoustic transponders are interrogated by an acoustic signal and will send a 25ms M-FSK (multi-frequency shift keying) reply. The 25ms M-FSK signal is a succession of ten monochromatic pulses (each 2.5ms long) of ten different frequencies ranging from 14.5 to 17.5 kHz. The order of frequencies during the pulse is determined by the M-FSK code. For best detection of the signal with Posidonia, codes 22 or 23 should

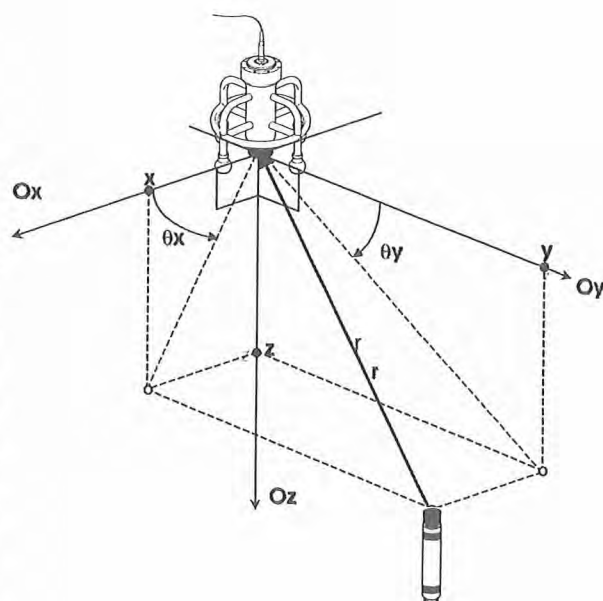


Figure 5.5.1 USBL Positioning principle

be selected. The four reception hydrophones of the array receive this signal that is then transmitted to the processing unit, which measure the phases of the signals and the time between interrogation and reply in order to deduce the relative position of the transponder and calculate its geographical position. The Posidonia positioning system runs in two different modes: the free mode and the towed-fish mode. In the free mode all four hydrophones are used to calculate the position of the transponder. In the towed-fish mode only the hydrophones aligned with the towed vehicle are used together with depth information provided by a built-in depth sensor. The free mode should be used when the transponder is located in a cone of 60 degrees underneath the ship, but this mode will work fine up to an angle of 120 degrees. Beyond this angle the towed-fish mode should be used.

The Posidonia array is deployed in the moon pool onboard FS Sonne. Although tightly fitted when locked, the antenna will not be in exactly the same position after each deployment. Following each deployment of the Posidonia antenna a calibration run for the Posidonia system is necessary. For this calibration a transponder of the ET/RT 8x1 series has to be deployed on the seafloor in 1000-2000 m of water depth (Fig. 5.5.2). Following deployment of the transducer, the ship will run a figure of eight above the transducer location in order to allow for communication between the antenna and the transducer from all sides of the ship and from all directions. A minimum total number of 400 data points should be registered during this calibration run. A special programme of the Posidonia system POSICAL will then allow to eliminate obviously erroneous data points and to calculate values for the bias on yaw, pitch and roll of the vessel.

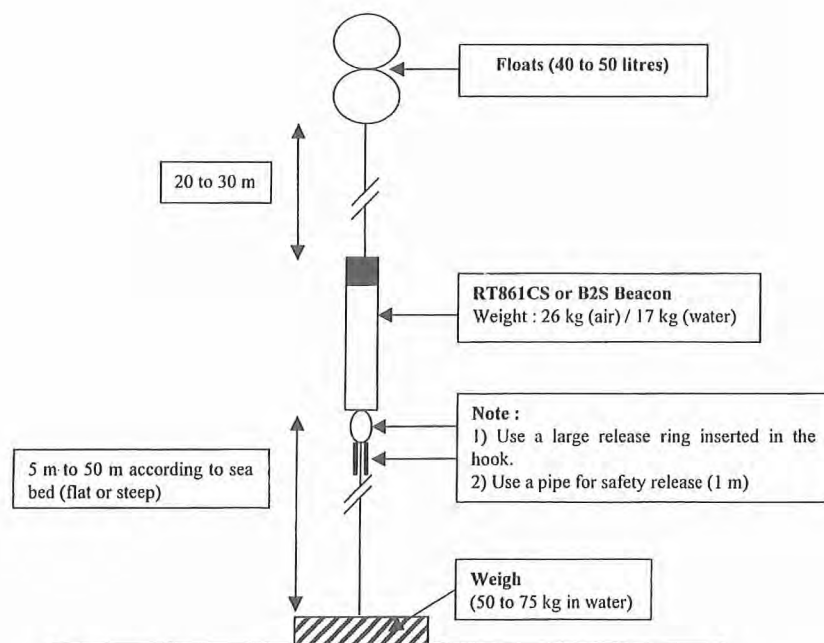


Figure 5.5.2 Mooring used for calibration

5.6 TOBI (Towed Ocean Bottom Instrument)

(D. Masson, I. Rouse, D. Matthew, V. Huehnerbach)

System Description

TOBI - Towed Ocean Bottom Instrument - is Southampton Oceanography Centre's deep towed vehicle. It is capable of operating in 6000 m of water (Fig. 5.6.1). The maximum water depth encountered during the TOBI surveys during this cruise was around 3000 m.

Although TOBI is primarily a sidescan sonar vehicle a number of other instruments are fitted to make use of the stable platform TOBI provides. For this cruise the instrument complement was:

1. 30 kHz sidescan sonar (Built by IOSDL)
2. 6-10 kHz chirp profiler sonar (Built by IOSDL/SOC)
3. Three-axis fluxgate magnetometer. (Ultra Electronics Magnetics Division MB5L)
4. CTD (Falmouth Scientific Instruments Micro-CTD)
5. Swath bathymetry capability
6. Light backscattering sensor (WET Labs Inc LBSS)
7. Pitch & Roll sensor (G + G Technics ag SSY0091)

A full specification of the TOBI instrumentation is given in **table 5.6.1**.

The TOBI system uses a two-bodied tow system to provide a highly stable platform for the on-board sonars (Fig. 5.6.2). The vehicle weighs two and a half tonnes in air but is made neutrally buoyant in water by using syntactic foam blocks. A neutrally buoyant umbilical connects the vehicle to the 600 kg depressor weight. This in turn is connected to the main armoured coaxial tow cable. All signals and power pass through this single conductor.

The deck electronic systems and the logging and monitoring systems were set up along the port side of the aft geophysical laboratory, giving both watchkeepers and scientists a clear view of the incoming data.

TOBI Deployment

The Sonne is equipped with a stern mounted hydraulic 'A' frame but it is too narrow to allow normal athwartships launch and recovery so the vehicle had to be launched in a fore-aft orientation. The main sheave was used for deploying and recovering both the TOBI vehicle and the depressor weight as well as towing during the survey. Launch and umbilical winches were supplied by SOC and mounted on the afterdeck using a custom designed bed plate originally built for Sonne 144 leg 2. The winches were run from a self-contained electrically-powered hydraulic power pack.

TOBI Watchkeeping

TOBI watchkeeping was split into three, four-hour watches repeating every 12 hours. Watchkeepers kept the TOBI vehicle flying at a height of ideally 350 to 400 m above the seabed by varying wire out and/or ship speed. Ship speed was usually kept at 2.5 kt over the ground with fine adjustments if the TOBI vehicle depth carried out by using the winch. As well as flying the vehicle and monitoring the instruments, watchkeepers also kept track of disk changes and course alterations.

The bathymetry charts of the work area were found to be very accurate which helped immensely when flying the vehicle. The ship's EM120 multibeam system provided a real time display of bathymetry.

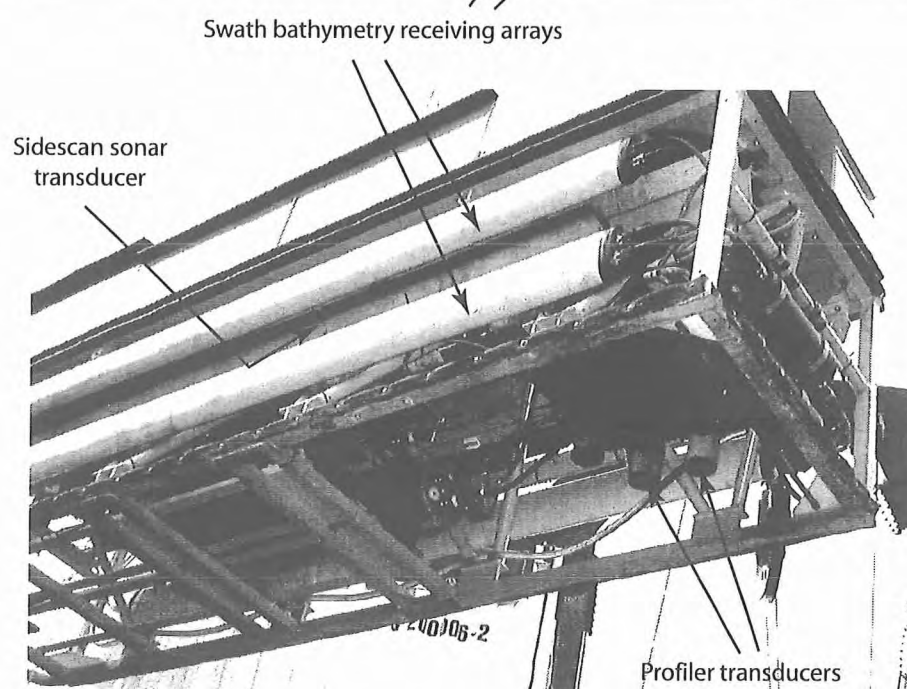
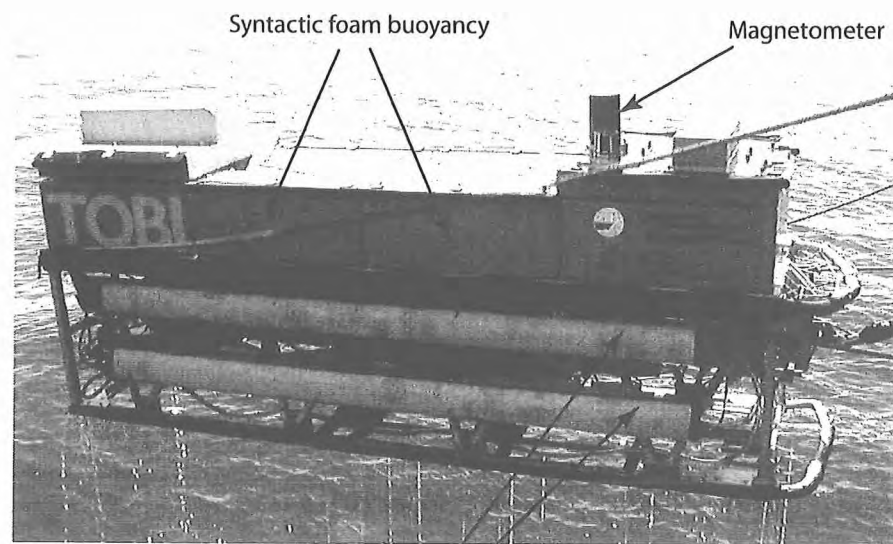


Figure 5.6.1 (top)

Figure 5.6.2 (bottom)

5.6.1. TOBI: Technical Specification

Mechanical

Towing method	Two bodied tow system using neutrally buoyant vehicle and 600kg depressor weight.
Size	4.5 m long, 1.5 m wide, 1.1 m high
Weight	2500 kg in air.
Tow cable	Up to 10 km armoured coax.
Umbilical	200 m long x 50 mm diameter, slightly buoyant.
Tow speed	1.5 to 3 knots (dependent on tow length).

Sonar Systems

Sidescan Sonar

Frequency	30.37 kHz (starboard) 32.15 kHz (port).
Pulse Length	2.8 ms.
Output Power	600 W each side.
Range	3000 m each side.
Beam Pattern	0.8 x 45 degree fan.

Bathymetry Sonar

Transmitter	Uses sidescan sonar.
Receiver	6 hydrophone arrays in 2 housings for each side.
Detection	Single and multi-phase.
Range	Up to 3000 m each side.

Profiler Sonar

Frequency	6 to 10 kHz Chirp.
Pulse Length	26 ms.
Output Power	1000 W
Range	>50 ms penetration over soft sediment.
Resolution	0.25 ms
Beam Pattern	25 degree cone.

STANDARD INSTRUMENTATION

Magnetometer	Ultra Electronics Magnetics Division MB5L.
Range	+/- 100,000 nT on each axis.
Resolution	0.2 nT.
Noise	+/- 0.4 nT.

CTD	Falmouth Scientific Instruments, Micro CTD.
Conductivity	
Range	0 to 65 mmho/cm.
Resolution	0.0002 mmho/cm.
Accuracy	+/- 0.005 mmho/cm.
Temperature	
Range	-2 to 32° C.
Resolution	0.0001° C.
Accuracy	+/- 0.005° C.
Depth	
Range	0 to 7000 dbar.
Resolution	0.02 dbar.
Accuracy	+/-0.12% F.S.

Extra Depth Sensor	AML Pressure Smart Sensor
Range	0 to 6000 dbar.
Resolution	0.1 dbar.
Accuracy	+/-0.05% F.S.
Light backscattering sensor	WET Labs Inc LBSS
Source	880 nm LEDS source and solar blind light detector
Range	0-10 $\mu\text{g l}^{-1}$ suspended particle concentration detection range
Resolution	0.01% FSD at 1 $\mu\text{g l}^{-1}$ resolution
Heading	S.G. Brown SGB 1000U gyrocompass.
Resolution	0.1 degrees.
Accuracy	Better than 1°, latitude < 70°.
Pitch/Roll	Dual Axis Electrolytic Inclinator.
Range	+/- 20 degrees.
Resolution	0.2 degrees.
Altitude	Taken from profiler sonar.
Range	1000 m.
Resolution	1 m.

5.6.2 TOBI Image Processing

Onboard processing equipment during this cruise consisted of a UNIX workstation (SUN ULTRA 10) with 36 Gb of disk space. Final maps containing side-scan sonar imagery were plotted on an A0 plotter. All data were also archived using CD-ROM (via a networked PC).

The ships navigation was recorded online on a UNIX server on the ship. The data was transferred on a daily basis and then tested for time continuity and abnormal speed values. Good navigation data is essential for processing, because the vehicle position is calculated from it. The winch data (wireout) were also recorded online and stored with the sidescan raw data. Errors in wireout values were corrected manually. The TOBI imagery was downloaded from the CD-ROM using a subsample and average factor of 8. This gives the pixel resolution of 6 m and a signal-to-noise ratio improvement by a factor of nearly three.

Each survey area was divided into smaller map blocks to facilitate processing. The approximate size of the blocks was 0.25° by 0.25°. As each area was completed the imagery was processed using the PRISM (v3.1) and ERDAS Imagine (v8.3) software suites to produce geographically registered imagery which can then be composed onto a series of mapsheets.

The processing of TOBI imagery has two main phases: Pre-processing and Mosaicing. The pre-processing stage is the correcting of the side-scan sonar characteristics, removal of sonar specific artefacts and geographical registration of each individual ping. This stage is solely composed of PRISM programs and runs from a graphical user interface.

For the pre-processing stage, the PRISM software uses a modular approach to 'correct' the imagery which is predefined by the user in a 'commands.cfg' file. For this data it was defined as:

```
suppress_tobi -i %1 -o %0
tobtvgr -i %1 -o %0 -a
mrgnav_inertia -i %1 -o %0 -t -u 260 -n navfile.veh_nav
tobtvgr -i %1 -o %0 -h -l 50
```

```
tobslr -i %1 -o %0 -r res , res
edge16 -i %1 -o %0 -m
drpout -i %1 -o %0 -u -f -p -k 201
drpout -i %1 -o %0 -u -f -p -k 51
shade_tobi -i %1 -o %0 -tl,4095
```

To explain this in sonar terms (in order):

- `suppress_tobi -i %1 -o %0`

Removal of any surface reflection (i.e. from TOBI to the sea surface and back) - only a problem in water depths <3000 m, where a bright stripe or line is seen semi-parallel to the ship's track. Removal is only done when the imagery is unambiguous, where the line is true artefact and not an actual seafloor feature. The result can sometimes be seen on the final imagery as a faint dark line.

- `tobtvgr -i %1 -o %0 -a`

Smoothing of the altitude of the vehicle above the seafloor. The altimeter sometimes cannot locate the seafloor, possibly due to very soft sediment reducing the strength of the profiler signal. Smoothing is done by a median filter of the given values, comparing this with the first return seen on the port and starboard sides, and applying a maximum threshold for altitude change if first return and altitude value differ. Generally first return values are used as that value will be used in the slant-range correction.

- `mrgnav_inertia -i %1 -o %0 -t -u 260 -n navfile.veh_nav`

Merging of ship navigation and cable data with the imagery and calculation of the TOBI position using an inertial navigation algorithm. The 'navfile.veh_nav' file contains ship position and cable values and an umbilical length of 200 m is assumed (default) plus an additional 60 m for the distance between the GPS receiver and the point where the cable enters the water. The recorded cable values in the TOBI data are used. Various defaults are assumed: the cable is assumed to be straight, the cable value is assumed to be accurate, and zero cable is set when the depressor enters the water.

- `tobtvgr -i %1 -o %0 -h -l 50`

Replaces the TOBI compass heading with the track heading calculated from the navigation data. A smoothing filter is applied. The heading values are used in the geographic registration process to angle each ping relative to the TOBI position.

- `tobslr -i %1 -o %0 -r res , res`

Slant-range correction assuming a flat bottom. This is a simple Pythagoras calculation assuming that the seafloor is horizontal across track and sound velocity is 1500 ms⁻¹. Each pixel is 8 ms and generally equates to 6 m resolution; any pixel gaps on the output file are filled by pixel replication.

- `edge16 -i %1 -o %0 -m`

Median filter to remove any high or bright speckle noise. A threshold is defined for the maximum deviation for adjoining pixels over a small area above which the pixel is replaced by a median value.

- `drpout -i %1 -o %0 -u -f -p -k 201`

Dropout removal. When the vehicle yaws excessively, it is possible for the 'transmit' and 'receive' phase of each ping to be angled apart. If this exceeds the beam sensitivity value (0.8°) little or no signal is received, creating a dark line on the imagery. The program detects the dropout lines and interpolates new pixel values. If more than 7 dropouts are present concurrently (time = 28 s) no interpolation is done.

- `drpout -i %1 -o %0 -u -f -p -k 51`

More dropout removal, but for smaller, partial line dropouts. If more than 7 partial dropouts are present concurrently (28 s) no interpolation is done.

• `shade_tobi -i %1 -o %0 -tl,4095`

Across-track equalisation of illumination on an equal range basis. This assumes that the backscatter from a particular range should average a given amount for each piece of data. The near-range pixels and far-range pixels are generally darker than mid-range pixels. This is due to the transducer's beam pattern and differences in seafloor backscatter response in terms of angle of incidence. This procedure typically amplifies the near and far-range pixels by about 1.5 and reduces the mid-range pixels by 0.8.

Once this has been done for a piece of data the individual pings are placed on a geographic map. To emulate beamspreading the pixels are smeared over a small angle (0.8°) if no other data is present in the area of those pixels. As survey tracks are designed to overlap the imagery at far range, any overlapping data pieces are placed on separate layers of the same map. This allows user intervention to define the join where one piece touches the other. If small pixel gaps are visible between the geographically mosaiced pings, these are filled with an interpolated value plus a random amount of noise (but having the same variance as the surrounding data pixels).

The second phase (mosaicing) allows the user to view all the 'layers' of data for an area. The software used is a commercial package named ERDAS Imagine (v8.3). Within this software the different layers can be displayed in different colours to distinguish the layers with data that overlaps. In order to merge the different layers and their data together polygons (Areas of Interest -or AOI) are drawn by the user to define the join lines between layers and then applied to create a single layer final image map. This procedure can also be used to remove shadow zones and areas of no data. The program that merges all data within selected AOIs into the final single layer image is called 'addstencil'. Several of these final images can then be mosaiced together into an overall image from which maps can be created in different projections and spheroids, including scales, co-ordinates and text. Also annotation such as ship track, vehicle track and dates and times can be added to the map. The map can then be plotted on the A0 plotter and/or converted into other format e.g. TIFF, JPEG, generic postscript etc. to be used for further analysis on PC, Macintosh or UNIX workstations.

5.7 DTS-1 deep-towed sidescan sonar

(I. Klaucke, C. Huguen, T. Schott, W. Weinrebe)

The GEOMAR DTS-1 deep-tow sidescan sonar (Fig. 5.7.1) was used during cruise SO163/1 in order to provide detailed acoustic imagery of the seafloor of a number of target selected on the basis of TOBI mid-range sidescan sonar data. Sidescan sonar mapping is a well established technique and a thorough summary of its theory and principles would be too long for this report. For more detailed information about sidescan sonar principles and data processing, please refer to Blondel and Murton (1998). The instrument used here allows imaging the backscatter intensity of the seafloor at high resolution due to 75 and 410 kHz sidescan transducers. This surface information can be integrated with very high-resolution subbottom information of the uppermost sedimentary layer, therefore allowing volume estimates of sedimentary units at the seafloor. During cruise SO163/1 a number of targets including zones of active mud volcanism, faulting, fluid venting and mass wasting have been investigated in detail.

5.7.1 Technical description of the instrument

5.7.1.1 Underwater set-up

The DTS-1 sidescan sonar (Fig. 5.7.1) is a EdgeTech Full-Spectrum (FS-DW) dual-frequency, chirp sidescan sonar working with 75 and 410 kHz centre frequencies. The 410 kHz sidescan sonar emits a pulse of 40 kHz bandwidth and 2.4 ms duration (giving a range resolution of 1.8 cm) and the 75 kHz sidescan sonar provides a choice between two pulses of 7.5 and 2 kHz bandwidth and 14 and 50 ms pulse length, respectively. They provide a maximum resolution of 10 cm. In addition to the sidescan sonar sensors, the DTS-1 contains a 2-16 kHz, chirp subbottom penetrator providing a choice of three different pulses of 20 ms pulse length each: a 2-10 kHz pulse, a 2-12 kHz pulse and a 2-15 kHz pulse giving nominal vertical resolution between 6 and 10 cm. The sidescan sonars and the subbottom penetrator can be run with different trigger modes: internal, external, coupled and gated triggers. Coupled and gated trigger modes also allow to specify trigger delays. The sonar electronics provide four serial ports (RS232) in order to attach up to four additional sensors. One of these ports is used for a Honeywell attitude sensor providing information on heading, roll and pitch. Finally, there is the possibility of recording data directly in the underwater unit through a mass-storage option with a total storage capacity of 2 x 30 Gbyte.

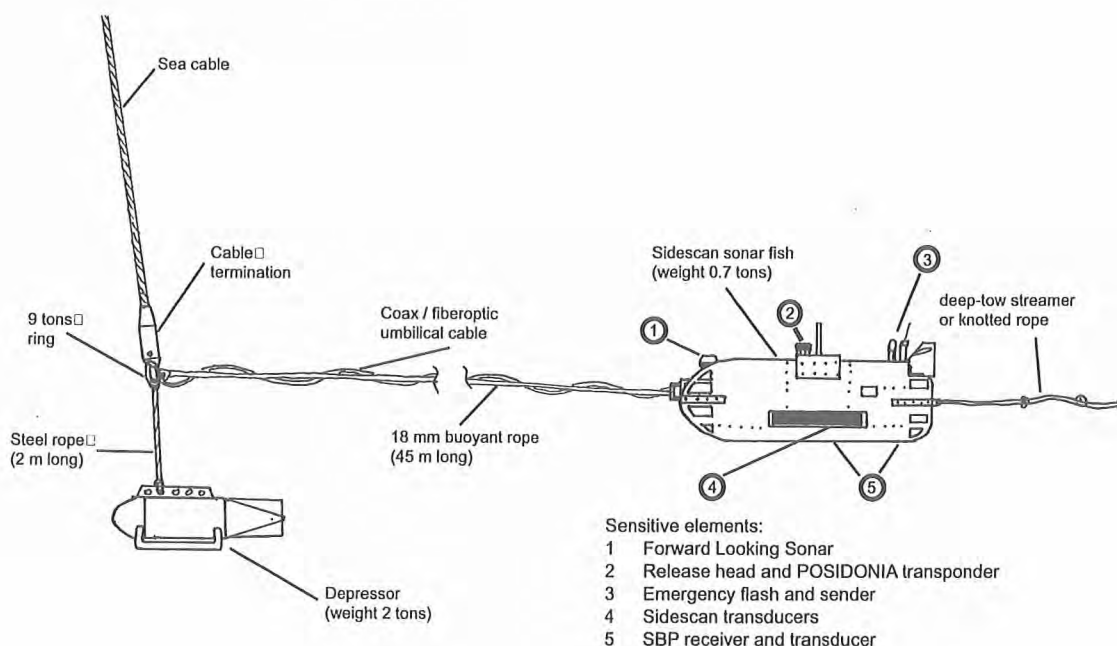


Figure 5.7.1 The DTS-1 towing configuration

The sonar electronic is housed in a titanium pressure vessel mounted on a towfish of 2.8m x 0.8m x 0.9m in dimension. The towfish houses a second titanium pressure vessel containing the wet-end of the SEND DSC-Link telemetry system and the bottom PC of the seismic streamer data acquisition system (see also section on deep-tow streamer). In addition, a Posidonia capable OCEANO releaser with separate receiver head and a NOVATECH emergency flash and sender are included in the towfish. Additional syntactic foam has also been placed in the front of the towfish in order to improve the fish's towing behaviour. For the same reason, the towfish has been fitted with a deflector at the rear. This deflector has five positions from 0 to -5 and is designed to reduce the pitch of the towfish.

The towfish is connected to the sea cable via the depressor (2 tons weight) through a 40 m long umbilical cable. The umbilical cable is tied to a buoyant rope that takes up the actual towing forces. An additional rope has been taped to the buoyant rope and serves to pull in the instrument.

5.7.1.2 Laboratory set-up

The laboratory set-up consists of four elements: the dry-end of the SEND DSC-Link telemetry, the topside PC of the streamer acquisition system, the Edgetech surface interface unit FS-IU and the topside unit running ELAC Hydrostar Online software (Fig. 5.7.2). In the basence of the deep-towed seismic streamer, the only function of the bottom and topside PC of the streamer acquisition system is to provide a serial link between the OCEANO releaser operating in responder mode and the Posidonia topside unit. Hydrostar Online allows general running of the sidescan sonar and subbottom penetrator operations as well as onscreen display of a subset of the acquired data. Unfortunately some additional settings such as the trigger mode or data window size can only be changed by accessing the underwater electronics directly via the FS-IU. The FS-IU also runs JStar, a diagnostic software tool, that also allows running some basic data acquisition and data display functions.

5.7.1.3 Software

The main operations of the DTS-1 sidescan sonar are essentially run using Hydrostar Online, a multibeam bathymetry software developed by ELAC Nautik GmbH and recently adapted to the acquisition of EdgeTech sidescan sonar data. For cruise SO-163 a new version of HydroStar Online (version 3.3.3) with improved

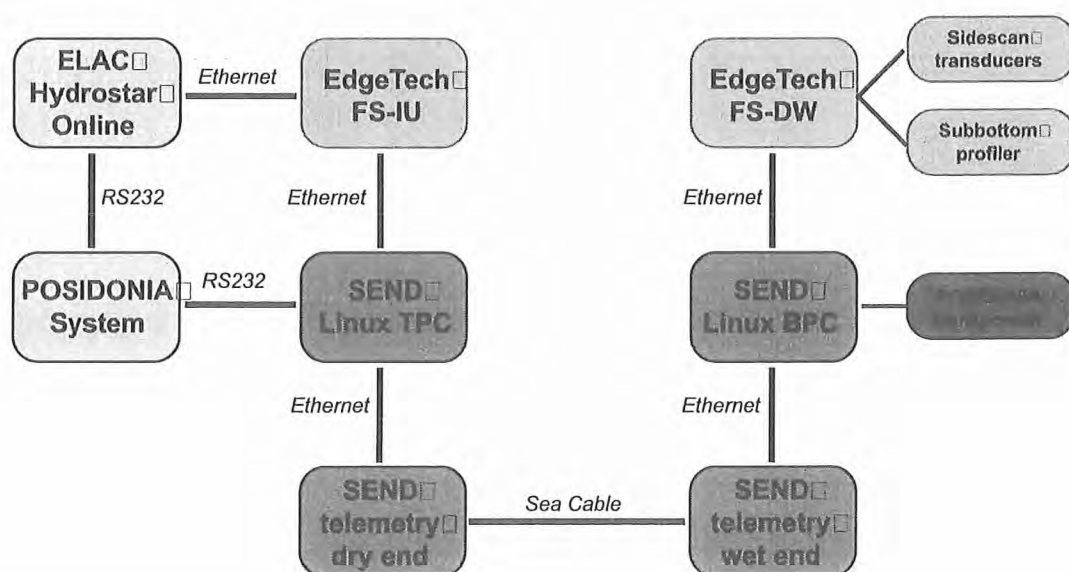


Figure 5.7.2 The DTS-1 electronics configuration

onscreen data representation and time synchronisation has been available. This software package allows onscreen representation of the data, of the fish's attitude, and of the towfish's navigation. It also allows setting some principle parameters of the sonar electronics, such as the selected pulse, the range, the power output, the gain, the ping rate, and the range of registered data. However, this version does not allow to set the trigger mode or the master subsystem in coupled trigger mode. HydroStar Online also allows to start and stop data storage either in XSE-format on the HydroStar Online computer or in JSF-format on the FS-DW. Simultaneous storage in both XSE and JSF-formats is also possible. HydroStar Online creates a new XSE-file when a file size of 10 Mb is reached, while a new JSF-file is created every 20 Mb. How fast this file size is reached depends on the amount of data generated, which in turn essentially depends on the use or not use of the high-frequency sidescan sonar. The amount of data generated is also a function of the sidescan sonar and subbottom pulses and of the data window that is specified in the sonar.ini file on the FS-DW. The data window specifies the range over which data are sampled. Proper selection of this parameter strongly depends on the selected range of the sidescan sonar systems in order to avoid good data to be cut-off, or to prevent too large amounts of unuseful data using up storage space.

Further handling of the data is still problematic as neither the XSE nor the JSF data format can be read directly by our sidescan sonar and seismic (for the subbottom penetrator) processing software, which require NetCDF and SEG-Y formats, respectively. Major efforts to achieve data conversion into formats for use with either PRISM or MB Systems are still going on onshore and will soon allow proper processing of the sidescan sonar data.

5.7.1.4 Deployment and recovery procedures

Operations for deployment and recovery of the sidescan sonar are a bit demanding and require relatively calm sea for handling that is safe for both crew and instrument. During cruise SO163-1 three seamen, the boatswain, the sidescan sonar technician and the responsible scientist were on deck for the operations. The sidescan sonar instrument should ideally be towed via the A-frame. With no speed made by the ship the kite tail is first thrown into the water and let to drift away. Then the sidescan towfish is heaved into the water (Fig. 5.7.3) and released with a special hook allowing to detach the crane cable. The sidescan fish then also drifts astern with minimal speed made by the ship. Meanwhile, the buoyant rope is secured. Then the depressor

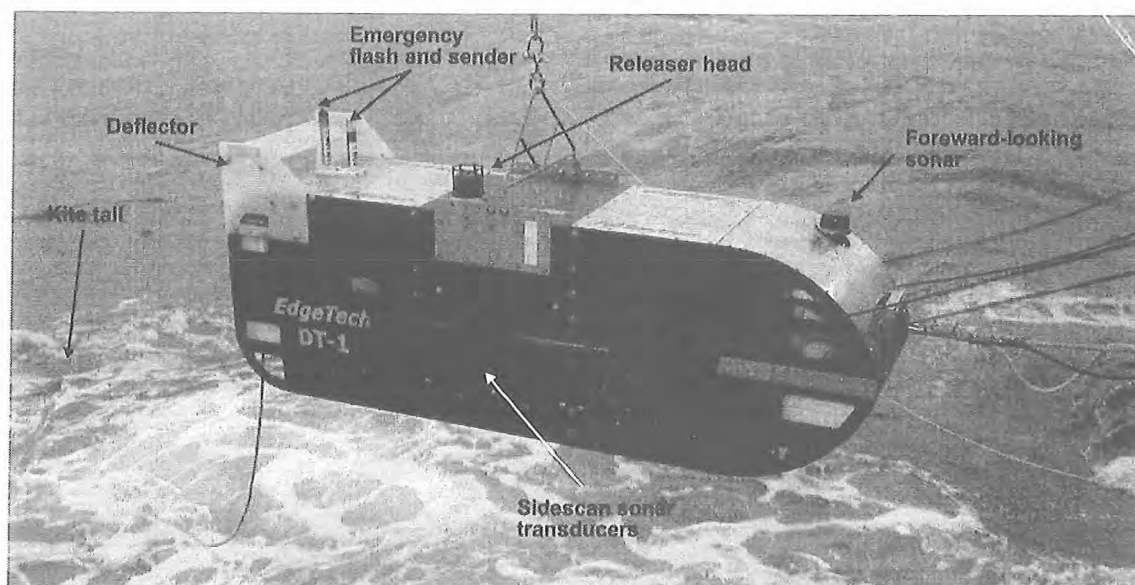


Figure 5.7.3 Picture of the DTS-1 towfish during deployment

5.8 Mini corer (MIC)

(S. Grandel)

Sediment samples were taken by means of a mini corer, which is a smaller version of the multiple-corer (Fig. 5.8) . The device is equipped with four sampling cores of about 50cm length, and was deployed together with the CTD-rosette unit, hanging approximately 20m beneath. Bottom contact was recorded by a pinger, and penetration of the cores into the sediment is controlled by deployment speed and weight of the device. While lifting the device, cores are closed by lids on the top and bottom. Collected sediment cores were approximately 25cm long, overlain by a clear bottom water column, indicating a sediment surface virtually undisturbed by sampling procedures (Barnett et al., 1984).

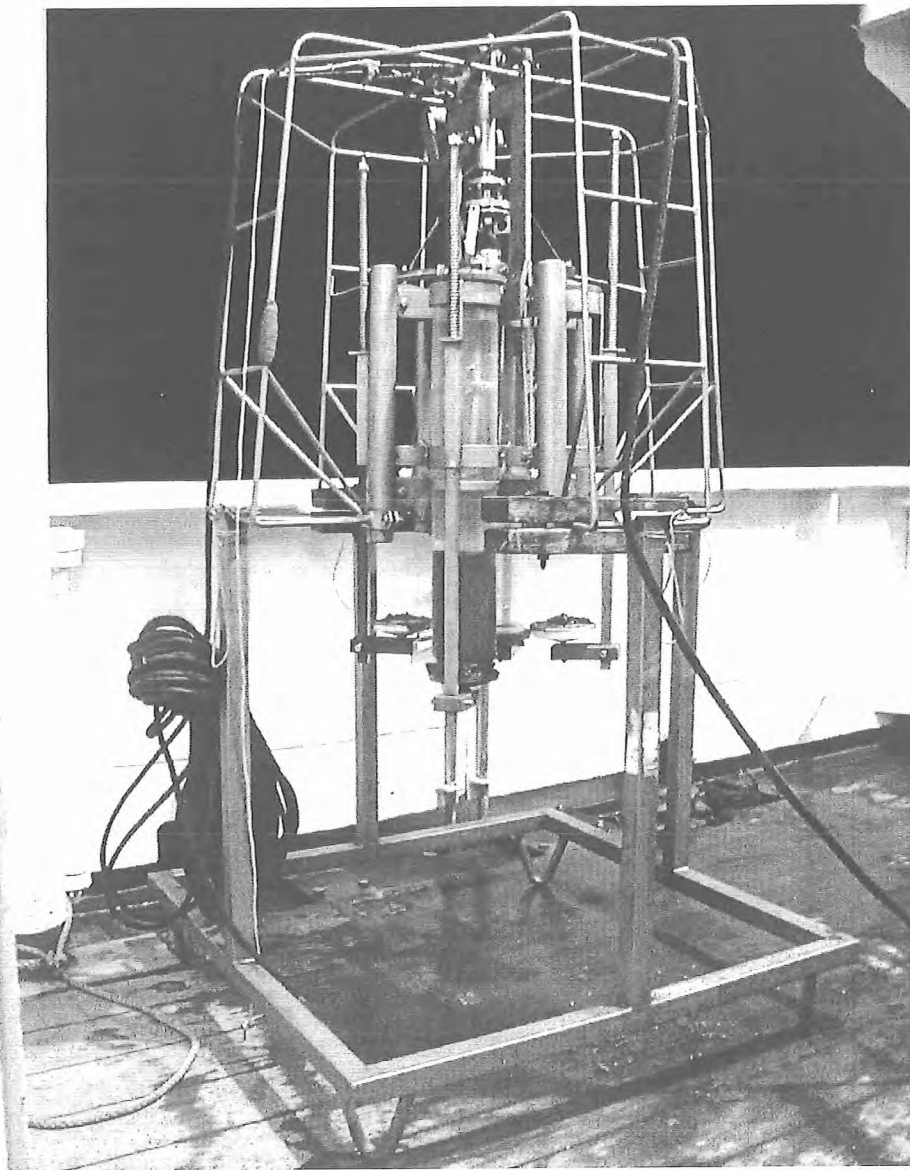


Figure 5.8: Mini corer used for the sampling of the upper sediment column.

6.1 Multibeam bathymetry

6.1.1 Multibeam bathymetry results from Sonne 163 leg 1

(C. R. Ranero, P. O. Thierer and watchstanders)

During Sonne 163 leg 1 multibeam bathymetry was collected during the entire leg in parallel with dedicated surveys for Tobi, DTS and OFOS. Time dedicated for the collection of only bathymetry was mainly available during down time of Tobi surveys for reparation of the instruments and some transits and therefore most of the bathymetric data are located in areas where there was existing bathymetric coverage. The only area not previously surveyed is west of the Nicoya Peninsula where bathymetric soundings filled some existing data gaps in the upper slope during Tobi repairs. In spite of covering areas with existing data the high performance of the new Simrad system produced high quality data that improved notably the resolution of the bathymetric coverage. The Simrad system, used for the first time in this margin where previously existing data had been collected with the hydrosweep system of R/V Sonne and R/V M. Ewing greatly outperforms the capabilities of the older systems and some of the detail surveys carried out during this leg have produced a brand new topographic map of areas that were previously mapped at lower resolution. Some of the description in this section will be dedicated to compare the results from the new system with older maps and to try to strengthen the great potential of the system if used in properly planned, dedicated surveys.

By the end of leg 1, the large majority of the bathymetric soundings had been cleaned using MBsystem version 5beta12 (Caress and Chayes, 1996) and only cleaned data were used for the illustrations in this section.

In the next points we will show a series of areas where bathymetric coverage with Simrad data has brought some interesting new results. The areas that will be discussed are indicated by the boxes in figure 6.1.1.1.

6.1.1.1 Mapping the ocean plate.

We run a small survey in an area of the ocean plate that had been previously covered with hydrosweep bathymetry. The difference between the results produced by both systems can be observed in figure 6.1.1.2. The purpose of the survey was to produce a detailed map to use for planning an OFOS study of the area where a cruise by R/V M. Ewing had found anomalously high heat flow values and indications of anomalous composition of fluids collected by piston coring. Those observations seem to point to vigorous deep penetrating fluid flow in the ocean lithosphere (Fisher et al., 2001).

Figure 6.1.1.2a shows the data collected with the Simrad system in a survey with two parallel tracks run at 5 knots and with a narrow aperture of the beams (4 km on either side) in order to obtain high quality data and a suitable coverage to grid the data at ~ 50 m in the area of interest. Figure 6.1.1.2b shows the compilation of soundings from previous cruises that could only be gridded at ~ 100 m to avoid spatial aliasing. The comparison of the images shows some straightforward differences. The hydrosweep data displays a considerable amount of noise that masks much of the fine scale topography of the plate. The noise of each individual track is particularly clear in this area because the seafloor topography is rather subdued. In contrast, the map constructed with the Simrad data shows a detailed image of a number of morphological features not observed in the other map. The new bathymetry shows that the heat flux measurements were carried out in an area where the ocean plate begins to be pervasively faulted due to bending as it approaches the trench. The faults run with a NW-SE strike with the hangingwall to the NE. In addition, a small graben structure about 200 meter wide is conspicuous in the SE of the surveyed area. The fault offsets, of the order of 50 meter (Figure 6.1.1.2c), and associated half grabens, are well displayed in the Simrad bathymetry, whereas only the graben is clearly visible in the SE part of the hydrosweep map. Although a more fair comparison could be done by removing some of the worse tracks (or beams) from the hydrosweep compilation, a detailed map like that of the Simrad system could most likely not be achieved.

The heat flux survey carried out with R/V M. Ewing is in between to main lineaments of faulting. The faults to the west of the measurements area smaller offset than to the east, where a SE striking fault with ~ 100 m throw is visible at the NE edge of the coverage. The faults have formed depressions (half grabens) that are filled in the deepest parts with a few tens of meters of strata (Figure 6.1.1.2d) that discordantly onlaps the basins and probably represent recent turbidites. The basal layers of the infilling strata are tilted forming wedge like geometries against the fault scarps indicating a syntectonic deposition.

All the above observations indicate that the heat flux data were collected in an extensional environment of currently tectonically active normal faults related to the bending of the ocean lithosphere. This could indicate that the anomalous fluid flow might be related to the extensional faulting. The lateral variability in the heat flux measurements (Figure 6.1.1.a) might indicate areas of inflow in the ocean plate (low values) and of outflow (local high values).

6.1.1.2 Ridge subduction and middle slope normal faulting and chemoherm mounds

An example of the middle slope structure offshore Nicoya peninsula is shown in Figure 6.1.1.3a. Much of area in the figure was covered during a DTS survey and parts during transit tracks. The area covered during the DTS survey is most of the center of the figure and can be easily identified because it shows smoother morphology lacking high frequency noise. The areas covered during transits are located to the SE and SW-W of the figure and two overlapping tracks leaving the area towards the NE. All those areas display high frequency noise that partially masks the real morphology and is close to aliasing at the ~ 50 m grid in the figure. A perspective view of the area is shown in figure 6.1.1.3b where the noise is also displayed.

Well above the noise level are a series of real morphological structures. In the SW of the images occurs a landslide with a concave up geometry and a clear semicircular headwall. The slide is in line with a small linear ridge in the ocean plate (figure 6.1.1.1) that is subducting beneath the upper plate (von Huene et al 2000). A seismic profile landward of the slide indicates that the ridge continues beneath the middle slope (Barckhausen et al., 2001) and the slide has probably been formed by the effect of the subducting ridge. The morphology within the slide is rougher than in the surrounding areas indicating blocks of slid material and a recently detached semicircular block is visible downslope of the headwall. Normal faults immediately upslope of the headwall and striking oblique to other faults in the areas probably are created by uplift due to the subducting ridge. In addition, two other sets of prominent structures are visible in the diagrams.

A series of roughly margin parallel offsets in the seafloor are the surface expression of normal faults. The faults are widespread in the area and have been mapped with bathymetry in other areas of the margin (e.g. Ranero et al., 2000b). Normal faults in other areas of the margin have been observed in seismic reflection records cutting deep into the slope sediment section and some of them possibly into the basement (McIntosh et al., 1993; Ranero and von Huene, 2000). Faults upslope of the slide and to the W seem to downthrow downslope, whereas some of the faults observed to the NE of the slide downthrow upslope.

A swath of mounds, few hundred meters to ~1 km wide and some tens to 100-200 meter high, spreads across the middle slope. These structures, initially interpreted as mud volcanoes (e.g. Shipley et al., 1990, Shipley et al., 1992) seem to be mounds formed essentially by carbonate chemohermes with abundant indications of active fluid flow (see OFOS chapter for a detail description; Kahn et al., 1996; Borhmann et al., in press). Although they cluster in a swath in the middle slope, two others occur upslope at the edge of the area of figures 6.1.1.3a and 6.1.1.3b (@ 85° 50' / 9° 44.5'). Is it interesting to note that the swath of carbonate mounds occurs in the area where, with the current resolution, also occurs the inception of widespread normal faulting. A more detailed image of the relationships between faulting and carbonate mounds is shown in the perspective diagrams of figures 6.1.1.3c and 6.1.1.3d that encompass the area of the detailed DTS survey where the coverage is the best allowing gridding at ~ 25 meter. A small portion of the area downslope in the figures is covered by transit tracks and aliasing has produced a poor quality image compared

to the ~50 meter grids of figures 6.1.1.3a and 6.1.1.3b. A preliminary interpretation of the images is that the carbonate mounds grow associated to fluids expelled along the faults, perhaps related to methane coming from deep in the sediment section. The sediment cover in this area is 1.5-2 km thick and bottom simulating reflections (indicating gas hydrates) occurring at ~400 m depth are widespread (Pecher et al., 1998; von Huene et al., 2000). Fluids either from sediment compaction or perhaps from greater depths may rise along the faults and locally disassociate hydrates releasing methane that will rise along faults. This might explain the spatial occurrence of most mounds and the inception of widespread faulting. Normal faulting may also occur upslope of our observations (e.g. McIntosh et al, 1993) where a few other mounds also occur but they might be below the resolution of our images. This in turn may indicate that faults in the middle slope are more frequently active favouring a more continuous fluid supply.

The fate of the mounds may also be governed by normal faulting. Faults offset is larger in the faults located downslope of the mound swath, and that may indicate that fault grow leads to the destruction of the mounds. Some of the mounds at base of the swath seem to be cut by the faults (e.g. @ -85° 58' / 9° 43') and no mounds are observed towards the lower slope (although our resolution is worse in that area, mounds several hundred meters wide should be visible). If this preliminary model were correct it would imply that the carbonate mounds develop for long periods of time. There is no age information on the faulting, but it is probably related to the collapse of the margin produced by basal tectonic erosion (Ranero and von Huene, 2000). This process leads to progressive subsidence of the margin in which the upper slope will eventually become a middle slope and even a lower slope at the end of the process in such a model, before being tectonically eroded away. The rates of erosion and subsidence are currently poorly constrained but it certainly implies millions of years. The top of the basement drilled in the ODP leg 170 (Kimura et al., 1997) was a beach rock about 17 Ma and it is currently beneath the lowermost slope (Vannuchhi et al., 2001).

6.1.1.3 Upper slope canyons, middle slope normal faulting and chemoherm mounds

A second area showing the relationships between faulting and carbonate mounds has been chosen to illustrate the great improvement obtained with the Simrad system. In addition, a spatial relationships among upper slope canyons, normal faulting and carbonate mounds is also starting to emerge from the new data (Figure 6.1.1.4). Much of figure 6.1.1.4 is in an area that had been previously surveyed with hydrosweep bathymetry, usually collected at speeds ranging from 8 to 12 knots. The quality of the previous data set can be appreciated in the lower left area of the figure where gridding at ~ 50 meter has yielded a poor image of the morphology and numerous small gaps in the coverage. In contrast, the rest of the figure shows a fairly new image of the slope morphology, in spite of the gaps among tracks in the uppermost slope. Much of the Simrad data in the figure was collected during a Tobi survey with speed of ~ 2.5 knots, but parts of the data were also collected at ~ 8 knots yielding comparably good data.

A complex pattern of canyons occurs in the upper slope running roughly downslope and with numerous tributary gullies that form a "herring bone" texture. At the mouth of the main canyons two areas show a fairly smooth morphology probably indicating recent sediment deposition and that the canyons are active (around -86° 20' / 10° 20' and -86° 18' / 10° 06'). Around these areas, numerous slope-parallel scarps break the seafloor indicating active faulting. The fact that the mouth of the canyons is associated to smoother topography may indicate that sediment supply is high enough to mask fault offsets. Most faults seems to occur downslope of the canyons but a few cut transverse across the them (e.g. @ -86° 15' / 10° 12').

Several mounds appear clearly displayed in the bathymetry. A cluster of 4-5 mounds occurs at -86° 24.5' / 10° 22' and an isolated large mound about 1 km wide and 200 m high occurs at -86° 18.5' / 10° 18' (surveyed with OFOS). Another cluster of mounds occurs at -86° 12' / 10° 1'. All those large mounds occur in the area of widespread extensional faulting inception, similarly to the mounds described in the previous section.

6.1.1.3 Uplift and slope failure related to seamount subduction: the Jaco scarp and uplift

Another example of the unprecedented resolution and a wealth of structural information obtained with Simrad multibeam bathymetry in the margin is the area of the Jaco scarp and uplift (Figure 6.1.1.4a and 6.1.1.4b). The uplift is produced by a subducting seamount (von Huene et al., 2000). The subducting seamount has uplifted the upper plate and produced numerous fractures that offset the seafloor and probably allow fluid flow (Ranero and von Huene, 2000). Large subducting seamounts, like those currently underthrusting central Costa Rica and causing the Jaco uplift break the entire upper plate where the plate thickness is less than ~ 10 km and may allow fluids from even the plate boundary to rise along the fractures (Ranero and von Huene, 2000).

In the wake of a subducting seamount forms a furrow parallel to the convergence vector (Figure 6.1.1.1; von Huene et al., 2000). The furrows indicate that material from the upper plate is removed during the seamount subduction process, and seismic images document that basal erosion along the underside of the upper plate is responsible for much of the material missing (Ranero and von Huene, 2000).

The new bathymetric data show that downslope from the edge of the scarp there is a terrace very similar to the seafloor surface of the uplift area (Figure 6.1.1.5c). This terrace probably corresponds to the last slid block that has moved a few hundred meters downslope as a coherent block. The block is dissected by several faults perpendicular to its strike and that seem to be the continuation of the faults observed in the uplift area. The fact that an elongated block, dissected by several faults, remains coherent after several hundred meters of displacement along the detachment plane might indicate creeping rather than a single catastrophic movement.

The edge of the uplifted area shows the largest amount of deformation with a network of fractures that break the upper plate in blocks similar to the block forming the terrace downslope of the scarp edge. The faults forming the lateral walls of the furrow seem to penetrate into the uplift area and crosscutting roughly perpendicular fractures dissect the terrain in blocks that later will slide down the scarp (Figure 6.1.1.5c). The area of maximum uplift, where faults have the largest offsets is located about 2 km upslope from the scarp and may indicate the position of the top of the seamount.

6.1.1.4 Summary

Bathymetric data collected with the new Simrad system and carefully cleaned with the Mbsystem has yielded a new, detailed image of the structure of the margin in areas where previous bathymetric sounding had a lower resolution. The use of dedicated surveys planned taking into account water depth and desired grid cell size to calculate ping aperture, and ship speed (ping rate) can produce high quality terrain data for maps of unprecedented high resolution.

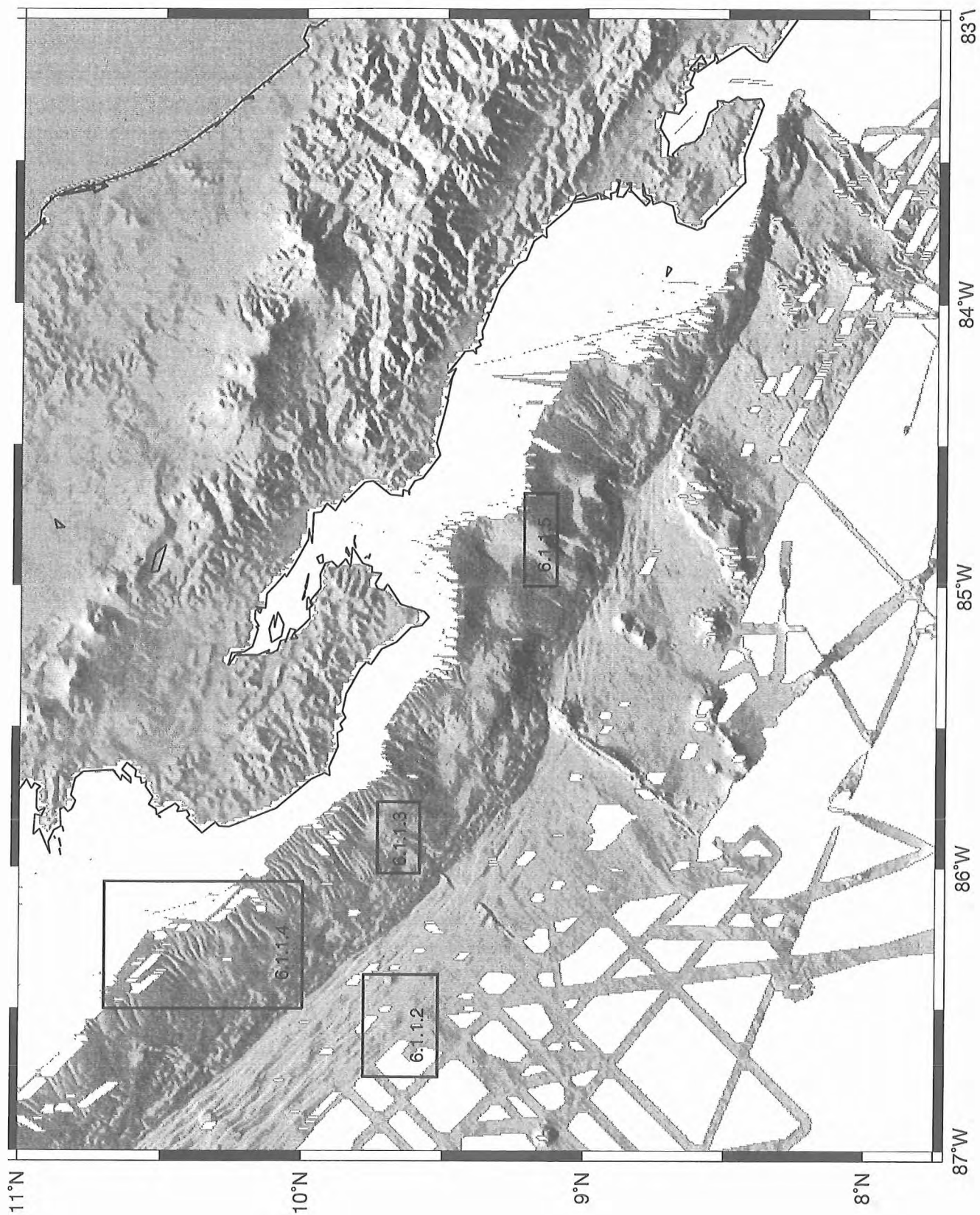


Figure 6.1.1.1. Compilation of multibeam bathymetry along the area studied during SONNE 163 leg 1. Boxes indicate location of areas discussed in detail in this chapter and numbers refer to figures.

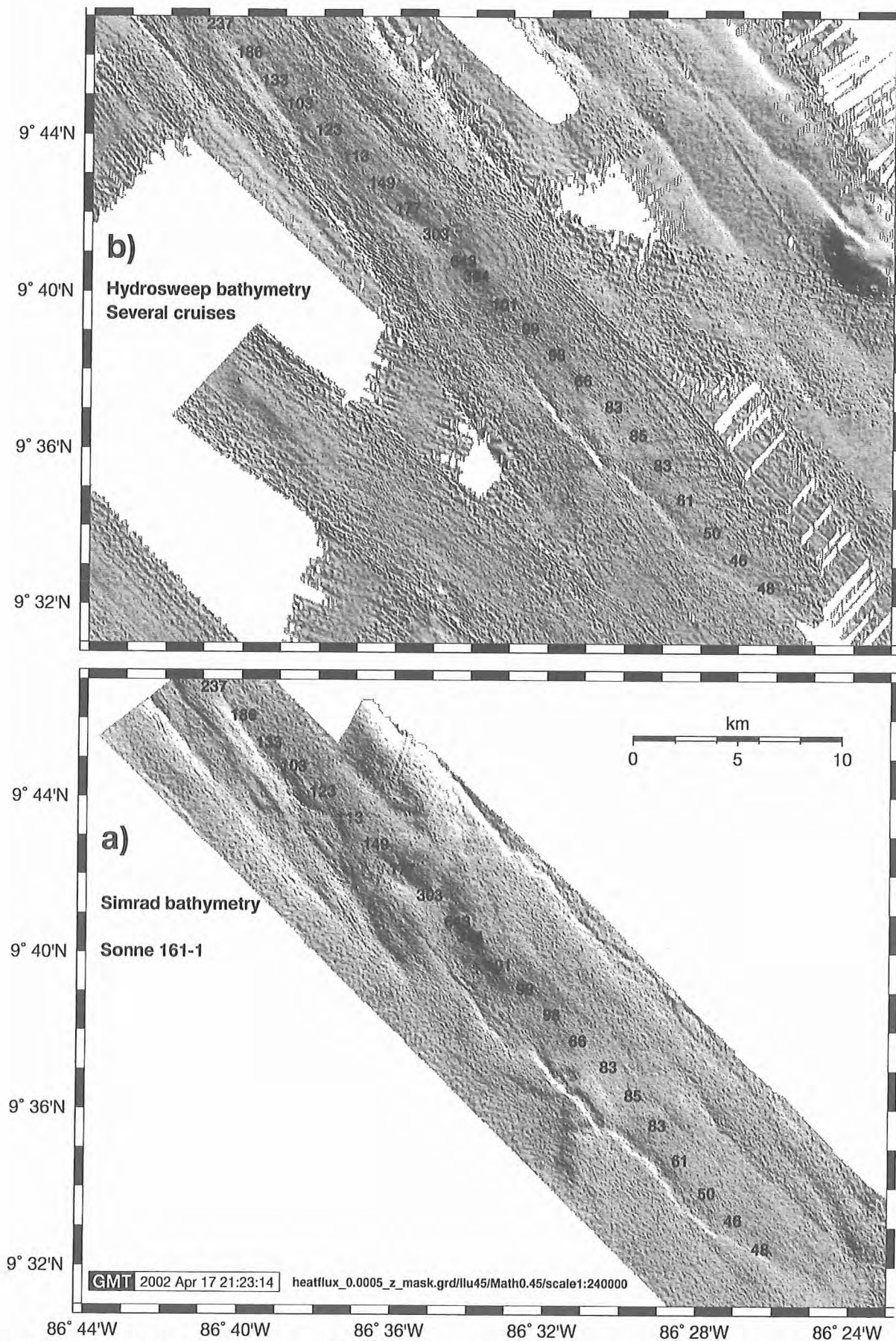


Figure 6.1.1.2. a) Shaded relief map of Simrad bathymetry on the ocean plate. Grid size is ~ 50 meter. Numbers along the bathymetric swath are heat flux values from Fisher et al. (2001) collected with R/V M. Ewing cruise 0104. Note some anomalously high values (303, 364, 643, 2370). **b)** Compilation of hydrosweep bathymetry from several cruises in the same area. Note the better resolution of Simrad data. Location on figure 6.1.1.1.

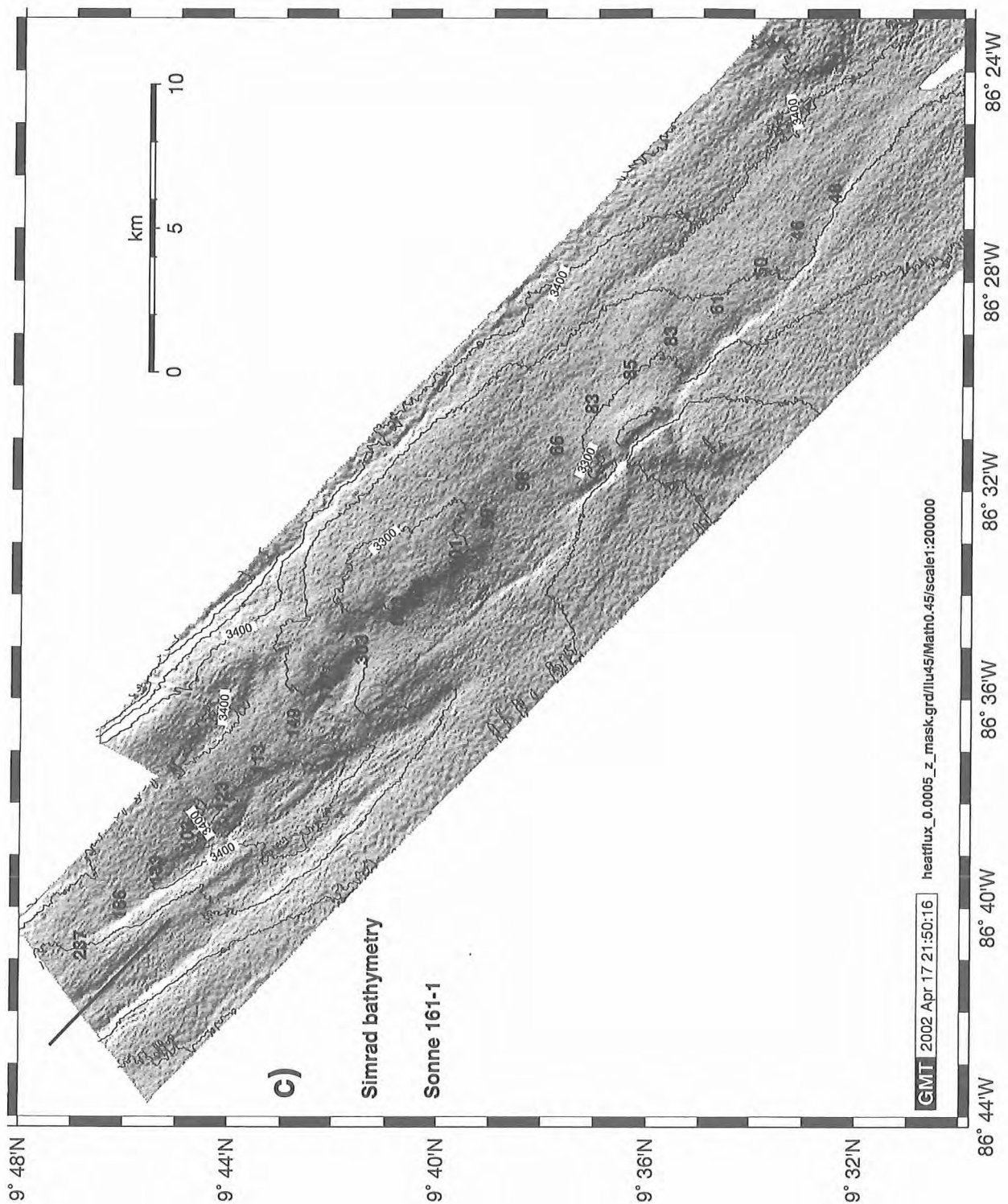


Figure 6.1.1.2. c) Shaded relief map of high resolution Simrad survey on the ocean plate. Grid size is ~ 50 meter. Illumination is from 45° and contours are every 50 meters. Numbers along the bathymetric swath are heat flux values from Fisher et al. (2001) collected with R/V M. Ewing cruise 0104. The black line is the parasound track of Figure 6.1.1.2d. Location on figure 6.1.1.1.

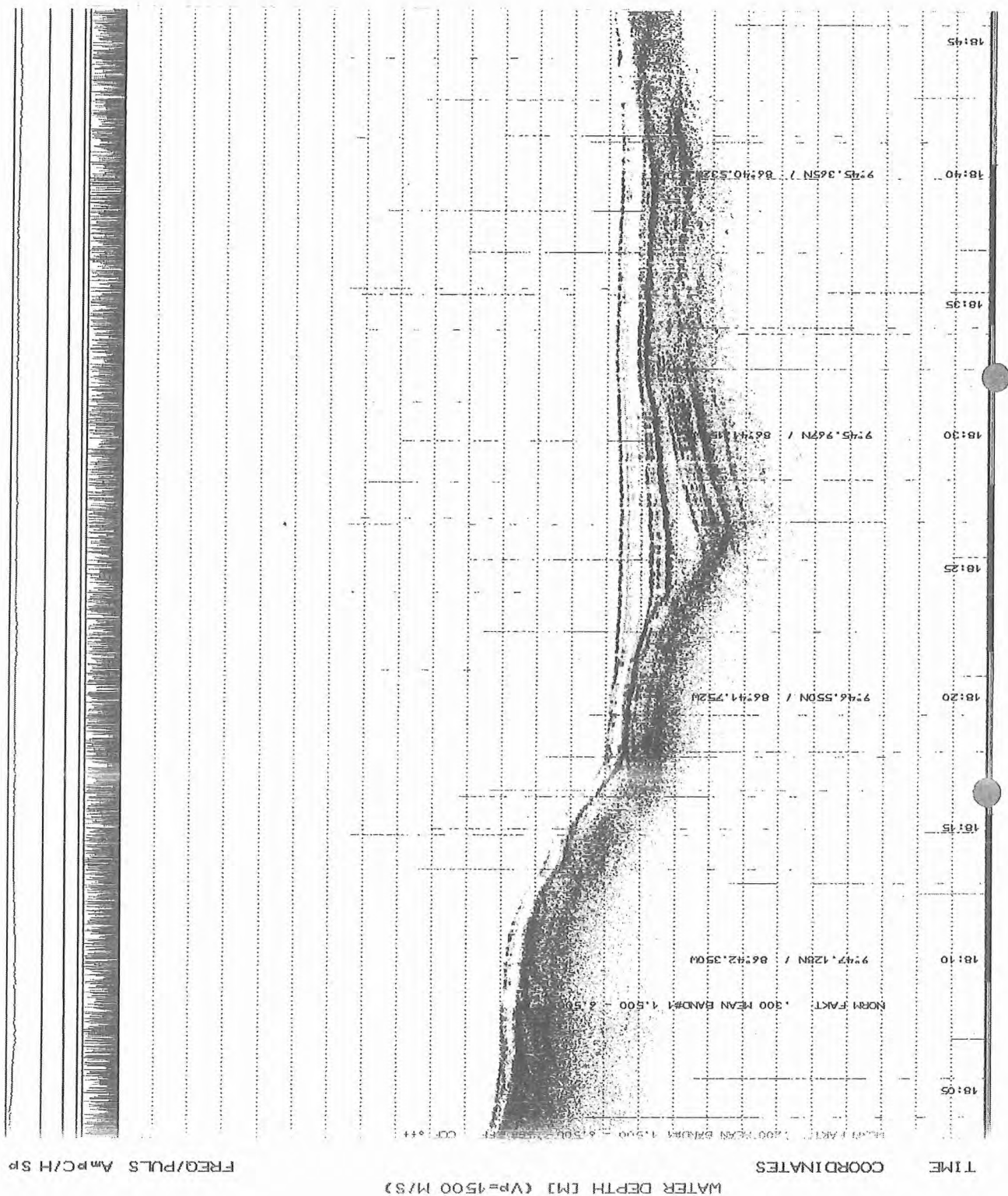


Figure 6.1.1.2. d) Parasound profile in the ocean plate showing tilting of deepest strata in the center of the figure and filling of basin with probably recent turbidites. Tilting and wedge shape units may indicate syntectonic deposit of some of the turbidites, although the geometry of the profile, parallel to the main faults, is inadequate for imaging of the true geometrical relationship between faults and sediment units. The location of the profile is shown in figure 6.1.1.2c.

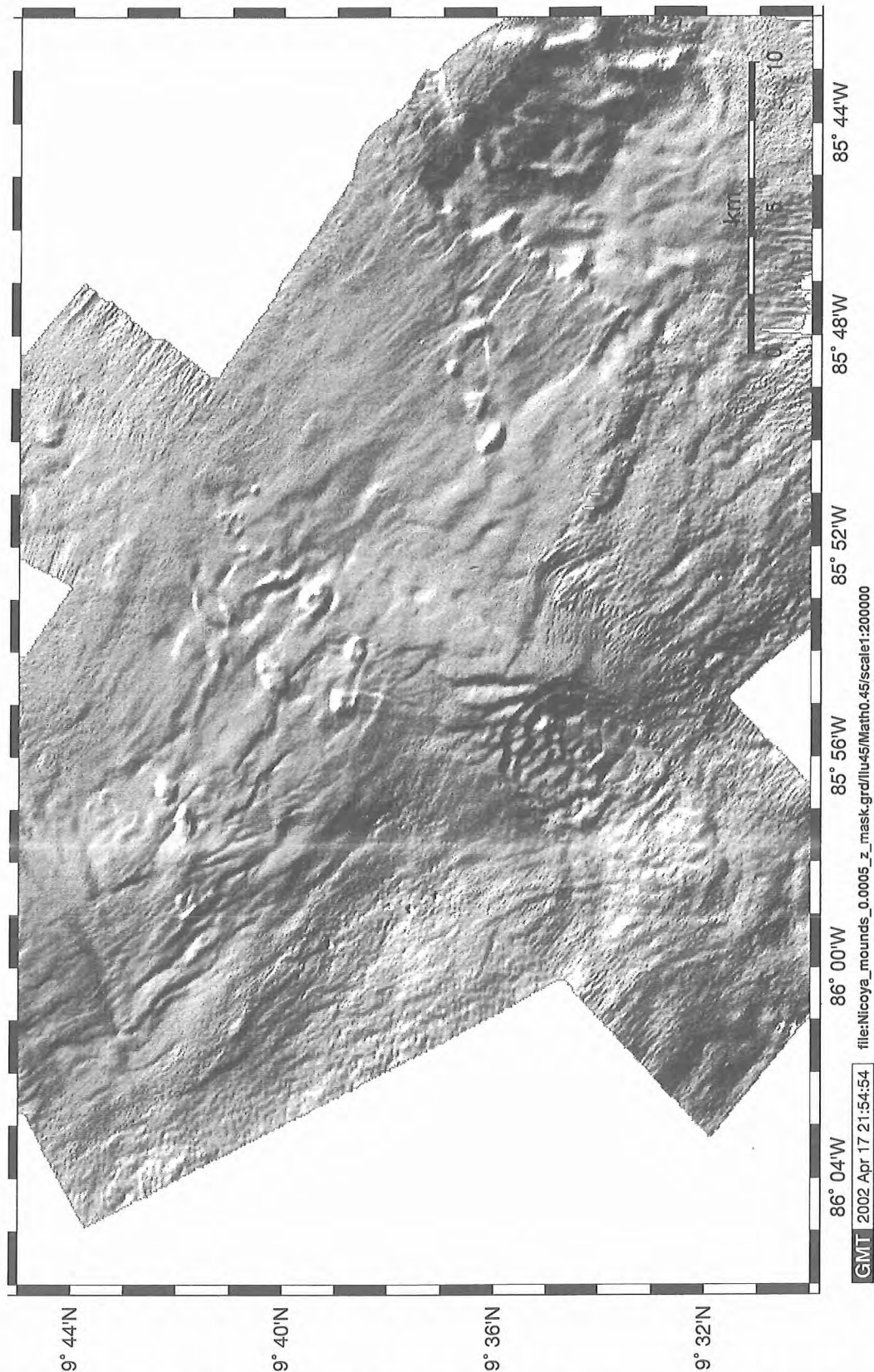


Figure 6.1.1.3. a) Shaded relief map of Simrad survey in the middle continental slope offshore Nicoya (see location in figure 6.1.1.1). The data are gridded at ~ 50 meters. The grid includes data collected during a detailed DTS survey in the center of the figure, with track spacing of ~ 1km, where the resolution is highest. Data to the S, SW and W on the figure and two tracks leaving the area to the NE were collected during transit lines and the faster ship speed of ~ 8-10 knots resulted in higher levels of noise. A large slide with a rough seafloor morphology produced by failed slope sediment is visible at ~ 85° 56' / 9° 34'. A swath of mounds runs along the middle slope. Linear, slope-parallel features are surface expression of normal faults.

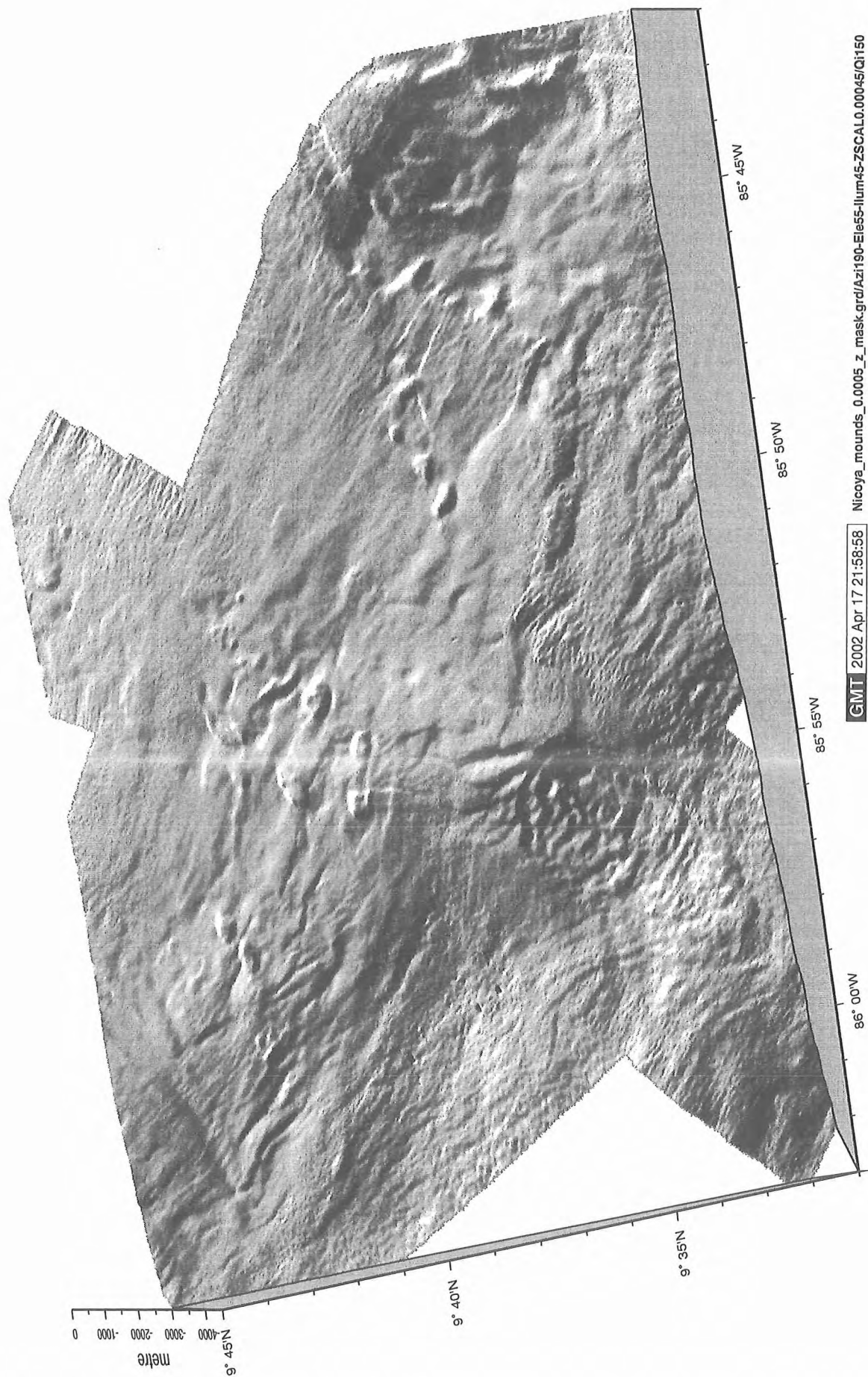


Figure 6.1.1.3. b) Shaded relief perspective view of Simrad survey in the middle continental slope offshore Nicoya (location in figure 6.1.1.1). See figure 6.1.1.3a for explanation and text for discussion.

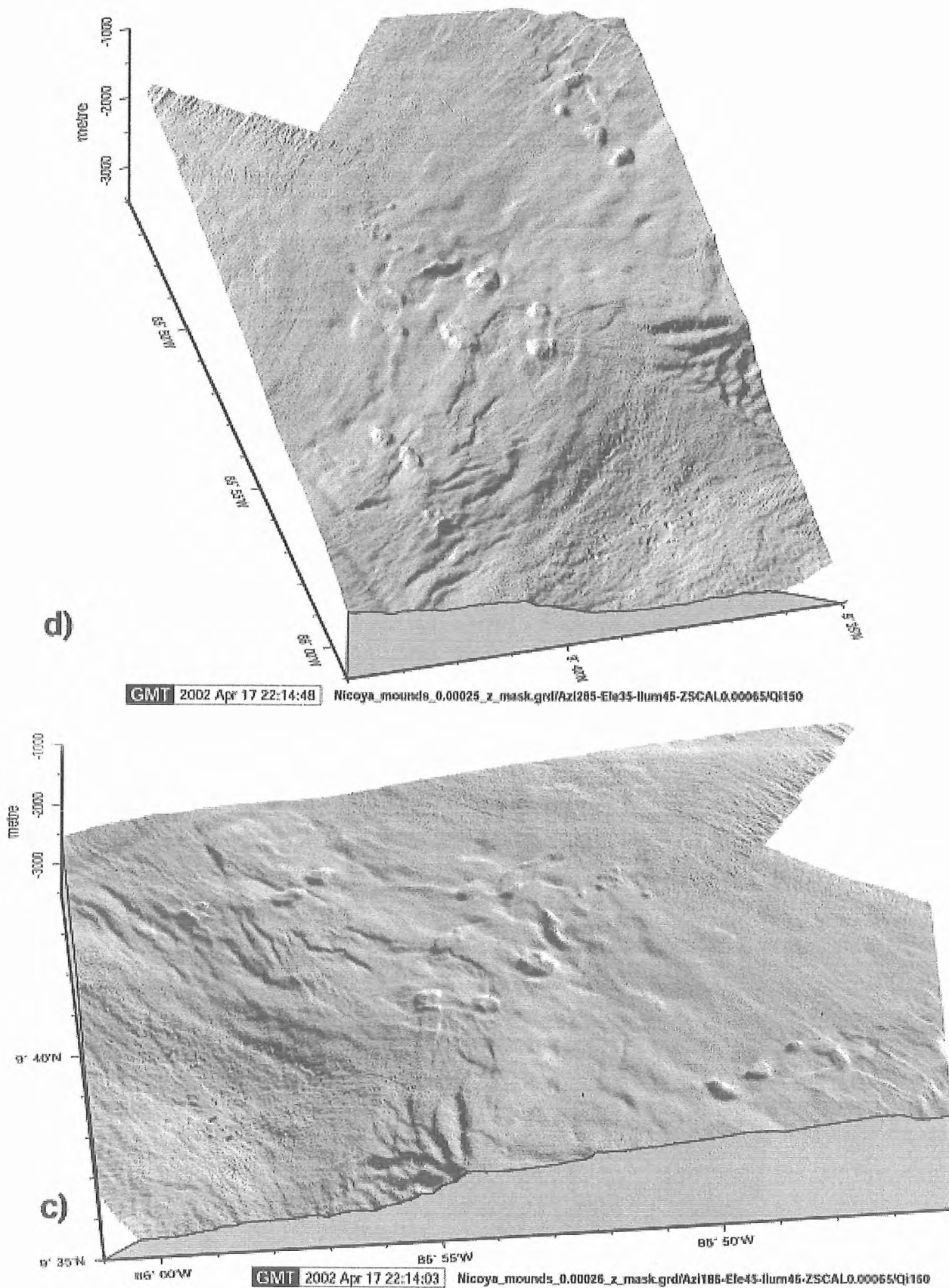


Figure 6.1.1.3. c) and d) Shaded relief perspective views of Simrad survey in the middle continental slope offshore Nicoya (location in figure 6.1.1.1). The grids were created using mainly the data of a detail DTS survey (see caption of Figure 6.1.1.3a) allowing gridding at ~ 25 meter. The two perspectives show the relationship between mounds and normal faulting. See text for discussion.

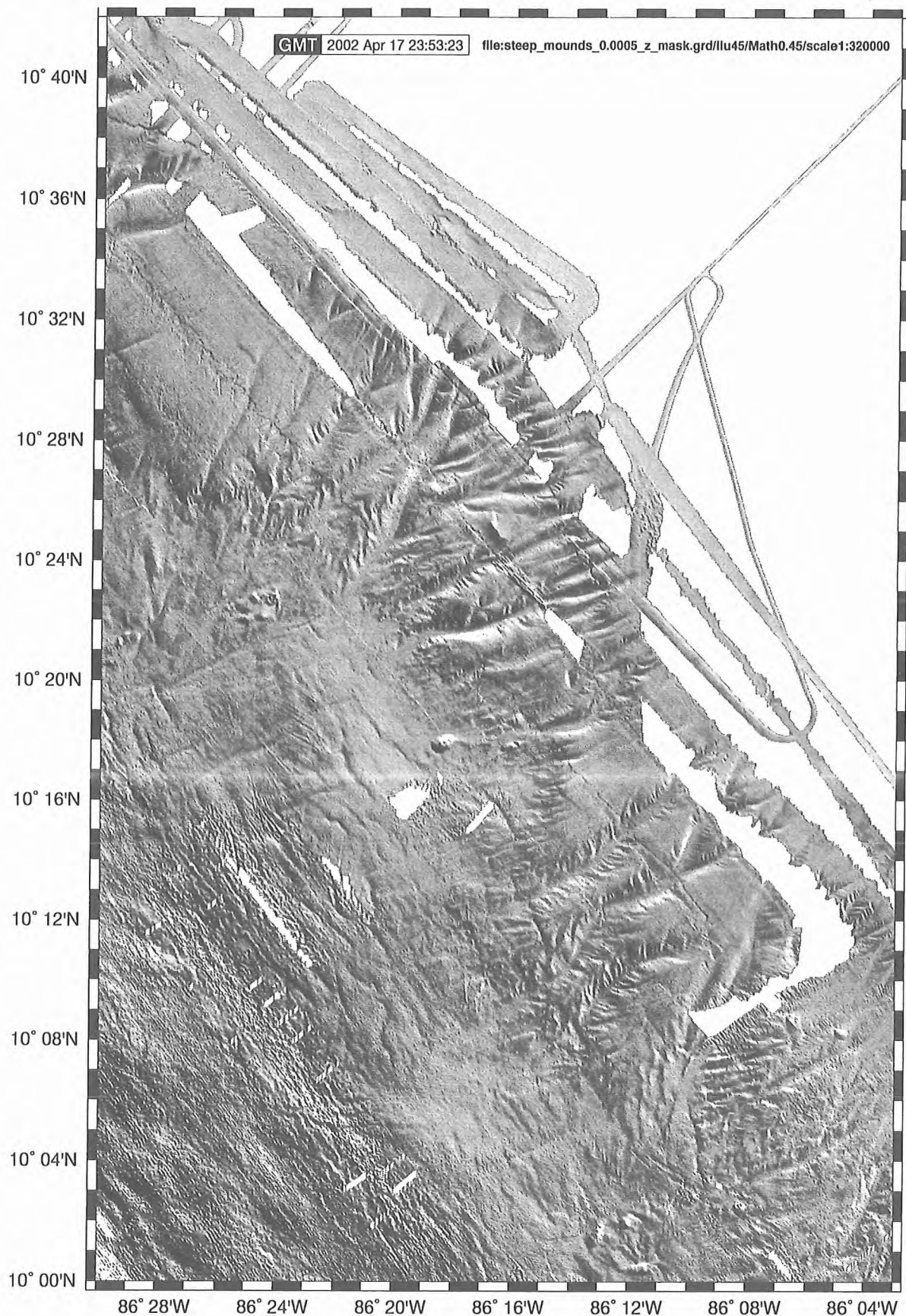


Figure 6.1.1.4. Shaded relief map of the compilation of Simrad data from So163 leg 1 and hydrosweep data from previous cruises. The area with no Simrad coverage is clearly visible in the SW corner of the figure where noise and data gaps largely obscure the morphology. The image shows the relationships among canyons in the upper slope, faulting and mounds in the middle slope. See text for discussion. Location in figure 6.1.1.1.

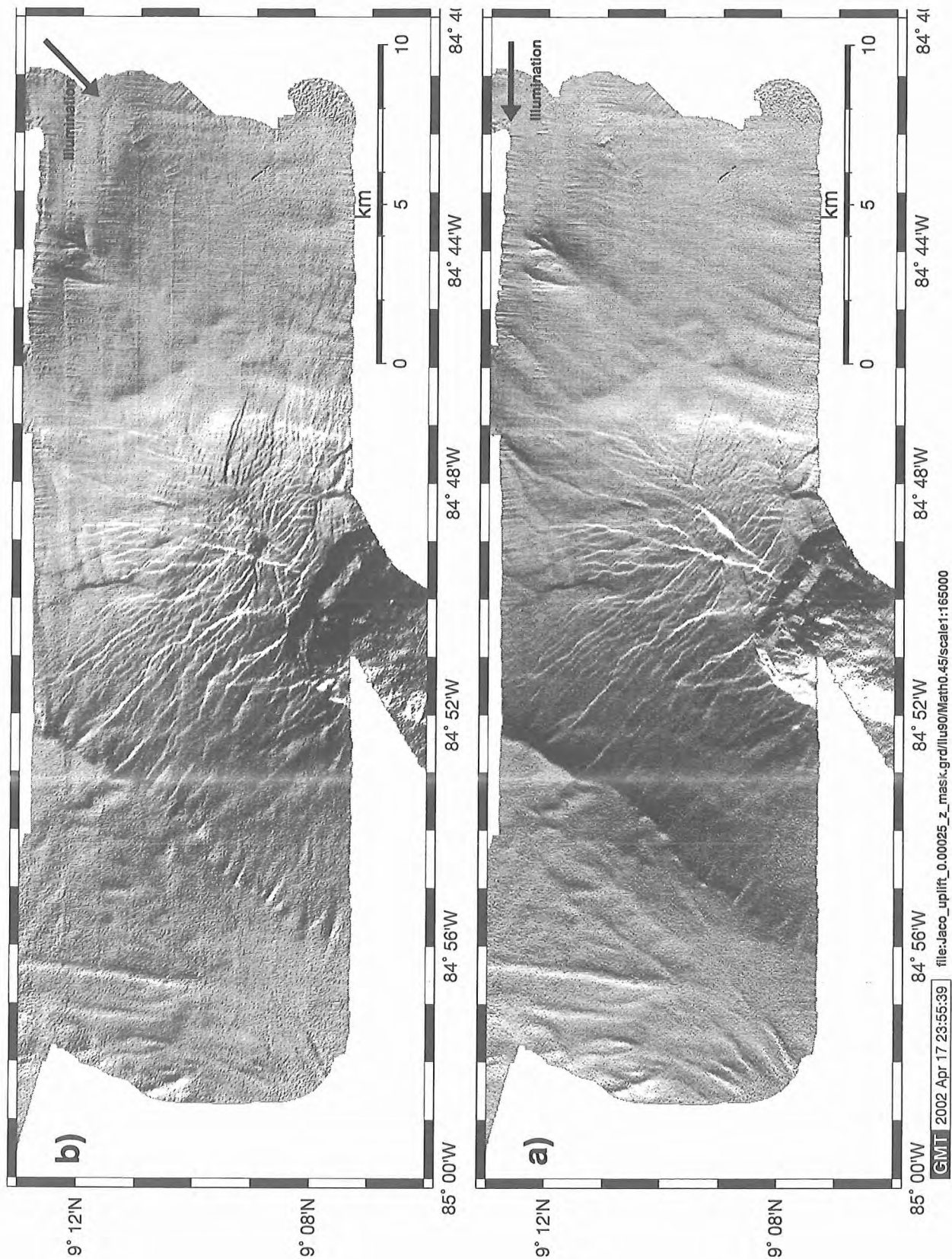


Figure 6.1.1.5. Shaded relief map of Simrad survey across the Jaco uplift and scarp. The uplift is produced by a subducting seamount in the ocean plate. The data were collected during a DTS survey allowing gridding at ~ 25 m. **a)** Data illuminated from 90° highlighting radial faults and headwall of the slide. **b)** Data illuminated from 45° highlighting faults roughly perpendicular to the convergence vector. Location in figure 6.1.1.1.

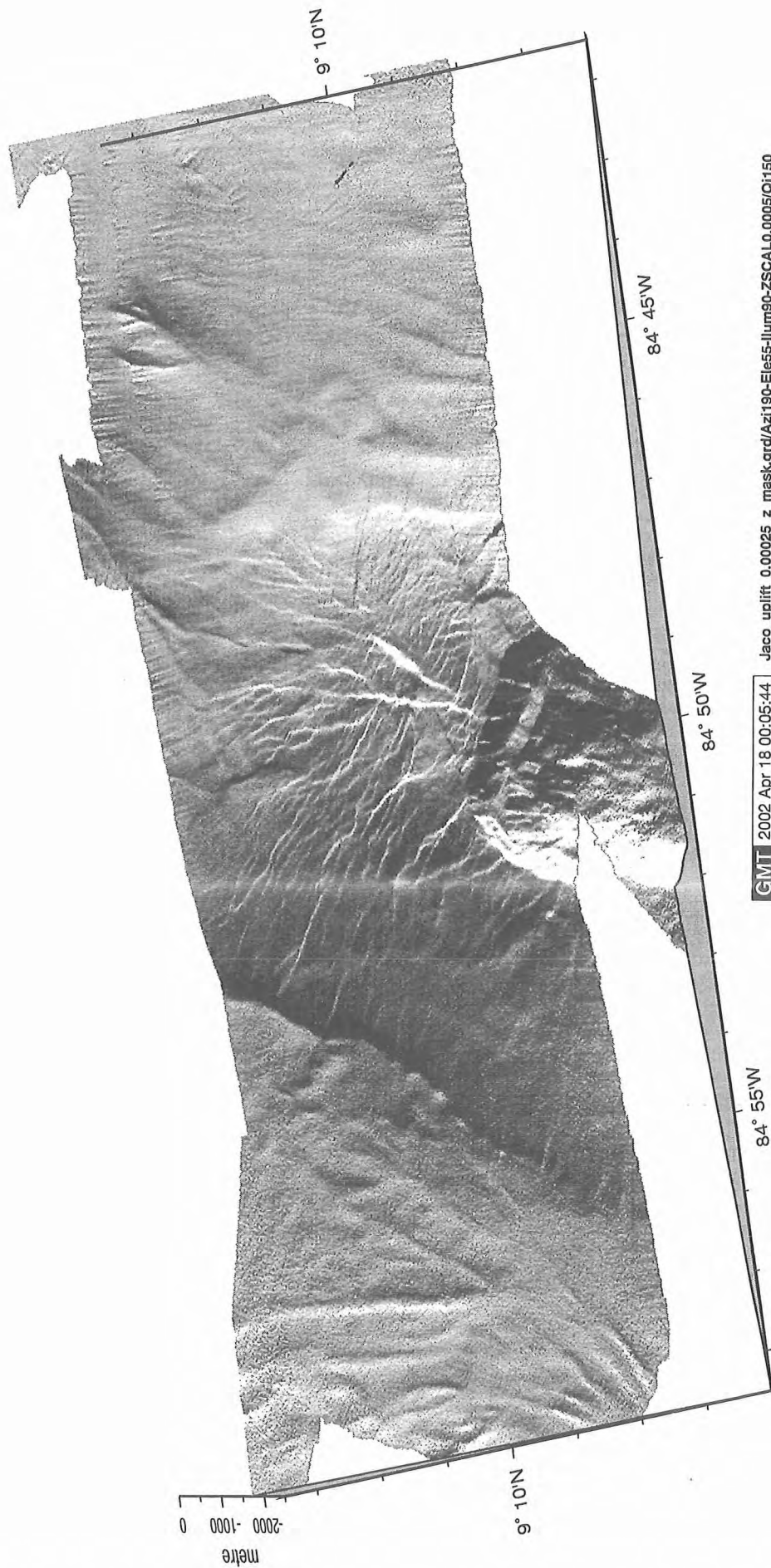


Figure 6.1.1.5. c) Shaded relief perspective view of Simrad survey across the Jaco uplift and upper scarp. Note that faults extend from the center of the uplift area laterally to the edge of the uplift region. Note terrace in headwall area from last slid block. Location in figure 6.1.1.1.

6.2. TOBI operation and initial data interpretation

(D.G. Masson, I.P. Rouse, V. Hühnerbach and D. Matthew)

6.2.1. Summary of TOBI operations

Operational Procedure

Mobilisation

The SOC TOBI system was transported to the Sonne in two 20' containers. One hardtop container was for electronics, computers, tools and spares, and one soft-top for the vehicle, winches, bedplate and depressor weight.

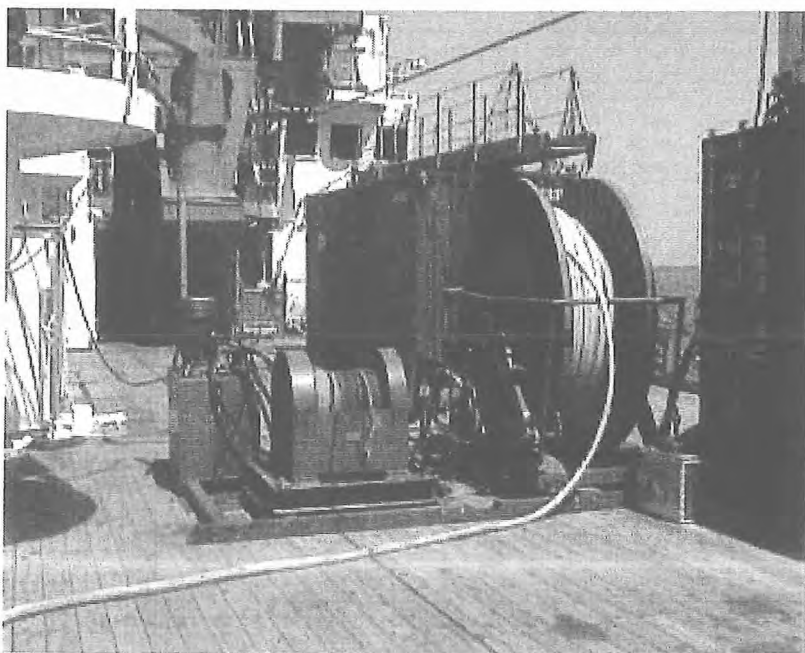


Figure 6.2.1 TOBI winch system. From left to right, power pack, launch winch, umbilical winch.



Figure 6.2.2 TOBI laboratory set-up.

As TOBI had previously been installed on the Sonne for SO-144, the mobilisation was quickly undertaken. The bedplate was assembled on the deck and bolted into position as before. The launch and umbilical winches and the power pack were bolted to the bedplate (Fig. 6.2.1). The power pack was connected to a 3-phase 15kW supply in the gun control shack aft of the geology laboratory.

The deck electronics systems were set up in the geology laboratory on the port side of the main deck. 8' x 4' x 18 mm plywood sheets were used to extend the bench space available on the middle benches in the laboratory. The electronics racks plus the TOBI replay system were mounted on these. It was found that the provision of electrical supply sockets above the benches provided a very neat installation (Fig. 6.2.2).

The GPS receiving aerial was mounted on a pole on the port side of the main deck. Experience on SO-144 showed that this position gave good reception and was conveniently close to the geology laboratory.

Approximately 1.5 days were spent setting up the equipment prior to the first test of the TOBI system on deck.

Pre-launch

Prior to the first launch of TOBI a number of essential tasks were carried out. The launch/recovery wire was wound onto the launch winch. A special link that had been manufactured in Southampton to connect the depressor weight to the main tow cable was offered up for fitting. It was found that it was too wide for the shackles that it was to fit and had to be machined down. The vehicle was rigged with four steadying lines plus the main recovery line for a fore-aft launch (Fig. 6.2.3). The launch wire was run through the main block and connected to the lifting strops on TOBI with the launch pin. The umbilical was connected to the vehicle and the emergency release mechanism cocked.

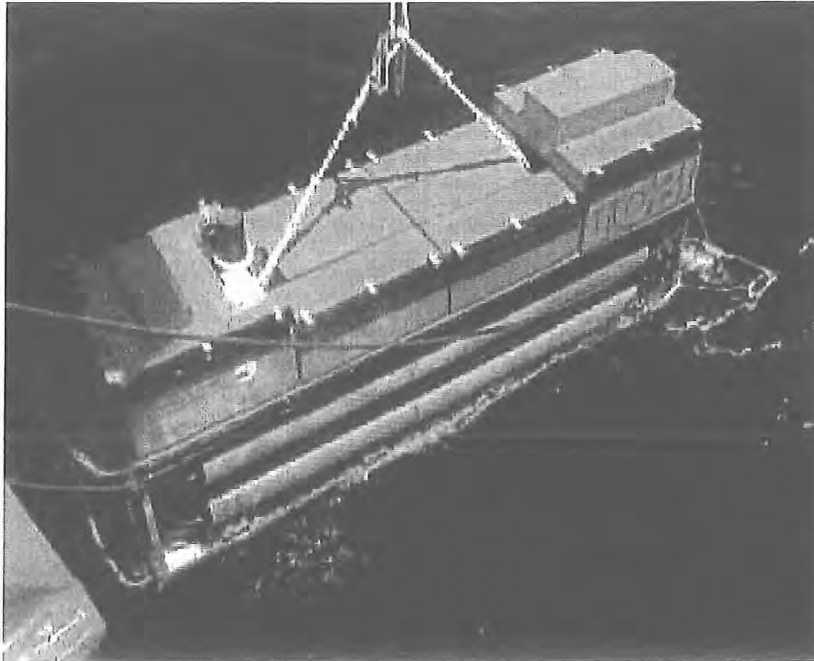


Figure 6.2.3 TOBI on recovery – note lifting strops and steadying lines.

A deck lead was run out to the vehicle for test purposes. The vehicle was powered up and all systems apart from the swath receivers worked. The problems with the swath system were worked on at length and although progress was made, it was clear that with the necessary hardware changes that had to be made to the system, that it was unlikely for it to be operation during the cruise.

A final check of the vehicle running the power through the umbilical revealed a short circuit. The umbilical was swapped for the spare prior to the first launch. The spare umbilical was found to have a crushed connector that had to be replaced.

Launch procedure

The TOBI vehicle is launched in a fore-aft position rather than the more preferred athwart ships position on the Sonne due to the narrow 'A' frame. The vehicle is lifted off the deck with the launch winch while being steadied by two lines to the fore of the vehicle. The 'A' frame is then extended over the stern of the ship and the launch winch paid out until the vehicle is in the water. The steadying lines are then let go, pulled back aboard and the launch pin pulled to release the vehicle. As the umbilical is being paid out the depressor weight is brought round under the 'A' frame and the main tow cable fed through the main block. The cable is mechanically connected to the depressor. The electrical termination bottles are attached to the main cable using clamps and cable ties. At the end of the umbilical a loop enables it to be tied off so that the free end can be mechanically connected to the depressor weight and electrically to the main cable. The vehicle is then powered up to check correct operation. If all is OK then the depressor is lifted into the water and survey commenced.

Recovery procedure

Recovery of the TOBI system commences with the recovery of the depressor weight. Making sure that the power to the vehicle is switched off, the depressor is swung aboard and landed in its cradle. The loop in the umbilical is grabbed and made fast to a cleat or eye with a rope. The umbilical is then disconnected from the depressor weight and main cable. The free end is attached to the rope pennant on the umbilical winch and is slowly recovered. During this time the depressor is disconnected from the main cable, and the main cable is replaced through the main block with the launch wire. When the vehicle is in boat hook range (approx 10 m) the recovery hoop is grabbed and the steady and recovery lines brought aboard. These are carefully

sorted out with the vehicle just astern of the ship. The recovery line is attached to the launch wire and once the steady lines are in position – two fore inside the ‘A’ frame, two aft outside the ‘A’ frame – the vehicle is brought out of the water (Fig. 6.2.3). Once clear of the stern, the ‘A’ frame is brought in and the vehicle landed on the deck.

Survey guidelines

As a general rule the survey speed for TOBI should not exceed 3 kts. Above this speed the stability of the vehicle is reduced, the wire out will become excessive and the along track resolution becomes large compared to that across track. During the cruise the tow speed was between 2.0 and 2.5 kts dependant on the wire out. It was only necessary to slow below 2.5 kts when surveying in water depths of 2,500 m or more.

TOBI is kept to between 350 and 400 m above the sea floor except on turns. Due to the amount of wire deployed it is essential that the ship does not make any sharp turns or speed fluctuations. A sharp turn or loss of speed will send TOBI plummeting whereas a rise in speed will take the vehicle above the most favourable altitude. A turn rate of 2° per minute has been found to be optimal for survey line spacing of 5 km and a ship speed of 2.5 kts. A maximum veer rate of 0.5 m s^{-1} and a maximum haul rate of 0.7 m s^{-1} are recommended so that no damage is done to the vehicle or cable due to too high a tension or too much slack wire causing instability.

TOBI Deployments

Run 1

The depressor was put into the water around 1810/076 GMT 12:10 local. Initially all look fine. Some speckle noise was observed on the starboard sidescan. Both sidescan channels seemed to be low on gain. The profiler was giving weak returns and the bottom-tracking algorithm couldn't lock onto the signal so hand altitude values had to be entered most of the time. The starboard sidescan suddenly increased in gain at around 2355/076 GMT and remained until 0335/077 GMT when it returned to its previous value. The digital data telemetry started giving corrupt readings at around 2025/076 GMT although the trigger and frame sync pulses were still being correctly detected. Substituting a spare card interface card in the display computer eventually cured the problem at around 2130/076 GMT. However the magnetometer didn't give reliable readings for the remainder of the run. At around 0530/077 GMT the system suffered a cable failure. The depressor weight was recovered at first light and the problem traced to the cable between the main tow cable termination and the TOBI connector bottle. This cable had rubbed through against the main cable. A test of the vehicle showed that there were no sonar transmissions so that a recovery had to be effected.

The lack of sonar transmissions was traced to a blown MOSFET in the sidescan power amplifier caused by the power cycling originated by the bad cable. This and a resistor were replaced. Also replaced was the magnetometer cable. A test on the sonar arrays showed that one of the port transducers was low in impedance ($\sim 20 \text{ Mohms}$ under test). A spare transducer replaced this. The vehicle was re-rigged ready for deployment.

Run 2

The depressor was put into the water around 2300/077 GMT. The starboard sidescan channel exhibited a low gain compared to the port channel. The magnetometer worked until a depth of 300 m was reached. All digital telemetry then gave corrupt data but again the trigger and frame sync pulses were unaffected. Eventually the digital data returned at 0230/078 at a depth of 430 m and subsequently the magnetometer started to work at a depth of 640 m at 0650/078. The magnetometer stayed correct until a depth of 420 m at 1550/078 when it stopped for the remainder of the run. The profiler was again giving weak returns and also lost its trigger intermittently. This improved with increasing depth and was solid for the remainder of the run. At 0330/078 the logging system was stopped to reset the day number. This had been incorrectly set before the first run. The run ended with the vehicle on deck at around 1500/081 GMT. The recovery was not good and the vehicle ran underneath the ship. This caused damage to the magnetometer mounting and to the rubber bumpers on the syntactic foam buoyancy blocks.

On deck the sidescan electronics was investigated and an intermittent contact was found on the starboard match pad. The whole of the starboard match pad was secured with heavy gauge wire to ensure connectivity. Also it was found that the port and starboard pre-amplifiers had been switched around. This was corrected and both match pad arrays checked for tuning. 21 dB of gain was added to the profiler receiver before the correlator. Repairs were done to the magnetometer mounts and the bumper fixings. The magnetometer was removed from its housing and two packages of silica gel added to try and combat any residual dampness. The vehicle was re-rigged ready for deployment.

Run 3

The depressor was put into the water at around 2015/082 GMT. A short on the cable became apparent almost immediately. The depressor was recovered and the fault traced to a connector. This was replaced with the vehicle still deployed and the depressor was re-launched. The sidescan and profiler sonar systems both gave excellent signals. After a brief spell during the descent the magnetometer worked throughout the run. Bursts of noise were observed on the sidescan but did not prove obtrusive. The vehicle was recovered at around 1630/084 GMT.

On deck all electronics cards were treated with conformal coating to improve their resistance to dampness. An extra 10dB of gain was added to the profiler. The swath pre-amp tube was isolated using blanking plugs to stop any possibility of it causing any noise. The vehicle was re-rigged ready for launch.

Run 4

The depressor was put into the water around 0130/096 GMT. Again a short period of corrupt digital data was observed during the descent. Also the magnetometer gave bad readings for a short period. These problems resolved themselves quickly at a depth of 630 m. Apart from some periods of low gain on both the sidescan and profiler all instruments worked fine throughout the run. Even on the lower gain the profiler gave a large enough signal for the bottom tracking to work correctly. The vehicle was recovered at around 1430/100 GMT.

On deck the cable connecting the hydro tube to the sidescan tube was replaced with the cable previously used on the swath pre-amps. The vehicle was re-rigged ready for launch.

Run 5

The depressor was put into the water at around 0048/102 GMT. All looked fine at the surface but at 0105/102 GMT the sonar systems stopped transmitting. The vehicle was recovered at around 0225/102 GMT. The problem was traced to the cable that had just been replaced. A spare was quickly installed and the vehicle re-rigged for deployment.

Run 6

The depressor was put into the water at around 0338/102 GMT. Apart from the magnetometers now obligatory funny five minutes on the descent all instruments worked well. Only in shallow water - less than 300 m on the rise and 400 m on the descent - did the magnetometer give corrupt readings. This points to a plug/cable problem. The light backscattering sensor failed at 1520/102 GMT. This was an instrument failure rather than a cable problem. The vehicle was recovered at around 2130/103 GMT.

Run Summary

The above detailed descriptions concentrate on the problems that were encountered during the TOBI surveys. This does not give a balanced picture as the system performed exceedingly well for the majority of the time deployed. The mosaics produced bear testament to this. The vehicle was deployed for a total of 12 days 14 hours and 32 minutes during the cruise. The only significant downtimes came as a result of the cable shorts during runs 1, 3 and 5. This totalled just over 8 hours or approximately 2.5% of the total time.

<u>Deployment</u>	<u>Start time/day</u>	<u>End time/day</u>	<u>Comments</u>
1	1810/076	1200/077	Short on cable.
2	2300/077	1500/081	Second run.
3	2015/082	1630/084	Third run.
4	0130/096	1430/100	Fourth run.
5	0048/102	0225/102	Aborted due to bad cable.
6	0338/102	2130/103	Sixth run.

TOBI Watchkeeping

TOBI watchkeeping was split into three, four-hour watches repeating every 12 hours.

Watchkeepers kept the TOBI vehicle flying at a height of ideally 350 to 400 m above the seabed by varying wire out and/or ship speed. Ship speed was usually kept at 2.5 kts over the ground with fine adjustments carried out by using the winch. In deeper water – greater than 2500 m – the speed was reduced to 2.0 kts to reduce the scope of wire deployed. As well as flying the vehicle and monitoring the instruments watchkeepers also kept track of disk changes and course alterations.

Data Management

Data from the TOBI vehicle is recorded onto 1.2 Gbyte magneto-optical (M-O) disks. One side of each disk gives approximately 16 hours 9 minutes of recording time. All data from the vehicle is recorded along with the ship position taken from the GPS receiver. Data was recorded using TOBI programme LOG. Data recorded on the M-O disks were copied onto CD-ROMs for archive and for importation into the on board image processing system.

As well as recording sidescan and digital telemetry data LOG displays real-time slant range corrected sidescan and logging system data, and outputs the sidescan to a Raytheon TDU850 thermal recorder. PROFDISP displays the chirp profiler signals and outputs them to a Raytheon TDU850. DIGIO9 displays the real-time telemetry from the vehicle – magnetometer, CTD, pitch and roll – plus derived data such as sound speed, heading, depth, vertical rate and salinity. LOG, PROFDISP and DIGIO9 are all run on separate computers, each having its own dedicated interface systems.

A program called HITSCOPY was used to strip off magnetic and attitude data from the raw TOBI data files and store it in ASCII format for direct importation into a spreadsheet.

The details of the times of the logged part of each TOBI run are given in Table 6.2.1

Instrument Performance

Most of the problems with individual instruments have been covered above in the details of each run. The relatively minor problems with the sonar systems were mostly easily fixed in the processing of the data. Whilst the individual instruments themselves performed exceedingly well the same cannot be said for the underwater connectors and cables. These were the source of almost all the problems encountered and it must be emphasised that these items must be regularly maintained and replaced if high reliability is to be ensured in the future.

M-O Number	File Name	Time/ Day START	Time/ Day STOP	Comments / Run #
830	TOBIB.DAT	1813/076	0523/077	Run 1 curtailed due to cable short. Day number corrected.
831	TOBI.DAT TOBIB.DAT TOBIC.DAT	2310/077 0334/078 0339/078	0329/078 0337/078 1526/078	Run 2 Day number corrected Day number correct
832	TOBI.DAT	1526/078	0733/079	
833	TOBI.DAT	0733/079	2340/079	
834	TOBI.DAT	2340/079	1548/080	
835	TOBI.DAT	1548/080	0755/081	
836	TOBI.DAT	0755/081	1419/081	End of run 2
837	TOBI.DAT	2120/082	1323/083	Start of run 3
838	TOBI.DAT	1323/083	0535/084	
839	TOBI.DAT	0535/084	1558/084	End of run 3
840	TOBI.DAT	0145/096	1752/096	Start of run 4
841	TOBI.DAT	1752/096	1000/097	
842	TOBI.DAT	1000/097	0207/098	
843	TOBI.DAT	0207/098	1815/098	
844	TOBI.DAT	1815/098	1022/099	
845	TOBI.DAT	1022/099	0229/100	
846	TOBI.DAT	0229/100	1355/100	End of run 4
847	TOBI.DAT TOBIA.DAT TOBIB.DAT	0052/102 0340/102 1016/102	0156/102 1014/102 1846/102	Run 5 Start of run 6 Reboot logging
848	TOBI.DAT	1846/102	1053/103	
849	TOBI.DAT	1053/103	2103/103	End of run 6

Table 6.2.1 Summary of logged TOBI data

Sonar artefacts

During the cruise a number of interesting artefacts appeared within the TOBI sidescan images from sources other than the direct sea floor. These are naturally occurring phenomena and are included here to enable users of TOBI imagery to recognise them.

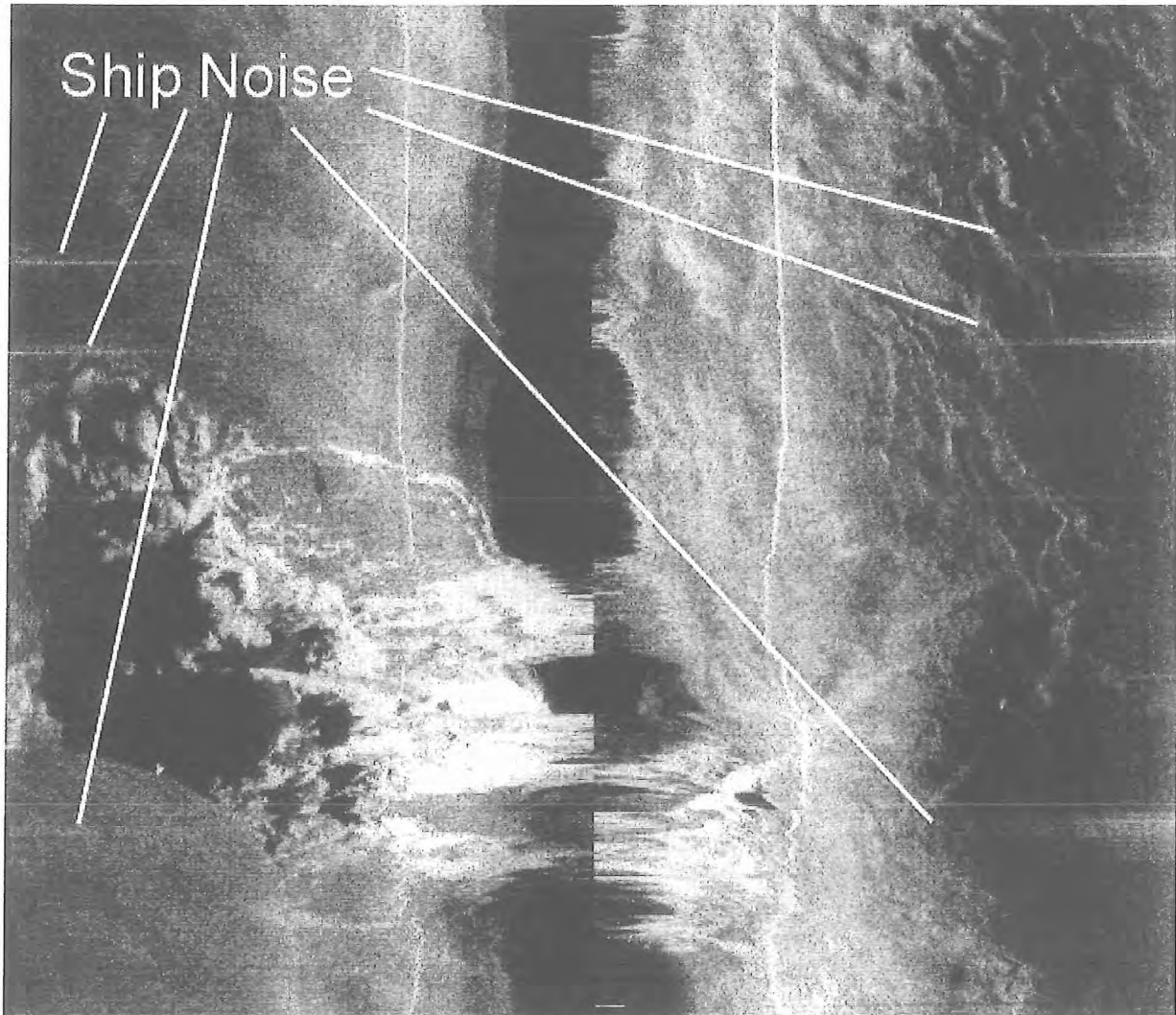


Figure 6.2.4 Passing ship noise on the TOBI sidescan record.

Ship noise (Fig 6.2.4) is usually seen when a passing ship is abeam of the TOBI vehicle position. Bubbles produced by the ship's propeller collapsing cause the noise that is quite intense but usually short-lived, a few minutes at the most. A look out of the porthole or radar is enough to identify the source.

Rain noise (Fig 6.2.5) is similar in appearance to ship noise. However it is not usually as intense and is longer lived, maybe over an hour. The heavy rainstorms experienced on the cruise were easily detected with TOBI as these usually give strong signals.

False sonar (Fig 6.2.6) occurs in shallow water, when TOBI is being flown at depths above the local sound speed minimum. This is especially true if the water is highly stratified. The acoustic impedance mismatches between the various layers or bodies of water are enough to create a sonar reflection. These images can look surprisingly like real geology so great care must be taken when interpreting images from these regions.

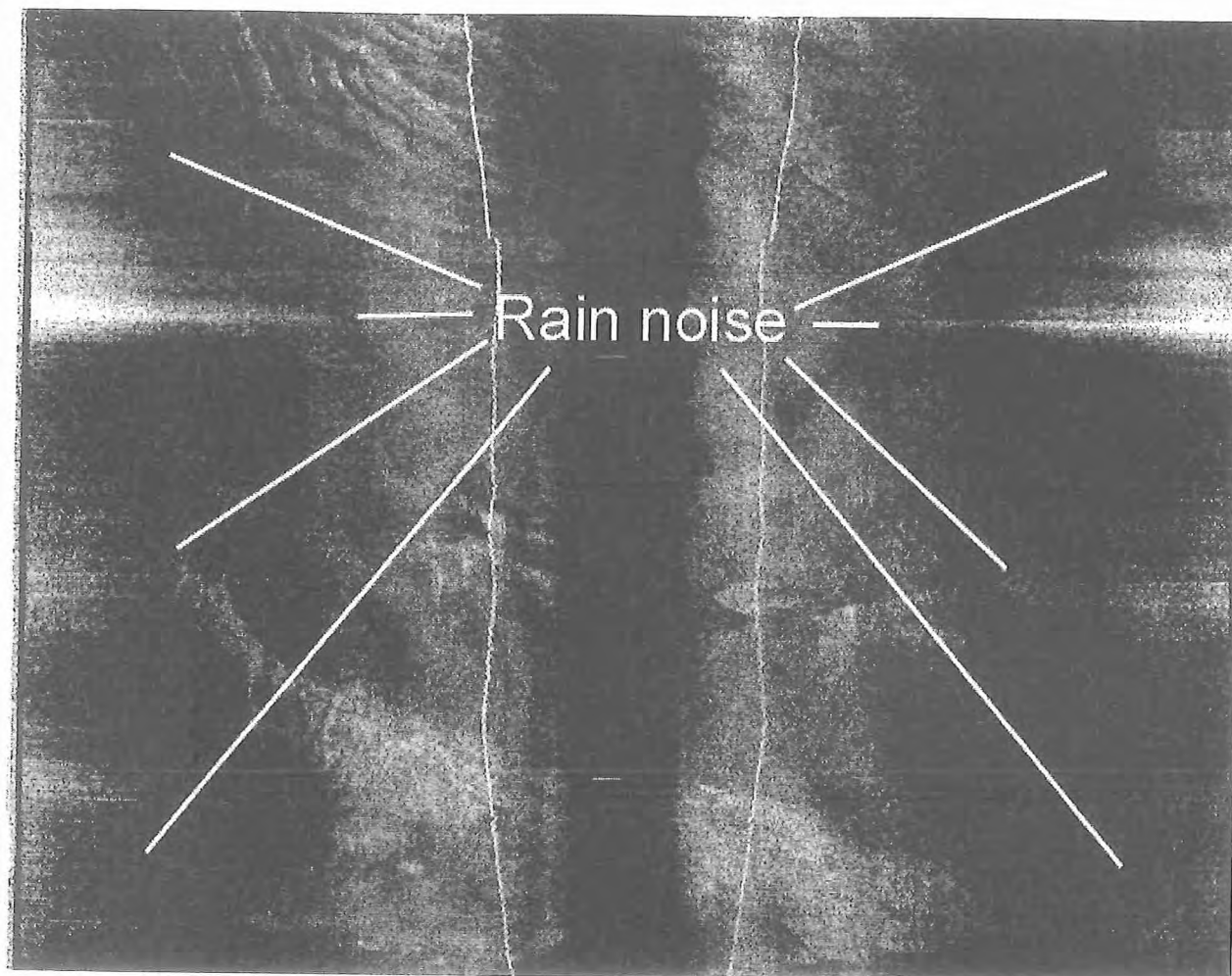


Figure 6.2.5 Rain noise on the TOBI sidescan record.

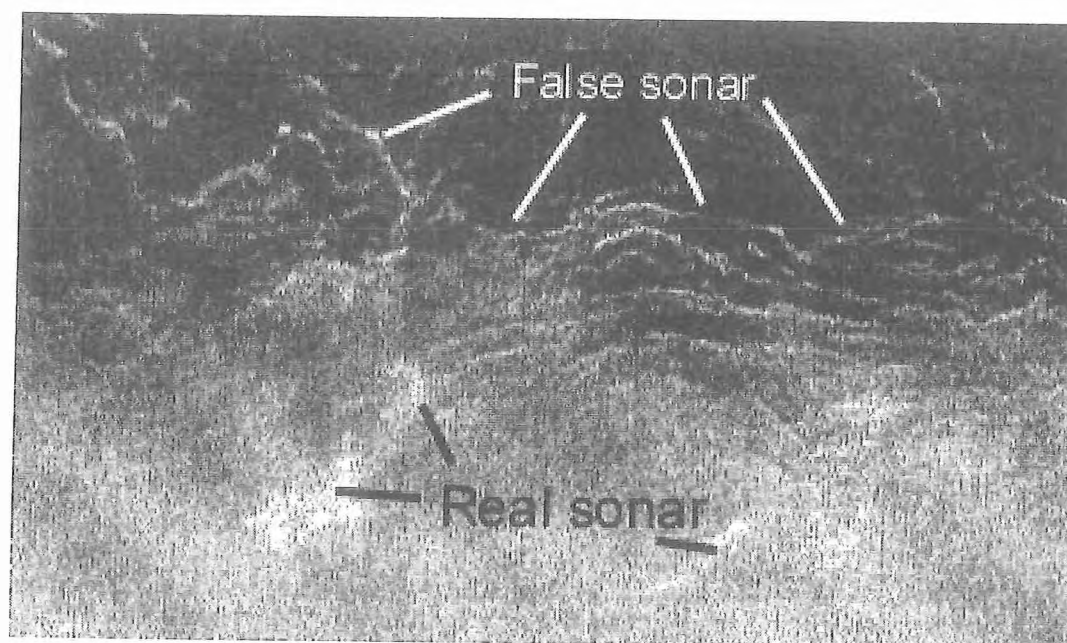


Figure 6.2.6 False sonar echoes on the TOBI sidescan record.

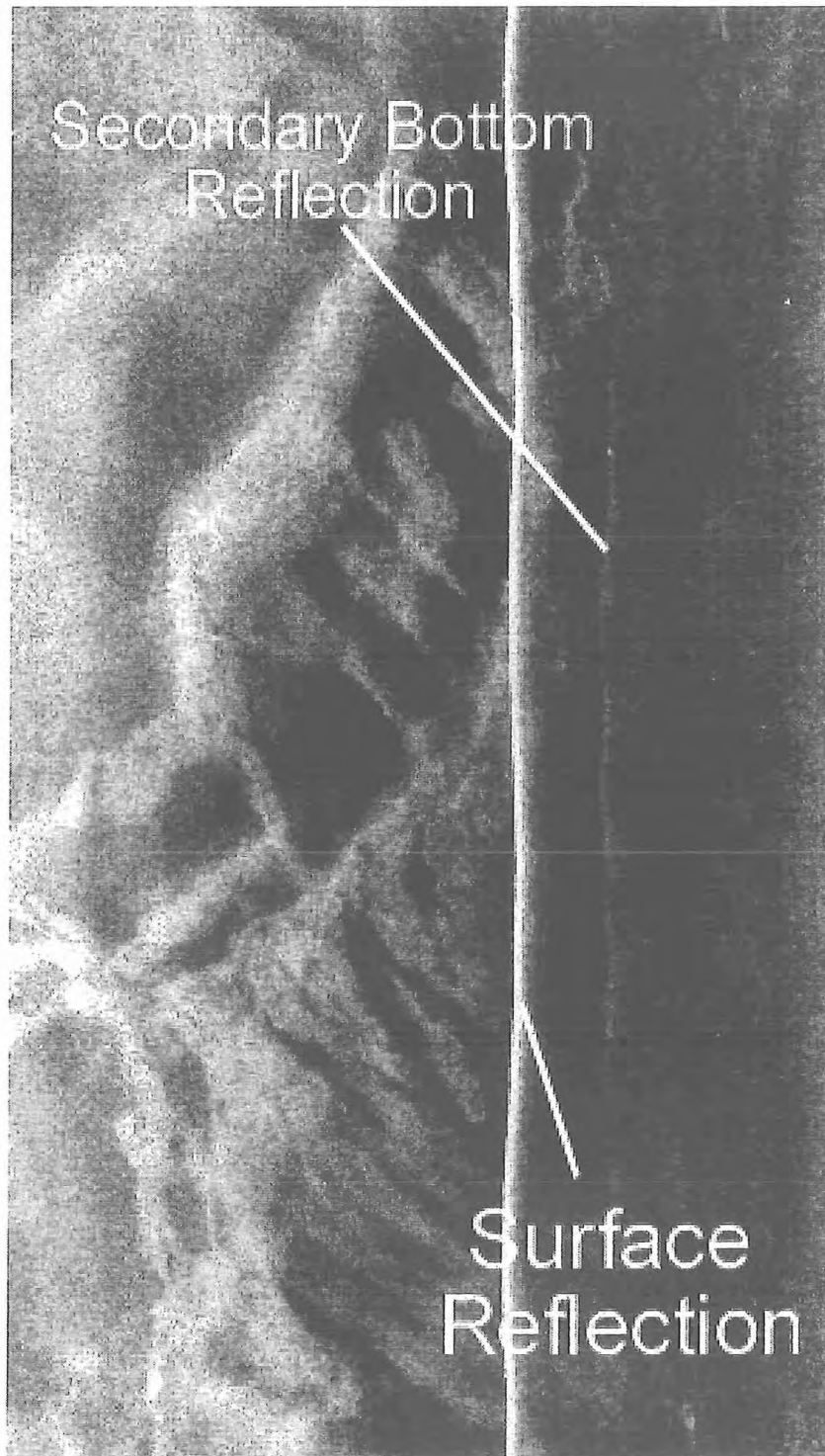


Figure 6.2.7 Secondary bottom reflection on the TOBI sidescan record.

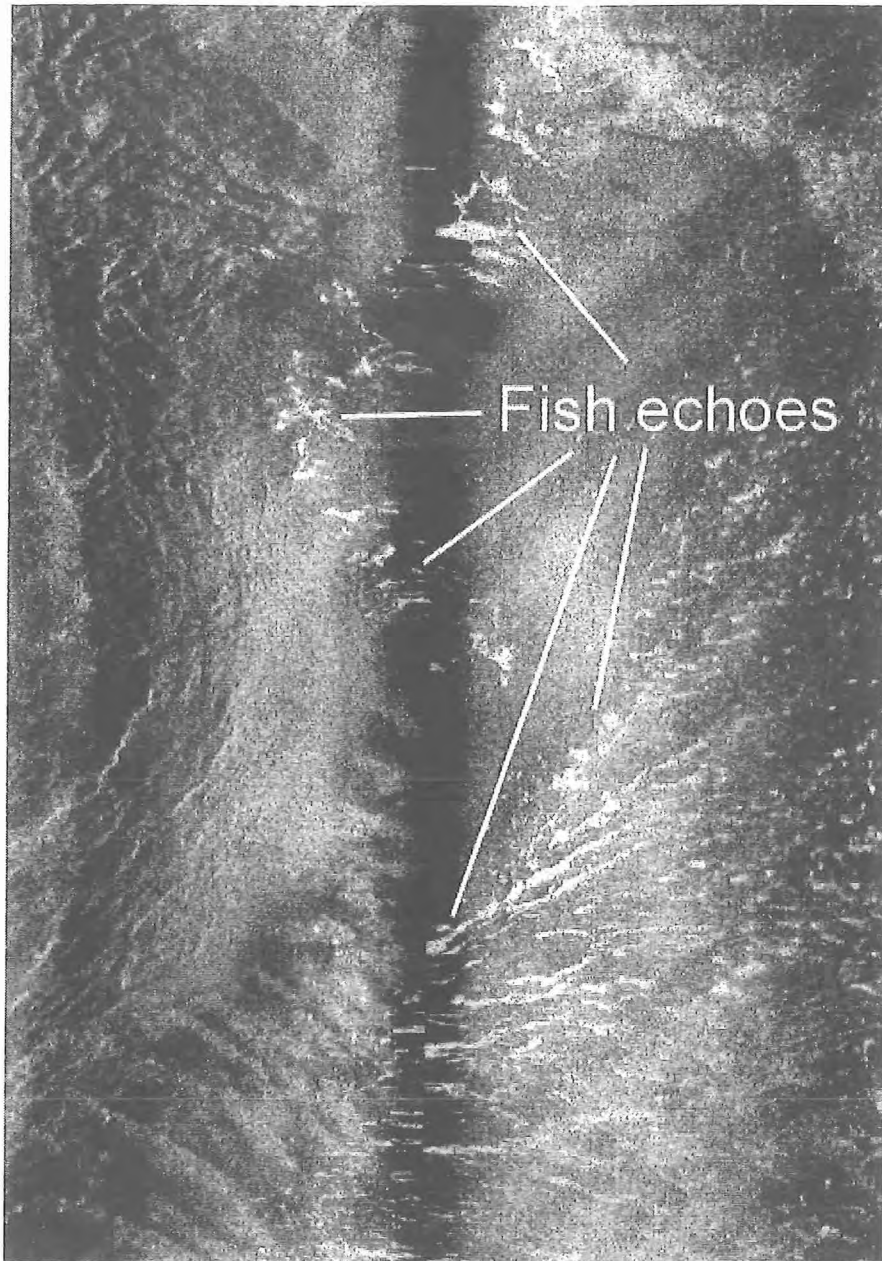


Figure 6.2.8 Echoes from shoals of fish on the TOBI sidescan record.

Secondary bottom (Fig 6.2.7) is where there is a strong reflector beneath the TOBI vehicle. Energy from the sonar is not only reflected back from the sea floor to the vehicle directly but also via the sea surface. This gives a second image of the sea floor at a distance of the water depth away from the original image. The secondary image is very weak compared to the direct image and is only seen over hard, rocky sea floor terrains. The secondary image disappears when the water depth exceeds 3 km due to the 4-second sweep rate of the TOBI vehicle.

Fish (Fig 6.2.8.) can generate sonar targets due to the air in their swim bladders. Shoals of fish can make quite large targets that can be imaged by TOBI. This was especially evident in one area where there seemed to be large shoals of fish over a shallow bottom during run 6.

6.2.2. Preliminary interpretation of TOBI sidescan sonar and profiler data

Two main areas were mapped with TOBI during Sonne 163. The first covered the middle and upper slope between 86° 06' and 86° 45' W and part of the middle slope between 85° 42' and 86° 06' W (Fig. 6.2.9). The main aim of this survey was to detect and characterise any mound-like structures or backscatter anomalies that could be related to fluid venting. The second area extends along the margin from 83° 15' to 84° 40' W. In the east, off the Osa Peninsula where the slope is very narrow, this survey covered the entire slope. Further west, it covered the middle and lower slope, including the deformation front at the base of the slope, with a limited area of coverage of the upper slope in the area of the Quepos landslide. The main aim of this survey was to examine the processes associated with the collision of seamounts and ridges, for example the Quepos plateau and a series of unnamed ridges on the Cocos plate, with the margin. As with the more western area, we also hoped to detect features associated with fluid venting.

The Costa Rican margin between 85° 42' and 86° 45' W (the 'western' TOBI survey area)

The main features seen on the TOBI images from this area include:

- (i) extensive submarine canyons
- (ii) seafloor mounds up to 1 km across and 100 m high
- (iii) complex patterns of faults which offset the present day seafloor
- (iv) erosional scars related to seamount and ridge subduction

Submarine canyons

Submarine canyons are the dominant seafloor feature along the entire upper slope area imaged with TOBI, between 86° 07' and 86° 45' W (Fig. 6.2.10). Typically they extend from the shelf edge to about 2500 m water depth, but die out further downslope. On the sidescan images, individual canyons have a typical 'fish bone' appearance, with the canyon floor imaged as a narrow, high-backscatter stripe and the canyon walls characterised by closely-spaced gullies which trend almost perpendicular to the main canyon axis (Fig. 6.2.10). Overall, the larger canyons tend to coalesce downslope in a dendritic pattern, so that only a few large canyons reach the middle slope at 2000 to 2500 m water depth. TOBI profile data confirm that the canyons usually have a maximum depth of <100 m (Fig. 6.2.11).

Seafloor mounds

A large number of seafloor mounds were identified using a combination of TOBI sidescan sonar images and Simrad EM120 bathymetry maps. Mounds are circular to elongate features a few hundred metres to 1 km in diameter and a few tens of metres high. On the sidescan images, mound relief is often severe enough to produce a clear shadow behind each mound (Fig. 6.2.12). Many mounds are characterised by high to very high backscatter, which may cover all or part of the mound (for further detailed mound images see section 6.4). The mound summits are often marked by small-scale rough topography, which can only just be resolved by the TOBI sidescan. Mounds often occur in groups. A series of OFOS transects across mounds established that, in general, the high backscatter corresponds to areas of carbonate or rocks outcropping at the seabed. Lower backscatter areas on the mounds are usually sediment covered. For further images and discussion of the mounds, see section 6.4.

Faults

Faults appear on the TOBI sidescan sonar images as distinct linear targets or shadows, dependent on whether they downthrow towards or away from the sonar (Fig. 6.2.13). Individual faults can rarely be traced for more than a few km, although zones of faulting can be traced for 10 km or more. Most faults are sub-parallel to the regional contours, except in the vicinity of seamount subduction scarps where more arcuate fault zones can occur (see below). Faults outcropping at the seabed are particularly common on the middle part of the slope in 2000 to 2500 m water depth. On the upper to mid-slope, faults predominantly downthrow in an upslope direction (Fig. 6.2.13). On the steeper lower slope, however, large faults which downthrow downslope become more common.

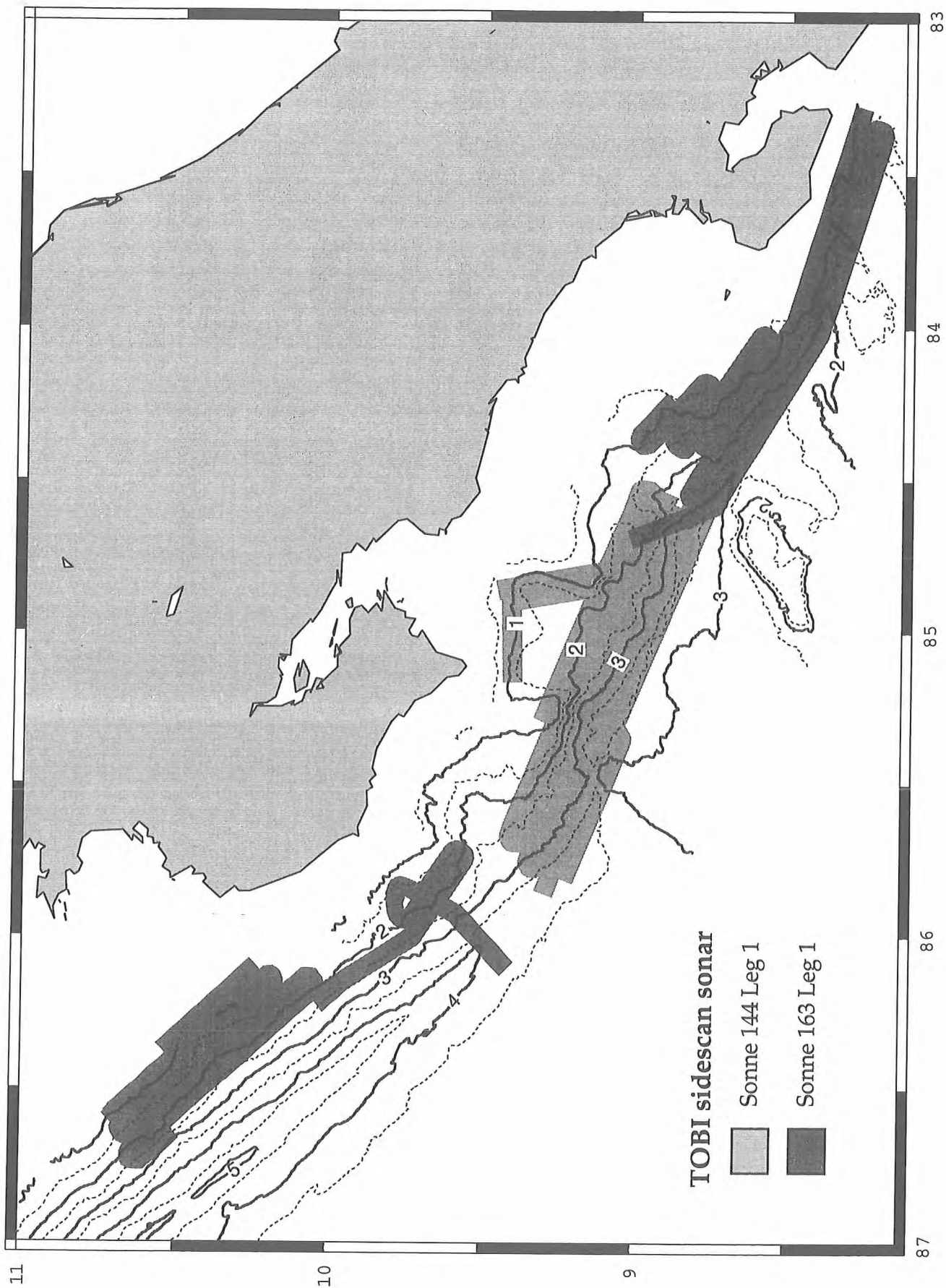


Figure 6.2.9 Map showing TOBI sidescan coverage obtained during S0 144 and 163. Contour interval 500 m.

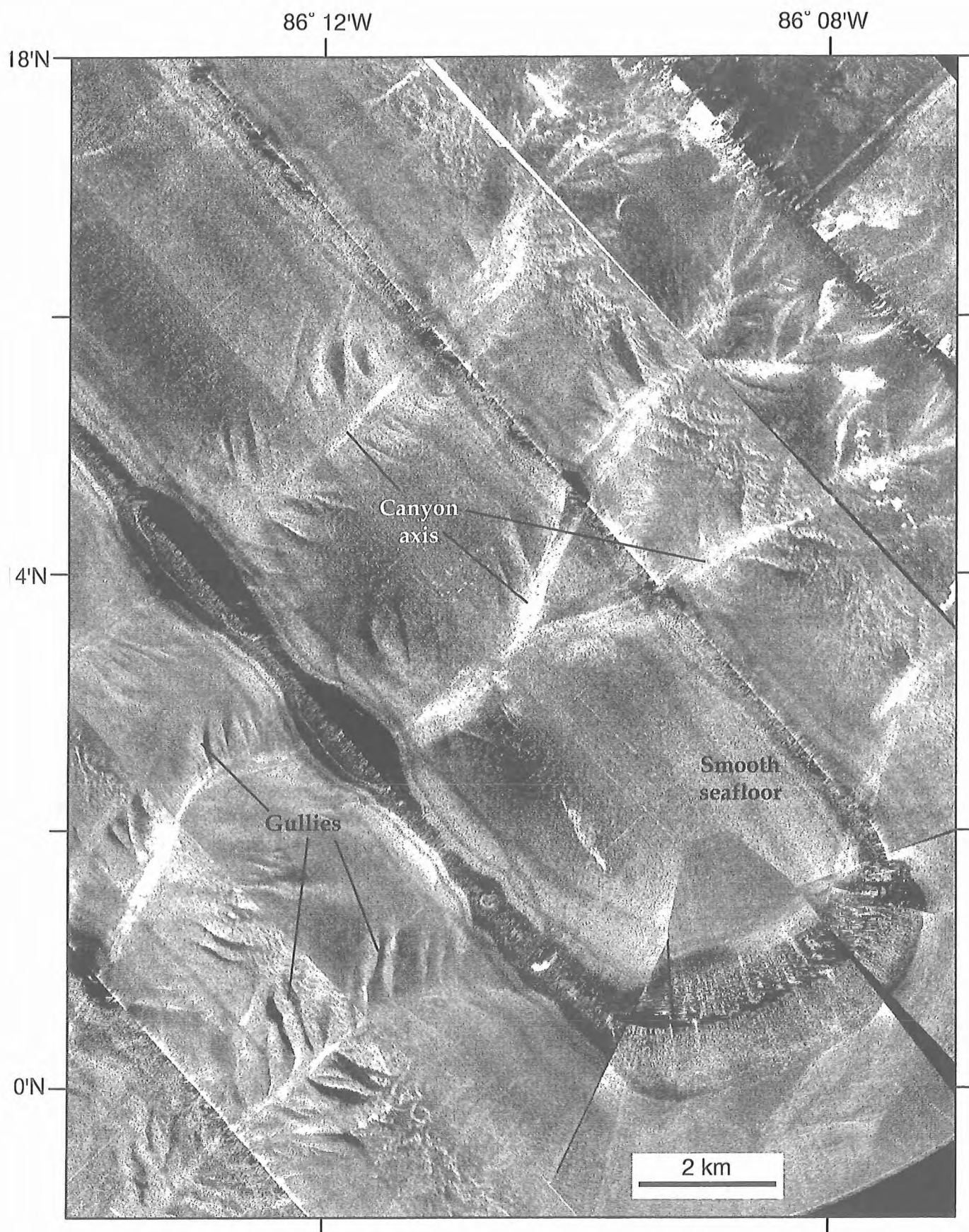
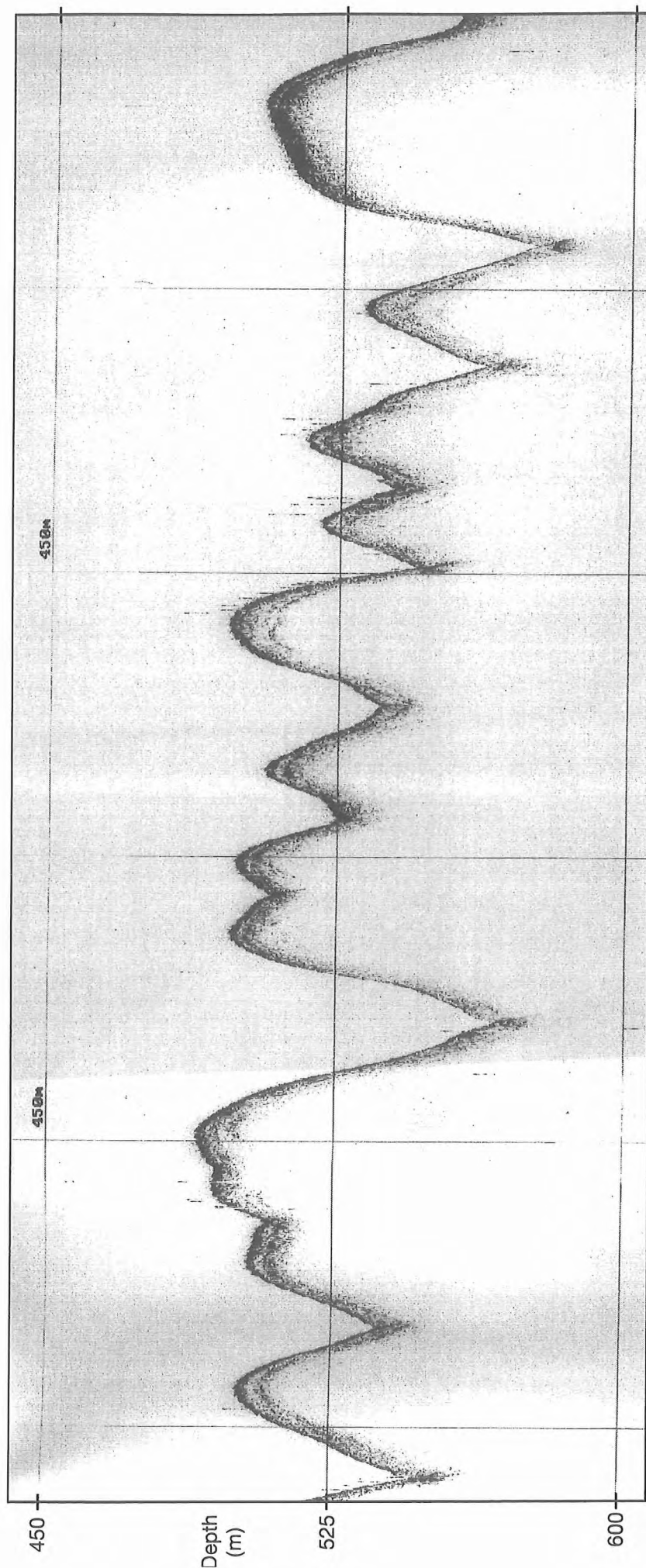


Figure 6.2.10 Sidescan sonar image of a dendritic canyon system in the western part of the survey area. Canyon axis are imaged as high backscatter stripes (white). Note the 'fish bone' pattern of the gullies on the walls of each canyon.



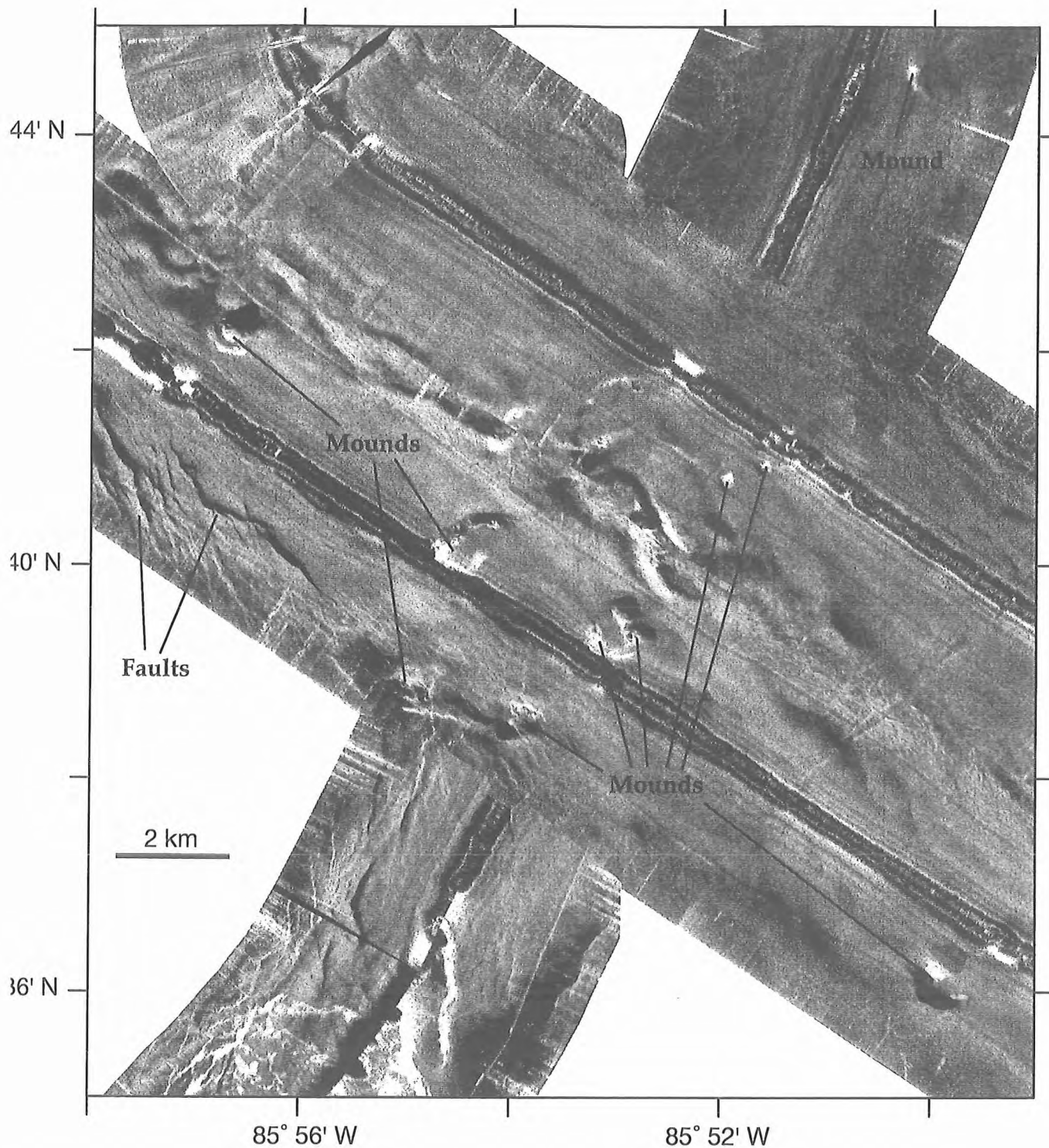


Figure 6.2.12 Sidescan sonar image of mounds in the western part of the survey area. For further detailed images of mounds see section 6.4.

Figure 6.2.11 (previous page): TOBI 7 kHz profile across the canyons on the upper slope in the western part of the survey area. Maximum canyon depth in this area is usually less than 100 m. Some sediment deposition can be seen on the interfluvies between canyons.

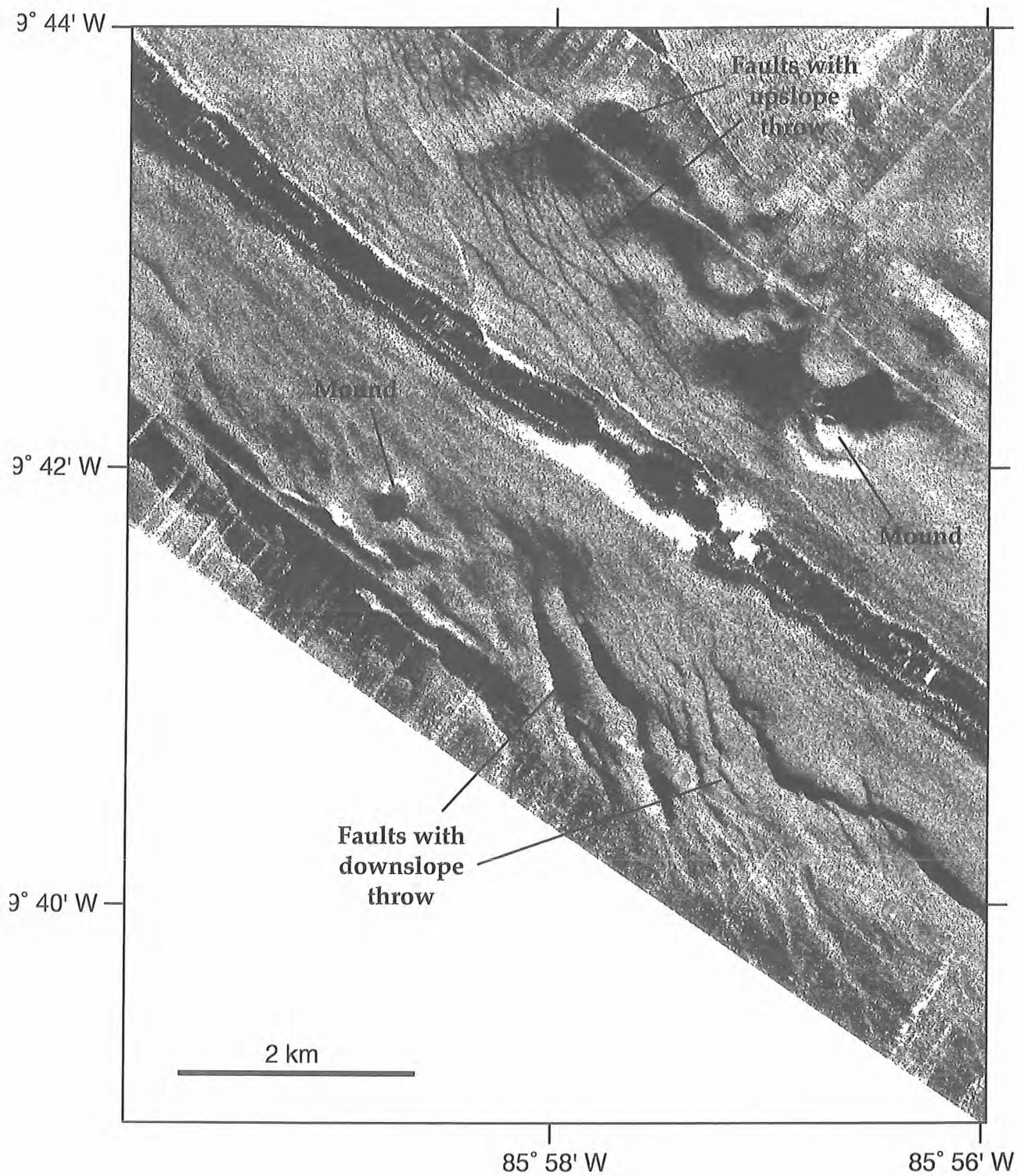


Figure 6.2.13 Sidescan sonar image showing the complexity of fault systems that are developed on some parts of the slope.

Seamount and ridge subduction

In the western survey area, a single example of an erosional scar related to seamount subduction was imaged (in part only) in the extreme northwest part of the area surveyed with TOBI. Toward the southeast limit of the western area, a single large oceanic basement ridge is subducted beneath the margin (Fig. 6.2.14). On the upper plate, the scar in the upper plate extends some 15 km landward of the deformation front. The material that partially fills the scar appears severely deformed, and probably consists largely of landslide debris. The headwall of the scar is marked by a pronounced scarp and several large arcuate faults

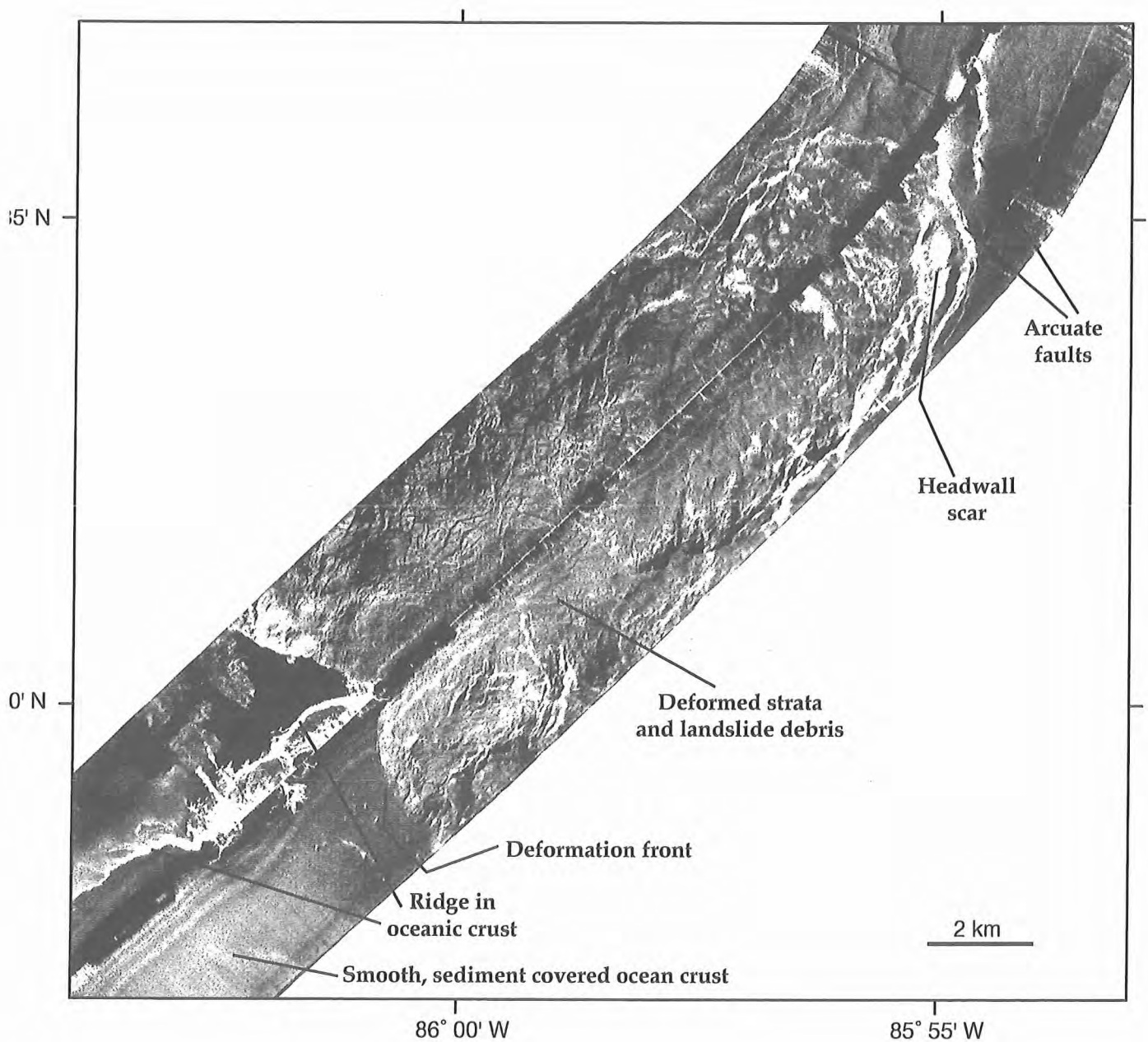


Figure 6.2.14 Sidescan sonar image showing an area where an oceanic basement ridge is being subducted.

The Costa Rican margin between 83° 15' and 84° 40' W

The main features seen on the TOBI images from this area include:

- (i) the deformation front between the subducting oceanic plate and the overriding continental margin
- (ii) several ridges and seamounts at various stages of subduction
- (iii) submarine canyons
- (iv) a variety of landslides
- (v) seafloor mounds

The deformation front

The deformation front, marking the outcrop of the plate boundary at the seabed, was imaged between 83° 20' and 84° 37' W. Except where it is affected by subducting seamounts and ridges, it forms a gently lobate feature separating smooth sediment covered ocean crust to the southwest from faulted and folded continental margin strata to the northeast (Fig. 6.2.15). On the basis of the sidescan imagery, we would suggest that very little sediment accretion occurs along this deformation front. Where accretion is suggested by the margin morphology, the accretionary wedge is rarely more than a few hundred metres wide (Fig. 6.2.15). The trend of normal faults in the ocean crust, which is slightly oblique to the deformation front, can be seen to extend into the overriding plate in several places, indicating that the subduction of 'steps' in the ocean crust can significantly modify the deformation front. There is some suggestion that accretion may occur primarily in the wake of such subducting steps, as the deformation front seeks to establish a new equilibrium profile. There is little evidence of significant sediment accumulation in the trench, with the exception of a few flat basins of limited extent that show low backscatter on the sidescan images; these are interpreted as ponded turbidites.

Seamounts and ridges

The deformation front is indented by several subducted seamounts and ridges of oceanic crust (Figs. 6.2.16, 6.2.17, 6.2.18). A small seamount that was been subducted at 83° 56' W is currently interpreted, on the basis of an uplift in the margin topography, to lie some 6 km landward of the deformation front (Fig. 6.2.16). Seaward of the uplift, a very steep scarp, generated by collapse of the uplift and landsliding as the support of the seamount progresses landward, is partially filled with debris avalanche material. The uplift itself is intensely fractured by arcuate faults, most of which throw downslope. This faulting extends well to the east of the scar cut by the seamount, and this, added to a second indentation of the deformation front at 83° 54' W, suggests that a second seamount may also have been subducted there (Fig. 6.2.16.). Between 83° 43' and 83° 48' W, three circular seamounts, each 2-3 km across at the base, are currently positioned adjacent to the deformation front (Fig. 6.2.17). One, at 83° 47' W, has just started to indent the deformation front; here, using OFOS, a small landslide scar was observed on the oversteepened slope. The other two seamounts have not yet collided with the margin, but they are of interest because of the unusual circular patterns which characterise their surfaces. Subducting ridges, such as that seen at 83° 42' W, create a high level of deformation in the overriding plate (Fig. 6.2.18). They drive deep re-entrants into the deformation front; these may migrate along the margin where the ridge is not exactly parallel to the convergence direction. Large fields of debris may be found at the base of slope and on the ocean plate where collapse of the oversteepened margin occurs as a result of ridge subduction (Fig. 6.2.18).

Submarine canyons

Submarine canyons occur in limited areas of the western part of the eastern survey area, where the slope is less steep than to the east. In contrast to the western area, canyons tend to be short, straight features of very limited length; dendritic systems are not seen (Fig. 6.2.19).

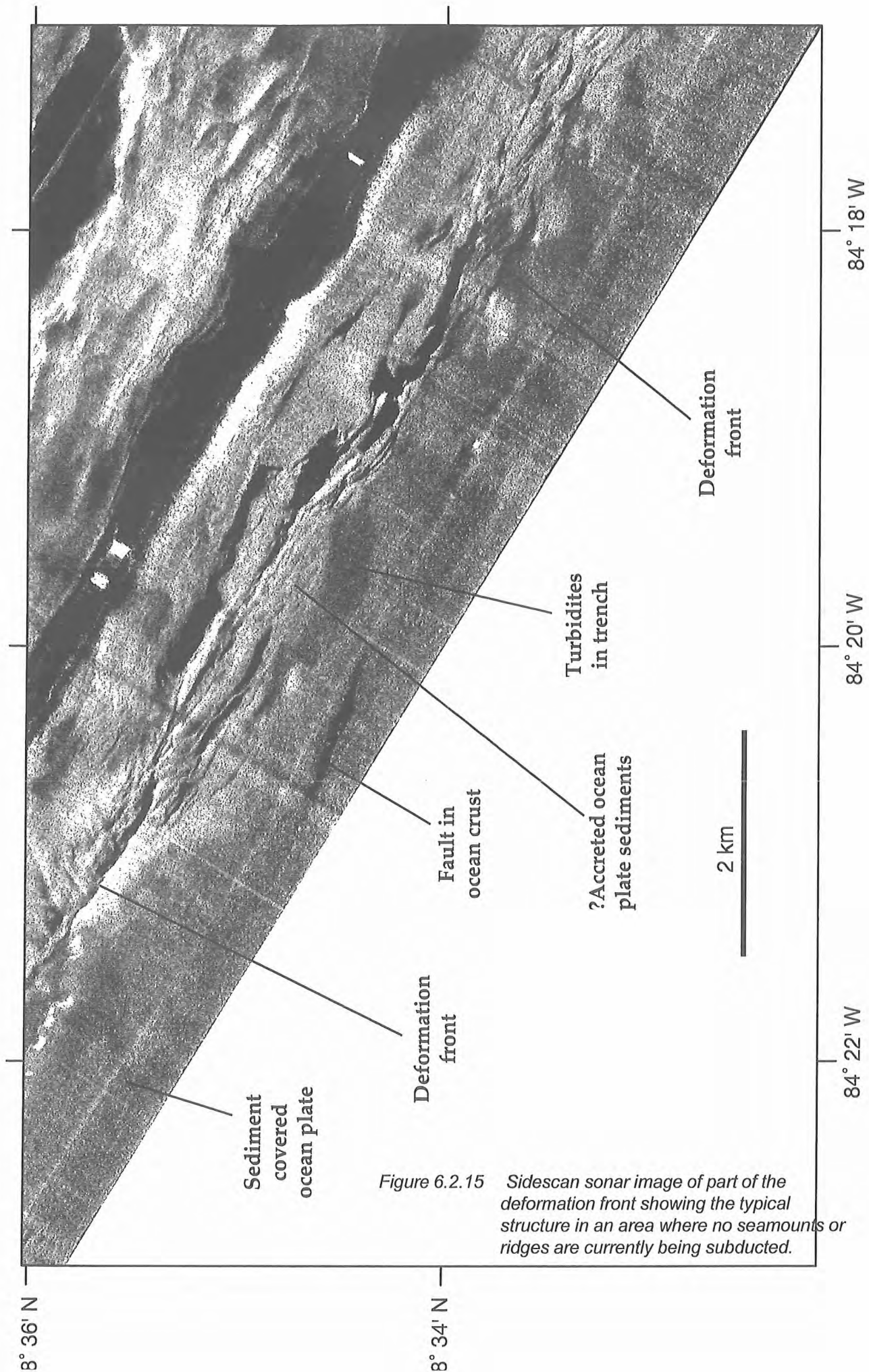


Figure 6.2.15 Sidescan sonar image of part of the deformation front showing the typical structure in an area where no seamounts or ridges are currently being subducted.

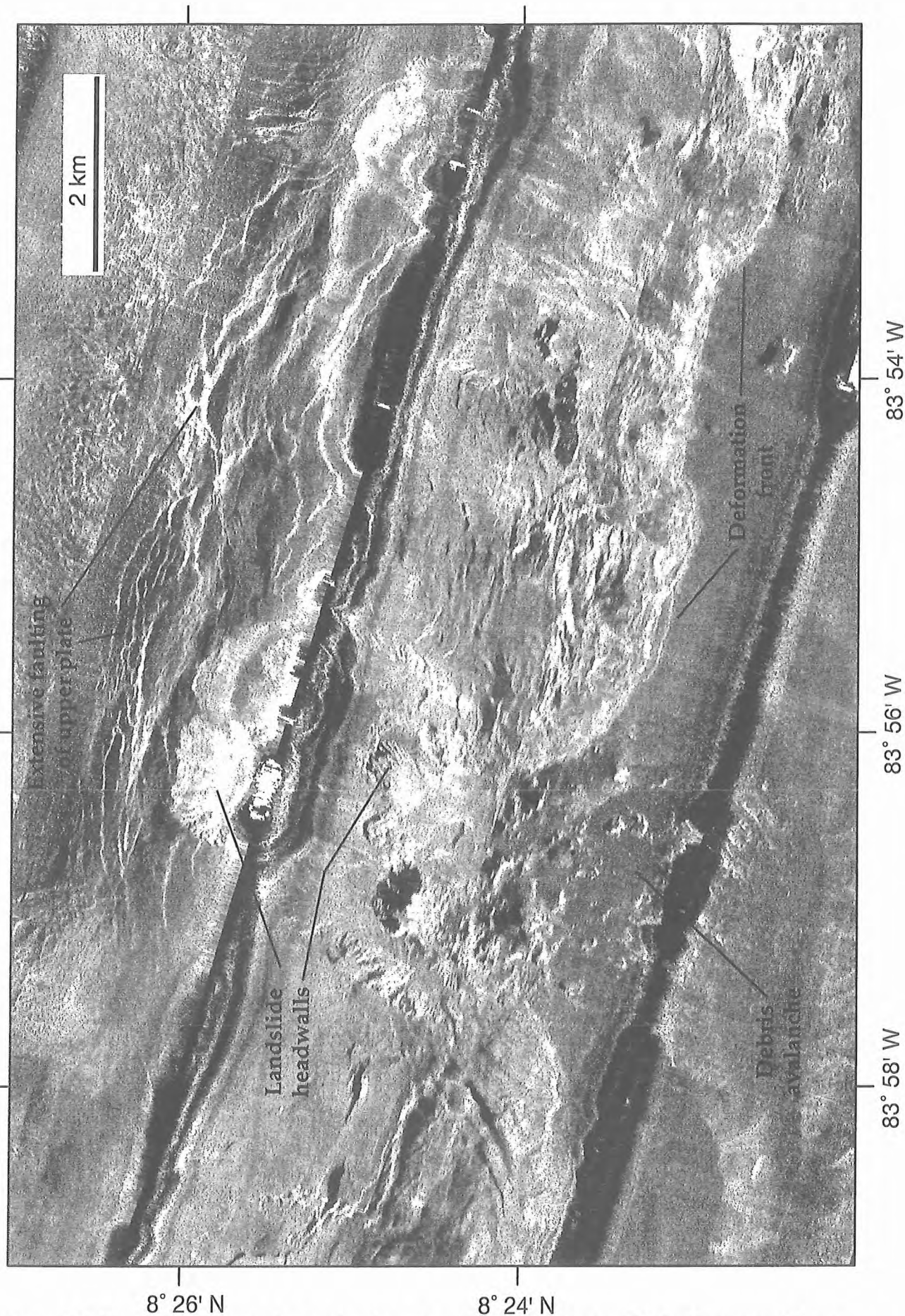


Figure 6.2.16 Sidescan sonar image showing an area where one (or possibly two) seamounts have been subducted. One scar, at $83^{\circ} 56' W$, is clearly defined by uplift and faulting of the upper plate above a scar cut through the plate by the seamount subduction process. A re-entrant is seen in the deformation front, and a large field of debris, derived from collapse of the steep headwall of the scar, covers the adjacent ocean plate. A second possible scar, at $83^{\circ} 54' W$, is marked by a small re-entrant and uplift and faulting of the upper plate, but

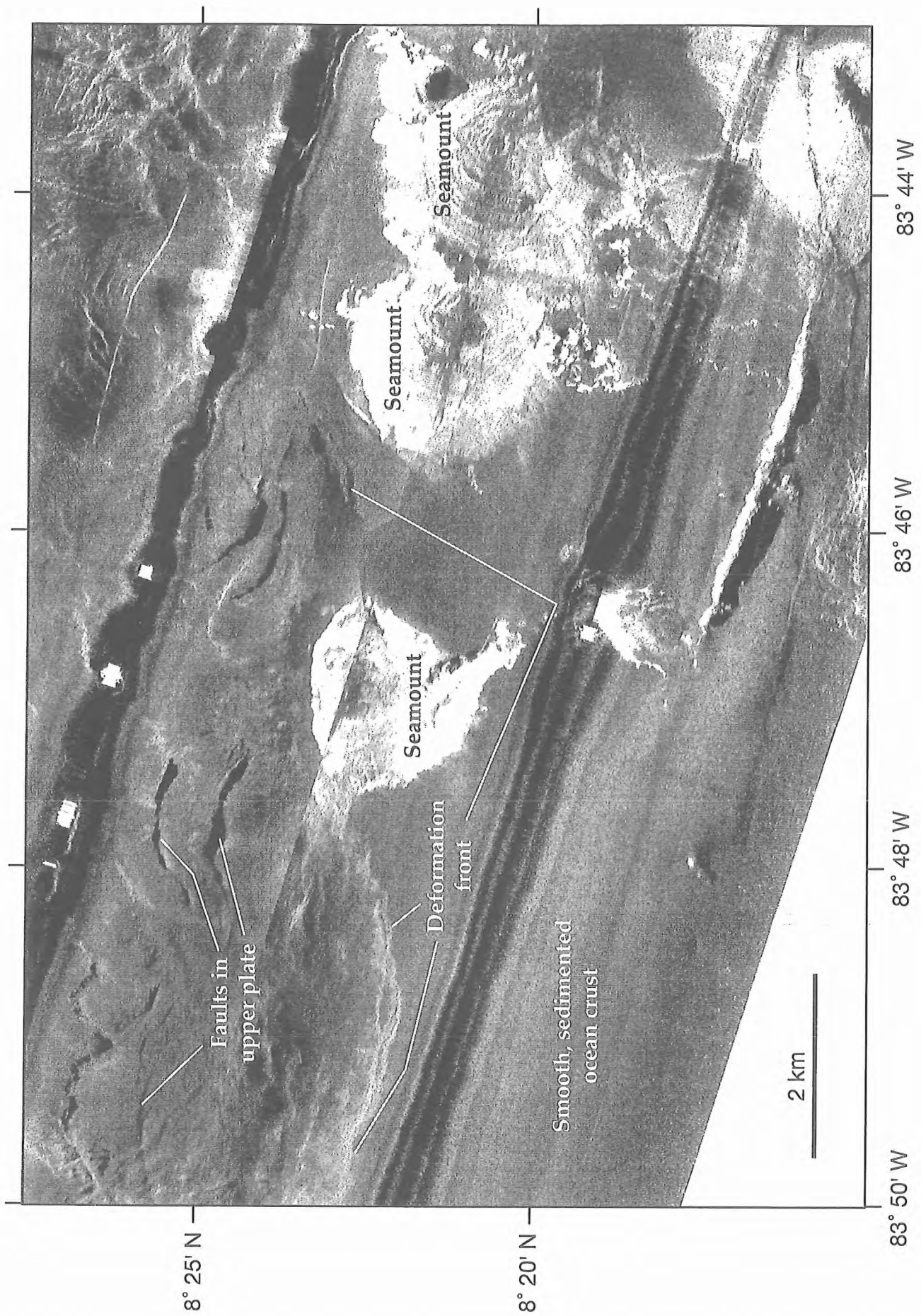


Figure 6.2.17 Sidescan sonar image showing three seamounts, defined by very high backscatter (white) which currently lie at or are approaching the deformation front.

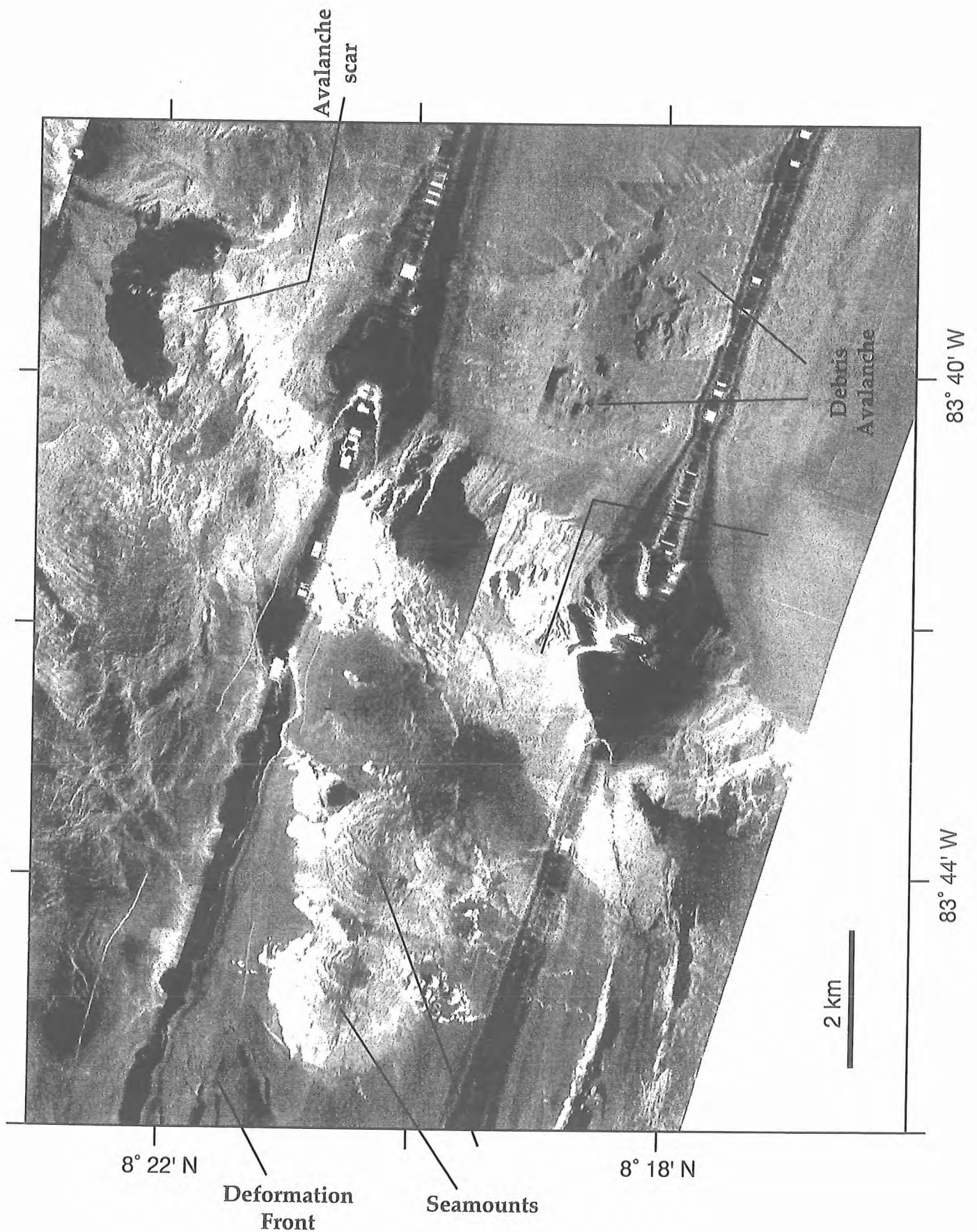
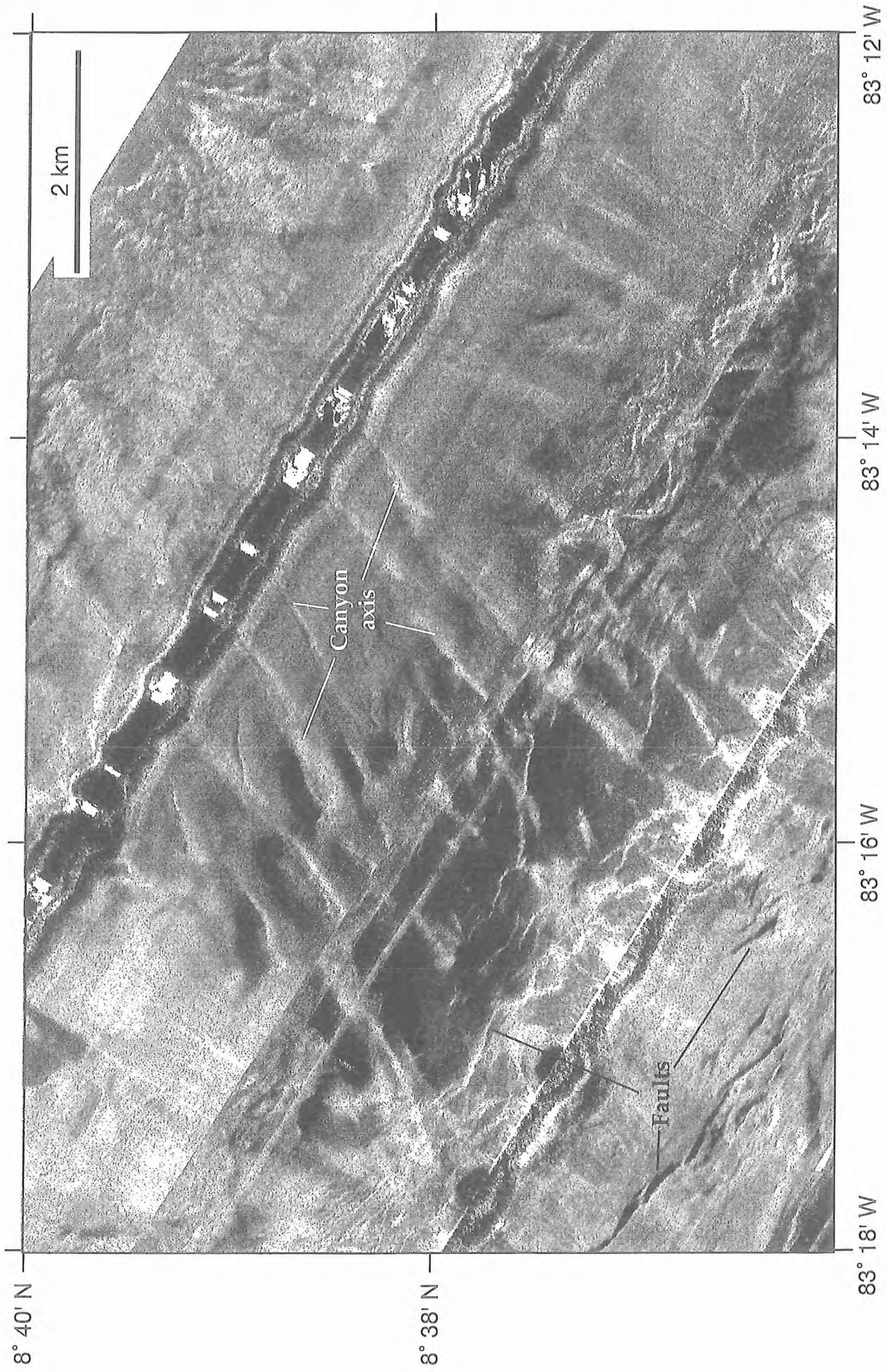


Figure 6.2.18 (above) Sidescan sonar image showing a large subducting ridge which indents the margin. Note the large field of debris derived from collapse of the oversteepened margin.

Figure 6.2.19 (next page) Sidescan sonar image of the typical submarine canyon morphology found in the eastern survey area.



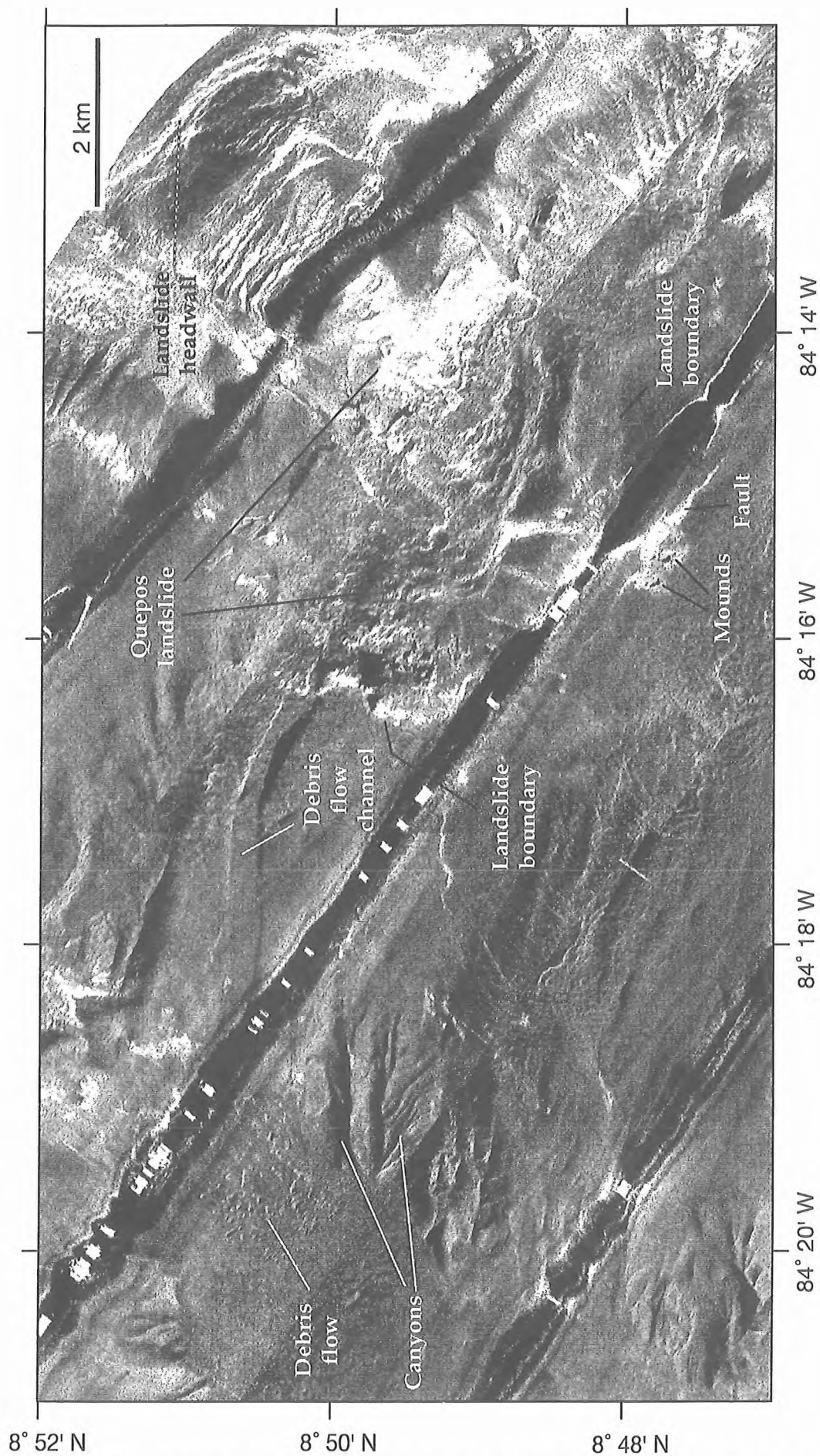


Figure 6.2.20 Sidescan sonar image of the Quepos landslide. The blocky and ridged topography suggests that slumping was the main process involved, with little disaggregation of the landslide material. Note the secondary debris flow channel and accumulation of debris to the west of the main landslide.



Figure 6.2.21 Sidescan sonar image of a linear group of mounds in the eastern survey area.

Landslides

A variety of landslide processes were observed in this area. As noted above, debris avalanches have occurred at several sites of seamount and ridge subduction (Figs. 6.2.16, 6.2.18). These contain angular blocks up to several hundred metres across. Near 8° 50' N, 83° 14' W, the Quepos landslide affects an area of some 8 x 4 km (Fig. 6.2.20). The morphology of this landslide suggests a slump and perhaps partial debris flow process. Immediately west of Quepos landslide, a broad, flat-floored channel or chute leads down to a accumulation of debris which fills a deeper valley. This is interpreted as a secondary debris flow triggered by the Quepos landslide.

Seafloor mounds

Mounds were less commonly observed in the eastern areas than either in the western area surveyed during Sonne 163 or the central area surveyed during Sonne 144 (see Fig. 6.2.9). A single group of four mounds, forming a linear chain, was discovered near 8° 56' N, 84° 18' W. This suggests that they are fault controlled. OFOS observations suggest that these were among the most active mounds examined in the area.

6.3 DTS1 operation and initial data interpretation

(C. Huguen, I. Klaucke)

Deep-towed side scan sonar deployments have been conducted during SO163-1 within the framework of subproject A1 of SFB 574. Subproject A1 is dedicated to continental margins processes and related changes in seafloor morphology. The EDGE-Tech DTS1 was deployed in the Central America subduction zone, offshore Costa Rica, to provide detailed surface (75 and 410 kHz side scan sonar) and subsurface (sub-bottom profiler) imaging of the continental slope. This deep-towed system has unfortunately not been operated in deep waters (oceanic plate) due to the length of the cable available on R/V SONNE. As a result, the survey areas focused on the continental slope (Fig. 6.3.1). They were chosen to assess a better understanding of both mud and fluid expulsions (DTS1-01, DTS1-02 and DTS1-05), as well as fault patterns and destabilisations related to seamount subduction (DTS1-03 and DTS1-04).

The GEOMAR deep-towed side scan sonar DTS1 has been successfully deployed five times during SO163-1 (Fig. 6.3.1), providing a total of **550 km** of low frequency side scan sonar (75 kHz) and **25.9 km** of high frequency (410 kHz) records. Both were run together with a 2-16 kHz sub-bottom profiler for **575.9 km** of high resolution seismic records. Details about the deployments are available on table 6.3.1.

Deployment	Average water depth	Track length	Mapped area	Time on sea-floor	Deployment	Recovering	Total
1	1300-2400	141.0 km	162.2 km ²	37 h	1 h 40	2 h	40 h 40
2	2500-3200	134.0 km	145.0 km ²	44 h	2 h 30	2 h 50	49 h 20
3	400-1900	172.4 km	175.2 km ²	44 h	35 min	2h	46 h 30
4	900-2000	120.4 km	161.0 km ²	30 h	2 h	2 h	34 h
5	1300-1500	9.0 km	1.8 km ²	2h	55 min	1h 05	4h

The *POSIDONIA* positioning system, installed on R/V SONNE during SO162 (*see* 5.5) was used to follow the position of the tow fish during deployments and to provide an accurate positioning on the sea floor. All data collected have been recorded by both *Hydrostar Online* (xse format) and the FS-DW unit directly (jsf format). Two backups of the xse files and one of the jsf were made on CD-ROM for the five deployments.

6.3.1 Deployment 1 (DTS1-01)

The first deployment of DTS1 during SO163 was conducted between 26.03.02, 6h16 and 27.03.02, 19h10, on an area of the continental slope, NW of Nicoya peninsula (Fig. 6.3.1). Here, small mound-like, circular features have been identified and initially interpreted as probable mud domes based on the multibeam bathymetry data (*see* 6.1.1) and the first TOBI deployment during SO163-1, which illustrates clear evidences of high backscatter patches (Fig. 6.3.2). During deployment 1 of DTS, 7 NW-SE trending parallel profiles have been run with a 1100 m track-spacing (Fig. 6.3.3), using the 75 kHz side scan sonar and the sub-bottom profiler. 162.2 km² of the sea surface have been mapped and 141 km of high resolution sub-bottom profiler data recorded over several of the inferred mud domes. A cross-cutting 10.4 km long track has been run with the high resolution (410 kHz) side scan sonar, along an E-W direction, over a small circular mud dome, located at 9°36 N - 85°49 W.

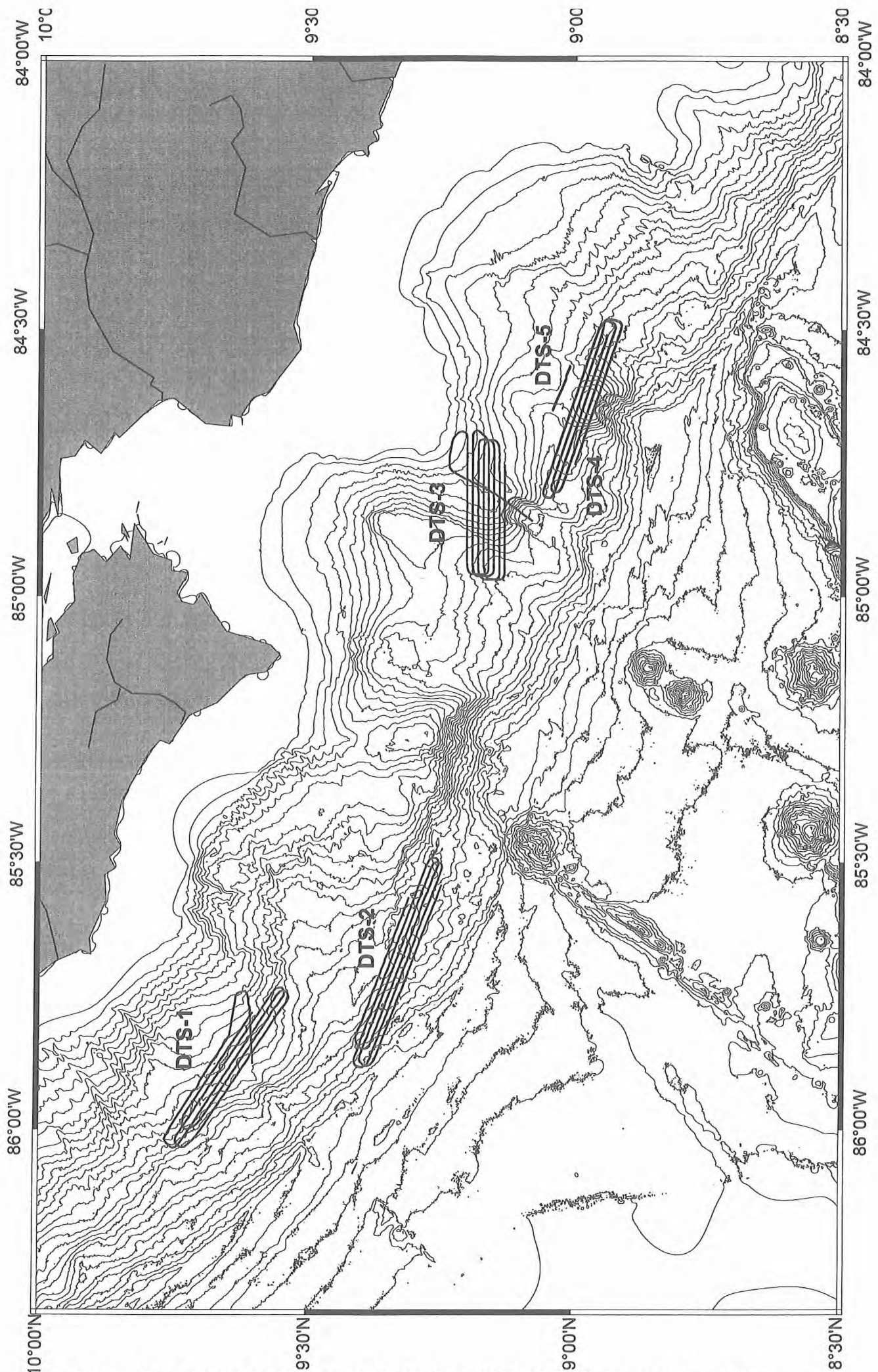


Figure 6 3 1: Map showing the ship's tracks for the five SO162-1 DTS deployments on a 100 m

During deployment 1 the range of the 75 kHz side scan sonar was set to 1500 m (0.49 Hz ping rate) and the data window size fixed to 26000 samples. The data window for the 2-10 kHz sub-bottom profiler, with its ping rate coupled to the 75 kHz side scan sonar, was set to 6000 samples, which corresponds to an average depth window of 250 m. For the last high resolution side scan sonar profile, the recording window size of the 410 kHz side scan sonar was fixed to 16000 samples corresponding to 100 m per side. The survey speed varied around 3.5 knots during all deployment. It was slightly increased during the turns (4.5 to 5 knots), in order to reduce them to less than 3 hours. During turns, enough cable was pulled in to keep the towed fish at a safe distance from the sea bottom.

Special remarks for deployment 1

- 1 - All signals were lost on *Hydrostar Online* for a few minutes between 12h58 and 13h16, on 26.03.02, and the winch stopped for safety reasons, while waiting for the signals to return.
- 2 - No response from the *Posidonia* fish transponder for more than 20 minutes between 13h17 and 14h39 on 26.03.02.
- 3 - Shortly after that, problems have been observed to find again the connection between Posidonia and the DTS acquisition system.
- 4 - DVS export data were not recorded between 15h55 and 19h54.
- 5 - DTS was crashed on a small mud mound, while using the high resolution side scan sonar (27.03.02, 19h05). The DTS frontal protection was slightly damaged during this crash, but the sensors were not affected.
- 6 - After recovering, the DTS titanium ring between the fish releaser and the umbilical cable (+ buoyant and pull in ropes) was broken. For the following deployments it has been replaced by a larger one for safer recovery.

6.3.2 Deployment 2 (DTS1-02)

The second DTS deployment during SO163-1 focused on an area downslope of the headwall of the prominent Nicoya slide (Fig. 6.3.1), where TOBI data recorded during SO 144-1 revealed scattered high reflectivity patches of yet unknown origin (Fig. 6.3.4).

This deployment lasted from 31.03.02, 10h34 to 02.04.02, 11h49. It consists of 6 WNW-ESE parallel lines (Fig. 6.3.5) along which the low resolution side scan sonar (75 kHz) and the sub-bottom profiler were run together, for a total length of 134 km and a 145 km² surveyed area. As during deployment 1, the range of the 75 kHz side scan sonar and sub-bottom profiler, was set to 1500 m (data window size 26000 samples) and 250 m (6000 samples), respectively. The ship's speed was kept at 3.5 knots and reduced to less than 3 knots when reaching 3000 m of water depth, in order to reduce the cable length and keep the tow fish within range of the sea-floor.

Special remarks for deployment 2

- 1 - The side scan sonar system worked properly during the whole deployment, except a 20 min break down on 02.04.03, between 22h32 and 22h53. The whole DTS acquisition system had to be restarted and the connection with *Posidonia* underwater navigation was lost after the restart of the DTS logging system.
- 2 - Two important *Posidonia* crashes also occurred during this deployment. A first one on 31.03.02 between 18h24 and 20h16, after 1818 cycles, and a second one on 01.04.02, between 9h16 and 12h51. No real explanations have been found for both crashes. Communication problems between the ship's information (DGPS, VHRU) and the *Posidonia* system have been suspected. Moreover, both communications losses between the antenna and the fish's transponder occurred in water depths between 2500 and 3000 m and could be related to large cable length (more than 6000m).

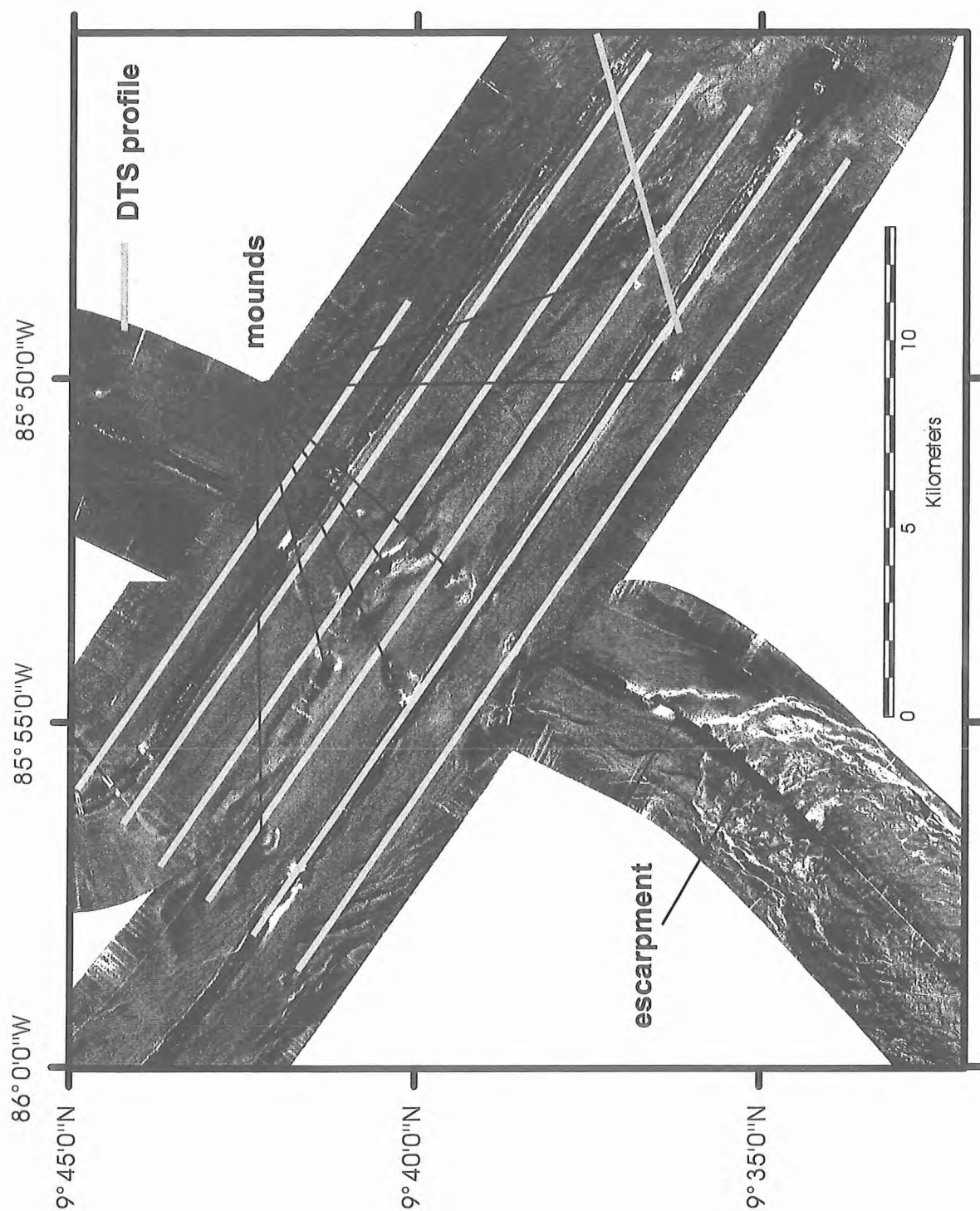


Figure 6.3.2: Positions of the first DTS deployment profiles superposed to TOBI imagery from SO163-1 (see 6.2).

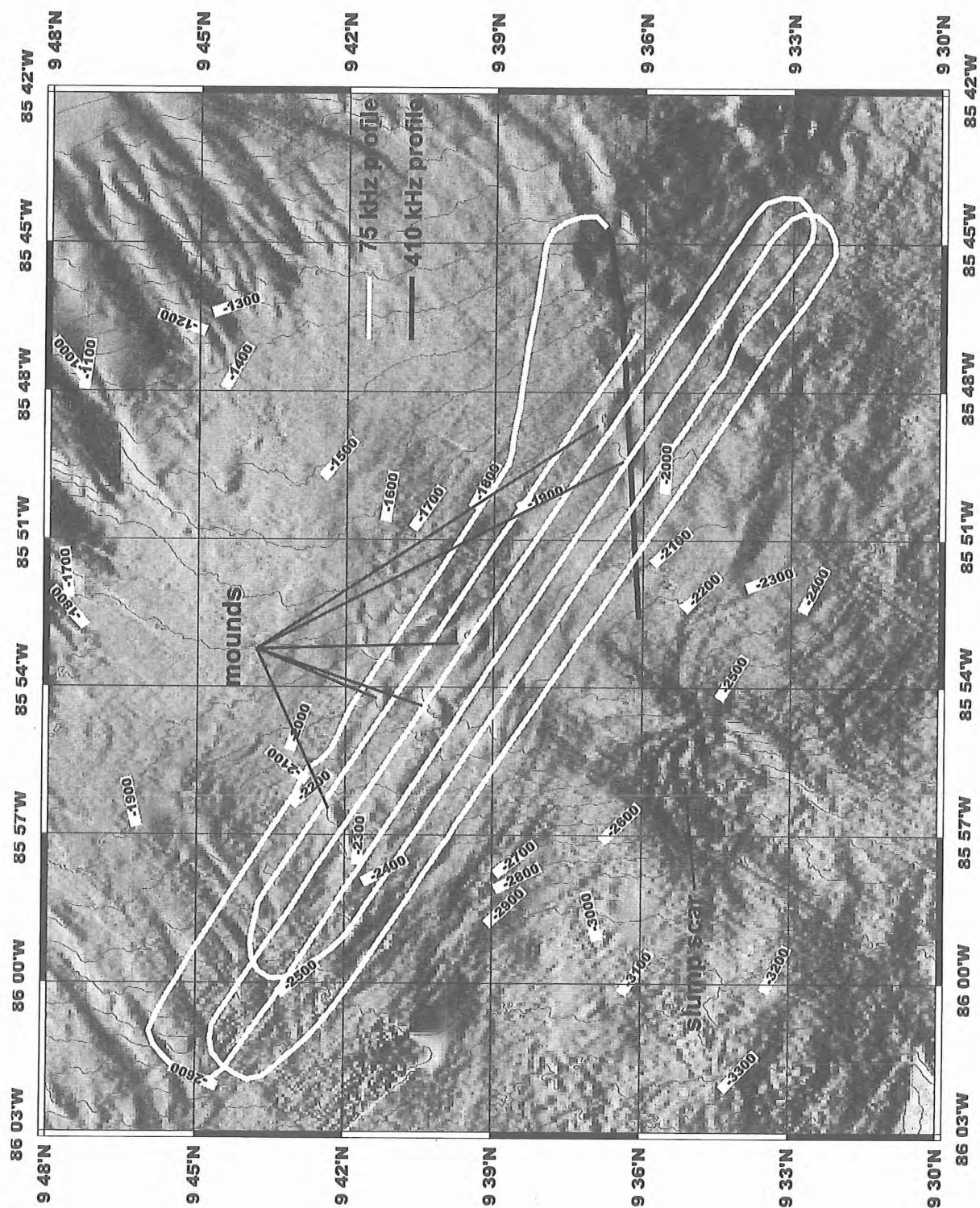


Figure 6.3.3: Map showing the ship's tracks during the first DTS deployment (DTS1-01) on the continental slope NW of Nicoya Peninsula.

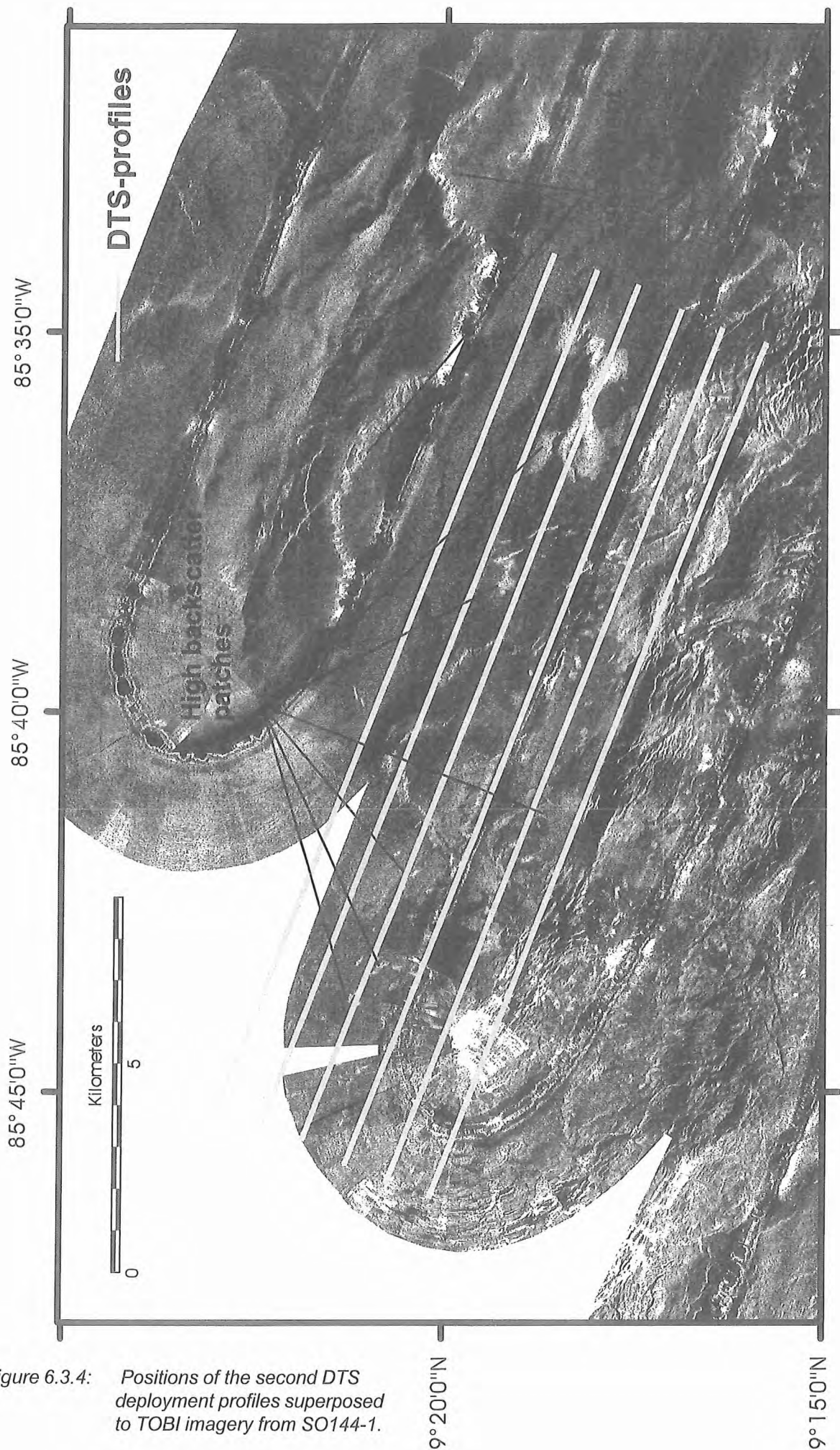


Figure 6.3.4: Positions of the second DTS deployment profiles superposed to TOBI imagery from SO144-1.

6.3.3 Deployment 3 (DTS1-03)

The third deployment of DTS1 was focused on the uplifted and faulted area upslope of Jaco Scar (Fig. 6.3.1) that was already surveyed with TOBI during SO144-1 (Fig. 6.3.6) and lasted for 46h30 between 03.04.02, 15h39 and 05.03.02, 13h54.

It consisted of 8 E-W trending parallel profiles upslope of Jaco Scar and a cross-cutting NE-SW profile ending at the base of Jaco escarpment (Fig. 6.3.7). A total of 44h of DTS records has been collected including 165.8 km of low resolution side scan sonar and sub-bottom profiler data covering 173.9 km² of the seafloor. In addition, a 6.52 km long high resolution side scan sonar profile was run, covering 1.3 km² of the seafloor. Starting with the same settings than for the two previous deployments the ranges have been switched to 750 m for the 75 kHz side scan sonar (0.98 Hz frequency) and 100 m for the high resolution side scan sonar from 05.04.02, 7h15 in order to optimise the along-track resolution. The ship speed was also held around 3.5 knots for most of the deployment and reduced to 2.8 knots during the high resolution profile for an improved along-track resolution.

Special remarks for deployment 3

- 1 - The contact with the sonar has been lost for 120 pings on 03.04.02, around 17h32 (no xse data recorded by *Hydrostar online*).
- 2 - Strong noise due to a sports fishing boat closeby has been observed on the sub-bottom profiler at 21h02 on 04.04.02. The sub-bottom profiler signals were strongly disturbed during more than 15 min, but the same noise has not been observed on the side scan sonar records. The DTS side scan sonar and profiler records have also been disturbed with some noise generated from another boat on 05.04.02, around 4h18.
- 3 - The *Posidonia* positioning system worked well during deployment 3. No crash or loss of connection with the fish's transponder have been observed. This probably results from smaller water depth and cable length.
- 4 - The connection with the *FS-DW* data disks has been lost on 05.04.02, 5h48. The fish storage disks D and E remained still not visible after reboot of the *FS-DW* unit. As a consequence, jsf data were not recorded until the end of the deployment. The *DTS* data were only recorded in xse format on the *Hydrostar Online* computer. After recovery of the system both storage disks were checked and they both worked properly. Therefore a faulty connector has been suspected as the cause of the problem. The connection between *Hydrostar Online* and the *FS-DW* unit was checked with DTS1 on deck before deployment 4. Both storage disks were visible and the jsf files recorded properly on the hard disk.

6.3.4 Deployment 4 (DTS1-04)

The fourth deployment of the DTS1 side scan sonar during SO163-1 focused on the area upslope the Parrita scar, located SW of the Jaco scar investigated during deployment 3 (Fig. 6.3.1). TOBI records available from SO144-1 (Fig. 6.3.8) at the front of the Parrita seamount, showed both intense faulting and large highly reflective patches, generating a complex morphostructural pattern.

This deployment lasted from 15.04.02, 17h14 to 17.04.02, 00h42 and consisted of 5 WNW-ESE parallel lines. As for all other deployments track-spacing was 1100 m (Fig. 6.3.9). As during deployment 2 only the 75 kHz and the sub-bottom profiler (2-10 kHz) were used along a total distance of 120.4 km providing 161 km² of total seafloor coverage. On the basis of the good quality data recorded during deployment 3 after the range modification, the same 750 m range (0.98 Hz frequency) has been been conserved for the 75 kHz side scan sonar during deployment 4. The ship speed averaged 3.5 knots during the whole deployment and had to be increased to 4.5 knots when crossing the major fault scarps, in order to maintain the fish at a safe distance from the seafloor. With some escarpments being over 700 m high, it proved important to keep cable length

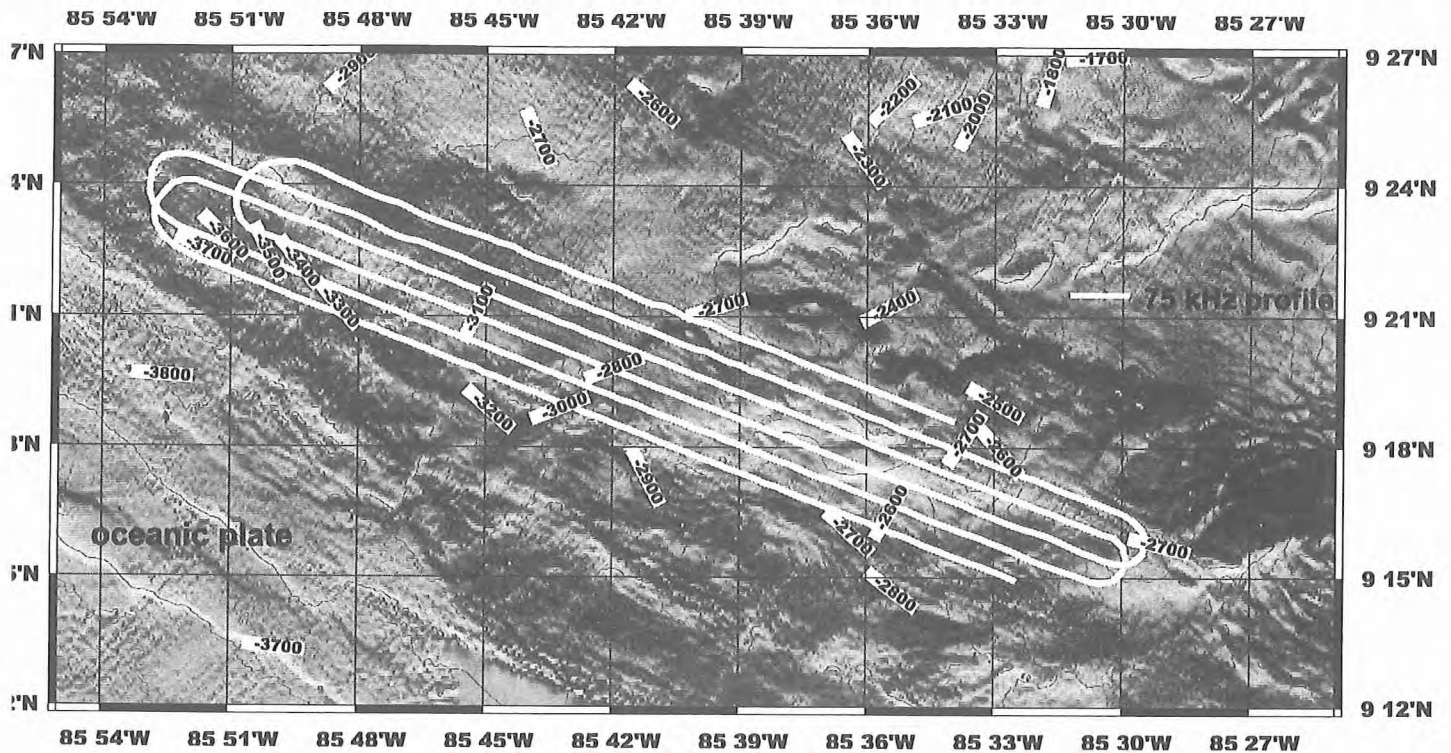


Figure 6.3.5: Map showing the ship's tracks during the second DTS deployment (DTS1-02) downslope of the headwall of the Nicoya slide.

at a minimum while approaching the escarpments. Reduced ship's speed of 3.0 knots allowed to have a shorter cable at the foot of the escarpment. Still, hauling at winch speeds of 1.0 m/s was necessary to keep the side scan fish off the seafloor.

Special remarks for deployment 4

- 1 - During this deployment the fiber optic cable has been used for the first time. After a first test with DTS on deck, the fiber optic modem initially set to a *low energy mode* was changed to the *high energy mode*, more appropriate for the 7 km long cable available on R/V SONNE. With such a configuration, the data and power signals are not modulated together as with the coaxial cable. As a consequence, significantly less noise was observed on the side scan sonar records during deployment 4.
- 2 - During deployment 4 serious problems have been observed with the *Posidonia* positioning system. Almost no position values were obtained from sensor H2 during the whole deployment and the calculated positions of the fish are strongly scattered around the ship. In some cases the fish's position had even been located ahead the ship... The only useful data available from the *Posidonia* system during this deployment are the depth sensor information. After recovering the DTS, the *Posidonia* acoustic array mounted below the ship hull has been checked and no real problem has been found. One possible explanation could be the presence of a fishing line around one of the sensors, lost when stopping the ship before pulling in the antenna or while pulling in the antenna. In any case, during the next deployment (DTS1-05), the *Posidonia* positioning system worked without problems.

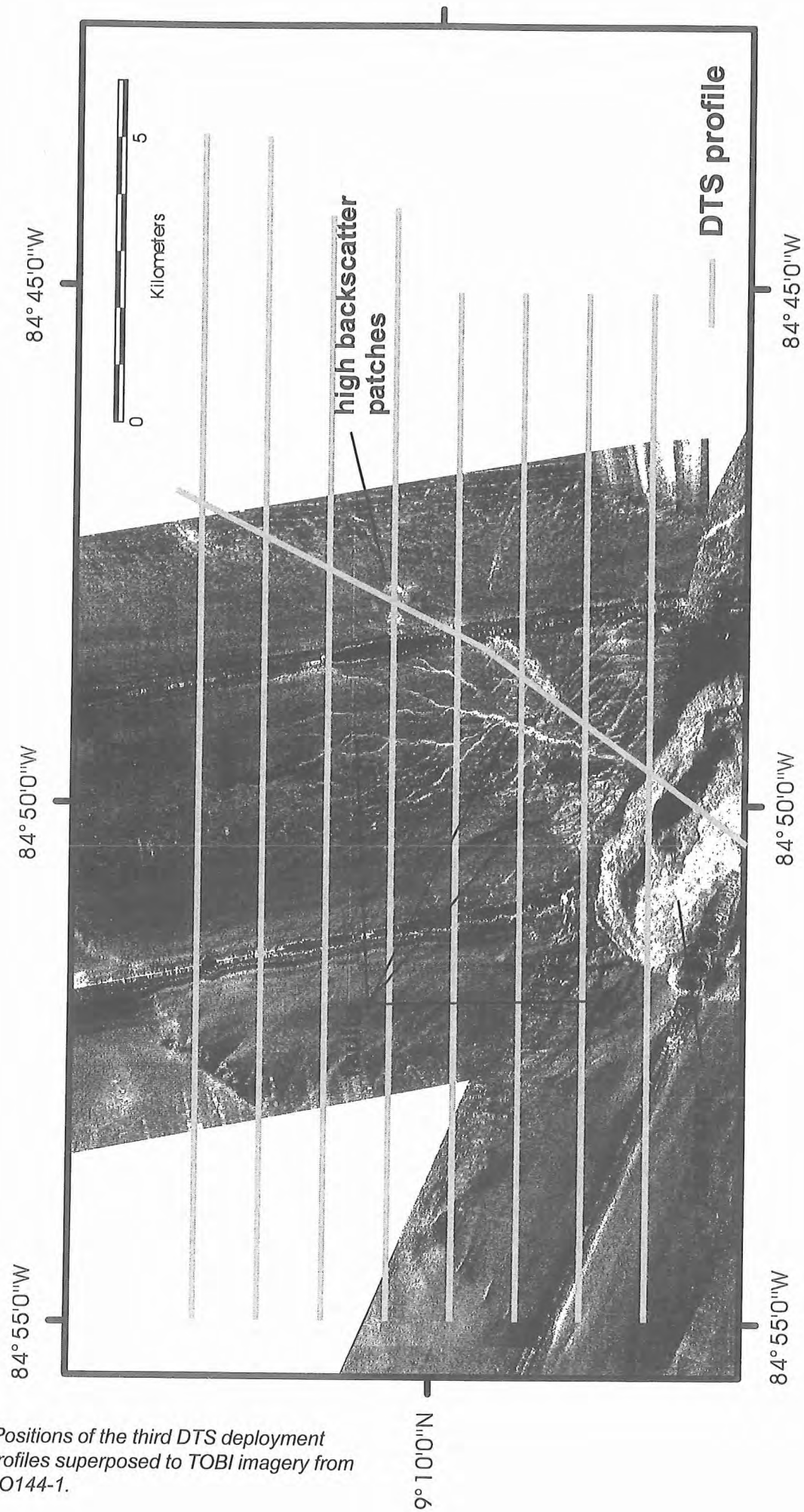


Figure 6.3.6: Positions of the third DTS deployment profiles superposed to TOBI imagery from SO144-1.

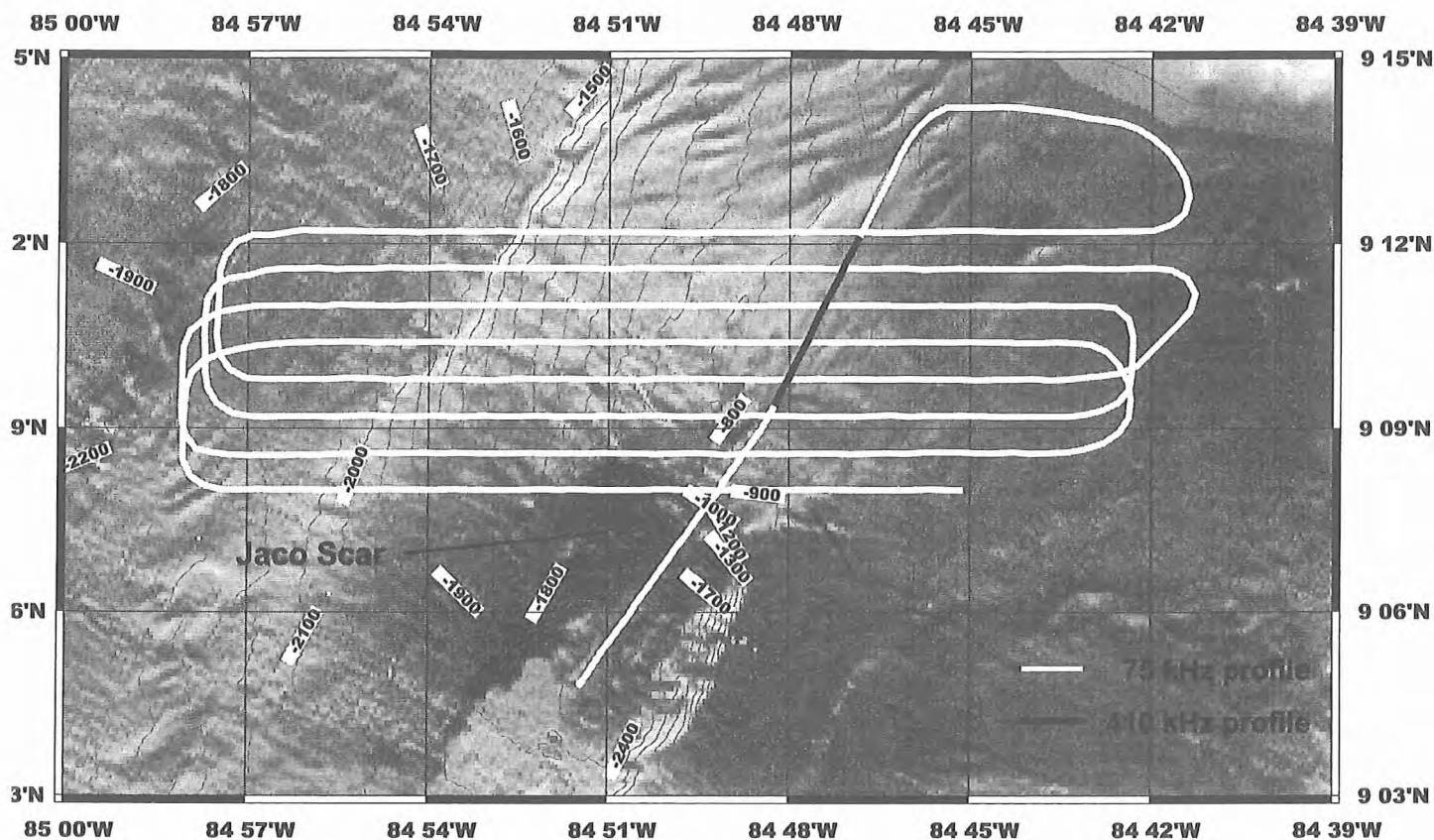


Figure 6.3.7: Map showing the ship's tracks during the third DTS deployment (DTS1-03) upslope of the Jaco Scar.

6.3.5 Deployment 5 (DTS1-05)

The fifth and last deployment of the DTS1 side scan sonar during SO163-1 was rather short on 17.04.02 between 18h10 and 22h20. It consisted of one high resolution (410 kHz) side scan sonar profile, located upslope Parrita scar, just north of the DTS1-04 low resolution profiles (Fig. 6.3.1, 6.3.8 and 6.3.9). This 9 km long profile has been run on the basis of the OFOS profiles obtained earlier this day over a small mound north of the Parrita scar (see 6.4), where intense activity has been observed.

During this deployment, the data window size was set to 10000 samples for the high resolution side scan sonar, which was run together with the 75 kHz side scan sonar (range 750 m, 26000 samples) and the sub-bottom profiler (range 250 m, 6000 samples). As during deployment 4 the fiber optic cable was used. The ship speed was kept around 2.5 knots in order to reduce the cable length for a more safe operation of the high resolution side scan sonar and to improve along-track resolution of the side scan sonar.

Special remarks for deployment 5

- 1- The connection to the fish was lost twice during deployment 5, for 1 or 2 minutes at 19h51 and for 100 pings at 20h27. No xse data have been recorded by *Hydrostar Online* during this time.
- 2- The quality of the sub-bottom profiler records was relatively poor during most of the profile, due to the relatively low speed of the ship resulting in a constantly important pitch for the fish. A slightly better display has been reached when the gain was reduced to values of 8.
- 3- The main objective of this deployment was to map a relatively small structure (no more than 200 m in diameter) which can easily be missed with the high resolution side scan sonar having a range of 100 m. The *Posidonia* positions were used to calculate the necessary offset of the ship to keep the fish exactly along the line and the structure was perfectly mapped.

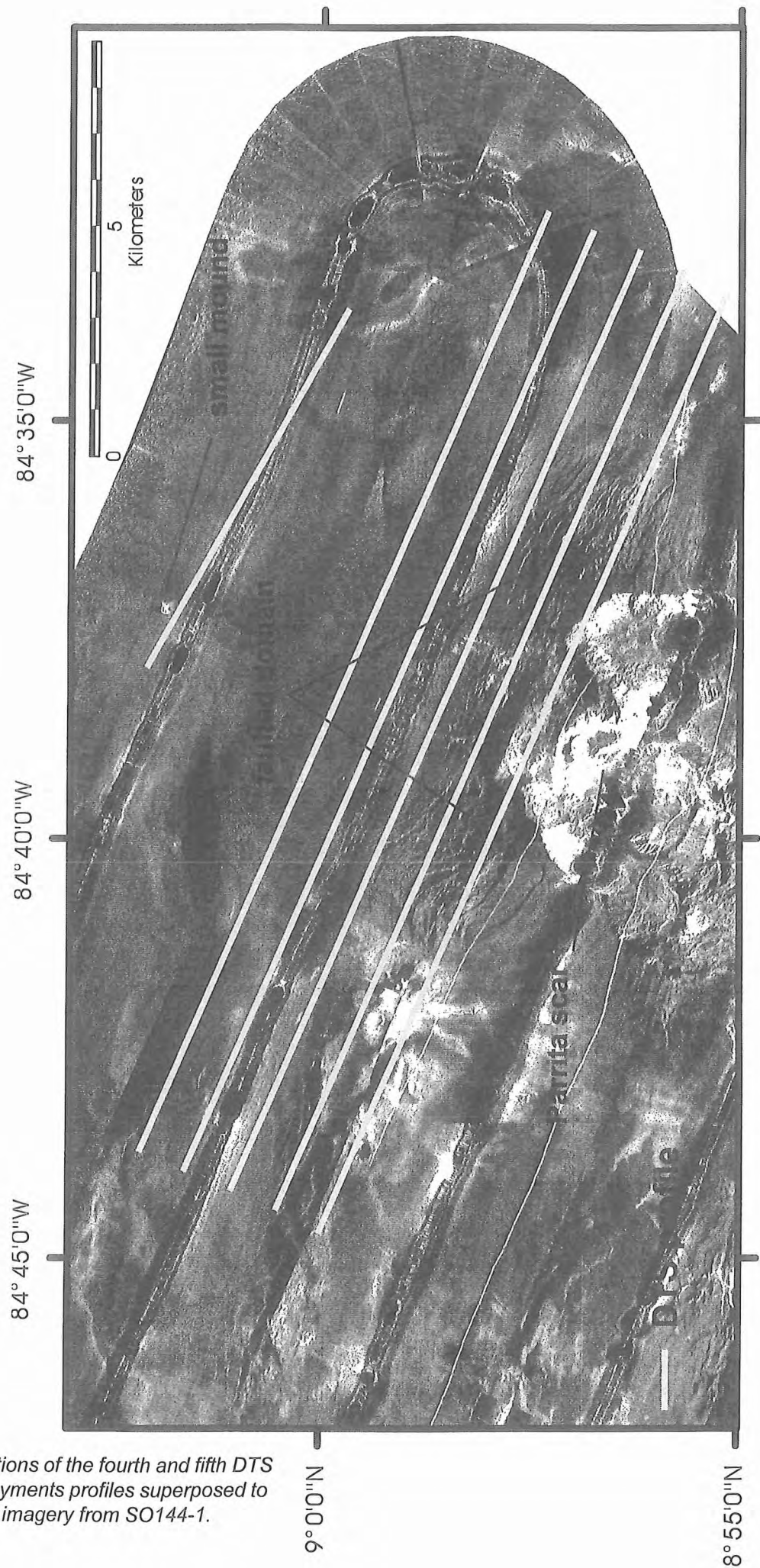


Figure 6.3.8: Positions of the fourth and fifth DTS deployments profiles superposed to TOBI imagery from SO144-1.

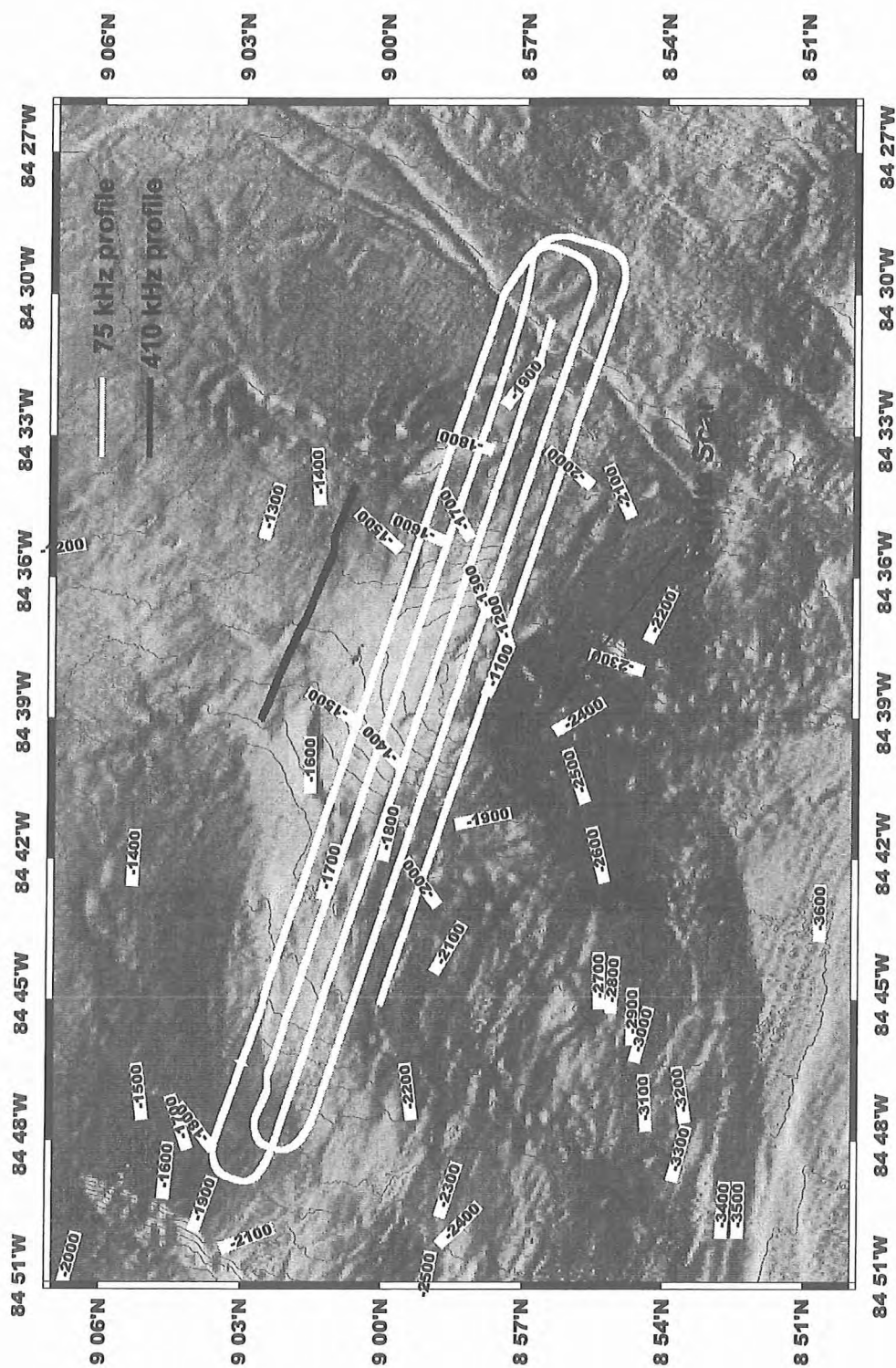


Figure 6.3.9: Map showing the ship's tracks during the fourth and fifth DTS deployments (DTS1-04 and DTS1-05) upslope of the Parrita Scar.

- 4- The *Posidonia* positioning system was run during this deployment without prior calibration after the *Posidonia* array had been pulled in after deployment 4. Yet, *Posidonia* underwater navigation data seemed to be sufficiently accurate to hit a small target.

6.3.6 General Remarks

a- Different types of noise observed

- During all deployments, regular interferences related to the *Posidonia* pinging sequence (5s pinging rate) has been clearly evidenced on the low resolution side scan sonar records and on the sub-bottom profiler.
- Some interference has been observed when running the high resolution side scan sonar (410 kHz) related to the sub-bottom profiler.
- Proximity of other boats always results in strong noise either on the side scan sonar or the sub-bottom profiler.
- Strong noise has also been observed at the beginning of all deployments on the low resolution side scan sonar. This noise decreases progressively during the first 30 minutes of DTS deployment and then completely disappears. This might be related to internal gain settings of the side scan data. As long as no bottom return is registered the gain is very high and amplifies noise. An alternative explanation might be that parasite energy level resulting from the start up sequence need some time to dissipate.
- Every day, between 20h03 and 21h50 UTC, a weak noise has been observed on the DTS data. All scientists involved agree to attribute this noise to some daily internal operations in the tow fish. The automatic backups done by the *Bottom PC* of the seismic streamer acquisition system that is housed in the telemetry pressure housing, has been switched off after deployment 3 and this noise completely disappeared.

b- A good bathymetry for safe deployments

- One of the most important remarks following those five deployments is the real need of a detailed bathymetric map for safe flying of the tow-fish, especially when running the high resolution mode with the fish only 20 m above the seafloor. A detailed bathymetry also helps to anticipate seafloor escarpments or other strong changes in relief and to keep the tow fish at a safe distance from the bottom by hauling in enough cable prior to bathymetric escarpments. During SO163, the available bathymetry was only a 100m grid and most of the fault scarps along the DTS profiles were not clearly shown. The result was two *emergency* situations during deployments 3 and 4 when cutting across the fault scarps upslope the Jaco and Parrita seamounts. Cable had to be pulled in at 1.0 m/s with ship speed increased from 3.5 to 4 knots during deployment 3 and at 1.5 m/s with a ship speed increased to 4.5 knots during deployment 3, in order to avoid bottom contact.

6.3.7 Preliminary observations and scientific results

Even though significant progress been in understanding and reading the *DTS* data format (thanks to W. Weinrebe during this cruise), we are still not yet able to process the side scan sonar data. Both subbottom profiler records and side scan sonar images are not georeferenced and corrected for towfish altitude variations, and we still have to find a way to properly filter and use the *Posidonia* underwater navigation.

As a result, we are not able to provide accurate scientific results or preliminary interpretations, but some interesting points can still be pointed out based on the online observation of the raw data during the deployments.

a- Structures related to mud and fluid escapes

One of the main interests of the DTS backscatter images is to provide information concerning the seafloor topography and lithology. An anomalous high backscatter signal could for example be tentatively interpreted as an evidence for a recent mud flow or a fluid escape structure. This information were obviously of high interest during DTS deployments 1, 2 and 5, focused on 3 areas where mud and fluid escape were supposed to take place.

- **DTS1-01:** During the first DTS deployment small bathymetric mounds have been investigated on the continental slope NW of the Nicoya Peninsula. The DTS low resolution side scan images illustrates two different types of anomalous high backscatter patches. **(1)** Some large highly reflective accumulations over several hundreds of meters, correlated with an echo-free surficial layer on the subbottom profiler which might correspond to mud accumulations. Some of them show scattered domains characterised by a more chaotic acoustic pattern, which could be interpreted as the acoustic expression of localised encrusted areas, covered by thick carbonate crusts. Clear relationships have been observed with tectonic features for several of these inferred mud flows. They could be observed on side scan images with small scale sigmoidal features like tension gashes or linear faults followed over hundreds of meters and clearly imaged on the subbottom profiler records when crossed by the fish. **(2)** Numerous small bright spots have been observed along all the lines. They might correspond to isolated fluid or gas escape centers like pockmarks structures. No special relationships with tectonic features have been observed so far.

- **DTS1-02:** During this deployment, several small reflective mounds which might correspond to mud diapirs have been observed, but no large highly reflective mud flow could be evidenced. The data quality of the subbottom profiler was really good during this deployment and illustrates a complex tectonic pattern with faults and tilted blocks clearly evidences on the profiles. This intense fracturation could constitute possible pathways for fluids and mud from depth to the surface.

- **DTS1-05:** The high resolution side scan sonar run during deployment five provides a very good quality backscatter image over a subcircular feature of about 200 m in diameter. Both the morphology and the backscatter strength of this topographic relief seems to indicate a mud construction. Moreover, some concentric circular rims have been clearly evidenced over the whole structure and could be the result of two different emplacement processes: **(a)** a progressive construction of the dome by successive flows, if an “extrusive” emplacement mode is evoqued, or **(b)** some circular cracks generated by a diapir-like “intrusive” emplacement. Some localised chaotic acoustic patterns have also been identified, and could be explained by the presence of carbonate crusts on top of the structure. Therefore, this relief could correspond to a relatively old structure and the circular rims be the result of a late subsidence of the whole structure, after the fluid escape activity has decreased.

During SO163-1 OFOS has been deployed on most of these possible mud structures (see 6.4) and provides the necessary ground-truthing for a meaningful interpretation of the DTS reflectivity data. Further detailed correlation between OFOS and DTS are definitely required to better understand the real signification of the acoustic patterns.

b- Fault patterns upslope Jaco and Parrita scars

Jaco and Parrita scars areas are well known since the SO144 expedition as complex tectonic features related to seamount subduction beneath the overriding plate, SE of the Nicoya Peninsula. The bathymetric highs observed upslope of both scars as well as their lateral walls are characterised by a complex morphostructural pattern shaped by crosscutting faults.

DTS deployments 3 and 4 have been run upslope Jaco and Parrita scars, respectively. During both deployments, the backscatter images from the low resolution side scan sonar clearly illustrates two main sets of faults: **(1)** radial normal faults, probably related to the seafloor uplift due to the seamount subduction and **(2)** linear faults trending perpendicular to the direction of subduction (N130-N140) over several kilometers. An well expressed surface deformation, mainly characterised by elongated short wavelength folding has been observed

on both structures and could be related to the combination of compressive stresses in front of the subducting seamounts and strike-slip motion on the lateral walls.

The different purposes of these two deployments were to understand the evolution and chronology of faulting, but also to image some possible fluid escape structures related to seamount subduction. Some noticeable differences have been observed between both scars. The radial normal faults seem much more developed upslope of Jaco scar and many sigmoidal structures might reveal strike-slip motions. The Parrita scar is mainly controlled upslope by folding and linear N140 faults. A large number of small scale highly reflective patches have been observed in both cases and could be attributed either to encrusted areas related to fluid escapes or to mud flows injected through the crosscutting faults which appears as the easiest way for fluid ascension. The compressive stresses at the front of the subducting seamounts would then provide the necessary overpressure to allow mud and fluid ascensions.

Several preliminary hypotheses could be proposed to explain the differences observed between both structures. (1) Parrita scar only 6 km upslope from the deformation front is a more recently subducted seamount covered by a thinner sedimentary sequence than Jaco scar. This could explain why Parrita scar shows less important uplift and brittle deformation. (2) The Parrita seamount could also be a smaller seamount more easily subducted. Both hypotheses also account for the presence of a larger amount of reflective patches around Jaco scar, where compressive tectonics seem to be more important. A further detailed analysis and interpretation of the side scan sonar mosaics and subbottom profiler lines should allow to better understand the tectonic mechanisms and associated fluid venting related to seamount subduction, as well as the evolution through time of these processes.

6.4 Results and discussion of OFOS and TVG deployments

(H. Sahling, T. Nadler, D. Masson, V. Hühnerbach, C. Ranero)

A major target of this cruise was to discover and sample seafloor areas that are influenced by venting of methane-rich fluids. It is generally accepted that the microbial-mediated process of anaerobic methane oxidation occurs, when these fluids come in contact with sulfate-rich waters at shallow sediment depths. This biogeochemical process causes, firstly, the precipitation of (chemoherm) carbonate and, secondly, the production of sulfide, that is then used by chemoautotrophic organisms. Carbonates and chemoautotrophic organisms are, thus, indicators of methane venting. One goal during this cruise was to collect these (geological and biological) indicators in order to confirm their methane-dependent origin.

It is generally thought that tectonic and mass wasting processes control fluid venting. Therefore, we concentrated our investigations on prominent geological structures; these are mounds (originally described as mud volcanoes), subducted seamounts and ridges, the accretionary wedge, and landslides. In addition, we surveyed a site on the oceanic plate that has revealed unusual high heatflux values and hydrothermally influenced porewater geochemistry.

A prerequisite for successful OFOS and TVG investigations are seafloor images based on multibeam echosoundings and sidescan sonar mapping. Wide areas were mapped during previous cruises or have been mapped during this cruise (compare SIMRAD and TOBI chapters). These images are shown in this chapter in order to draw a comprehensive picture of the seafloor processes.

During this cruise 26 OFOS deployments and 9 TVG deployments were successfully accomplished (Fig. 6.4.1, Fig. 6.4.2, Tab. 6.4.1).

6.4.1 Mounds

Circular structures tens of meters high that were originally described as mud volcanoes, occur along the Costa Rican continental margin (Shipley et al., 1992). At these structures luxuriant chemosynthetic communities as well as chemoherm-like carbonates were observed by DSS Alvin and OFOS (Bohrmann et al., in press; Kahn et al., 1996). During the first OFOS deployments it became evident that these structures are indeed fluid venting related: at the top or the flanks of the mounds we found massive chemoherm-like carbonates outcropping and, at some of these structures, vesicomid, mytilid or solemyid bivalves and pogonophoran tubeworms occur. However, there was no evidence for mud expulsion. Furthermore, at the top of some of these mounds, described in detail below, massive carbonate crop out, suggesting that these structures constitute partly, if not entirely, of carbonate. These structures may be mounds build up by chemoherm-like carbonates. This is further corroborated by the occurrence of chemosynthetic communities on top and on the flanks of the mounds, which indicate fluid flow at present.

Along the continental margin of Costa Rica several mounds were discovered using TOBI. Most mounds are characterised by high backscatter relative to the surrounding sediment covered slope. This can be attributed mainly to the occurrence of carbonate (or carbonate-cemented sediment?) although mound topography may also contribute to elevated backscatter levels on slopes facing the sidescan. Backscatter levels show some variability both within and between mounds, probably reflecting variation in sediment cover on the mound. Most mounds are readily identified in the bathymetric map. For clarity, numbers were given for each mound that were surveyed with OFOS during this cruise (Tab. 6.4.1.1, Fig. 6.4.1.1).

Mound # 1 (Mound Culebra) OFOS 1 & 2, TVG 1 & 2

Mound # 1 is located at 1600-m water depth and is recognisable in the bathymetry. It is a circular structure of about 0.8 nautical miles in diameter and rises about 100 m above the surrounding seafloor. The mound was surveyed by two TOBI lines, which showed high reflectivity throughout the mound (Fig. 6.4.1.2). Due to ocean currents and wind direction we were forced to tow OFOS from west to east (Fig. 6.4.1.3). To avoid going up the steep western slope we deployed both OFOS close to the top and surveyed the highest parts and the flanks on the E and NE side. "Bumps" about 30 m high largely structure the top of the mud volcano (Fig. 6.4.1.4); these can also be seen on the TOBI image. On top of these "bumps" massive carbonate is exposed. Carbonate also outcrops at the flanks suggesting that the whole structure is cemented. We found no evidence

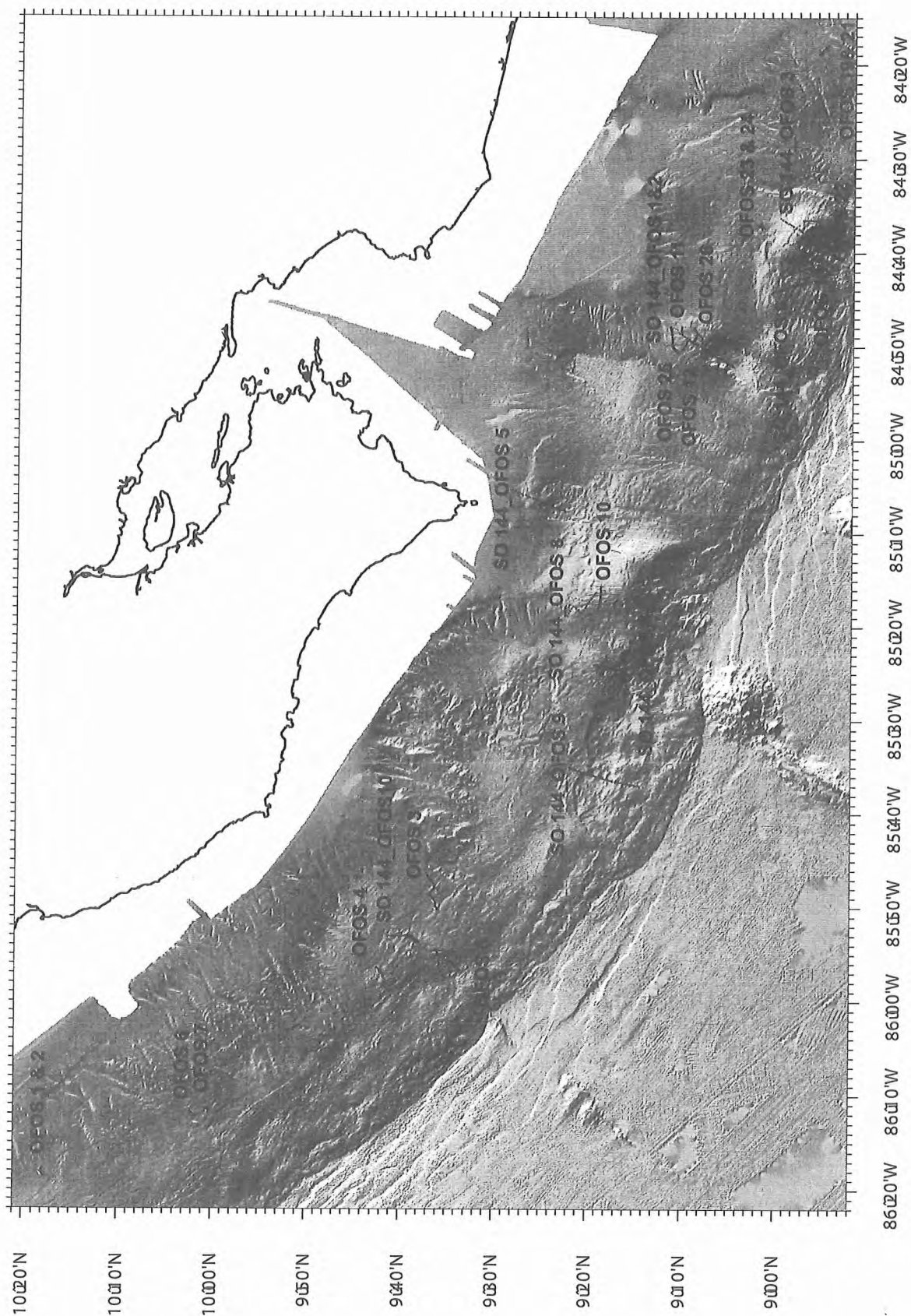


Figure 6.4.1: Overview of all OFOS deployments during SO 163 and SO 144 on the Western continental margin off Costa Rica.

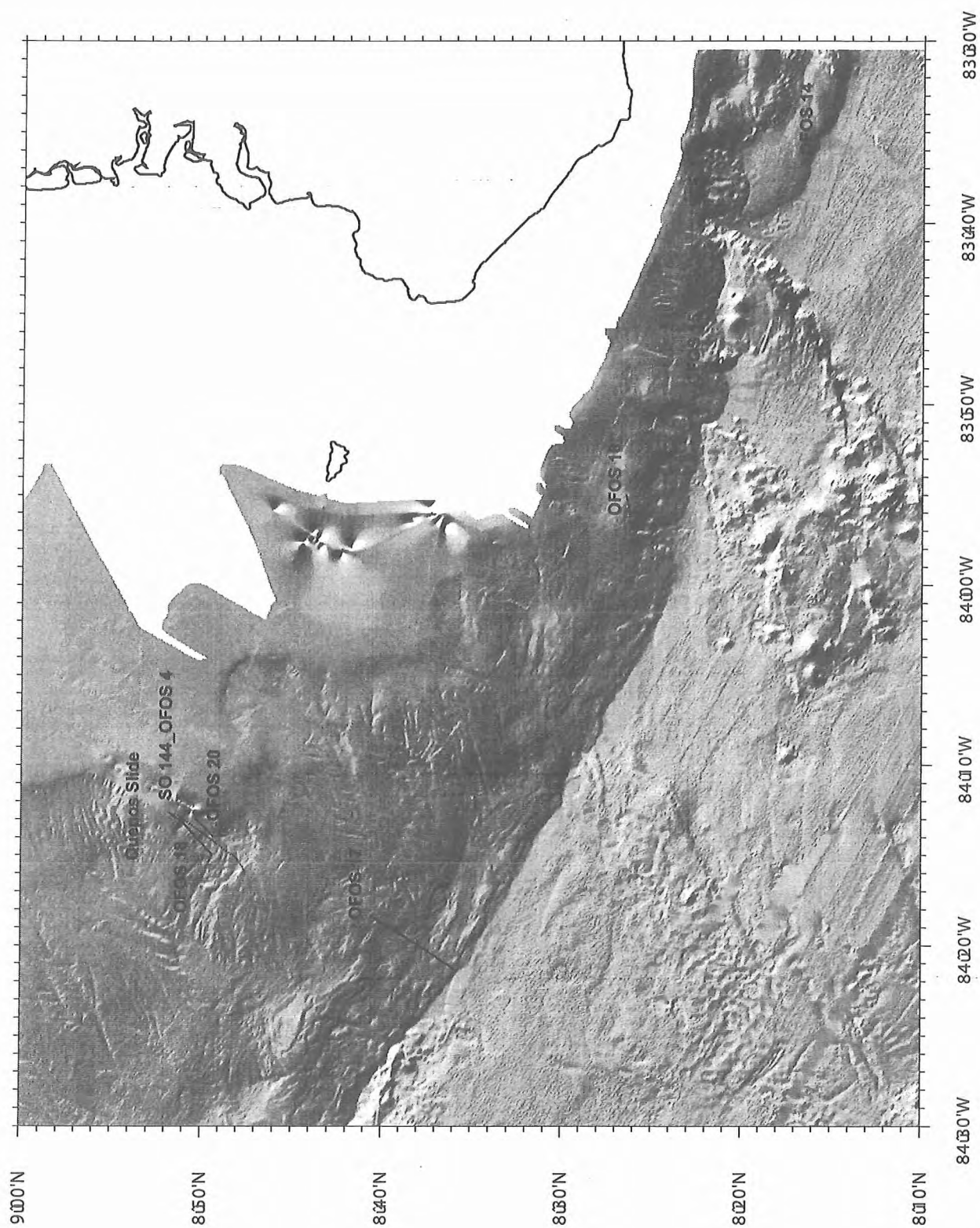


Figure 6.4.2: Overview of OFOS deployments during SO 163 on the Eastern Costa Rican margin.

Tab. 6.4.1:

Summary of OFOS and TVG deployments during SO 163-1.

Station	Area	Date	Time	at seafloor			off seafloor / grab location			Observation / recovery	
				Lat	Long	Depth	Time	Lat	Long		Depth
Mounds											
OFOS 1	Mound # 1	22.03.2002	17:41:48	10° 17.86	86° 18.37	1540	19:41:49	10° 18.01	86° 17.99	1570	Carbonates, clams and pogonophorans, No SSBL Position
OFOS 2	Mound # 1	22.03.2002	21:29:10	10° 17.84	86° 18.56	1510	23:21:10	10° 18.14	86° 18.07	1550	Carbonates, clams and pogonophorans, SSBL 50 m above OFOS
TVG 1	Mound # 1	23.03.2002	01:59:50	10° 17.88	86° 18.40	1500	02:07:00	10° 17.923	86° 18.32	1516	2 carbonate pieces on top of grab
TVG 2	Mound # 1	23.03.2002	03:26:00	10° 17.93	86° 18.33	1490	03:36:00	10° 17.987	86° 18.29	1473	Several carbonate pieces and clamshell fragments
OFOS 3	Mound # 2-6	28.03.2002	00:10:30	9° 37.12	85° 46.53	1588	08:25:26	9° 35.95	85° 50.15	1989	Crossed 4 mounts from NE to SW, some with carbonates, pogonophorans, clams; Camera B problems
OFOS 4	Mound # 7, 8	18.03.2002	11:32:36	9° 42.53	85° 56.30	2203	15:17:14	9° 41.98	85° 56.90	2240	Small aera with carbonates, pogonophorans and clams
OFOS 6	Mound # 9	29.03.2002	05:01:30	10° 02.16	86° 11.80	2240	07:24:10	10° 01.56	86° 12.23	2342	No seep indication
OFOS 7	Mound # 10	29.03.2002	09:24:10	10° 00.73	86° 11.35	2305	11:36:10	10° 00.65	86° 11.50	2341	Abundant vent indication; clams, pogonophorans, carbonate
OFOS 13	"Bright Spot W of Parrita"	05.04.2002	20:19:50	9° 59.70	84° 42.50		23:35:00	8° 58.00	84° 41.65		Beside scattered clamshells no seep indications
OFOS 19	Mound # 11, 12	14.04.2002	05:44:20	8° 54.94	84° 17.91	1023	08:14:32	8° 55.95	84° 18.83	1005	Extensive carbonates, clams, mussels and pogonophorans
OFOS 21	Mound # 12	14.04.2002	15:09:30	8° 55.94	84° 18.53	998	18:40:16	8° 55.73	84° 18.95	1013	Extensive carbonates, clams, mussels and pogonophorans; Video on 16:32
TVG 4	Mound # 12	14.04.2002	22:29:35	8° 55.79	84° 18.74	1002	23:14:45	8° 55.81	84° 18.62	992	Piece of outcropping carbonate
TVG 5	Mound # 12	15.04.2002	00:16:50	8° 55.79	84° 18.75	1002	00:41:00	8° 55.80	84° 18.74	1005	Complete seep ecosystem: Acharax, Bathymodiolus, Lamellibrachia, galatheid crabs and carbonates
OFOS 23	Mound # 13	17.04.2002	05:12:40	9° 02.20	84° 37.24	1451	12:28:39	9° 01.78	84° 37.06	1403	Most extensive field of carbonates, pogonophorans, bivalves
OFOS 24	Mound # 13	17.04.2002	14:27:44	9° 01.95	84° 37.22	1406	16:41:50	9° 01.95	84° 37.13	1397	Mapping most extensive seep area
TVG 8	Mound # 13	18.04.2002	16:55:10	9° 01.96	84° 37.22	1408	17:06:40	9° 01.97	84° 37.22	1408	Carbonates
TVG 9	Mound # 13	18.04.2002	18:44:15	9° 01.96	84° 37.25	1408	18:47:35	9° 01.96	84° 37.24	1408	Vent fauna, carbonates

Station	Area	Date	Time	at seafloor			off seafloor / grab location			Observation / recovery	
				Lat	Long	Depth	Time	Lat	Long		Depth
High heatflow area											
OFOS 8	Basement high	30.03.2002	04:43:17	9° 40.25	86° 33.76	3275	08:20:00	9° 41.01	86° 33.17	3267	No indication for hydrothermal venting
OFOS 9	Fault	30.03.2002	10:58:04	9° 46.48	86° 40.78	3315	13:40:51	9° 46.89	86° 40.41	3346	No indication for hydrothermal venting
seamount scars											
OFOS 5	Scarp of fracture ridge trace	28.03.2002	17:42:39	9° 36.31	85° 54.28	2194	23:00:40	9° 33.94	85° 56.53	3145	very few seep indication or other geological structures
OFOS 10	Rio Bongo Scarp	02.04.2002	18:39:39	9° 18.86	85° 15.54	470	23:27:30	9° 19.16	85° 18.31	772	Carbonated top, few vent fauna
OFOS 11	Jaco Scarp	03.04.2002	04:52:39	9° 10.43	84° 47.75	731	07:16:42	9° 10.40	84° 48.66	779	Carbonates, bivalve shells
OFOS 12	Jaco Scarp	03.04.2002	09:03:40	9° 10.40	84° 50.34	928	13:15:33	9° 08.74	84° 48.69	865	Chemoherm-like carbonates, vesicomyid clamfield
TVG 3	Jaco Scarp	05.04.2002	15:44:10	9° 10.42	84° 48.25	763	15:50:57	9° 10.43	84° 48.24	763	Chemoherm-like carbonates, living solemyid bivalves, bivalve shells
OFOS 25	Jaco Scarp	18.04.2002	00:09:48	9° 11.60	84° 47.80	774	05:03:20	9° 10.78	84° 50.62	996	No vent indication
OFOS 26	Jaco Scarp	18.04.2002	06:22:55	9° 08.24	84° 49.64	953	09:08:31	9° 09.14	84° 48.04	810	No vent indication
TVG 7	Jaco Scarp	18.04.2002	14:01:32	9° 09.03	84° 49.18	834	14:32:00	9° 09.01	84° 49.13	824	Carbonates
OFOS 15	"Frontal collision"	11.04.2002	00:02:00	8° 22.20	83° 47.684	1260	02:52:24	8° 21.182	83° 48.013	1633	Steep morphology, beside clamshells no seep indication
OFOS 16	"Small seamount scarp"	11.04.2002	05:05:00	8° 26.48	83° 54.981	701	07:48:14	8° 25.893	83° 56.002	1095	No vent indication
OFOS 22	Parrita	15.04.2002	09:15:58	8° 57.55	84° 36.51	1228	14:12:57	8° 55.17	84° 37.89	2338	Beside scattered clamshells no seep indications

Station	Area	Date	Time	at seafloor			off seafloor / grab location				
				Lat	Long	Depth	Time	Lat	Long	Depth	Observation / recovery
Accretionary wedge											
OFOS 14	"Crest of Cocos Ridge"	10.04.2002	19:21:24	8° 15.572	83° 34.185	1710	20:41:58	8° 15.343	83° 34.803	1902	No seep indication
OFOS 17	"Seismic Line SO 81 #5"	11.04.2002	12:34:30	8° 40.56	84° 18.35	2000	20:57:25	8° 35.93	84° 21.30	2500	No seep indication
Landslide											
OFOS 18	Quepos Landslide	13.04.2002	22:56:41	8° 51.87	84° 12.60	179	03:48:43	8° 49.50	84° 15.51	831	Larg bacterial mats at headwall, few seep indications in slide mass
OFOS 20	Quepos Landslide	14.04.2002	10:16:00	8° 50.30	84° 13.00	438	13:06:00	8° 48.61	84° 13.90	740	Beside one bacterial mat no seep indication at slide mass

Mound No.	Latitude	Longitude	Depth	DTS	SO 163	Results
1	10°17.90 N	86°18.40 W	1600	--	OFOS 1, 2; TVG 1, 2	carbonates, clams and pogonophorans
2	9°36.50 N	85°47.40 W	1760	75 kHz	OFOS 3	sediment covered
3	9°36.40 N	85°47.80 W	1760	75 kHz	OFOS 3	carbonates, clamshells
4	9°36.80 N	85°48.60 W	1800	75 kHz	OFOS 3	carbonates, clamshells
5	9°36.30 N	85°49.30 W	1800	75 kHz	OFOS 3	few carbonates, largely sediment covered
6	9°36.10 N	85°50.00 W	1950	75 kHz	OFOS 3	sediment covered
7	9°42.20 N	85°56.60 W	2200	75 kHz	OFOS 4	no seep indication
8	9°41.90 N	85°57.10 W	2200	75 kHz	OFOS 4	small area with carbonates, pogonophorans and clams
9	10°02.00 N	86°12.00 W	2300	75 kHz	OFOS 6	no seep indication
10	10°00.40 N	86°11.50 W	2300	75 kHz	OFOS 7	abundant vent indication; clams, pogonophorans, carbonate
11	8°55.33 N	84°18.26 W	1000	--	OFOS 19	carbonates
12	8°55.80 N	84°18.70 W	980	--	OFOS 19, 21; TVG 4, 5	extensive field of carbonates, clams, mussels and pogonophorans
13	9°01.90 N	84°37.23 W	1400	410 kHz	OFOS 23, 24	most extensive field of carbonates, pogonophorans, bivalves

Tab. 6.4.1.1: List of mounds surveyed during SO 163 cruise along the Cost Rican continental margin.

of mudflows. At the top of the mound chemoautotrophic species, vesicomid clams and pogonophoran tubeworms indicate that venting still occurs. We deployed two TVG in order to sample the carbonates. Both grabs felt over during the closing of the jaws. However, we retrieved some pieces of rocks, sediment with sulfide odour and shell fragments. The calcareous cemented sediments are probably formed in the course of methane venting.

Mound # 2-6 OFOS 3

With OFOS 3 we surveyed fault-like structures and 5 mounds that were imaged by TOBI (Fig. 6.4.1.5). The OFOS survey revealed that the spotted backscatter on the slope between the beginning and a (OFOS-CTD) depth of 1750 m correspond to hardground that is hardly covered by soft sediment. At a depth of 1730 m big slabs of several m indicate tectonic activity that is probably linked to a fault imaged by TOBI. Mound # 2 was completely sediment covered, and only a few single clamshells were observed. The track crossed Mound # 3 and followed the southern slope, crossed the eastern flank of Mound # 4, where the direction of the sled turned so that it crossed, then, the top of Mound # 4 (Fig. 6.4.1.6, Fig. 6.4.1.7). At the summits of both mounds as well as the steep S flanks massive carbonate crops out. At the top of Mound # 3 the carbonate appears chemoherm-like and indicates venting activity. Additionally, clamshells were observed at the top that further confirms venting processes at present. The origin of the outcrops on the flanks is unknown, although we speculate that the entire mounds are made of carbonate originating from methane venting. In the following we have crossed the top area of Mound # 5 twice. This mound as well as Mound # 6 is largely sediment covered. The area Northeast of Mound # 6 is characterised by lower backscatter than the surrounding seafloor. The OFOS depth profile indicates an apparently flat seafloor (area 8' - 9'), suggesting a sediment filled depression.

In general, methane venting in this area appears to be largely inactive. The massive carbonate outcrops indicate considerable methane venting in the past.

Future investigation: To gain better understanding of the origin of carbonate we suggest dredge deployments in this area.

Additional note: Around 6:51 h (UTC) OFOS passed through a cloud of benthic-pelagic animals of very high densities. These may have been juvenile crustaceans.

Mound # 7, 8 OFOS 4

OFOS 4 covered two mounds that are largely sediment covered (Fig. 6.4.1.8, Fig. 6.4.1.9, Fig. 6.4.1.10). On top of Mound # 8 several boulders of unweathered chemoherm carbonates, a few bushes of vestimentiferan tubeworms and vesicomid clam clusters are present in a very small area, which we crossed twice with OFOS. Due to the small extension of venting this area may not be a high priority target for future investigations.

Mound # 9 OFOS 6

This mound was completely sediment covered and does not show any signs of carbonates, or biological communities associated with fluid venting (Fig. 6.4.1.11, Fig. 6.4.1.12, Fig. 6.4.1.13).

Additional note: A dense aggregation of several dozens large spider crabs occurred at the beginning of the track.

Mound # 10 OFOS 7 and OFOS 6

This mound is associated with crosscutting, fault patterns that are clearly seen upslope of the mound on the TOBI image. We found abundant vent indications in the top area and the on the higher northern flank. Here, unweathered chemoherm carbonate boulders, vestimentiferan tubeworms and vesicomid clam occur. Clamshells are scattered over wide areas but only a few clusters of living clams were seen. However, one clam cluster several meters wide and very long suggests, firstly, venting at present and, secondly, that faults control their distribution. Again, we found no evidence for mudflows. Due to the extensive vent areas this site should be a subject for future geochemical investigations.

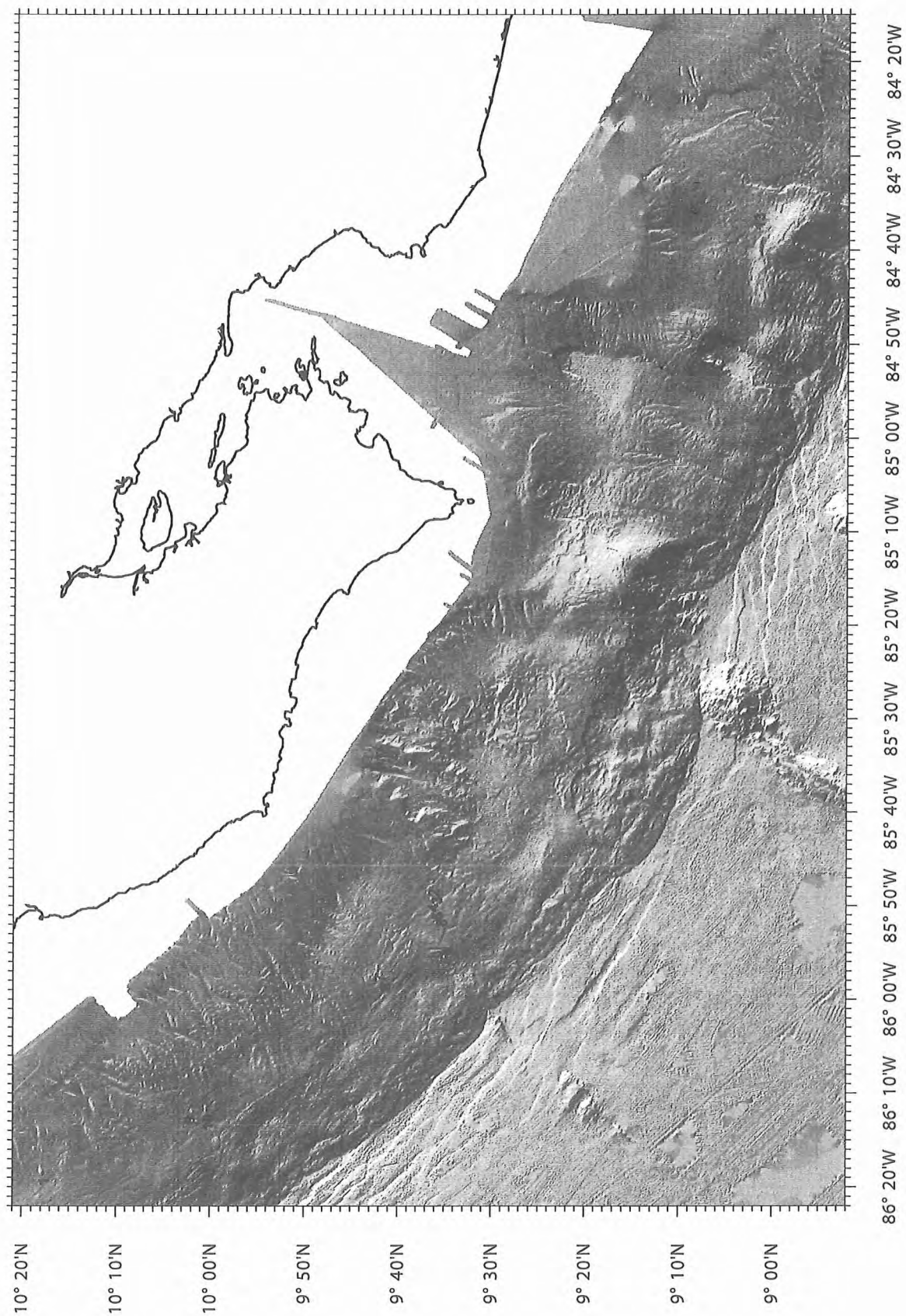


Figure 6.4.1.1: Location of mounds surveyed during SO 163 cruise along the continental margin off Costa Rica. The numbers indicate mounds listed in Tab. 6.4.1.1.

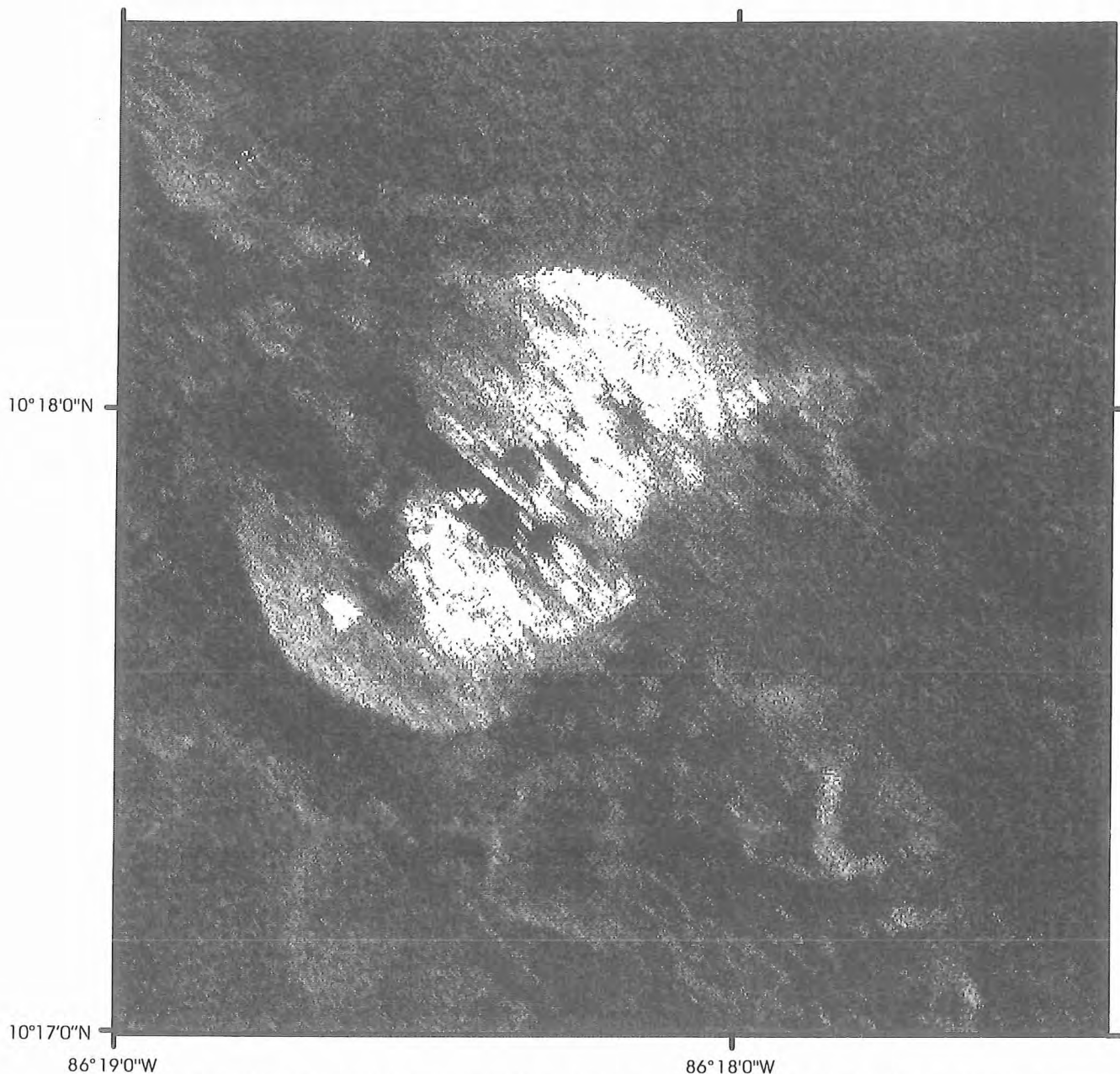


Figure 6.4.1.2: TOBI backscatter image of Mound # 1. White = high backscatter, black = shadows. Rough areas on the mound crest correlate with elevated carbonate outcrops seen on the OFOS transact.

Mound # 11, 12 OFOS 19, 21, TVG 4, 5

Mound # 11 and 12 were imaged by TOBI (see Fig. 6.2.21). With OFOS 19 we surveyed both mounds (Fig. 6.4.1.14). At Mound # 11 outcrops were seen, at Mound # 12 outcropping chemoherm-like carbonates and, in addition, abundant seep fauna. With OFOS 21 we surveyed on a very small scale the distribution of carbonates and organisms (Fig. 6.4.1.15). The abundance and diversity of pogonophorans and bivalves were high. At some sites the sediment between carbonates appeared blackish and very thin white bacterial mats occur. We found an area at which the mytilid bivalve *Bathymodiolus* occurs in cluster of up to about 5 meter in diameter. Although we do not know the relation between geochemical processes and the occurrence of seep organisms, we speculate that the high abundance of bivalves indicate the sites of most active methane seepage. With TVG 4 we cut a piece of outcropping carbonate, that has brecciated appearance. With TVG 5 we sampled a whole small ecosystem, that consisted of a living pogonophoran tubeworms (cf. *Lamellibrachia* sp.) and mytilid bivalves (*Bathymodiolus* sp.) in addition to chemoherm-like carbonates.

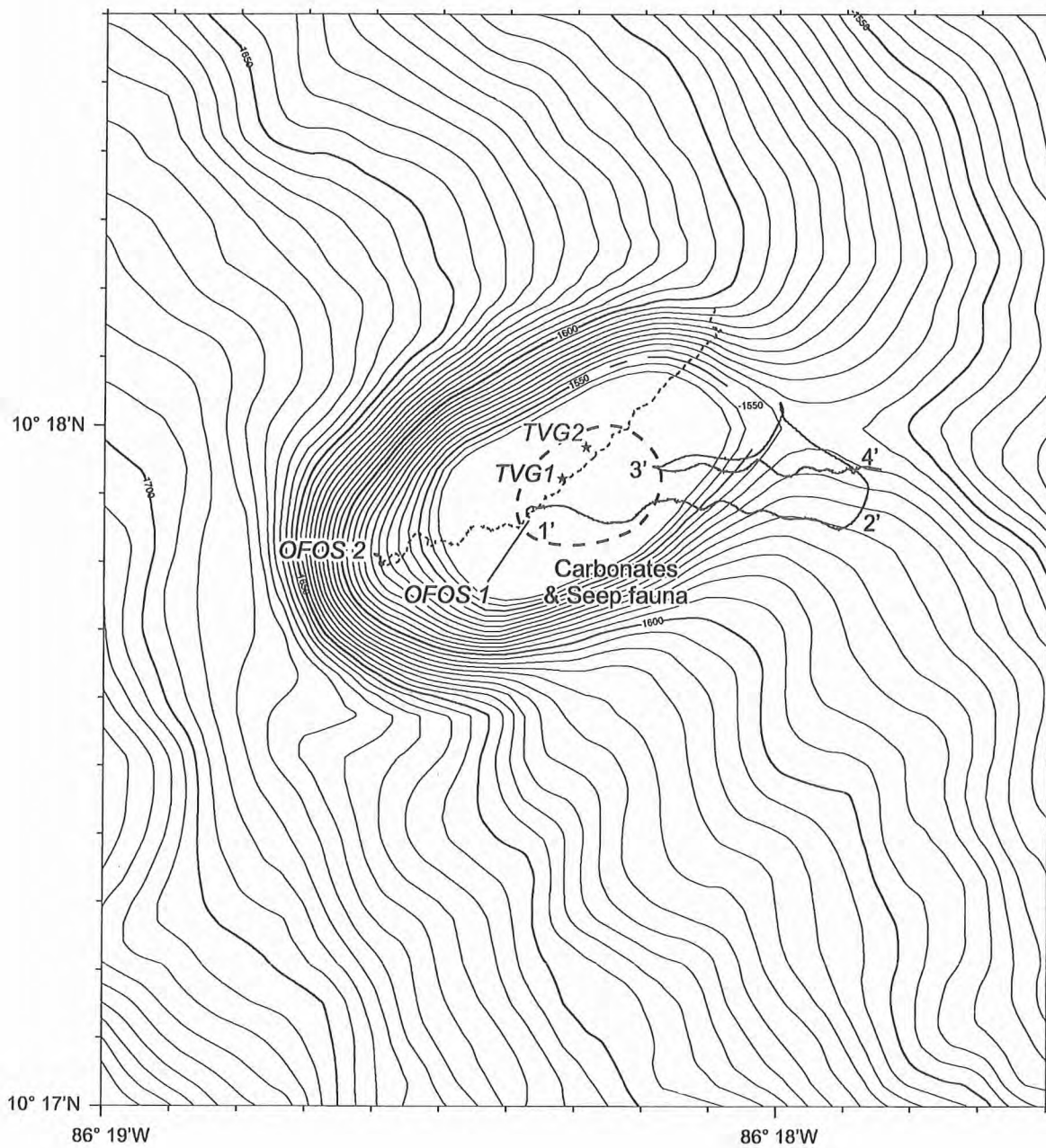


Figure 6.4.1.3: OFOS and TVG deployments at Mound # 1 (Mound Culebra). Tracks indicate OFOS 1, TVG 2 (ship position), OFOS 2, TVG 1 (SSBL positions). The bathymetry is based on Simrad multibeam obtained during this cruise. The position of major areas with outcropping carbonates and seep fauna is indicated.

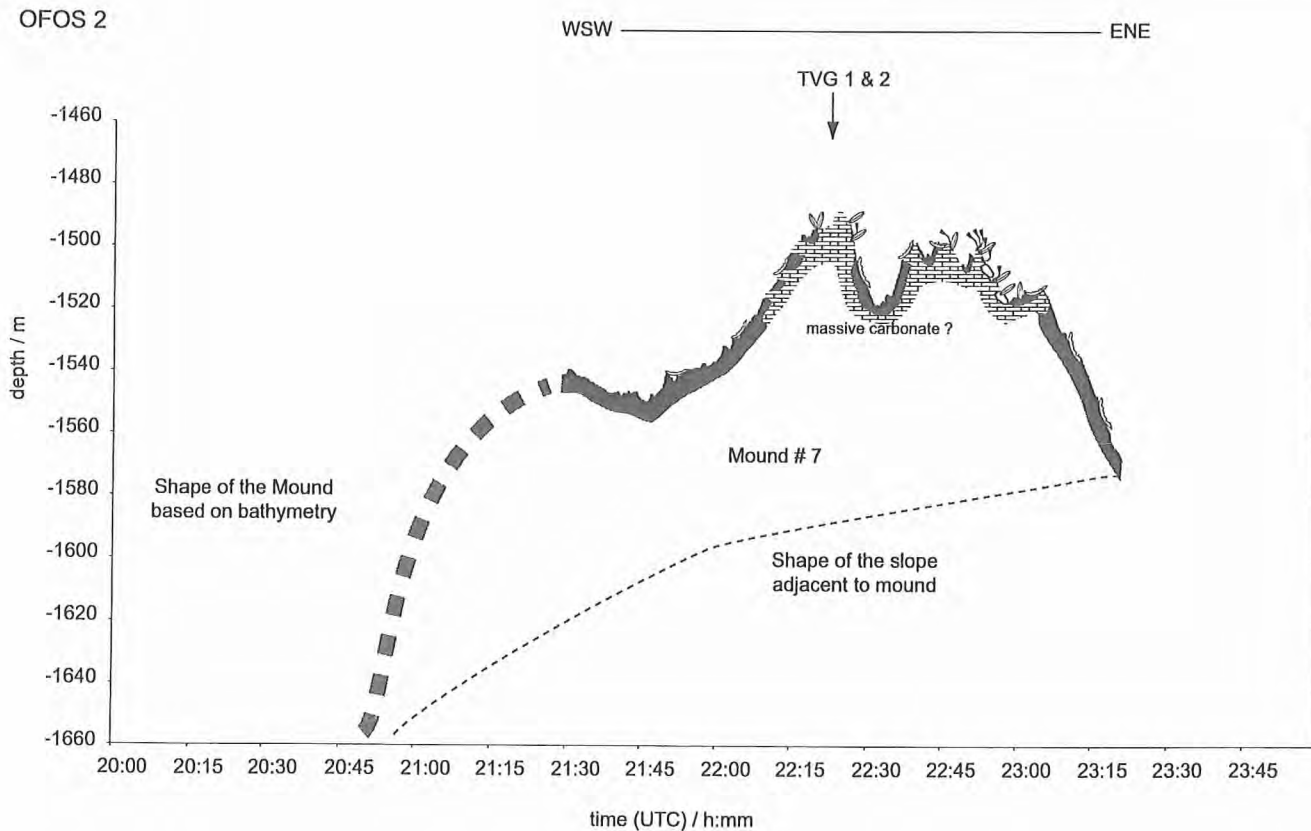
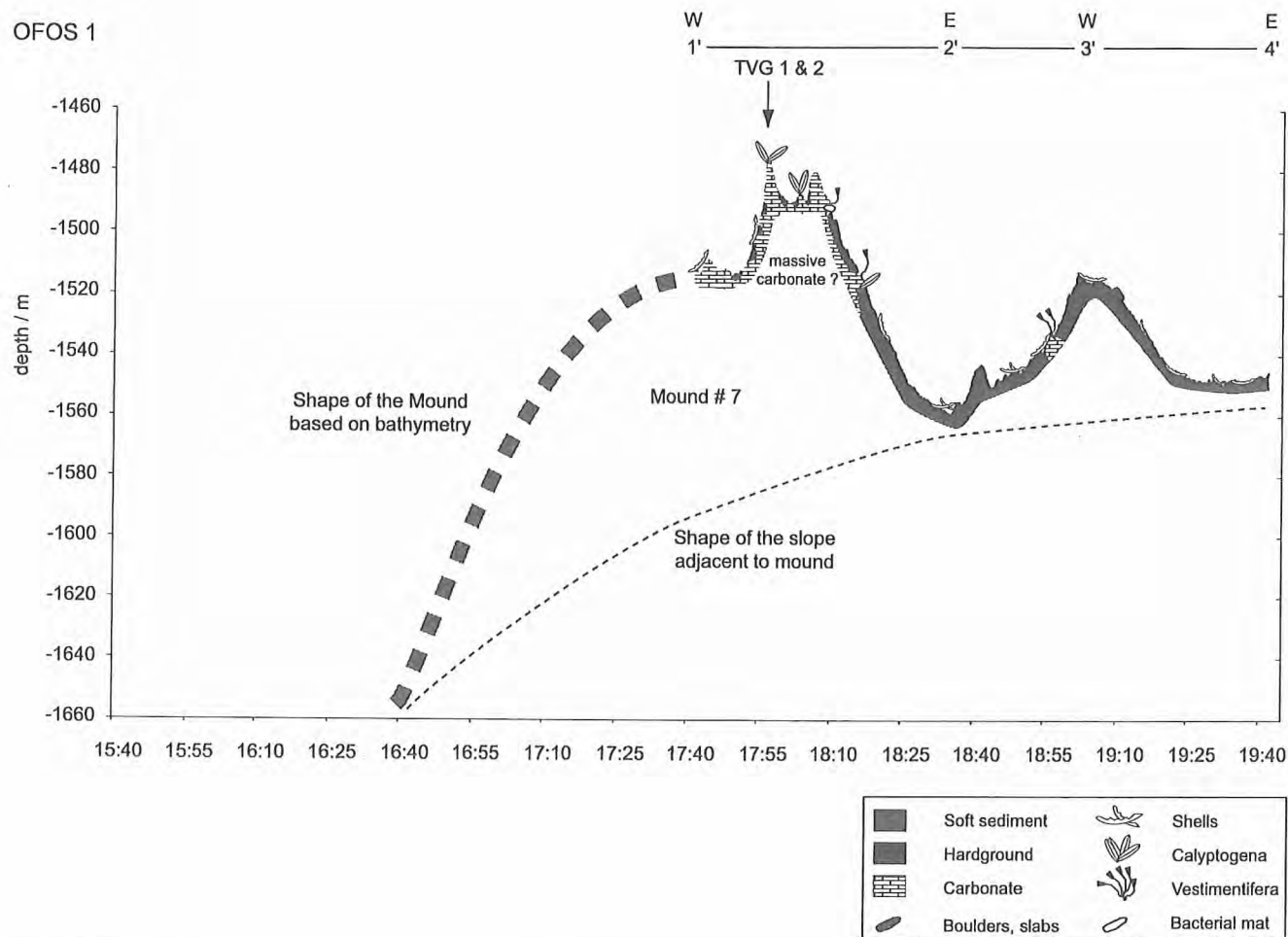


Figure 6.4.1.4: OFOS depth profiles at Mound # 1. The outcropping carbonates at the top and on the flanks of the mound suggest that the whole summit, if not the entire mound, consist of carbonate.

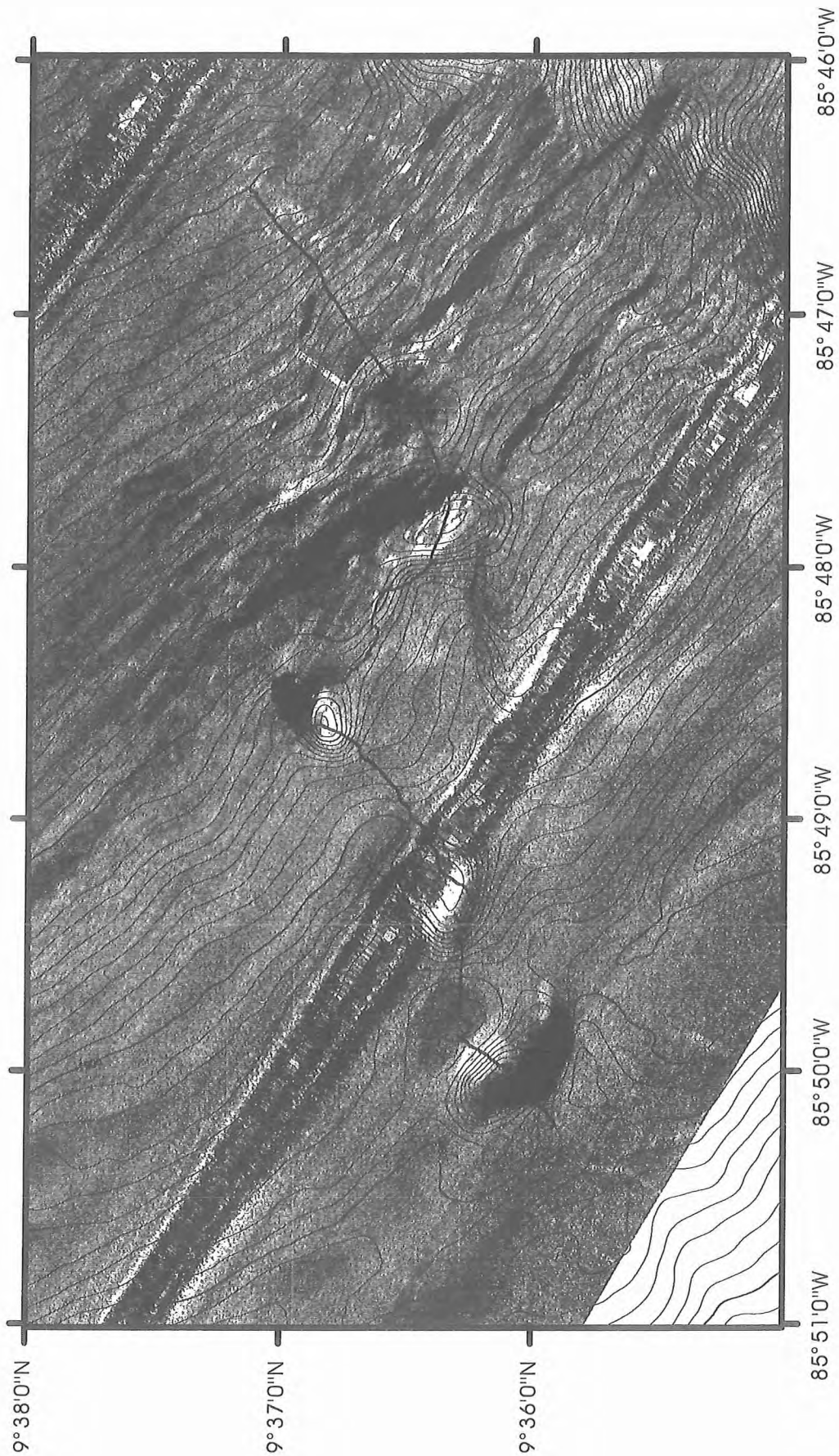


Figure 6.4.1.5: TOBI image with overlaid bathymetry and the track of OFOS 3 in the area of Mound # 2-6. This map is on the same scale as Figure 6.4.1.6 . High backscatter (white) is found at mounds. Linear high and low backscatter features to the east of the two central mounds are interpreted as outcropping faults.

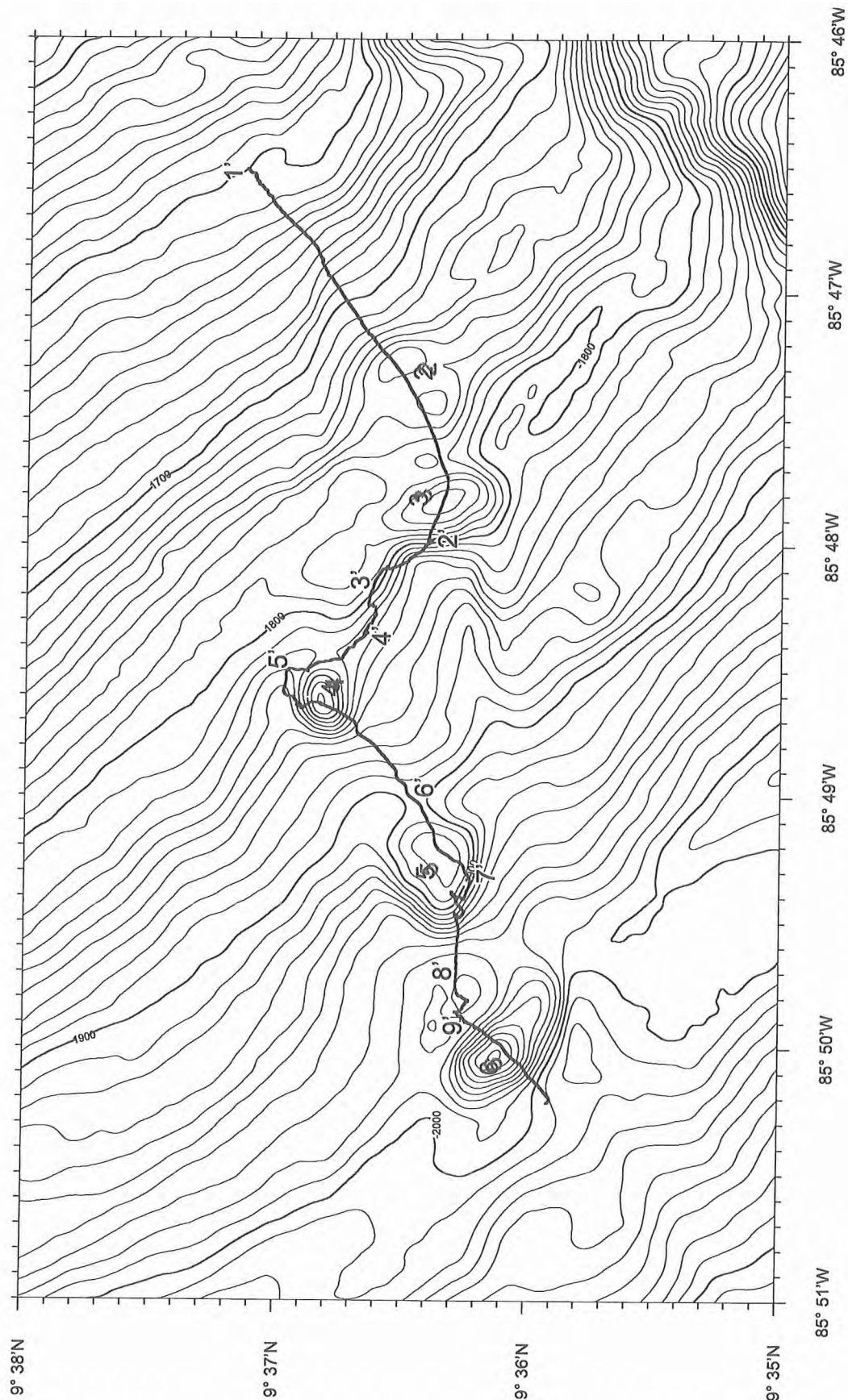
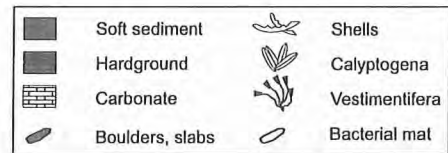
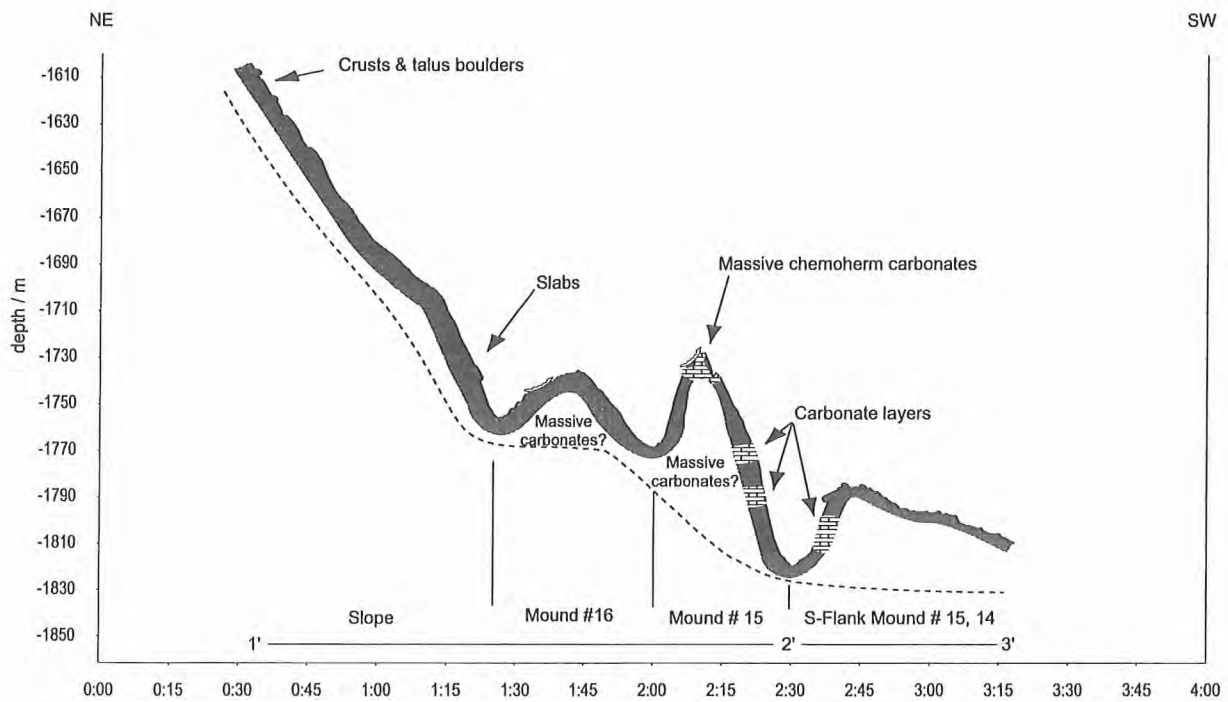


Figure 6.4.1.6: OFOS 3 at Mound # 2-6. The survey revealed that: (1) the spotted high backscatter at the beginning ('1') corresponds to highly consolidated or lithified sediments, (2) slabs probably indicate the existence of a fault east of Mound # 2, (3) the sloping flanks of the mounds cause the pronounced high and low backscatter values, (4) layered carbonates crop out at the south flanks of Mound # 3 and 4, (5) low backscatter east of Mound # 6 (8'-9') caused by a difference in sediment texture or thickness (possibly recent sediment deposition in the lee of the mound).

OFOS 3 - Profile 1



OFOS 3 - Profile 2

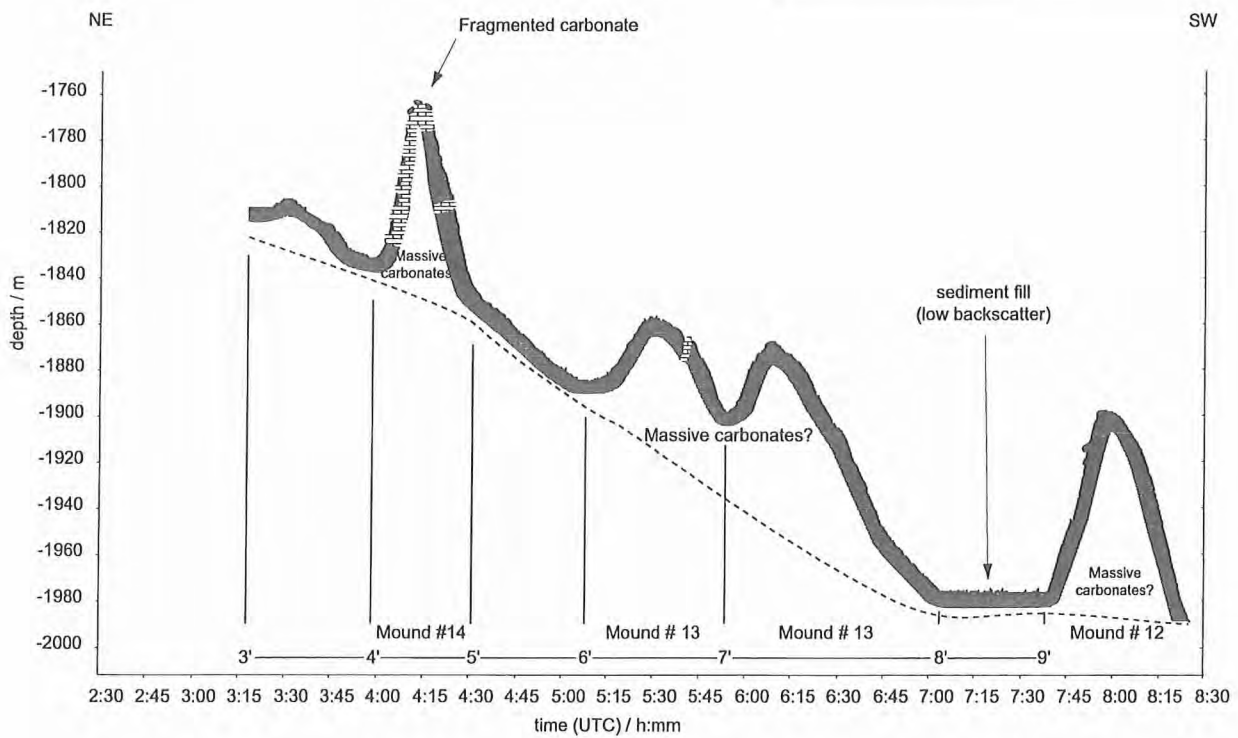


Figure 6.4.1.7: OFOS 3 depth profiles at Mound # 2-6. The profile shows the shape of the mounds, the dotted line indicates the profile of the adjacent "mound-free" slope.

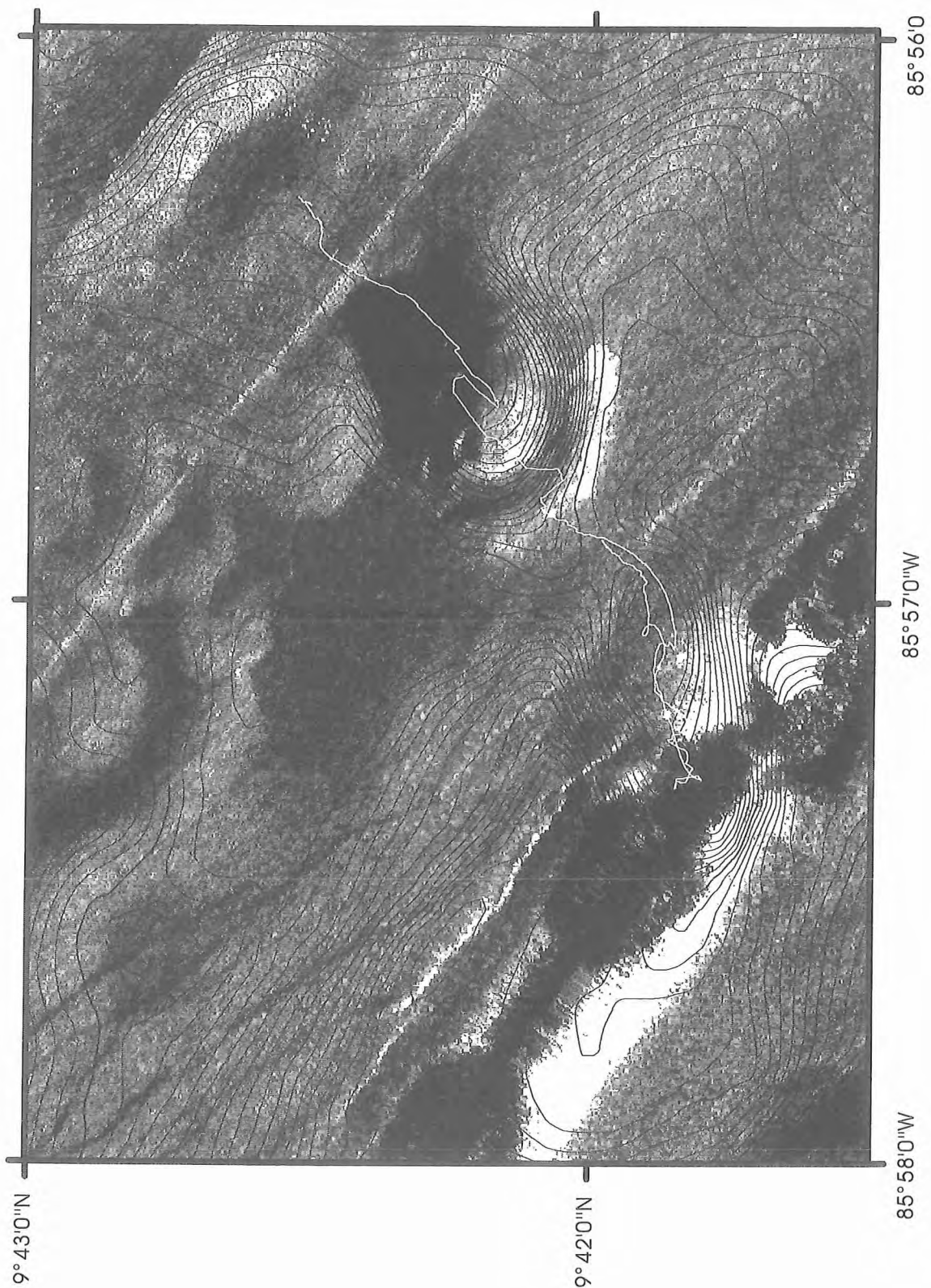


Figure 6.4.1.8: TOBI image showing the same area as Figure 6.4.1.9. Linear, high and low backscatter features are faults. Mound 8 was poorly imaged as it lies directly beneath the TOBI track, but nevertheless still shows high backscatter.

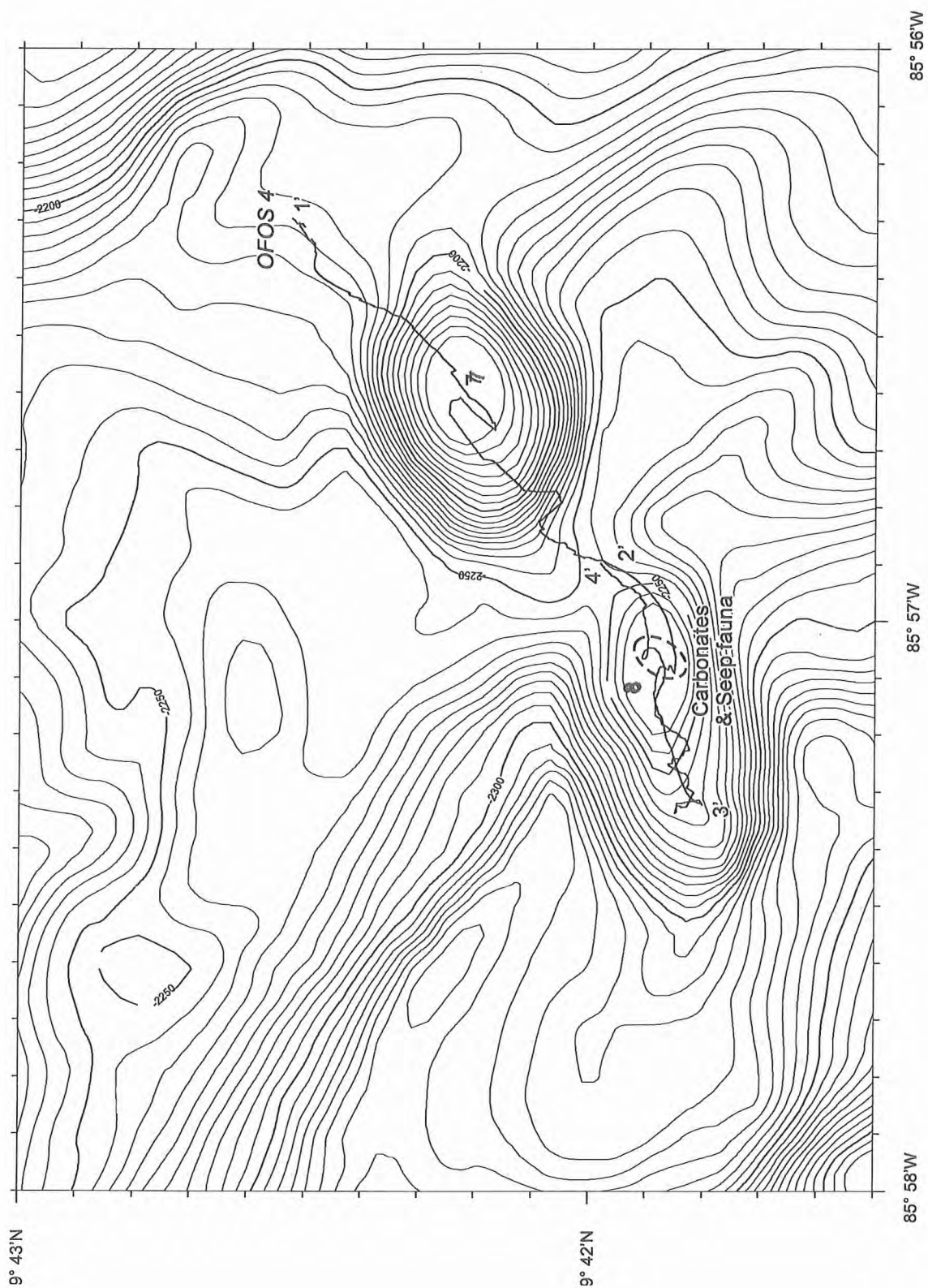


Figure 6.4.1.9: OFOS 4 surveys at Mound # 7 and 8. Chemoherm-like carbonates, vestimentiferan tubeworms and clams were seen on top of Mound # 8.

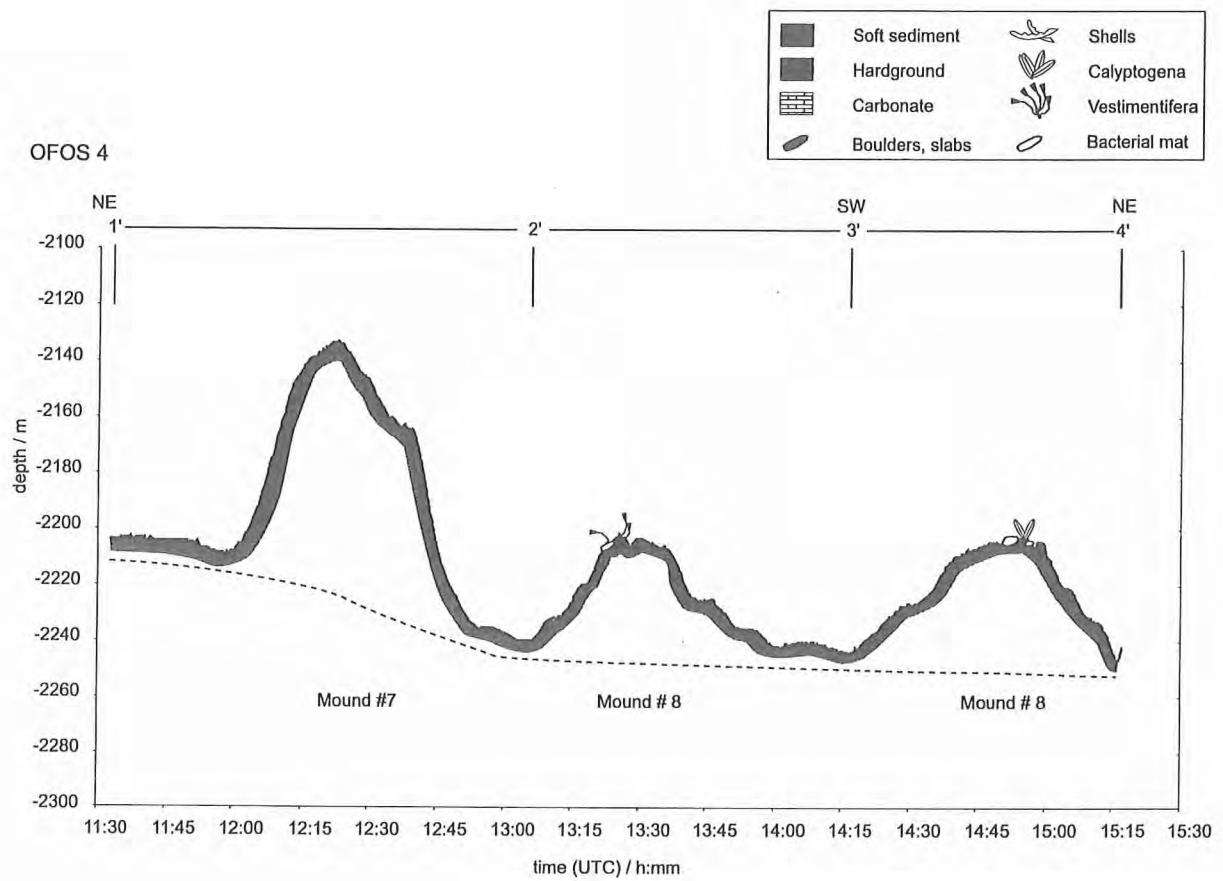


Figure 6.4.1.10: OFOS 4 depth profile.

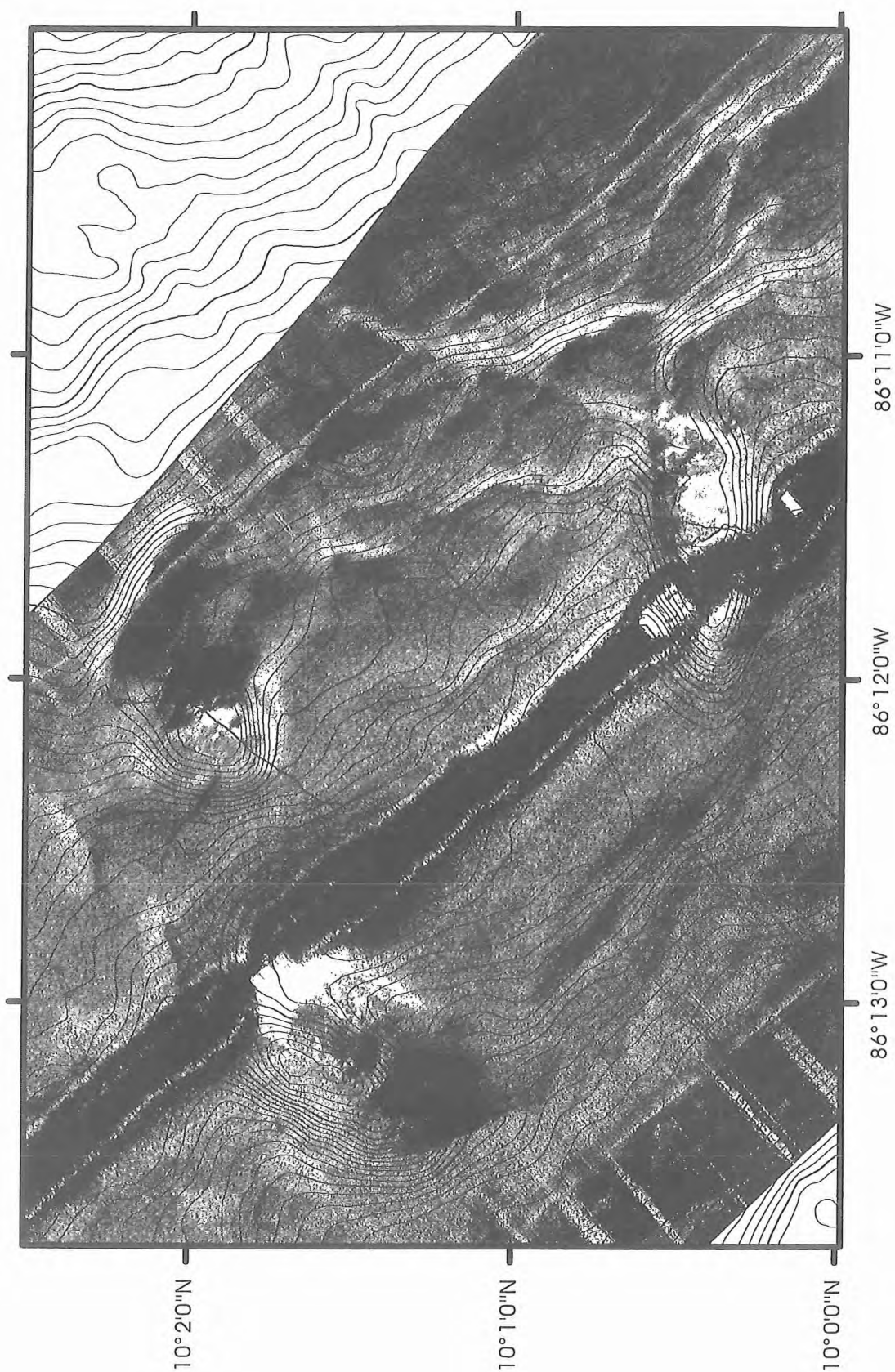


Figure 6.4.1.11: TOBI image showing the same area as Figure 6.4.1.12. A complex fault pattern is shown by the pattern of linear features (mainly high backscatter) crossing the image, especially to the northeast of mound 10.

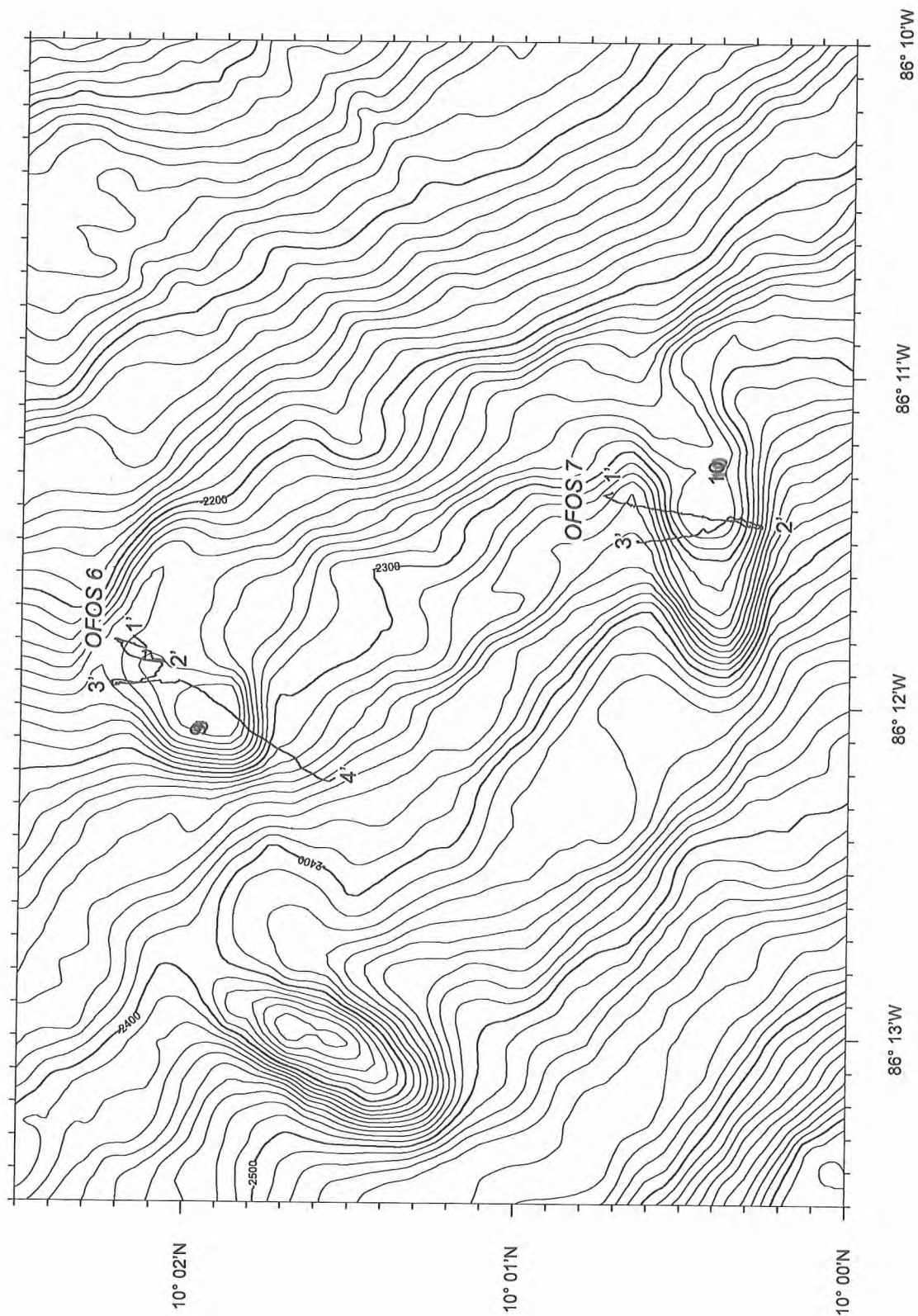


Figure 6.4.1.12: OFOS profiles at mound # 9 and 10. High backscatter characterises the mounds (Figure 6.4.1.11). OFOS 6 surveyed Mound # 9 but did not reveal any signs of venting, i.e. carbonates or chemoautotrophic fauna. Mound # 10 is associated with faults (OFOS 7). At the northern flank and the top of this mound vigorous venting was indicated by an extensive clamfield and chemoherm carbonates.

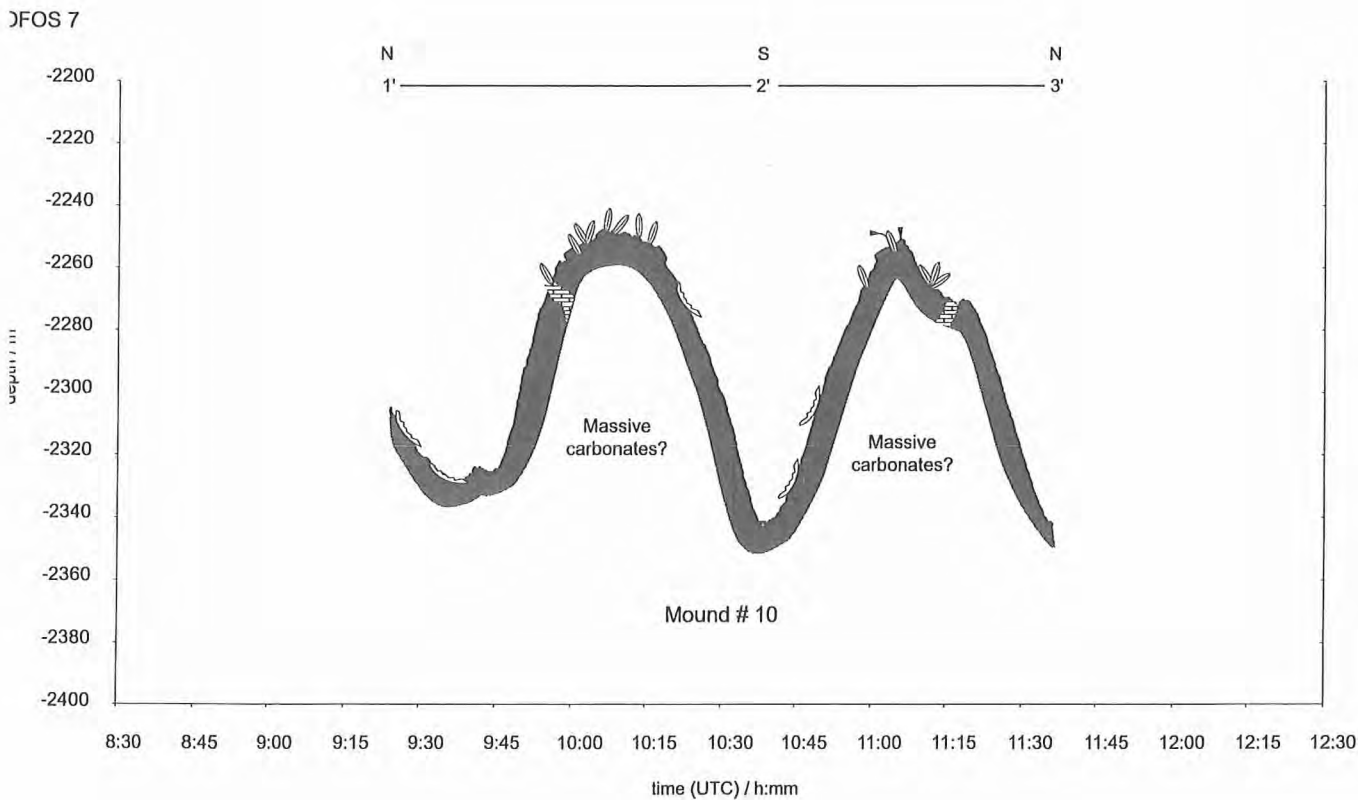
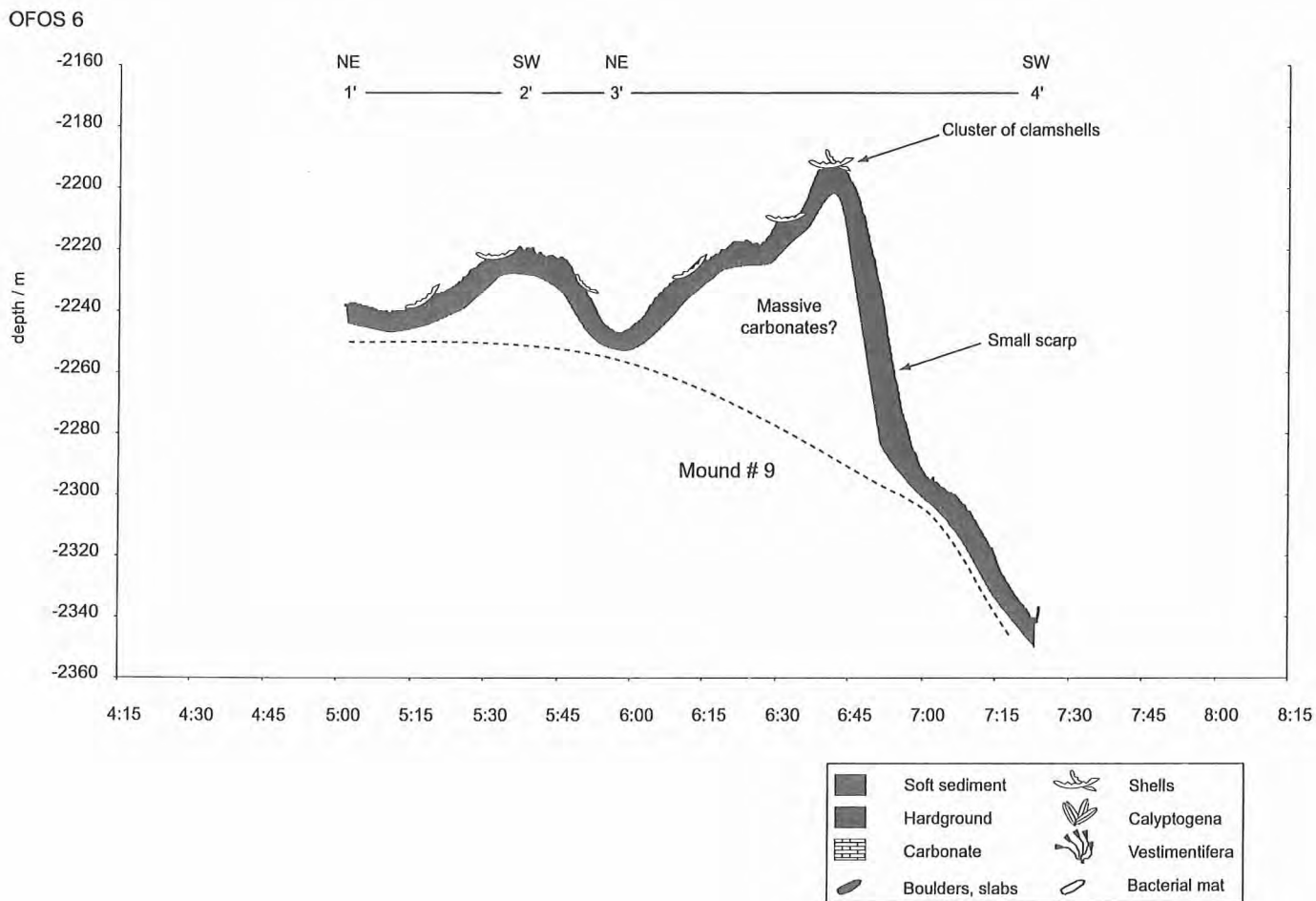


Figure 6.4.1.13: Depth profiles of OFOS 6 at mound # 9 and OFOS 7 at Mound # 10.

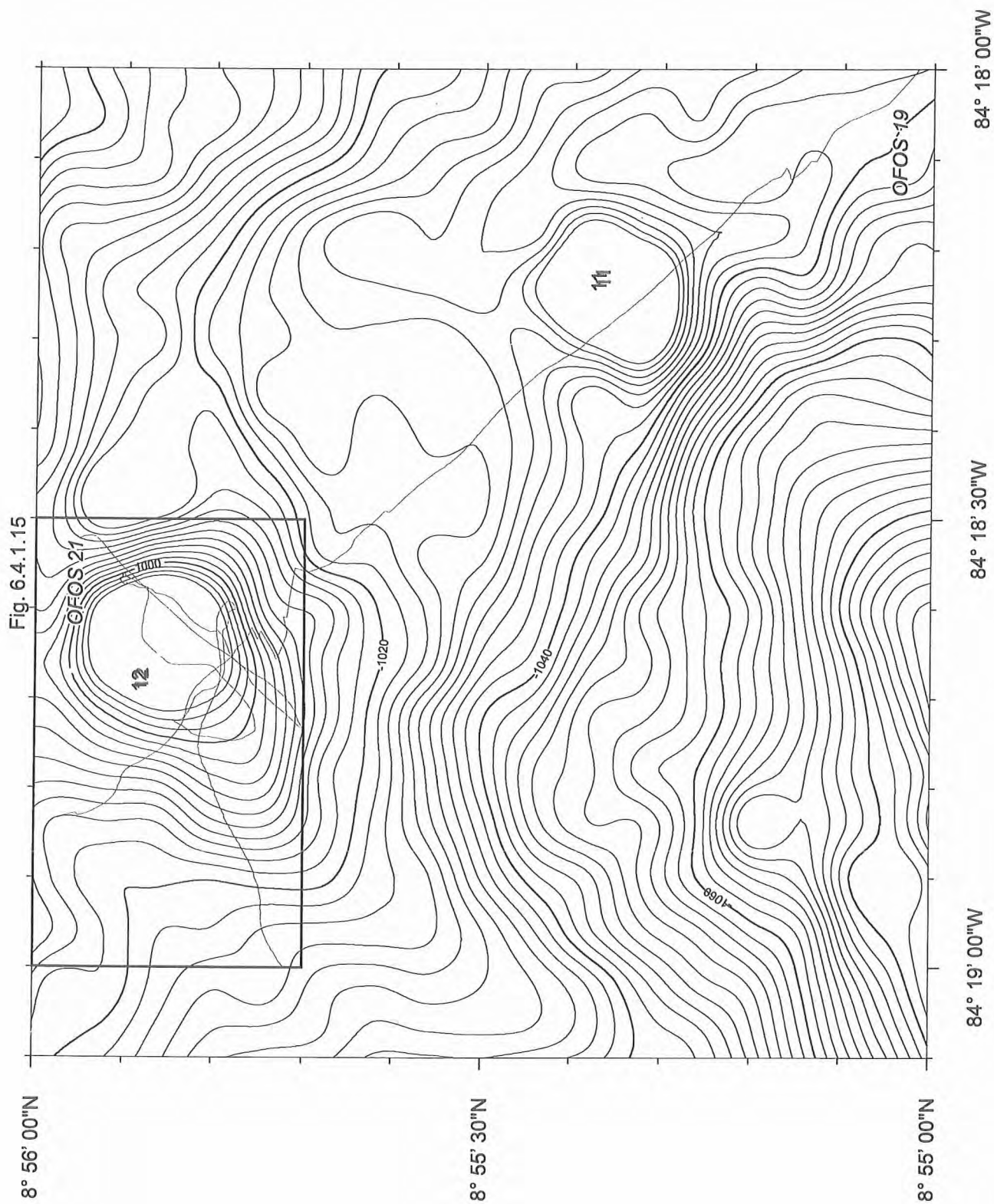
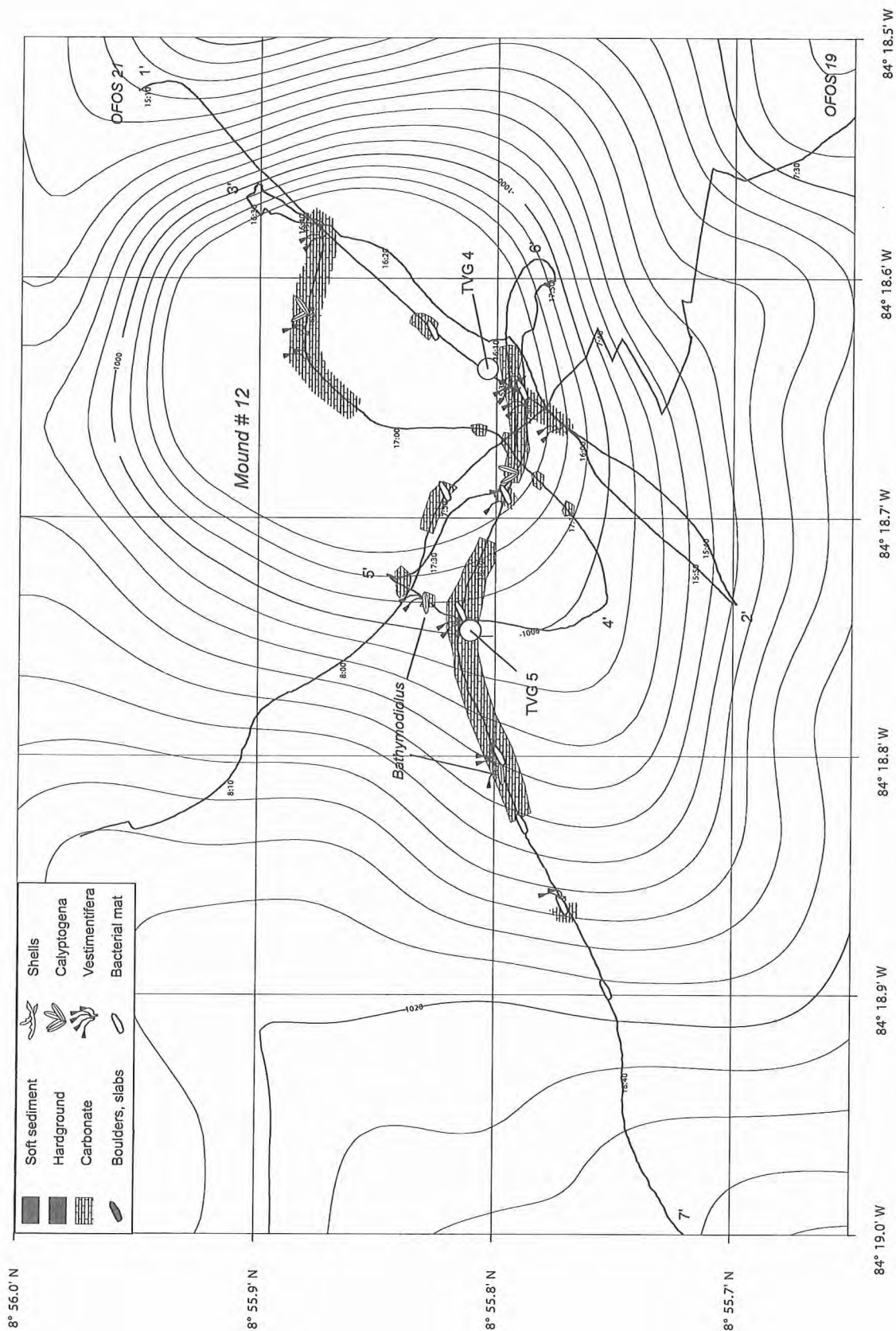
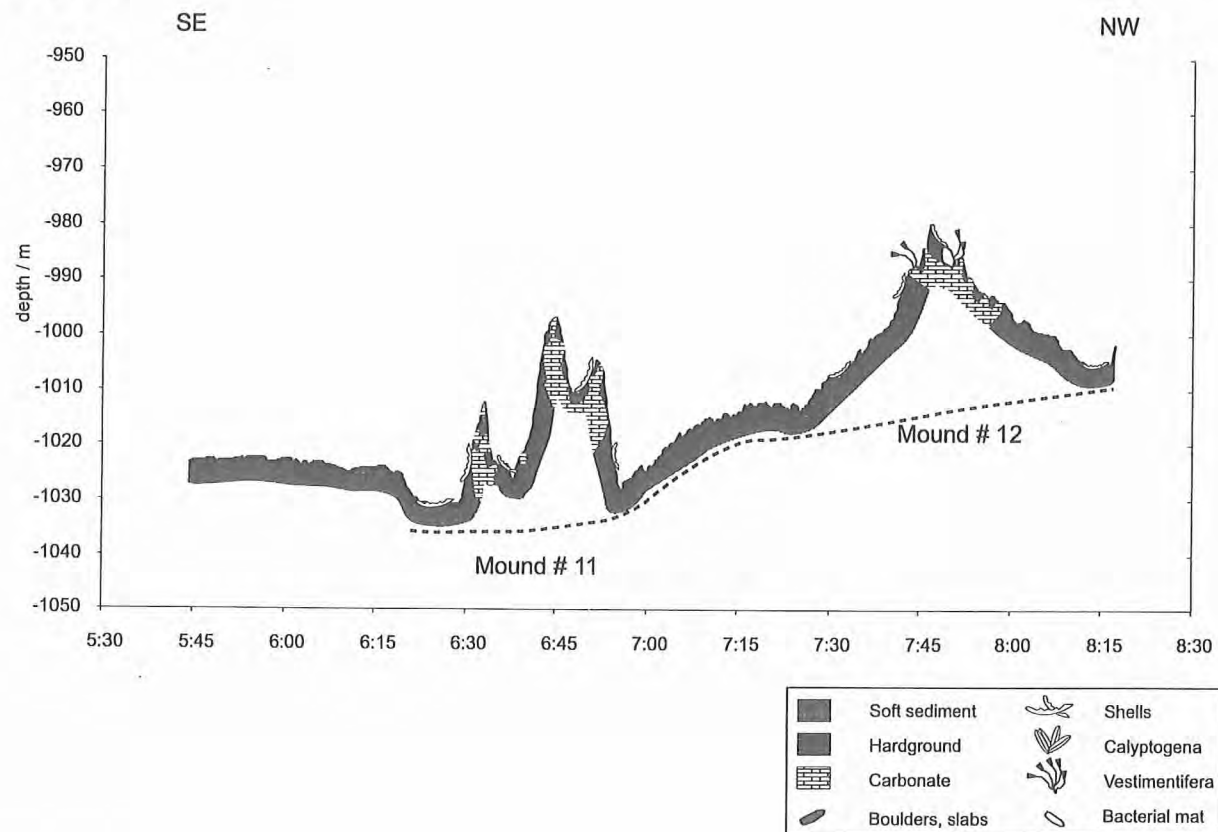


Figure 6.4.1.14: Location of OFOS 19 and 21 at Mound # 11 and 12.



OFOS 19



OFOS 21

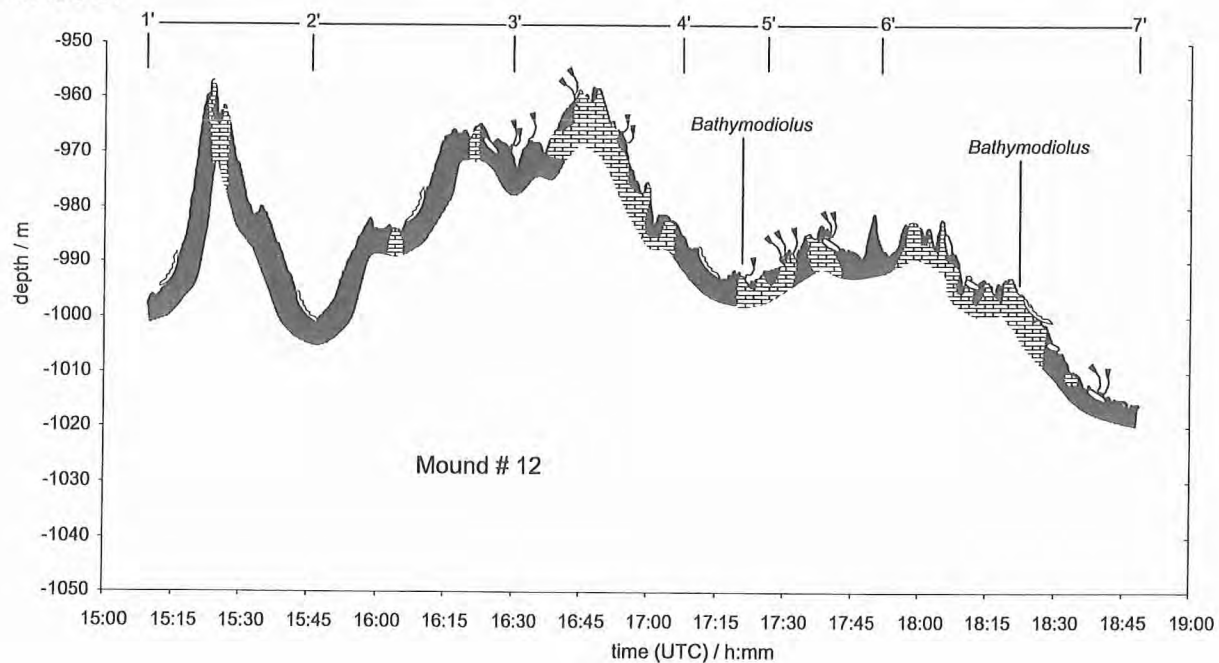


Figure 6.4.1.16: Depth profiles of OFOS 19 and 21 at Mound # 11 and 12.

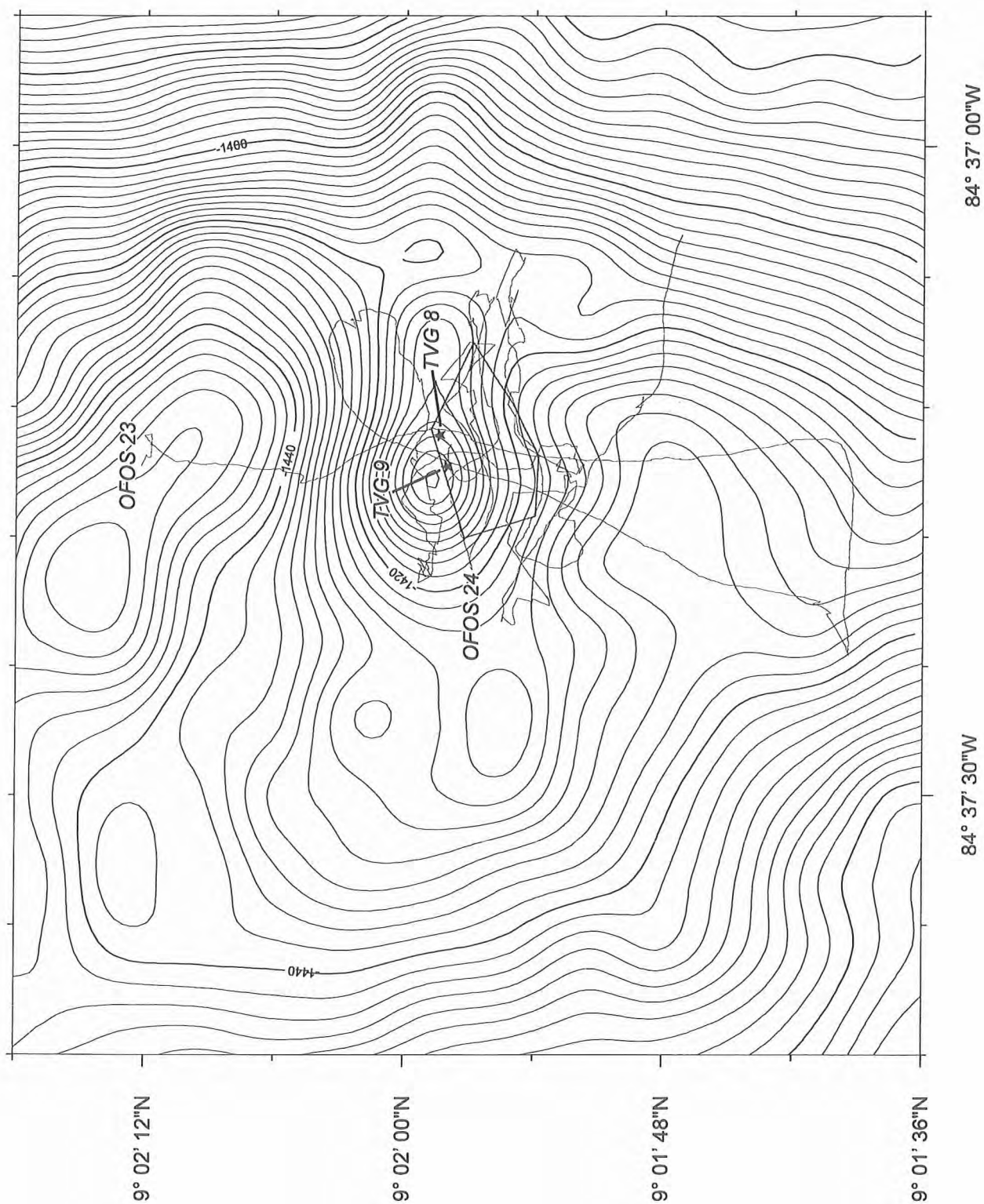


Figure 6.4.1.17: Location of OFOS 23, 24 and TVG 8, 9 at "Ranero Mound". The polygon roughly indicates the extent of the vent field, i.e. the occurrence of carbonate, pogonophorans, mytilid and vesicomid bivalves.

In general, venting activity at Mound # 12 appears to be high, sustaining a luxuriant seep community. Further investigations should concentrate on this mound, which is among the most active ones surveyed during this cruise. Only at Mound # 13 the seep organisms occur in even larger areas.

Mound # 13 (Quepos Mound) OFOS 23, 24, TVG 8, 9

Mound # 13 were discovered a sampled during the last days of the cruise. The area influenced by methane seepage at this mound is by far the most extensive we have observed during this cruise.

6.4.2 High heatflux area

During the TicoFlux I expedition with R/V Maurice Ewing A. Fisher and the Shipboard Party conducted heatflux measurements as well as investigations on the porewater of sediments obtained by gravity- or pistoncorer. They report heatflux values ($> 600 \text{ mW m}^{-2}$) and porewater compositions (high Ca concentrations) which strongly indicate hydrothermal venting. Based on these findings we surveyed this area located on the ocean plate with OFOS 8 and 9 (Fig. 6.4.2.1). The plan was to survey the two sites, at which the highest heatflux values were found, in a North-South transect by OFOS 8 and, then, to turn East in order to survey the flank of the seafloor elevation, which is the surface expression of a basement high (Fig. 6.4.2.2). Due to wind and currents we did not actually reached the (most Northward) point with the highest heatflux values. Furthermore, we had to heave OFOS for a while in order to allow the ship to adjust position. However, the track did not reveal any indications for fluid venting. A second OFOS profile (OFOS 9) were conducted at a second site with high heatflux values ($> 300 \text{ mW m}^{-2}$) further to the North. Again, no indications for fluid venting were found.

6.4.3 Seamount subduction

The subduction of seamounts and ridges on the lower (oceanic) plate cause deformation of the upper plate. High resolution bathymetry and TOBI images show clearly the uplift and faulting processes of the overriding upper plate (dome), in addition to mass wasting processes in the wake of the subducted seamount or ridge (scarp). We have surveyed with OFOS several of these structures in order to, first, find indications for fluid venting and, second, reveal the surface expressions of tectonic processes.

Scarp of Fracture Zone Trace OFOS 5

A narrow ridge mark the boundary between the lithosphere formed at the East Pacific Rise and the Cocos-Nasca spreading center on the oceanic plate named Fracture Zone Trace (Barckhausen et al., 2001). The subduction of this fracture zone and its associated morphological structures has caused uplift of the upper plate sediments and landsliding in the wake of it, a feature we call here the Scarp of the Fracture Zone Trace. The TOBI image revealed high backscatter in the area of the multiple headwalls as well as in the deposit area. OFOS 5 surveyed the scarp from the headwalls between 2210 m and 2700 m and the deposit areas down to a depth of 3100 m (Fig. 6.4.3.1). In contrast to the seamount scarps surveyed during SO 144 cruise, this scarp does not show abundant indications for venting. The OFOS depth profile over the time (Fig. 6.4.3.2) traces well the tilted blocks. We found only few segments of the headwalls which show bedding. The talus blocks areas are also limited. Few scattered empty shells were observed. No vent indications were recorded at the deposit area where high backscatter was found. Mass wasting seems to be minor at present. Most structures are sediment covered.

Uplifted area above Christmas Ridge (Rio Bongo Scarp) OFOS 10

The Rio Bongo Scarp clearly indicates the location of the subducted Christmas Ridge. Due to the subduction of the Christmas Ridge the upper plate sediments are uplifted, forming a dome. The TOBI image, which covers only the area south of $9^{\circ}19.5' \text{ N}$, indicated high reflectivity at the top of the dome. A survey (OFOS 8) during SO 144 revealed the existence of carbonates and clams at some sites on the NW Flank of the uplifted areas. With OFOS 10 we surveyed the top area of the dome and the eastern flank (Fig. 6.4.3.3). The

survey revealed that the whole top area at water depth above 550 m consists of hardground, that is probably carbonate. Although the origin of this carbonate is largely unknown, we speculate that tectonic processes in the course of ridge subduction have caused the carbonate exposure. Along the track we observed a few sites at which the hardground was cutted. There, NE-SW striking bedding was seen. We found no clear evidence for fluid venting. At some sites the outcropping rocks were very bright or had coating of whitish material, that may or may not be bacterial mats. Only at one location, around 20:17 h (UTC), whitish material on rocks as well as in depressions appeared to be bacterial mats. Here, rocks with sharp edges were seen, that could probably be chemoherm carbonates. Additionally, at 20:27:24 something was observed that could possibly have been rising gas bubbles. In general, we are not convinced that these structures are related to venting. The OFOS line was interrupted for a while because the captain was afraid that we caught a fishing line, thus, OFOS was heaved to the surface and lowered to the seafloor again. The following track surveyed faults that were largely sediment covered and did not show any indication for venting activity. In general, this structure is probably not a high priority target for future investigations.

Jaco Scarp OFOS 11, 12, TVG 3

A detailed image of Jaco Scarp was obtained by TOBI during SO 144. At that cruise OFOS tracks 1&2 surveyed the uplifted areas and the scarp. We used a variety of OFOS surveys to further understand tectonic processes and fluid venting patterns at this feature, which was also imaged by DTS # 3. That DTS covered along WE profiles the uplifted areas. The regular survey was followed by a high resolution (410 kHz) line, which was planed to cover the circular structures seen in the TOBI, but the DTS passed these structures to the south. The last line then followed the scarp downslope.

OFOS 11 was deployed in order to survey the circular structures with high backscatter. We crossed these with an east to west profile. Due to wind and currents the ships speed was very low. In the area of high backscatter the seafloor appears to be continuously paved by carbonates below a very thin sediment cover in which very scattered bivalves occur. At some areas, donut-shaped carbonates occur which are surrounded by clamshells. We deployed TVG 3 in order to sample these carbonates. Again, this grab felt to the we were able to sample this carbonates, 89 living specimens of the solemyid bivalve *Acharax johnsoni*, shells of vesicomylid and thyasirid clams and sediments with a weak sulfidic odour. The high abundance of *Acharax*, that were approximately the same size (46.2 ± 8.7 mm), indicate that there is a considerable but low supply of sulfide. This sulfide is possibly produced by anaerobic methane oxidation and, thus, we conclude that methane venting does occur at present.

Additional OFOS surveys were conducted towards the end of the cruise (OFOS 25, 26). These did not reveal any further indications for fluid venting. With TVG # 7 we were able to sample carbonate rocks.

Parrita Scarp OFOS 13, 22

Two OFOS surveys during the SO 144 cruise and this leg revealed that only few indications for venting exist in the uplifted area of Parrita Scarp. Additionally, during both cruises the area West of Parrita Scarp was surveyed. There, the TOBI image showed high backscatter reflectivities. Both OFOS runs showed only infrequent clamshells. Later during the cruise the area was surveyed by DTS and revealed the existence of very small mounds, which were hardly seen on the TOBI image and do not show up in the bathymetry. This small mounds were not surveyed by OFOS.

“Frontal collision” OFOS 15

With OFOS 15 we surveyed the collision zone between the accretionary wedge and a seamount. The area at the frontal wedge is characterised by very steep morphology. Beside few scattered clamshells we found no indication for methane seepage.

“Small Seamount Scarp” OFOS 16

With OFOS 16 we surveyed the uplifted area of a seamount, that is probably smaller than the seamounts subducted at Jaco or Parrita Scarp. We did not find any indications for methane seepage.

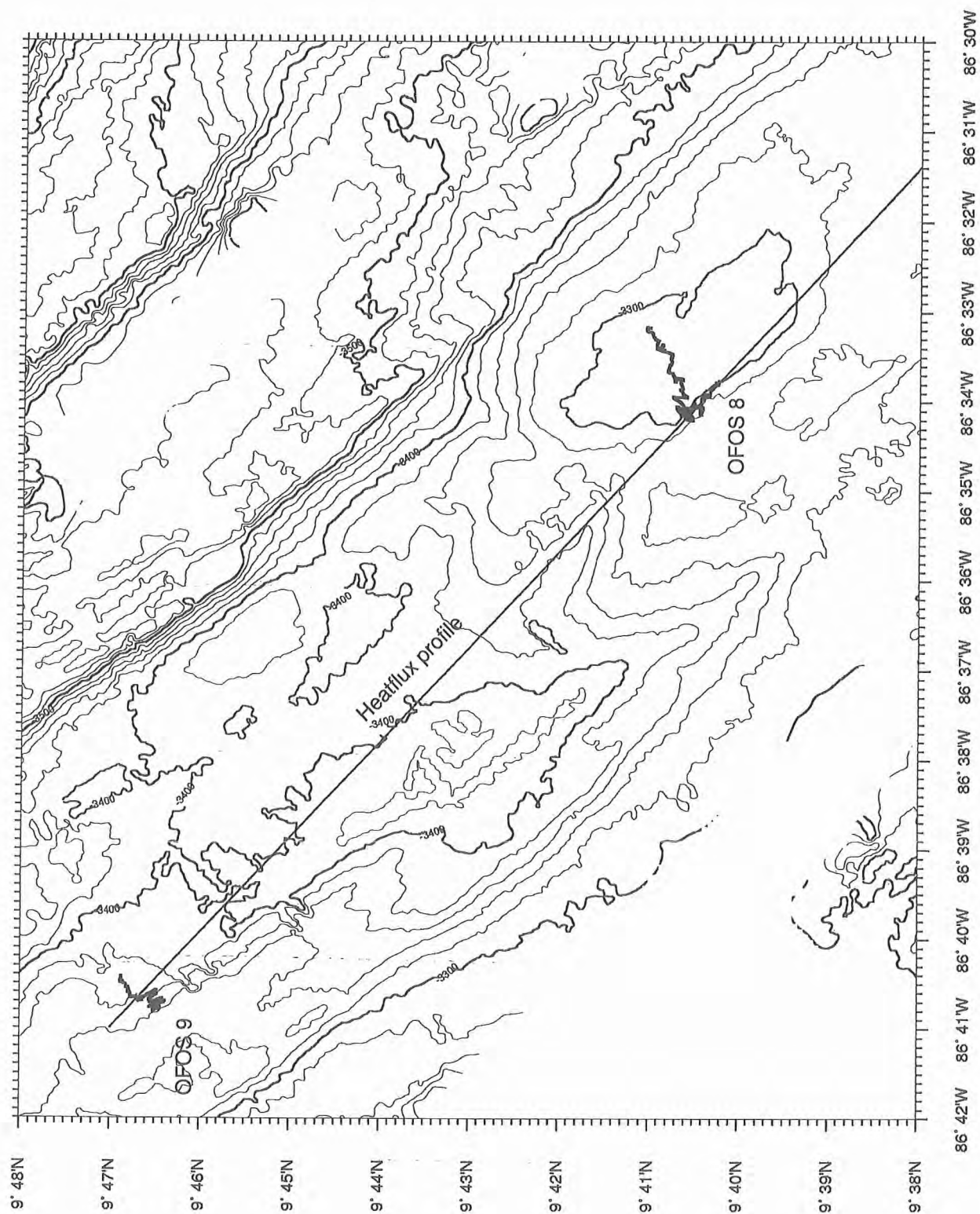
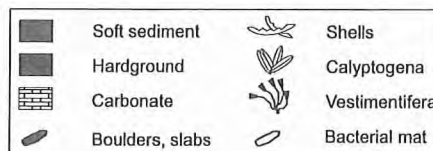
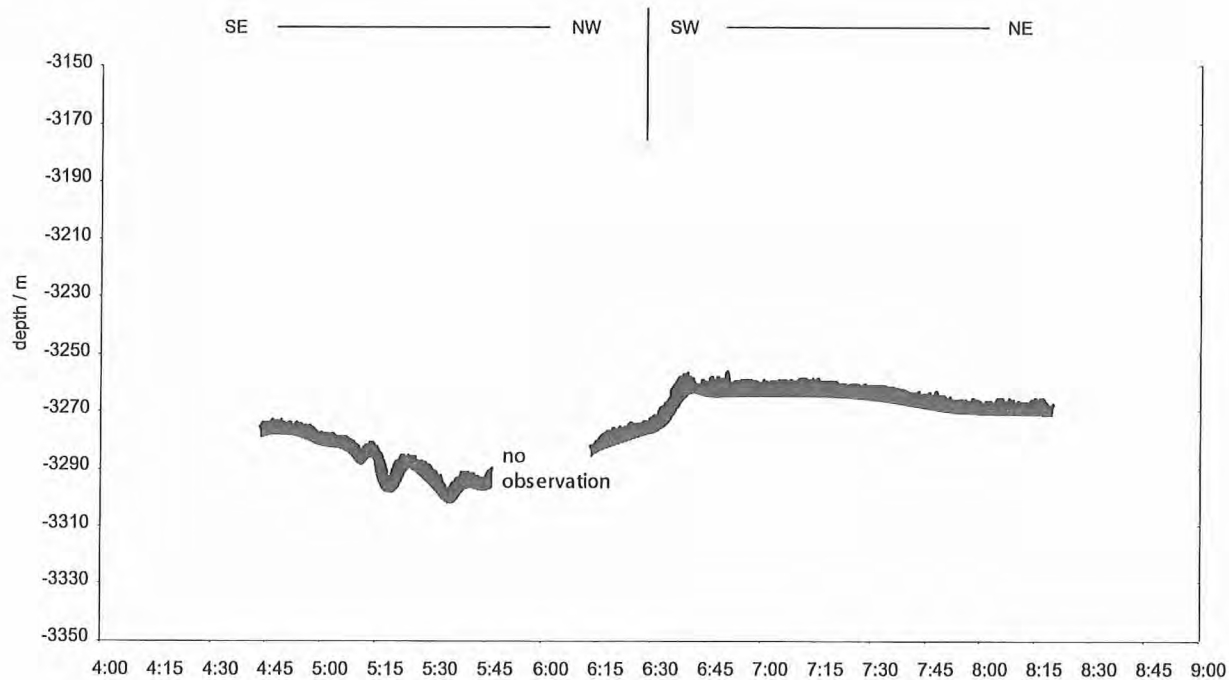


Figure 6.4.2.1: Location of OFOS 8 and 9 profiles at the high heatflux sites on the ocean plate.

OFOS 8



OFOS 9

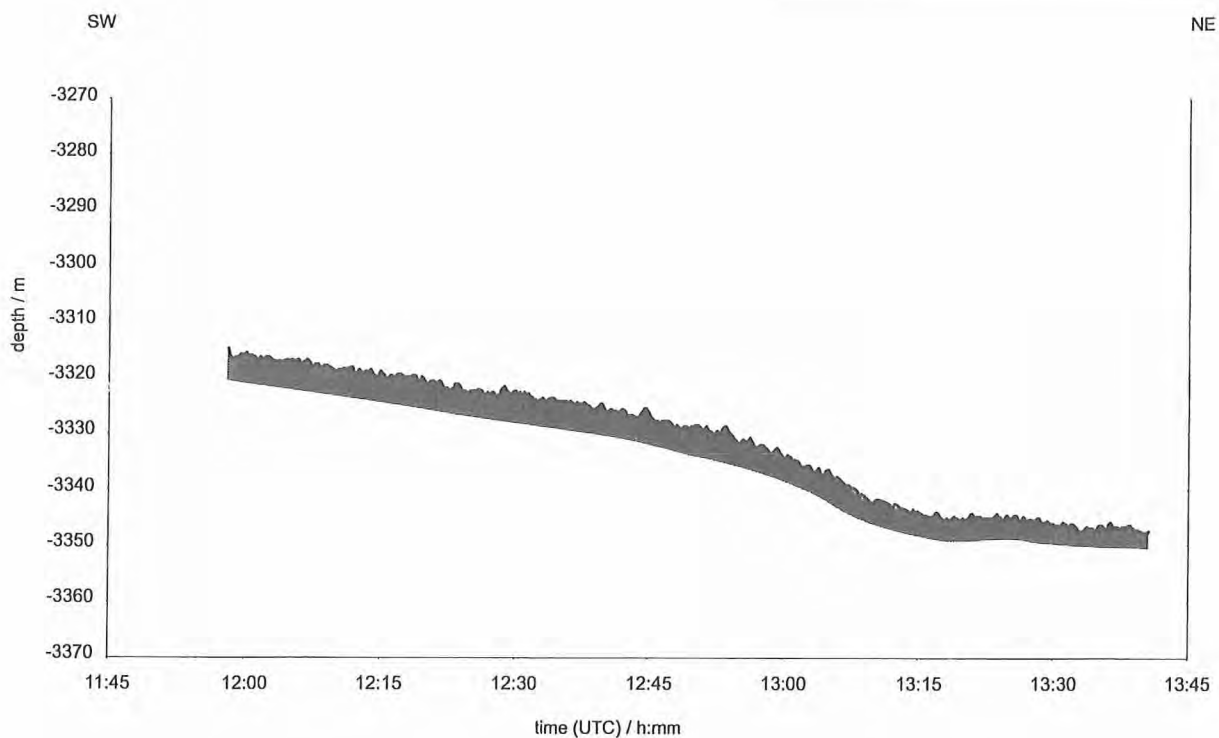


Figure 6.4.2.2: Depth profiles of OFOS 8 and 9 at the high heatflux site on the ocean plate. We found no evidence for hydrothermal venting.

6.4.4 Landslides

Quepos Landslide OFOS 18, 20, TVG 6

An OFOS survey conducted during SO 144 at the Quepos landslide (OFOS 4) revealed that enormous bacterial mats cover the sediments at and below the headwall of this landslide. Additionally, bacterial mats and clams were observed in the area of the slide tail. The occurrence of bacteria in the headwall area can be explained either by anaerobic methane oxidation or by anaerobic organic matter oxidation. We surveyed the area with two OFOS profiles (18 & 20), that run parallel to the SO 144 OFOS. OFOS 18 confirmed the existence of the bacterial mats covering the sediments below the headwall. Below, at the slide mass, we observed with OFOS 18 and 20 very few bacterial mats and no living clams or clamshells that are typical for methane venting. With TVG 6 we sampled sediments covered by bacterial mats just below the headwall. Just before the grab touched the ground we saw in the cameras, that thin Beggiatoa-like bacterial filaments were looking through the sediment surface. When the grab came close to the sea surface we saw bubbles rising from the grab. On deck, gas was released from the sediment. The grab was completely full of sediment that was still increasing in volume due to degassing. The entire sediment was mousse-au-chocolate-like, highly porous, although at the same time considerable “dry”. It had a considerable sulfidic odour. Several 2-3 cm thick layers of cemented sediments were found throughout the sediment. The sediment surface did not remain intact and, thus, we did not find any bacterial mats. In conclusion, we believe that methane venting rather than anaerobic organic matter oxidation cause sulfidic conditions. Porewater studies are needed to confirm which geochemical pathway causes the sulfidic conditions. The area covered by this bacteria are large enough to be sample by conventional equipment, i.e. multicorer or gravitycorer.

6.4.5 Accretionary wedge

Accretionary wedge at graben at crest of Cocos Ridge (off Osa Peninsula)

The wedge where the graben (formed at the crest of the Cocos Ridge) collides with the upper plate is characterised by a largely undisturbed accretionary sequence. At the very frontal slope of these sequence linear patterns of alternating high and low backscatter imaged by TOBI exists. Based on geophysical data, that show the greatest porosity reduction at the frontal prism, this area should be influenced by fluid escape. With OFOS 14 we surveyed the front down to the oceanic plate, but could not detect any indication for fluid venting. The seafloor was covered by very soft sediment, thus, no geological features associated with deformation processes could be observed.

Accretionary wedge along seismic line SO 81 # 05

The accretionary wedge along the seismic line SO 81 # 05 (see (Huene et al., 2000) for location) was surveyed by OFOS 17 in order to search for evidence of fluid venting. This area is, as shown by the seismic line, highly influenced by deep-seated faults. The seafloor expressions of these faults are probably small displacements seen on the OFOS depth profile. We did not discover living chemoautotrophic species or carbonate precipitations during this deployment. The seafloor was uniformly covered by soft sediment. However, we found shells that probably originate from vesicomid clams that are indicative for methane venting processes.

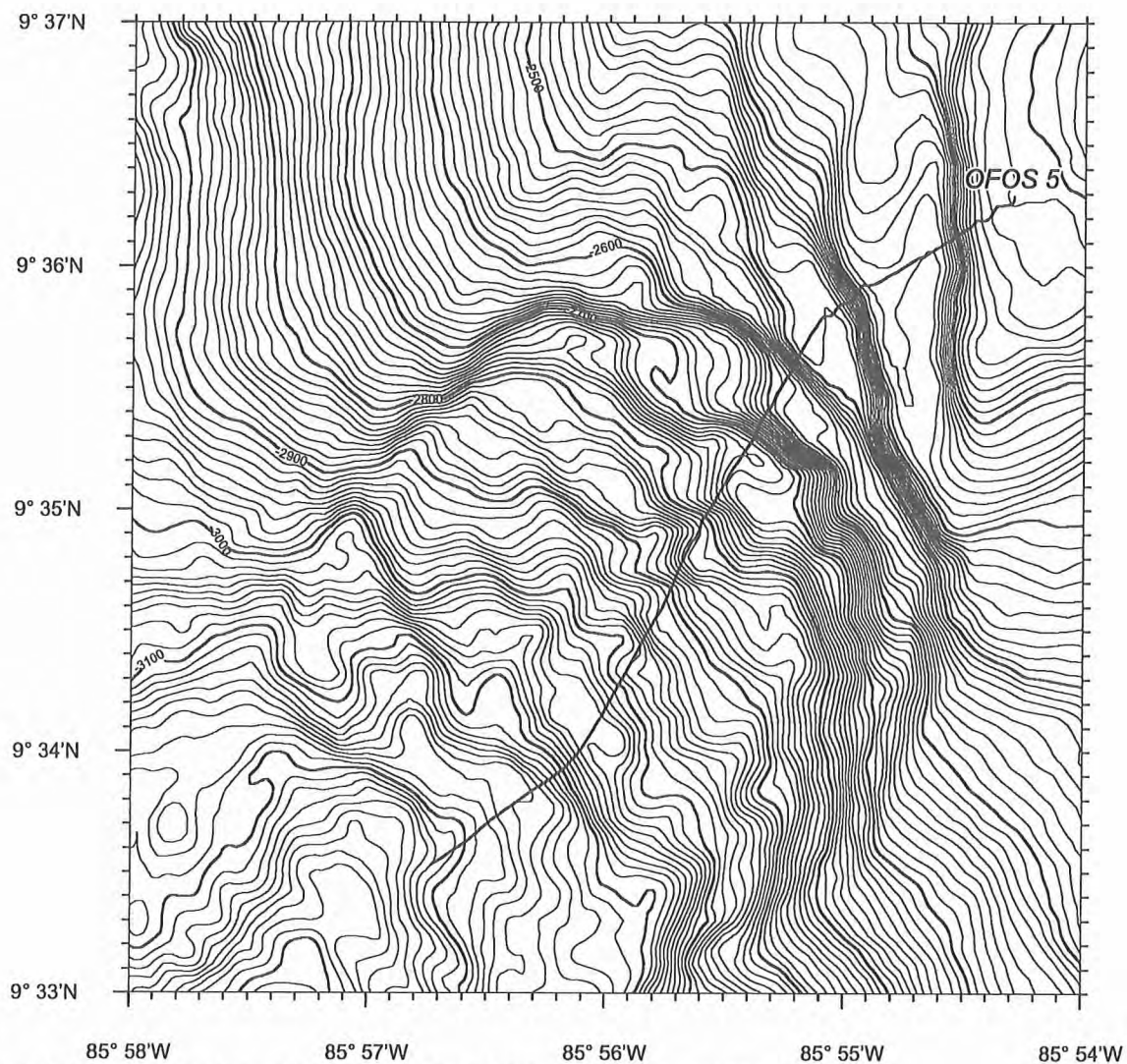


Figure 6.4.3.1: Location of OFOS 5 profile at Scarp of Fracture Zone Trace.

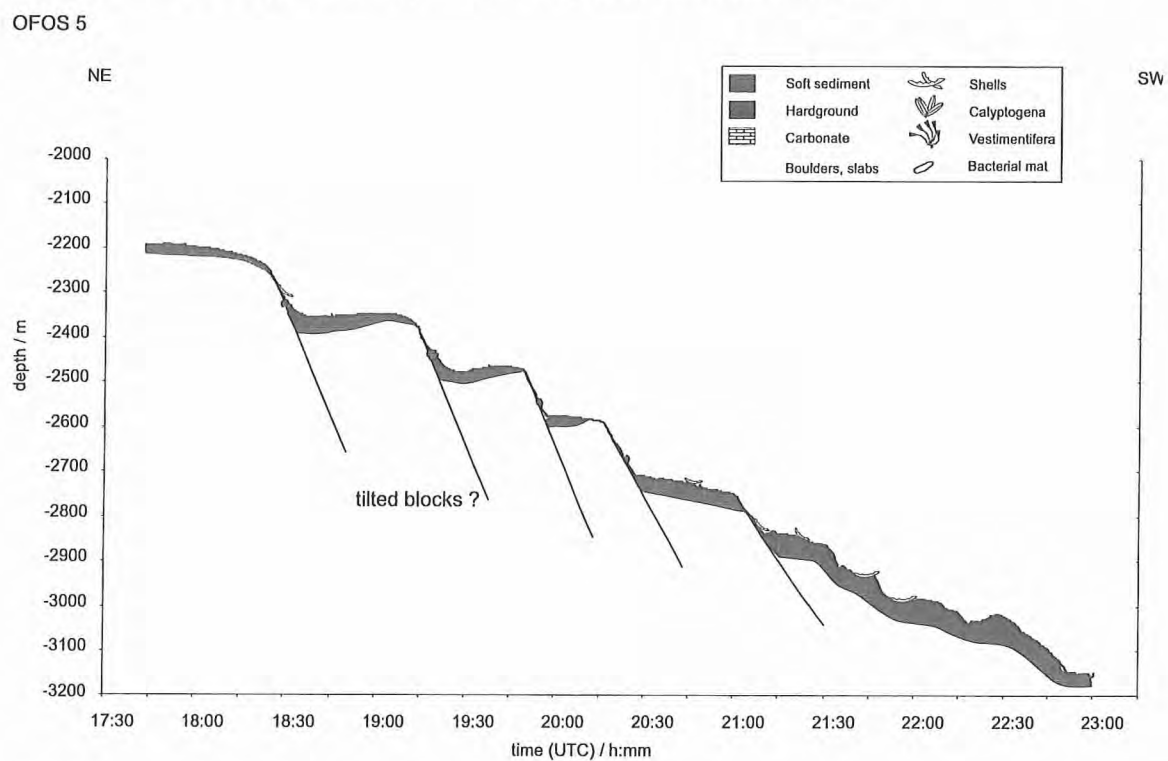


Figure 6.4.3.2: OFOS 5 depth profile at Scarp of Fracture Zone Trace. Beside very few scattered clamshells we found no evidence for fluid venting.

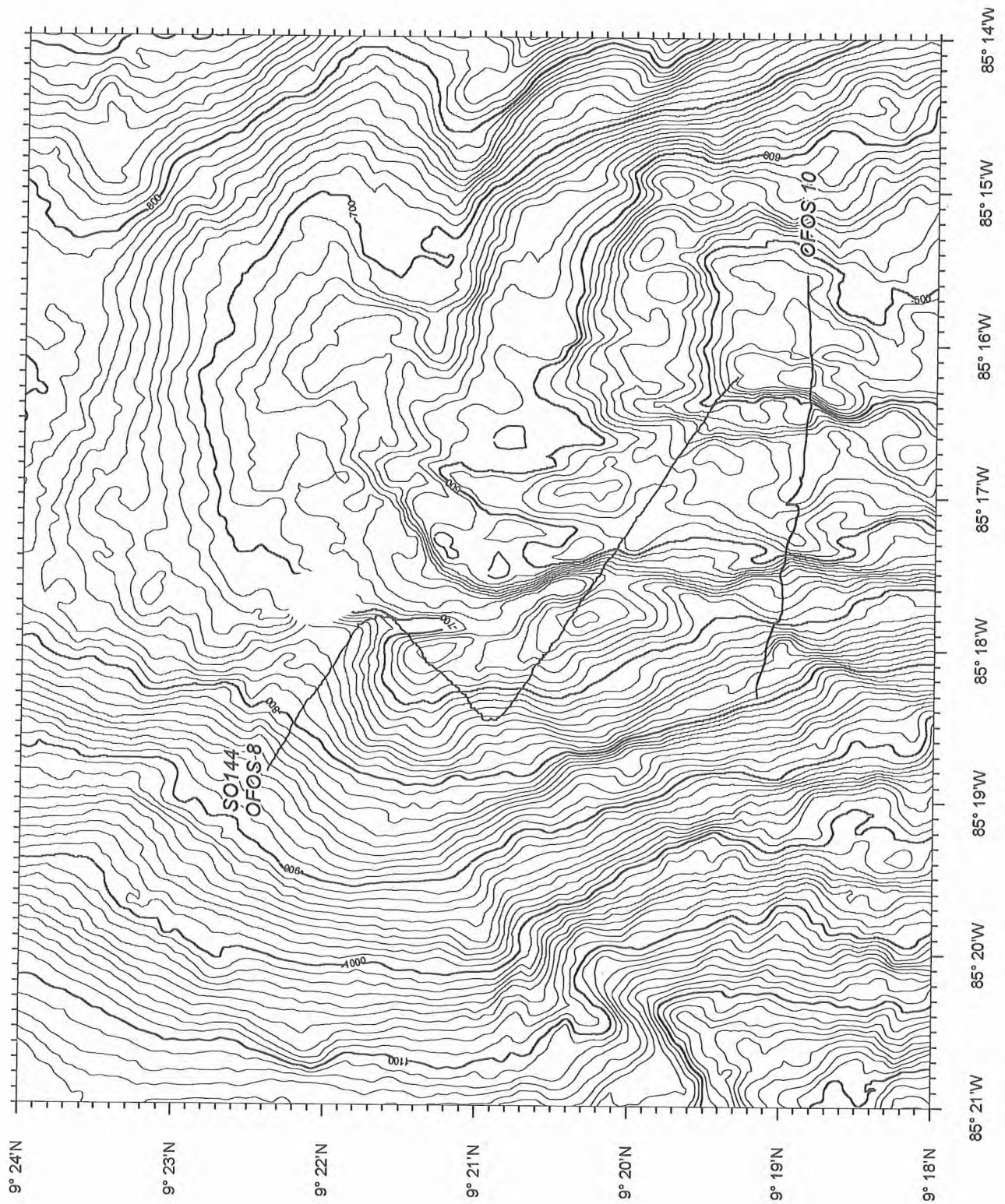


Figure 6.4.3.3: Location of OFOS profiles at uplifted area of Rio Bongo Scarp.

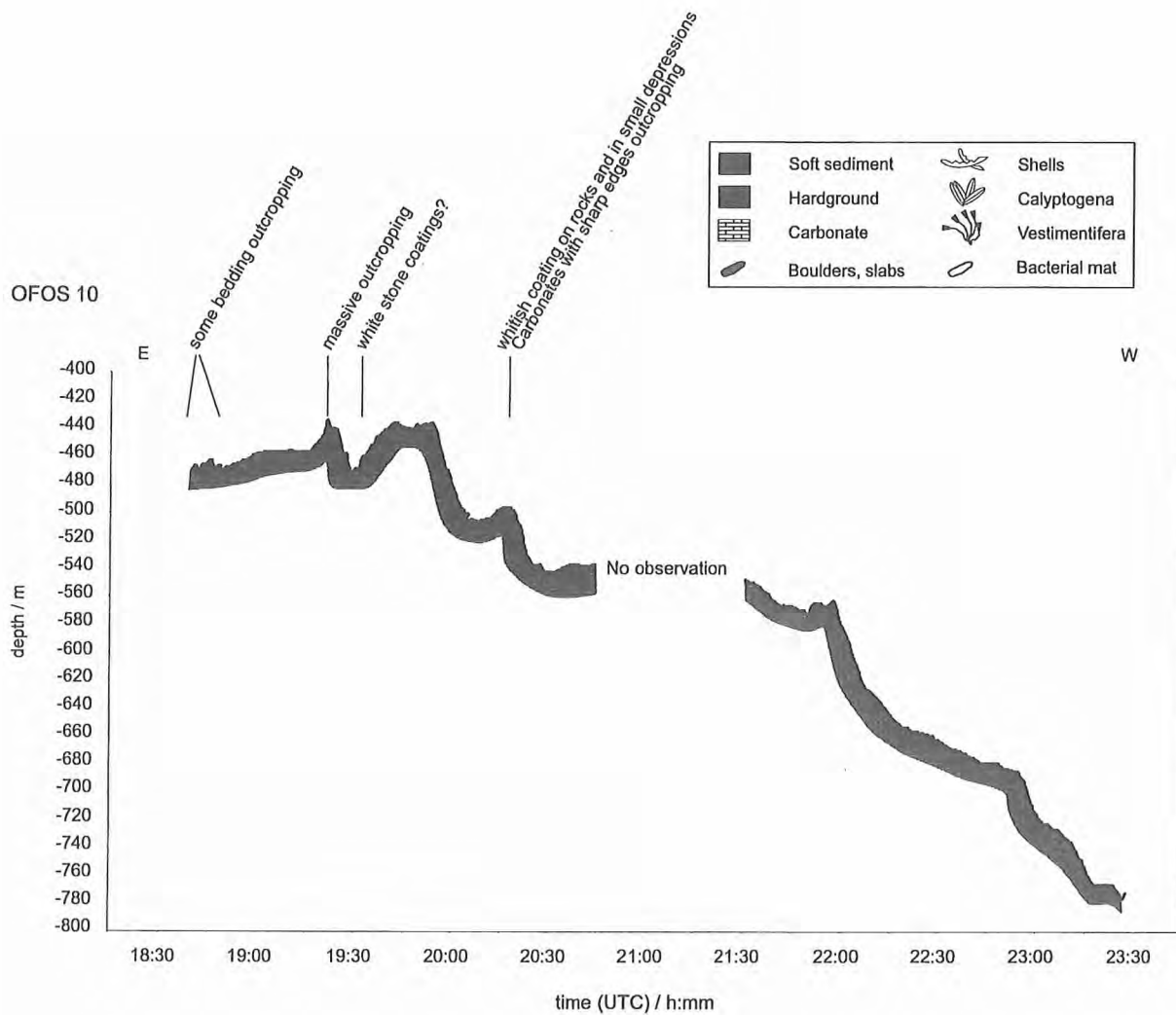


Figure 6.4.3.4: OFOS 10 depth profile at uplifted area on Rio Bongo Scarp. The entire top area above 550-m water depth consists of hardground. This is probably outcropping massive carbonate forming a uniform flat plateau. The hardground is cutted at places, exposing bedding that is NE-SW striking. We found no indication for fluid venting beside whitish coatings on rocks and in depressions that can possibly be bacterial mats and chemoherm-like carbonates.

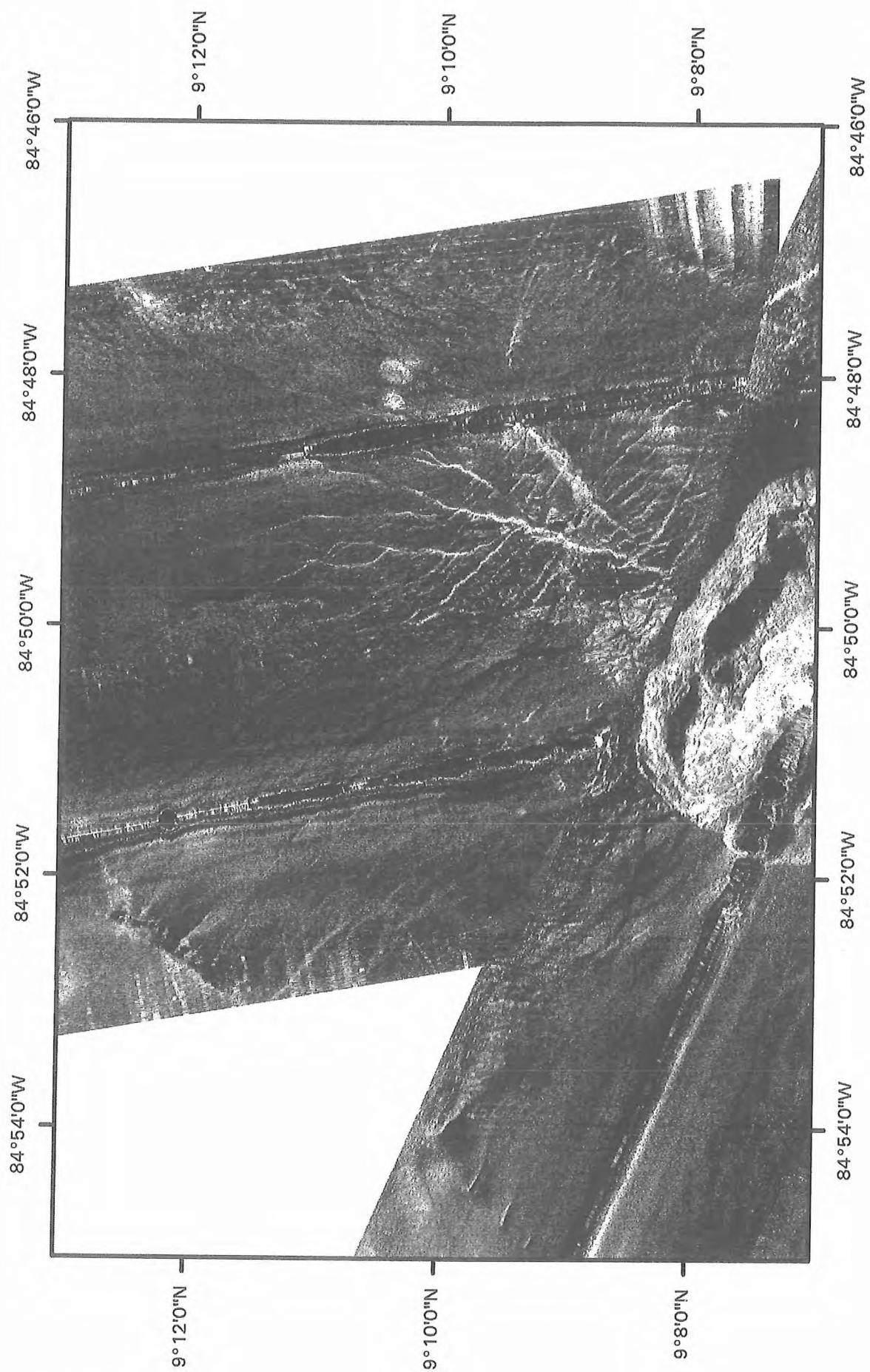


Figure 6.4.3.5: TOBI image of Jaco Scarp.

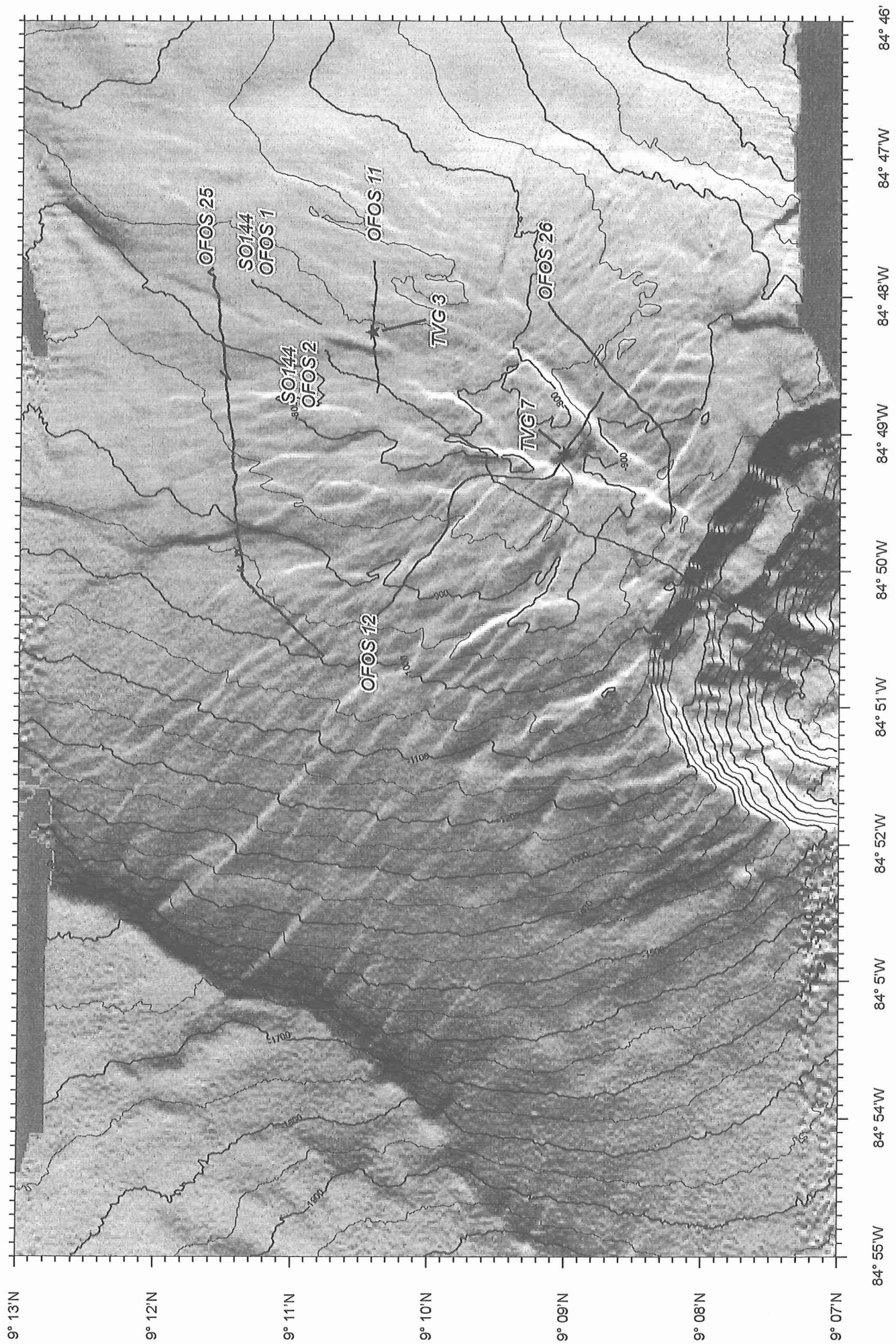
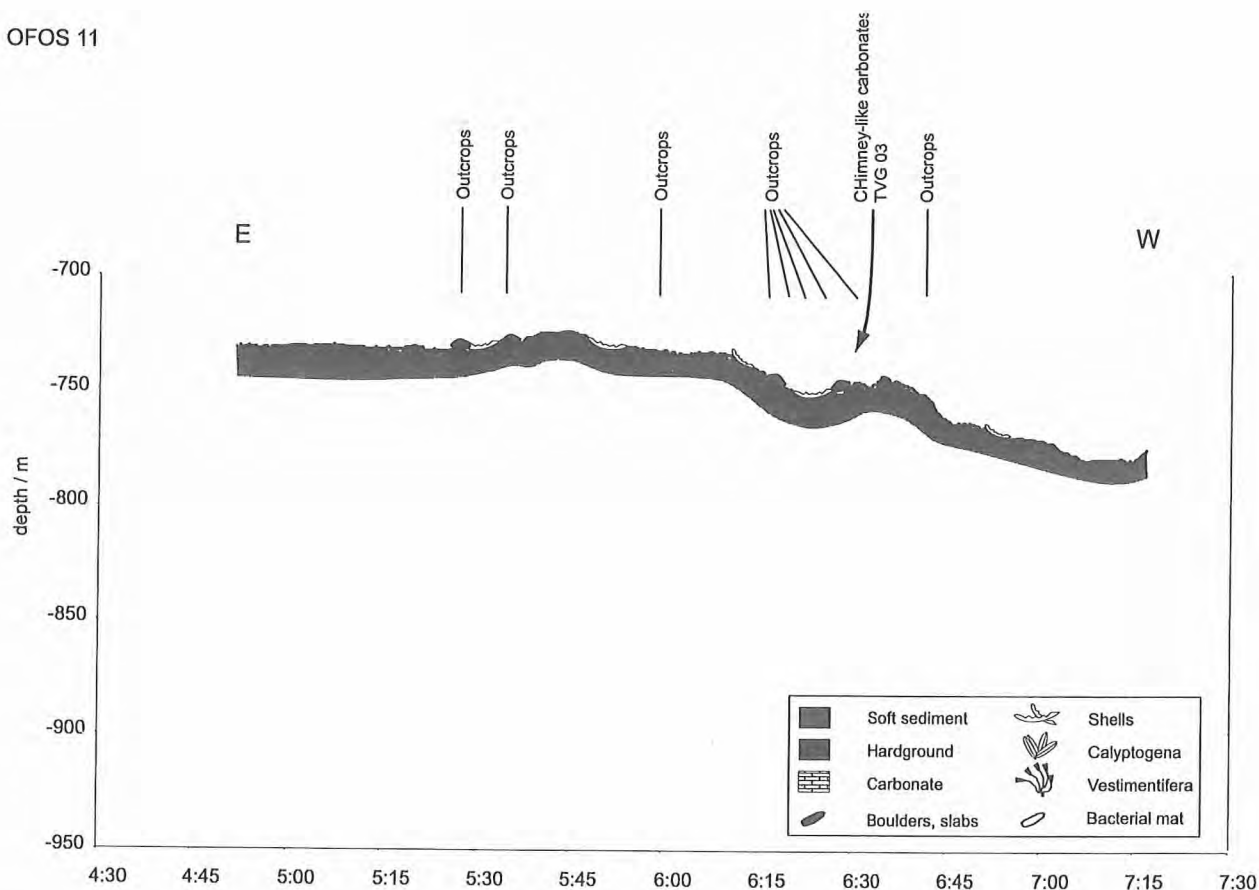


Figure 6.4.3.6: Locations of OFOS tracks and TVG sampling stations at Jaco Scarp. The high-resolution bathymetry was obtained by the Simrad system during this cruise.

OFOS 11



OFOS 12

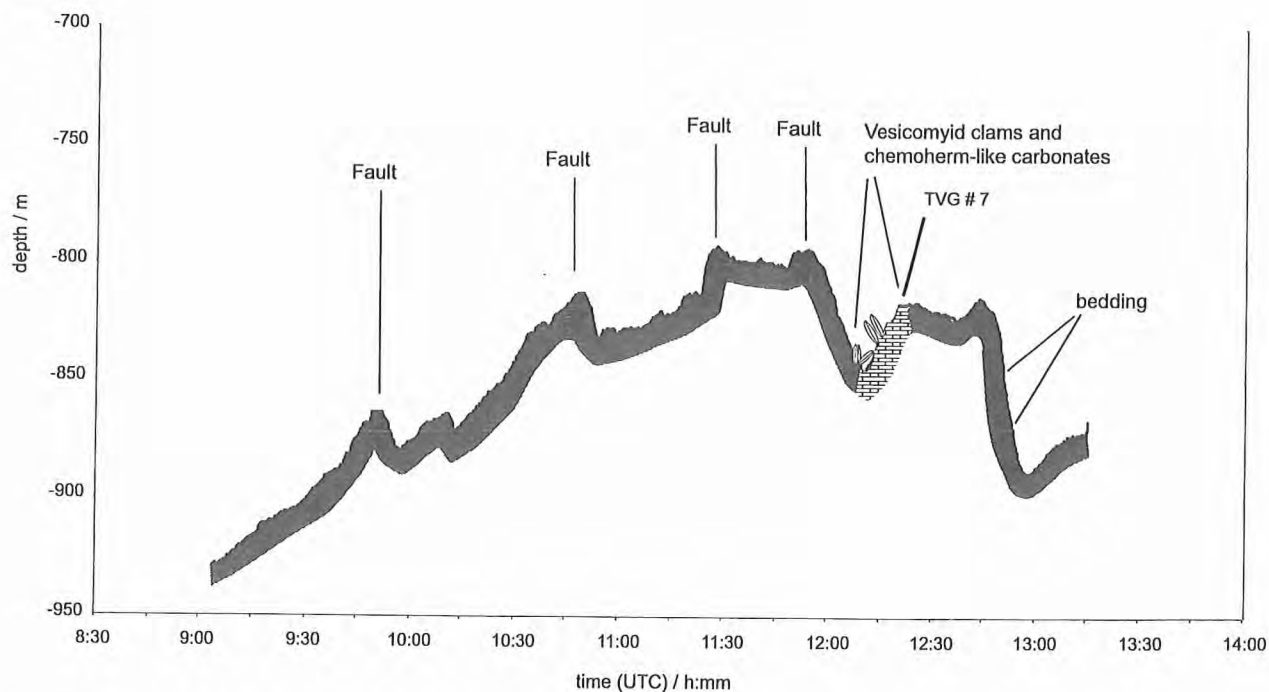


Figure 6.4.3.7: OFOS 11 and 12 depth profiles at dome of Jaco Scarp. Circular high backscatter features were surveyed by OFOS 11. The locations of these features correspond to areas with outcrops. In addition, hard ground exists in a large area, in which scattered clamshells are abundant. Chimney or donut-like carbonates were observed around 6:25 (UTC) and sampled by TVG 03. These carbonates as well as abundant living solemyid bivalves indicate methane venting at this site. OFOS 12 surveyed the radial fault patterns. We crossed several faults at which outcrops were seen. We found indications for methane

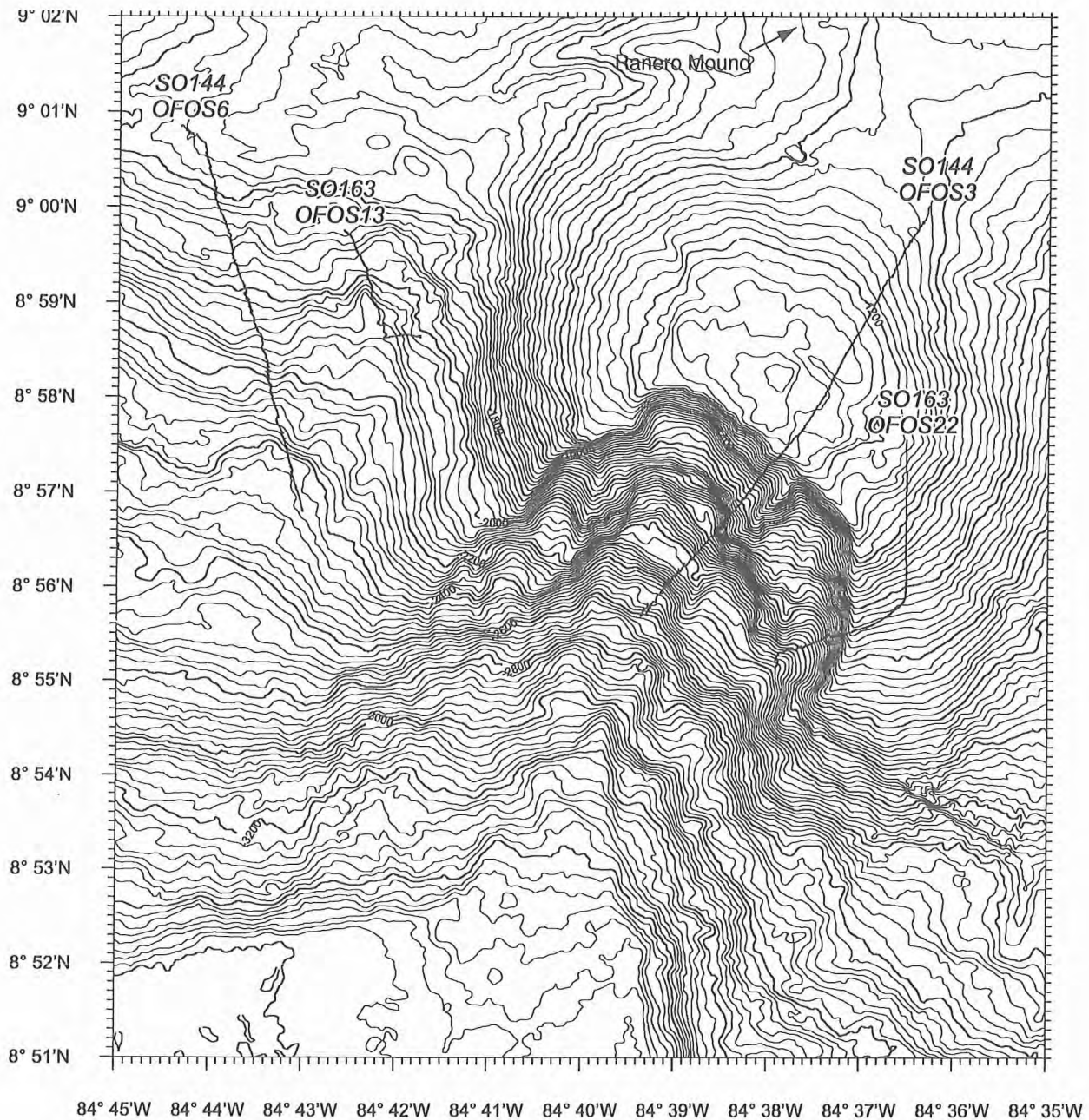


Figure 6.4.3.8: Location of OFOS tracks at and near Parrita Scarp. SO 144 OFOS 3 and SO 163 OFOS 22 surveyed the uplifted area of Parrita Scarp. SO 144 OFOS 6 and SO 163 OFOS 13 were deployed in order to survey the seafloor at sites which revealed high backscatter in the TOBI image. Note that we discovered with DTS after the OFOS deployments the existence of very small features that are probably mounds. We did not survey these mounds with OFOS.

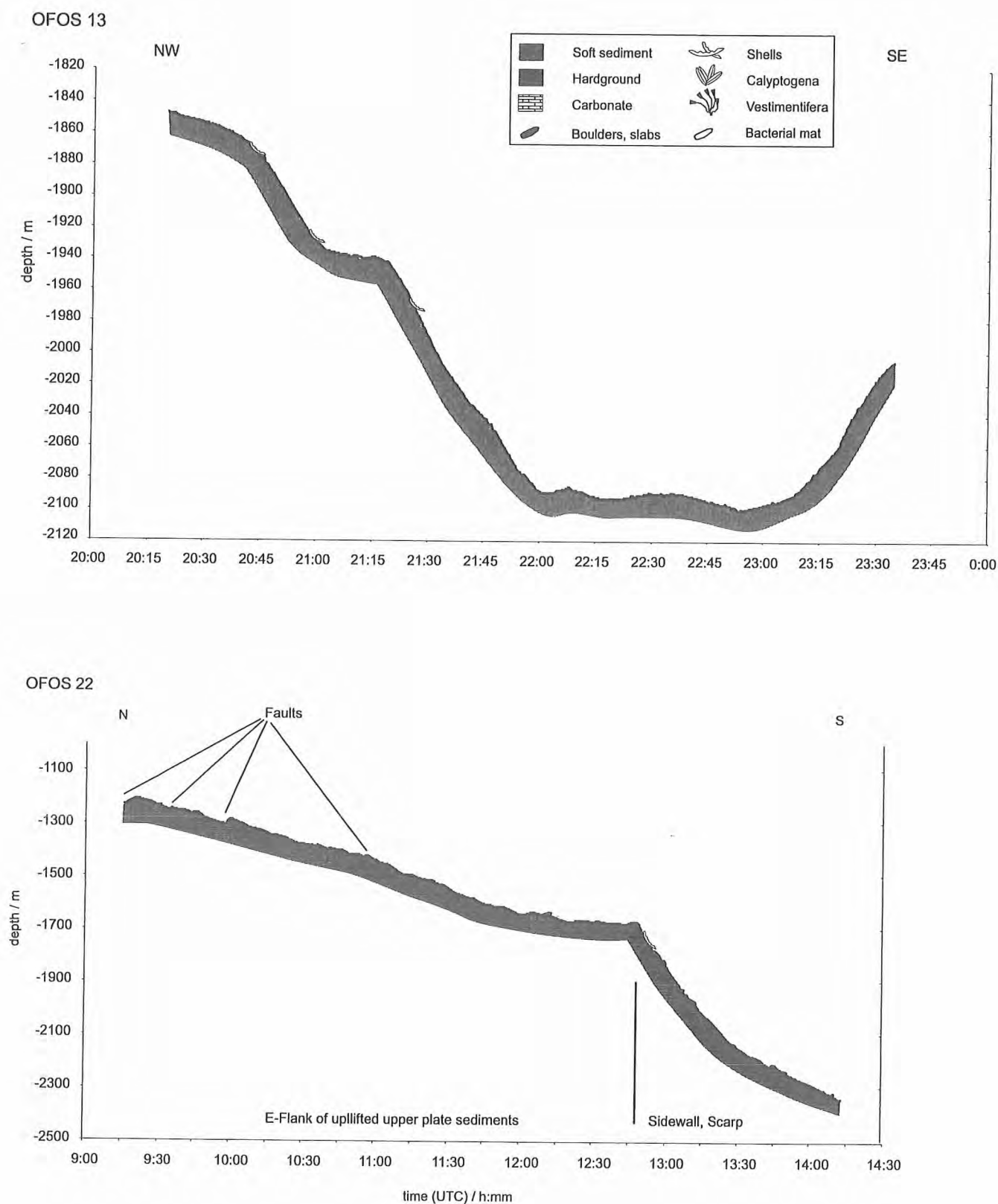


Figure 6.4.3.9: OFOS 13 and 22 depth profile. OFOS 13 was designed to survey an area of high reflectivity E of Parrita Scarp. During the DTS profile it turned out that in this area small mounds are present, that we have just missed. OFOS 22 surveyed the E side of the uplifted area at Parrita Scarp but did not reveal any signs of venting.

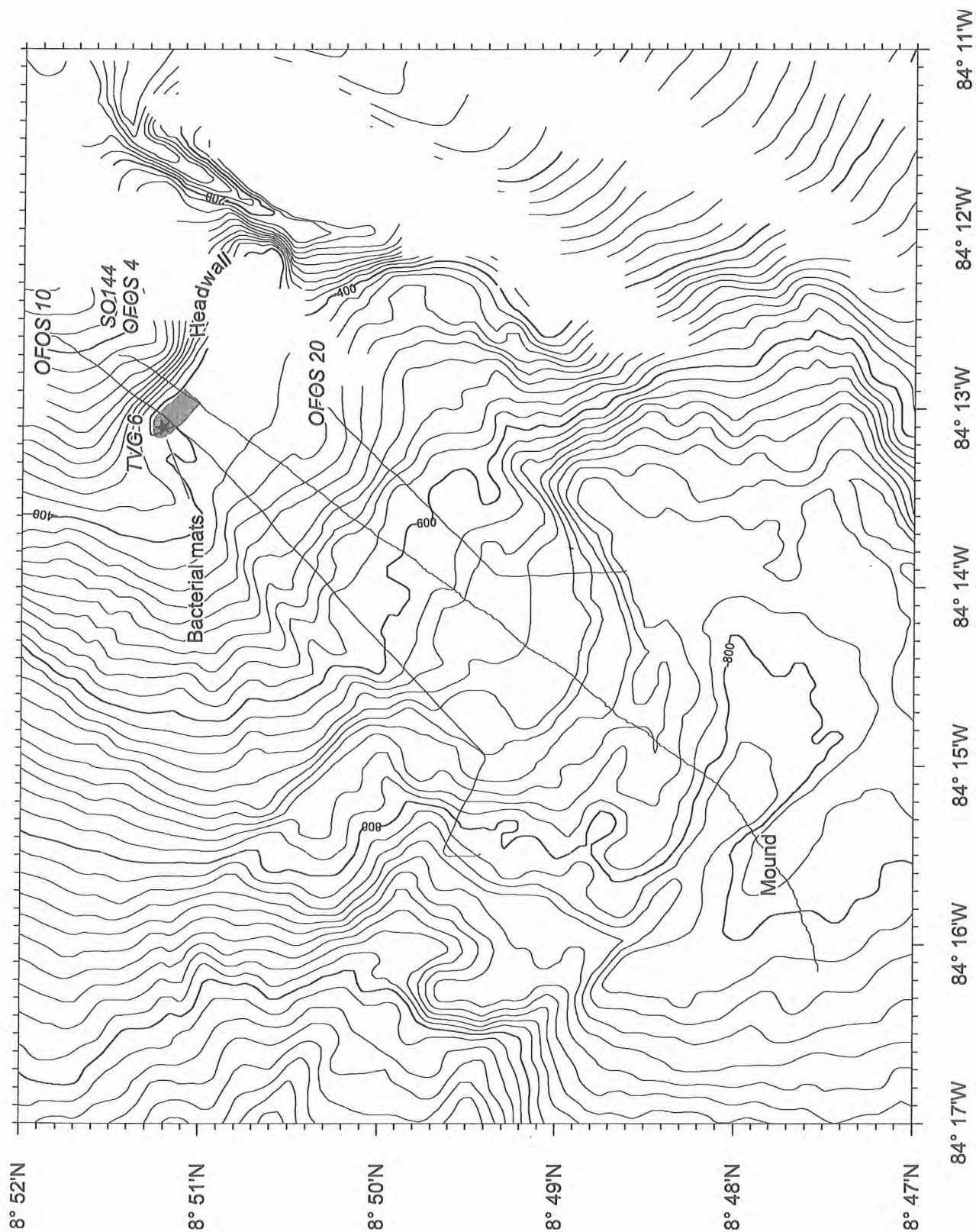
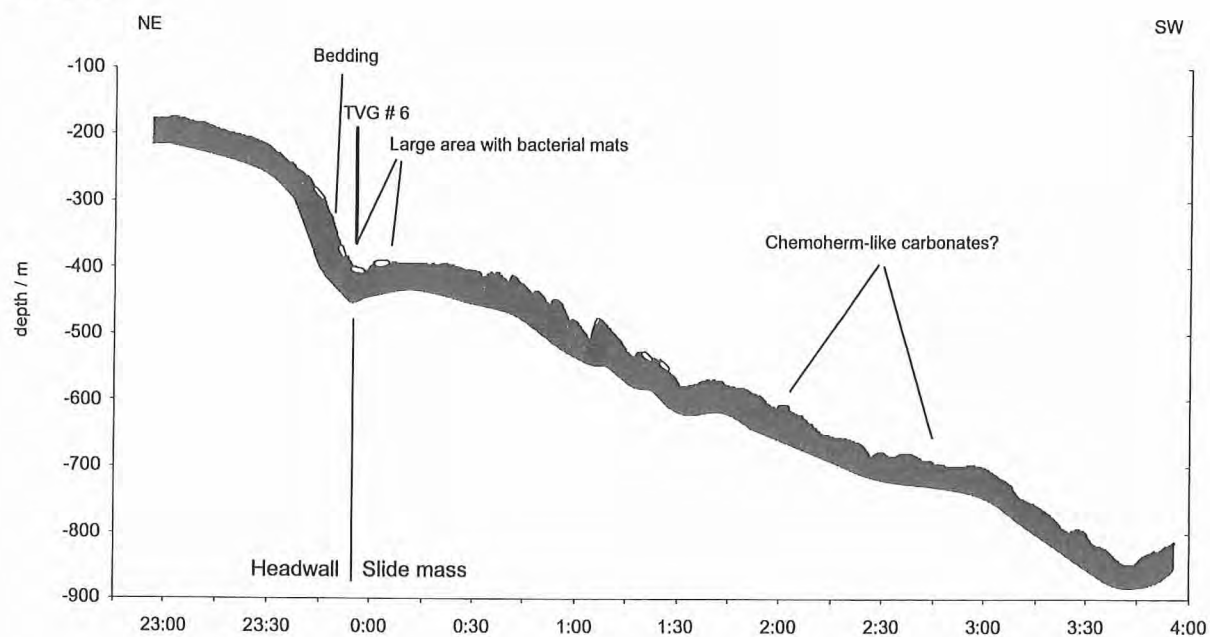


Figure 6.4.4.1: The locations of OFOS profiles, TVG deployment and major bacterial mat occurrence at Quepos landslide. Notice that with SO 144 OFOS 4 by chance carbonates at a mound were surveyed.

OFOS 18



OFOS 20

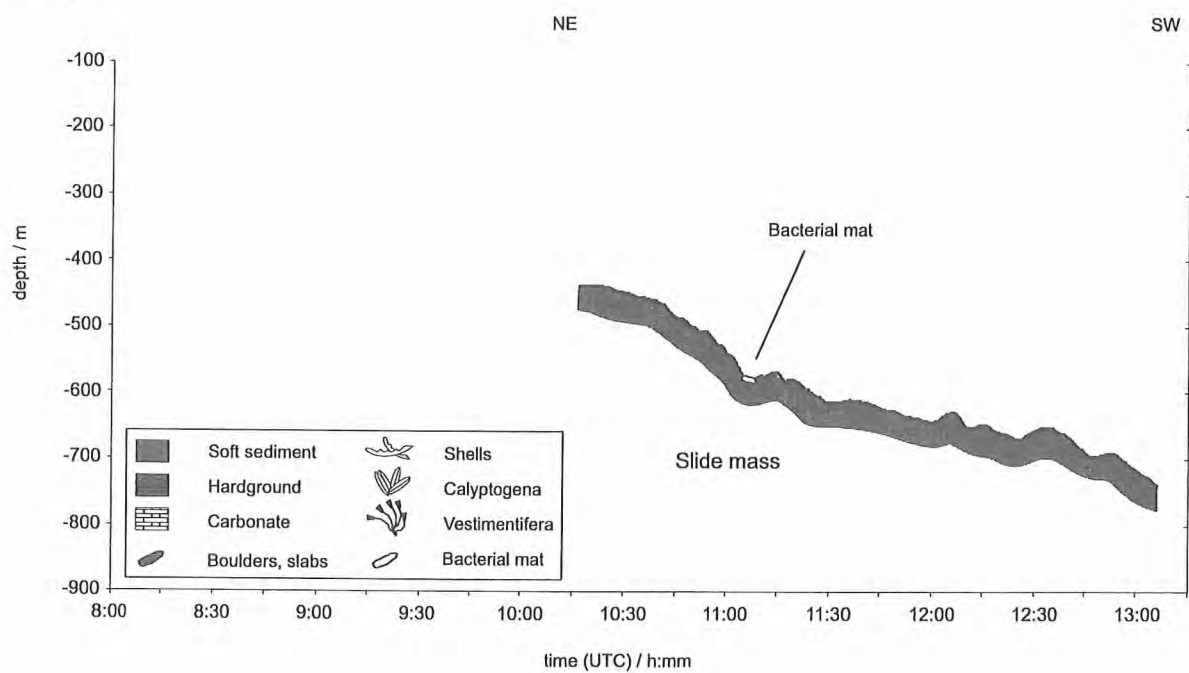


Figure 6.4.4.2: OFOS 18 and 20 depth profiles. Below the headwall, the most extensive bacterial mats are present. In contrast to SO 144 OFOS 4 only few bacterial mats and no seep-typical living clams or clamshells were observed in the slide mass. The most extensive bacterial mats were observed between the points: 8°51.21'N, 84°13.10'W and 8°51.1'N, 84°13.10'W.

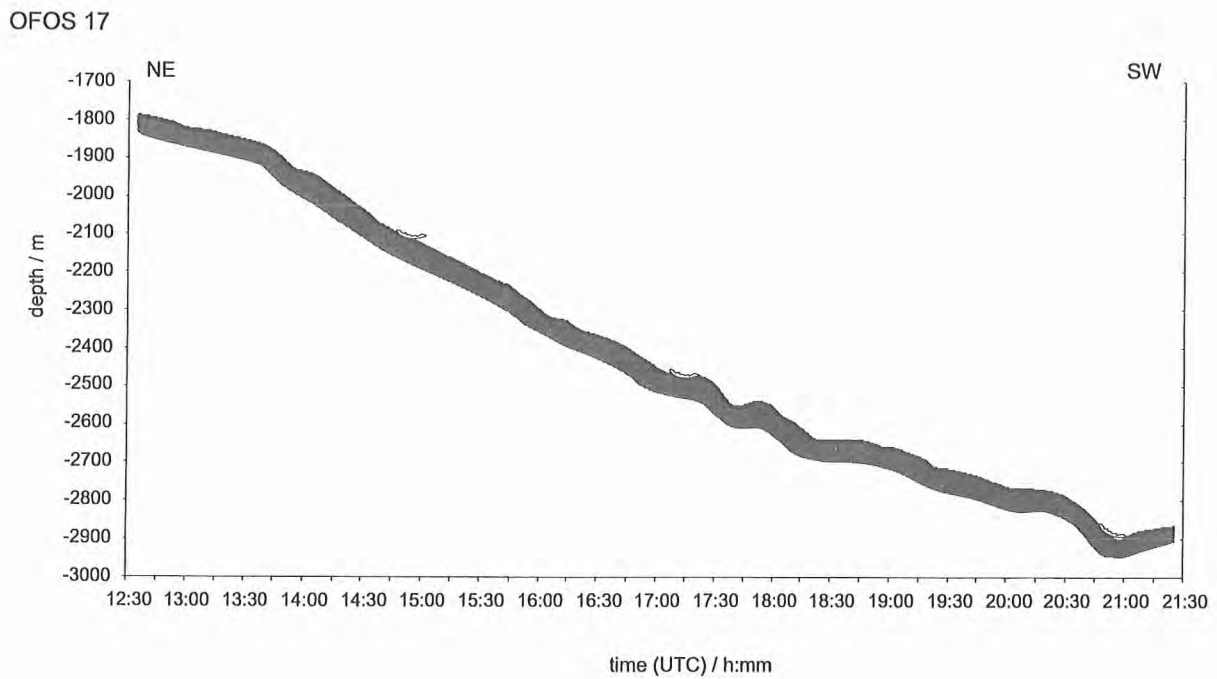
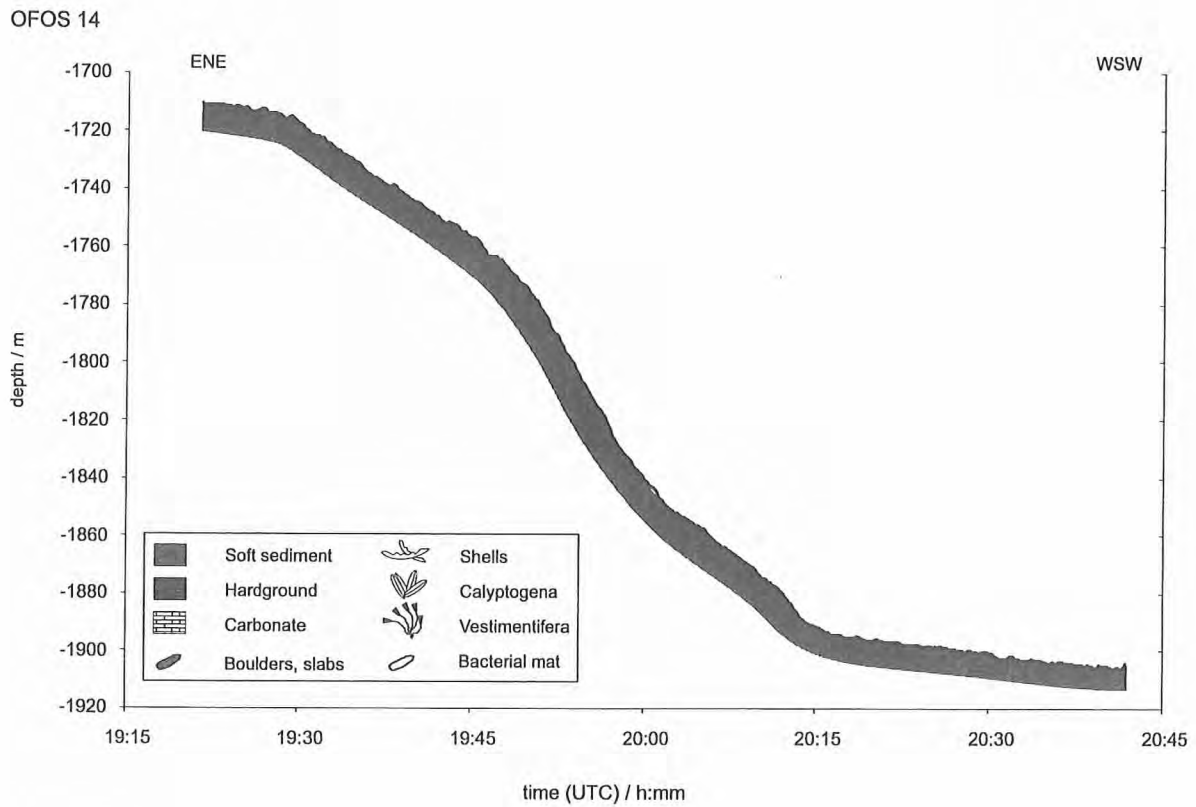


Figure 6.4.5.1: OFOS 14 and 17 depth profiles at the accretionary wedge. For locations see Figure 6.4.2. Soft sediment prevails throughout the profiles.

6.5 Magnetic Data

(P.O. Thierer)

During the cruise SO163-1 of R/V SONNE, in total 16 magnetic profiles were collected. The magnetometer was deployed during SIMRAD^{EM} 120 bathymetry profiling and on transit between different working areas. Due to the technical parameters of the GeoMetrics G801/3 Marine Proton Magnetometer system, its maximum working depth is limited to 70 meters below seafloor.

The measurement of the natural magnetic field requires that the influence of other, artificial magnetic fields is minimised. In consequence a distance of at least three times ship length between the vessel and the sensor is required; in case of R/V SONNE 300 meters. To keep the depth of the sensor above 70 meters below seafloor, the velocity of the ship is forced to be at least 4 nms per hour. Therefore magnetic profiling could not be done while TOBI or DTS-1 Side Scan Sonar profiles were done.

Interpretations of the magnetic data has been done on the total field magnetic raw data. No corrections as e.g. daily variations or latitude correction was done. Therefore, interpretations are of preliminary character.

The magnetic profiles obtained are shown in Figure 6.5.1 to 6.5.12. All profiles are situated on the Costa Rica continental margin.

Profile 1 is striking about 120 degrees and crosses the lower part of the slope. The magnetic data show variations in the magnetic field strength of about 400 nT, they amount to a relative maximum toward the continent. Furthermore, we can observe variations which correlate to the observed topography on the margin.

Profile 2 was carried out during SIMRAD^{EM} 120 bathymetric profiling on the slope. We can observe variations in the total magnetic field strength and a local maximum in the outermost north of this survey area. Variations along the track might correlate to topographic features.

The magnetic profiles 3, 4 and 5 are only of short length and trends (except profile 5) parallel to the slope. They all show a constant magnetic field strength. Probably the wavelength of the magnetic variations is too long in relation to our profiles.

Near by the Nicoya Slide, the magnetic profile 6,7,8 and 9 were obtained. At this place, a seamount is subducting. The magnetic profiles show a relative homogeneous northern part and a distinct magnetic anomaly to the south in the near offset area of slide. This magnetic anomaly fits to the subducting seamount at this place and clearly evidences its influence to the change of the total magnetic field at this place.

Profile 10 is situated in the lower part of the slope and reaches near by point "A" the oceanic plate. We can observe a little change in the magnetic field strength which might be part of the magnetic anomaly pattern of the oceanic plate. Toward the slope, magnetic variations are only very little. They might be linked to the normally very homogeneous magnetic behaviour of continental slopes and the fact that the seafloor topography in this area is very smooth.

Profile 11 instead is located in higher position on the slope and furthermore to the north than profile 10. However, magnetic variations are not very distinct too; they might be explained as it is done for profile number 10.

Magnetic line 12 is situated on the lower slope, its geographical position is between profile 10 and 11. In contrast to these lines, magnetic line 12 shows distinct magnetic anomalies. The influence of the normal oceanic plate magnetic anomaly pattern to the total magnetic field strength could be observed. Furthermore, we can observe variations of the seafloor topography as shown on the bathymetric map. Probably, a portion of the observed magnetic anomalies might be caused by some of those small and

already subducted topographic features which might influence the magnetic anomaly pattern as shown in line 6 to 9.

During SO163-2 only four magnetic profiles were collected, two during transit and two while shooting on the oceanic plate offshore Nicoya peninsula.

The profiles are shown in Figures 6.5.13 to 6.5.16.

Profile 13 is located off Nicoya Peninsula and crosses the Pacific Costa Rica Margin, striking NE - SW. The total field strength is very constant along the profile. Variations are very weak, both, on the margin and in the deep sea basin. Furthermore, no significant correlation between topography and total magnetic field strength can be observed.

Profile 14 and 15 are located off the northern part of Nicoya Peninsula. Profile 14 is coincident with the northern part of profile 15. Along the profile, the bathymetry is very flat and not influenced by e.g. seamounts or other topographic features. The total magnetic field strength ranges about 35500 nT in the northern part; in the southern part of magnetic line 15, a weak and negative magnetic variation can be observed.

Magnetic line 16 was recorded during transit towards Panama. The line starts near to Jaco Scar and stops at the Costa Rica - Panama frontier, off Burica Peninsula. The line, runs on the ocean plate and crosses several topographic elements like Cocos Ridge and several seamounts. The total magnetic field strength shows many variations along the profile. In the outermost east at 83.1° W, a smooth shaped magnetic anomaly was found which has no correspondence in the bathymetric profile and can be interpreted as a part of the magnetic anomaly pattern of the oceanic plate. The eastern part of the profile can be divided into two parts, west and east of latitude 84.0° W respectively.

In the western part, the bathymetric and magnetic profiles are trending nearly parallel; bathymetric highs are corresponding to regions of high magnetic field strength. The eastern portion of the profile shows the opposite tendency. Bathymetric highs are related to relative minimum zones of the magnetic field strength.

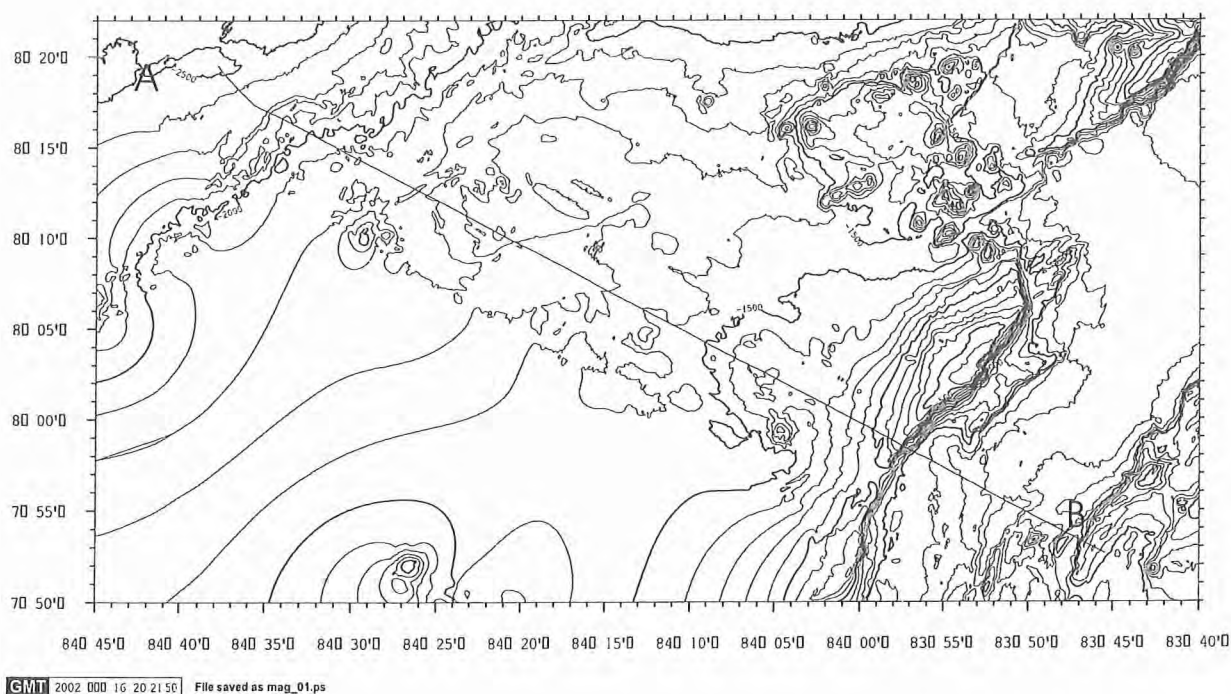
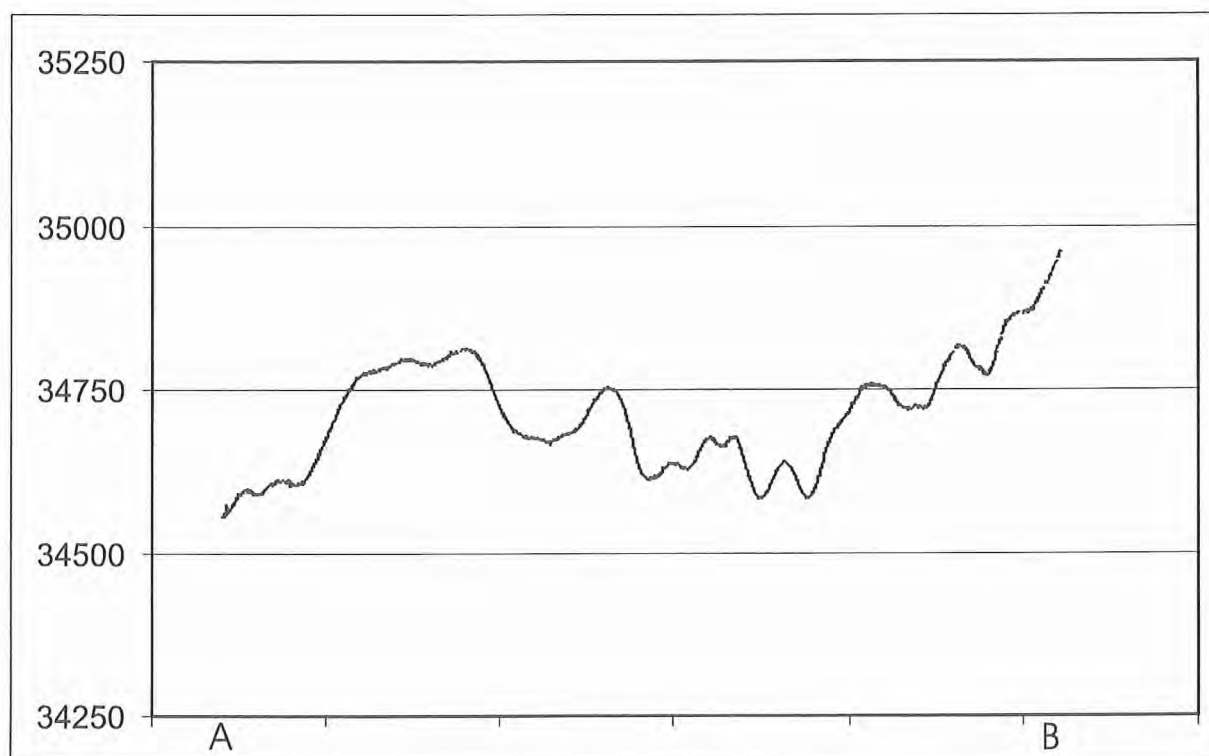
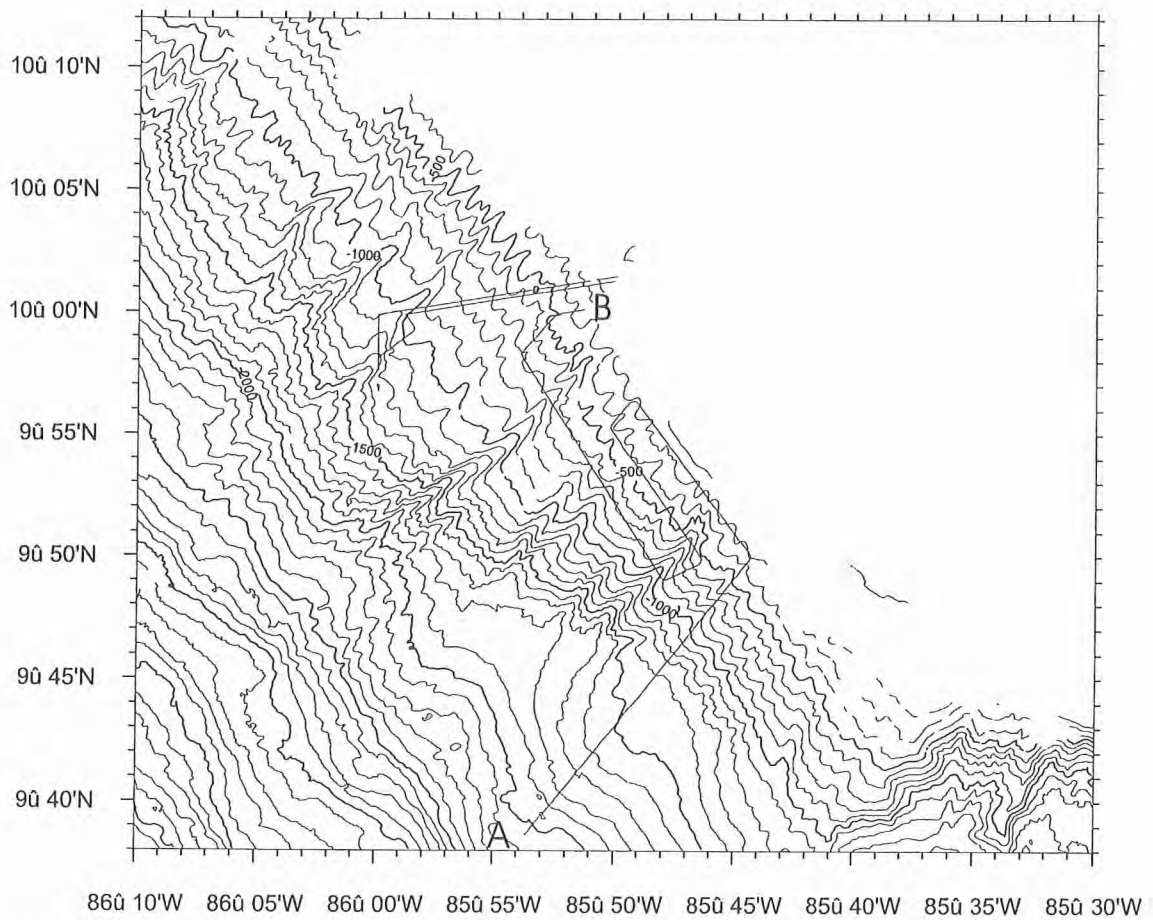
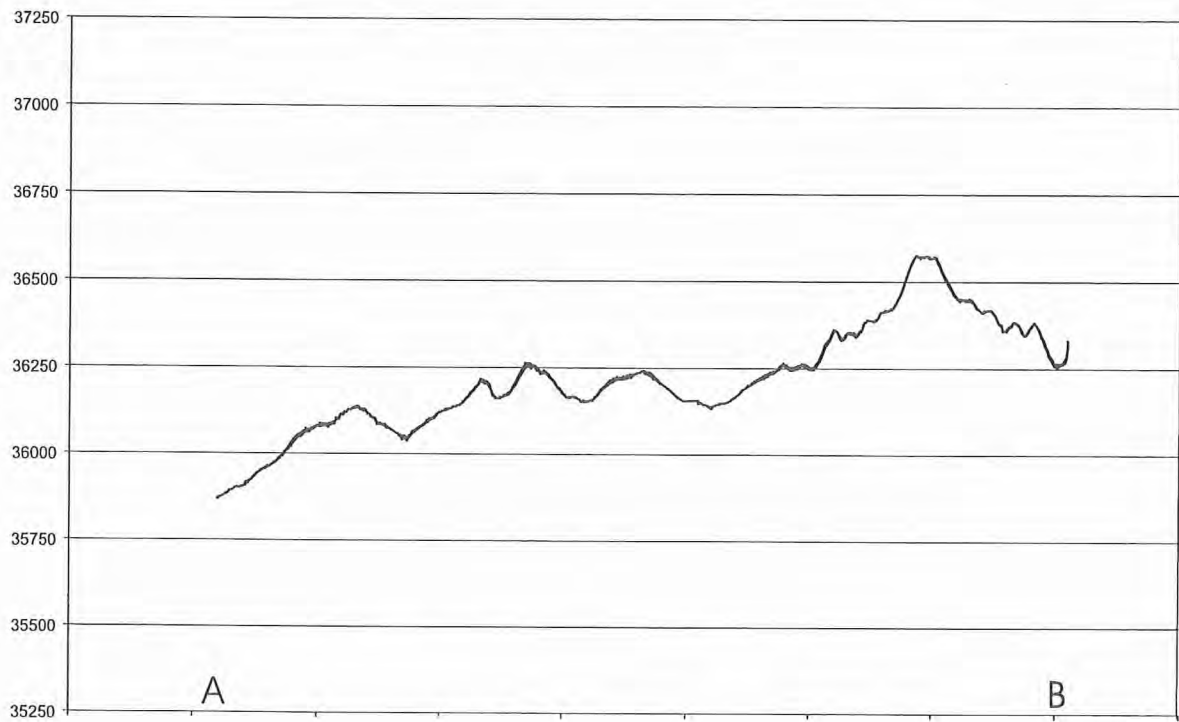
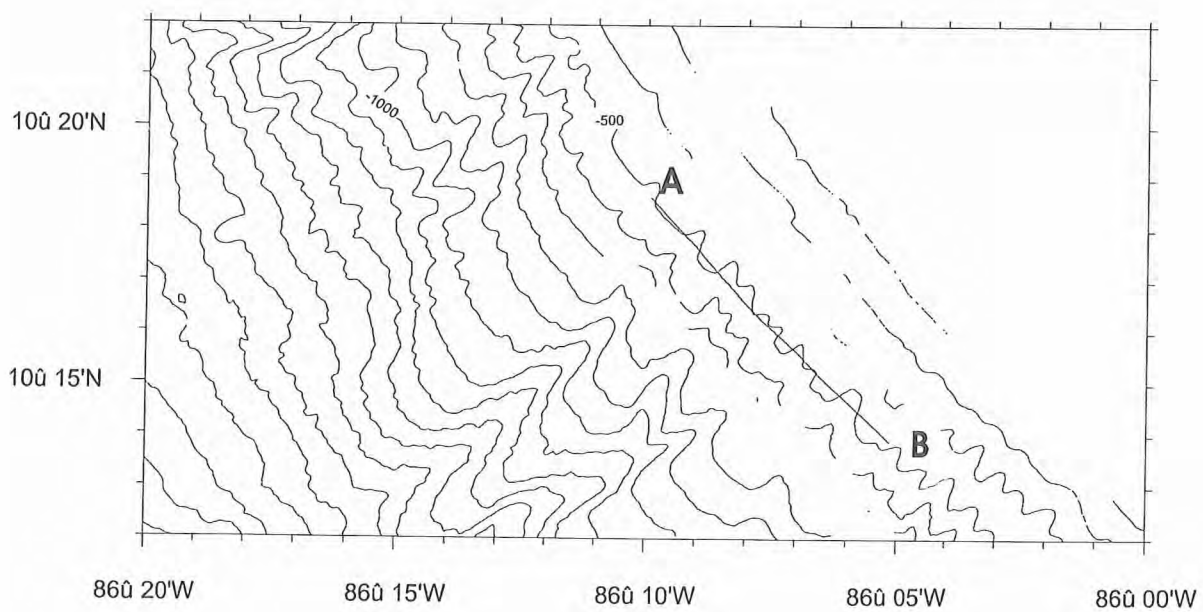
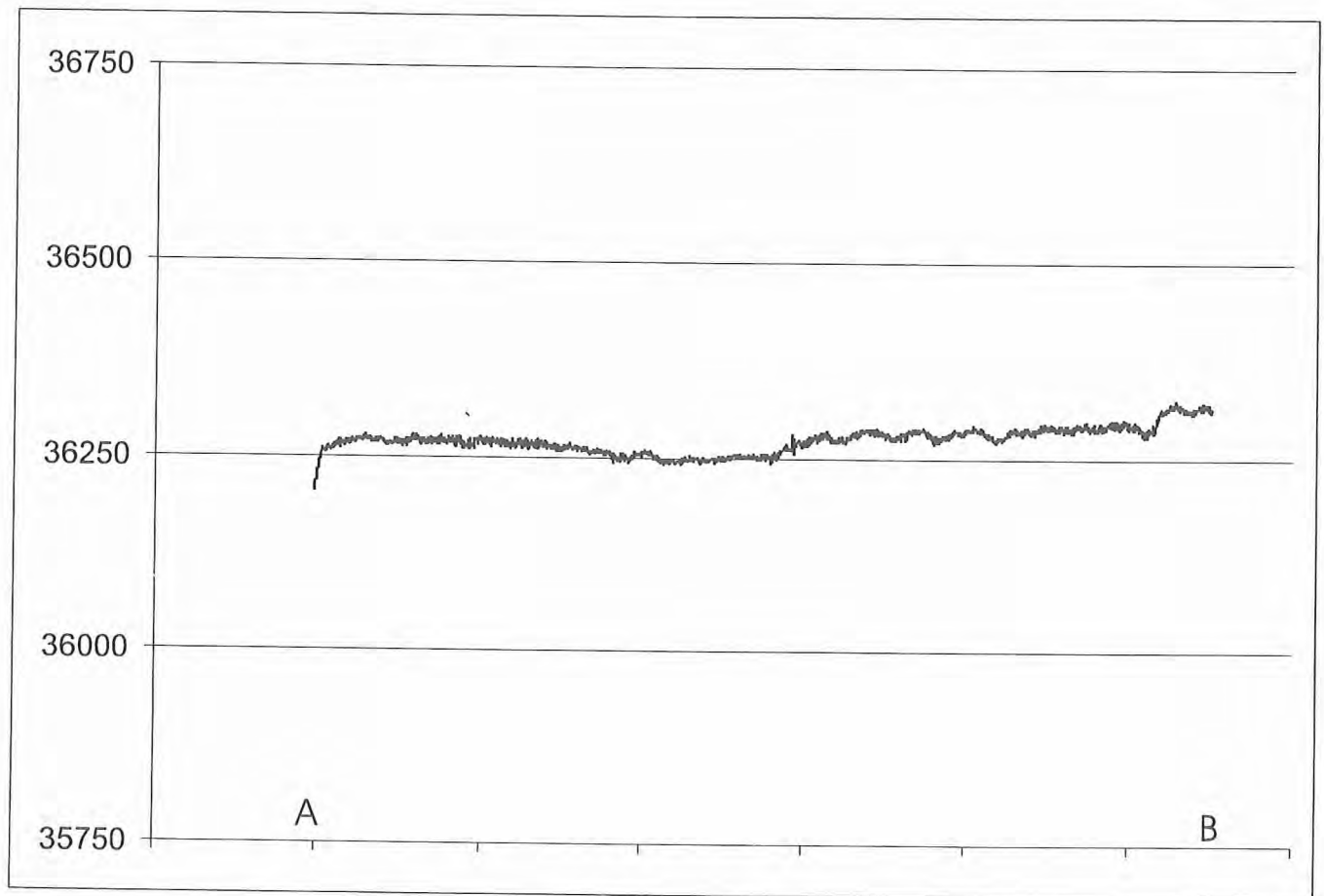


Figure 6.5.1:
Magnetic field strength [nT] and track line of magnetic profile 01, SO 163-1



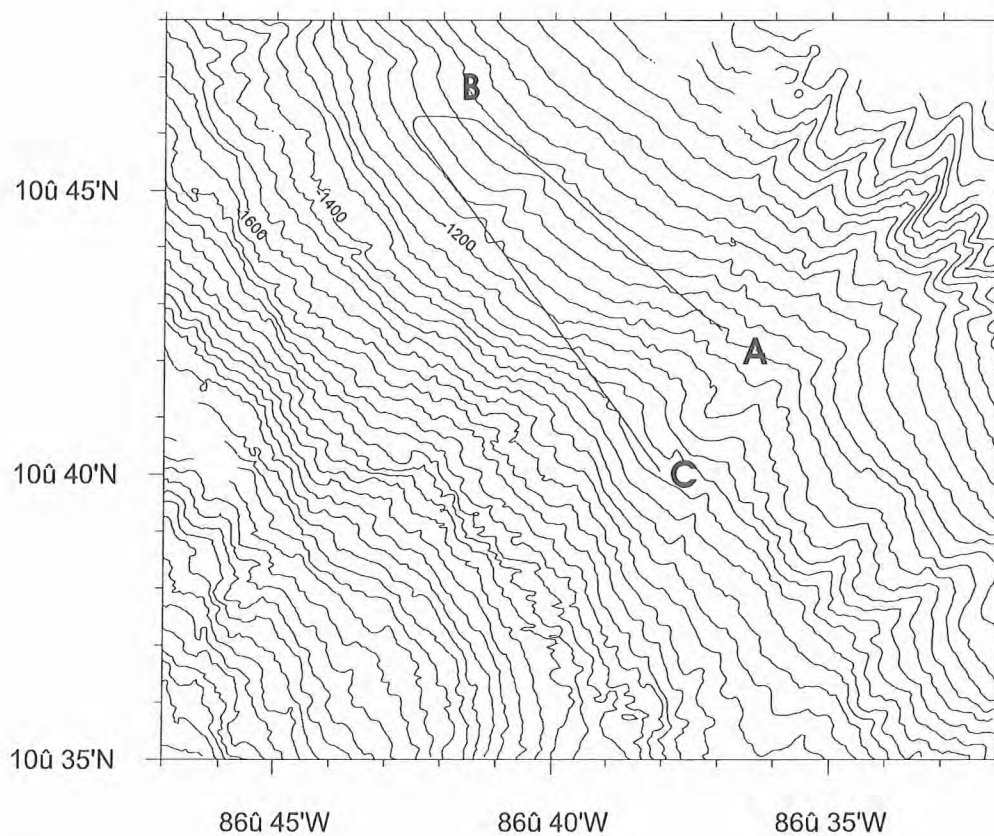
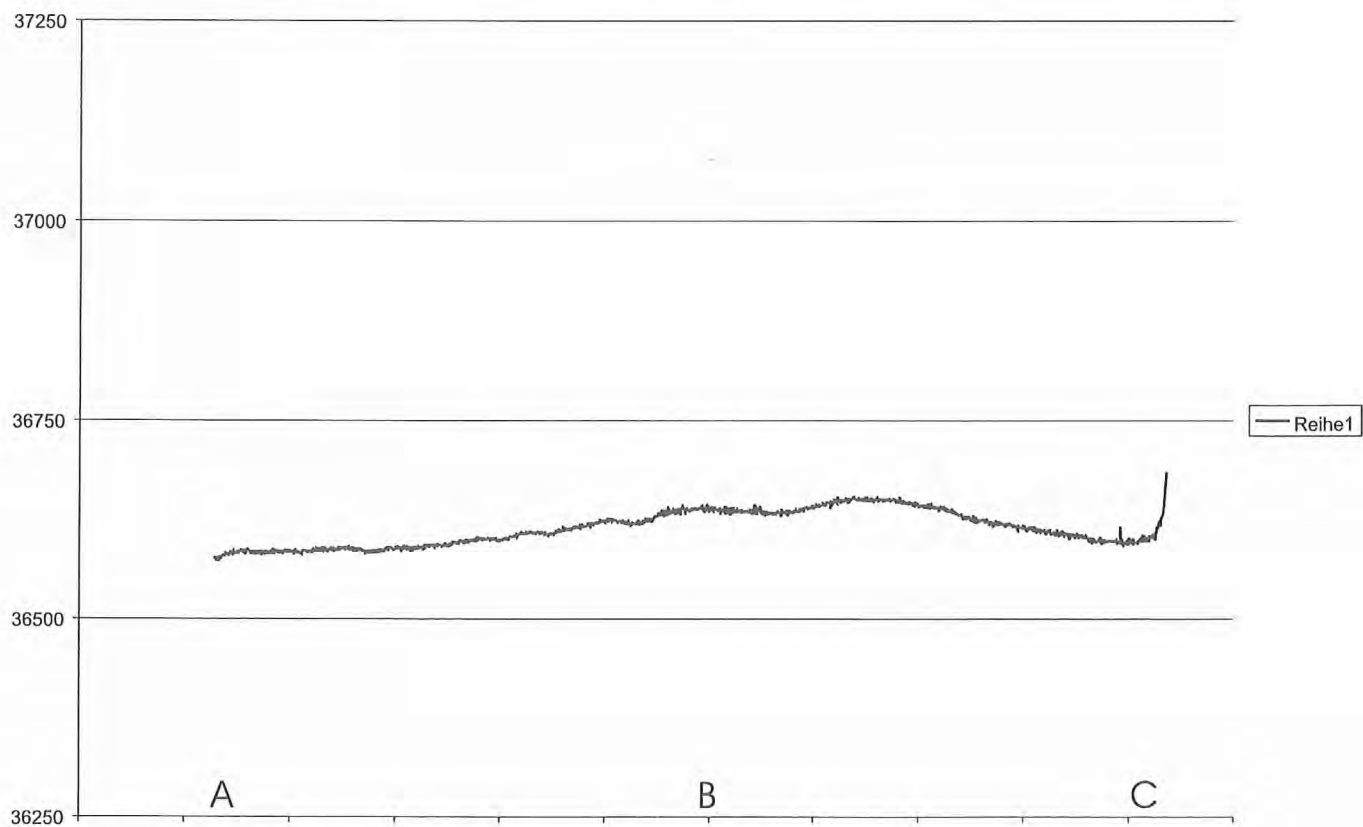
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Figure 6.5.2:
Magnetic field strength [nT] and track line of magnetic profile 02, SO 163-1



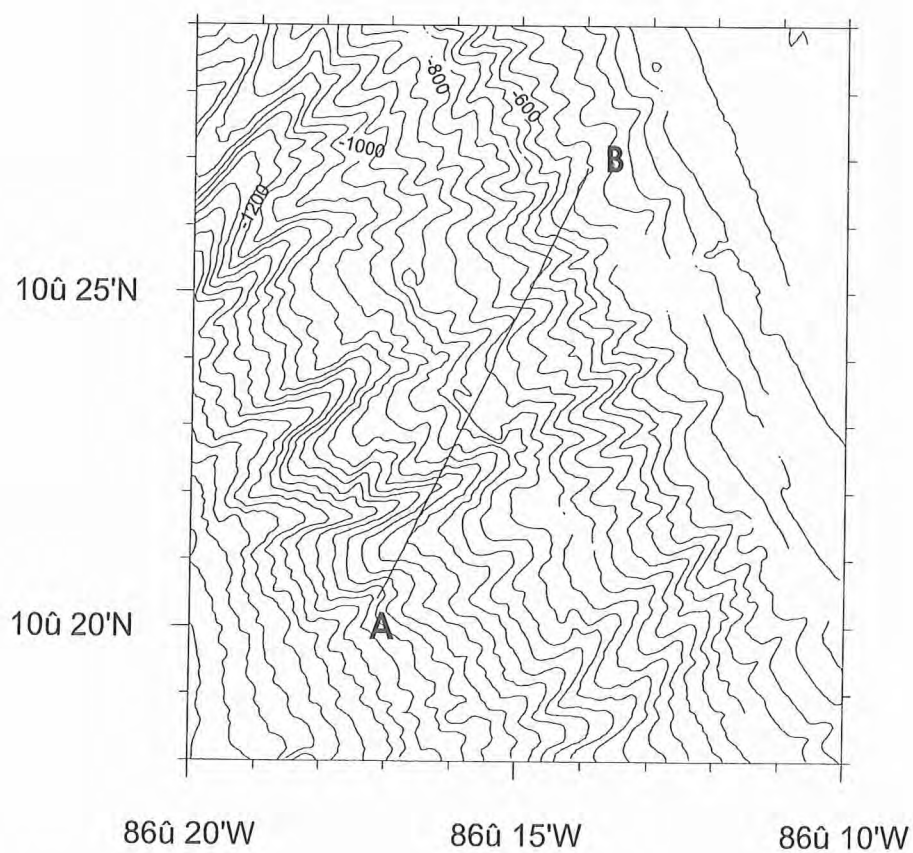
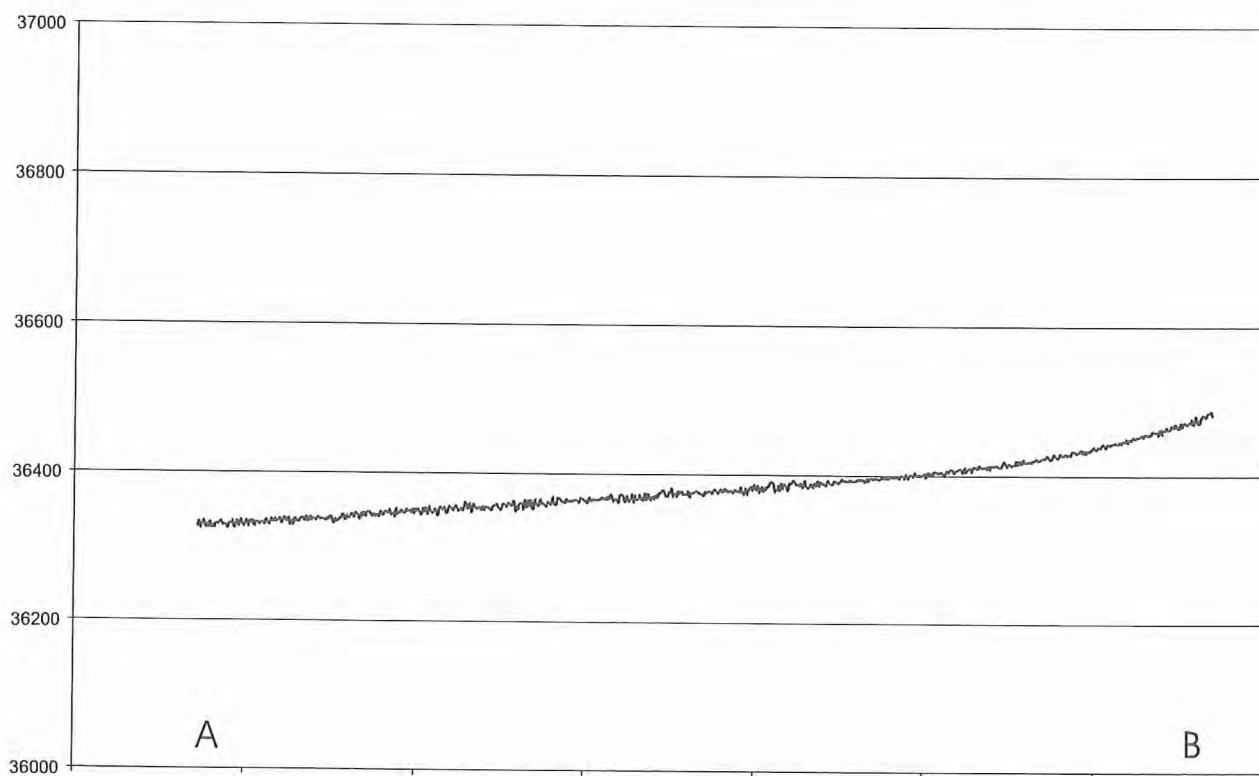
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Figure 6.5.3:
Magnetic field strength [nT] and track line of magnetic profile 03, SO 163-1



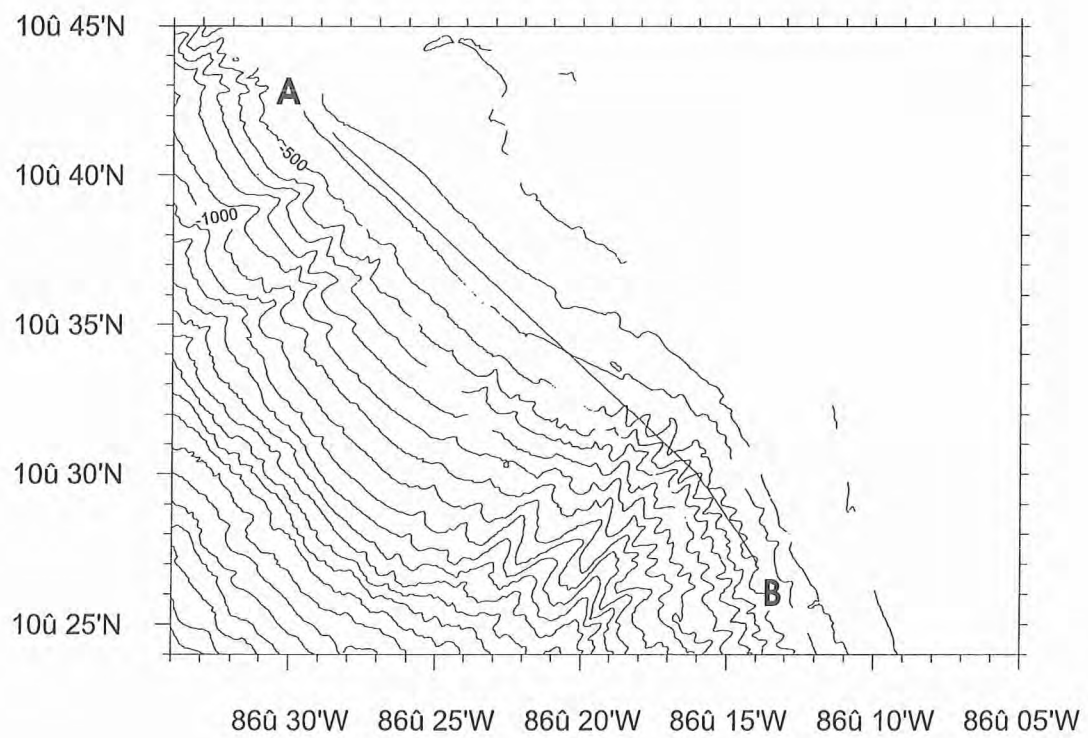
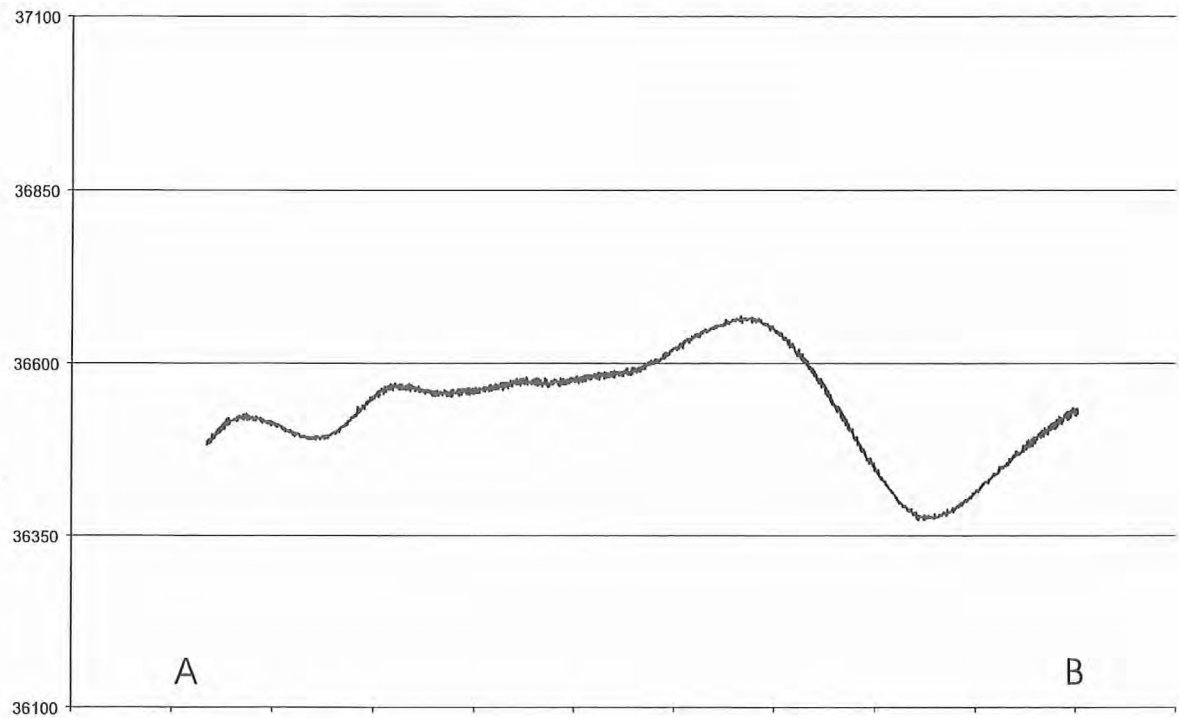
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Figure 6.5.4:
Magnetic field strength [nT] and track line of magnetic profile 04, SO 163-1



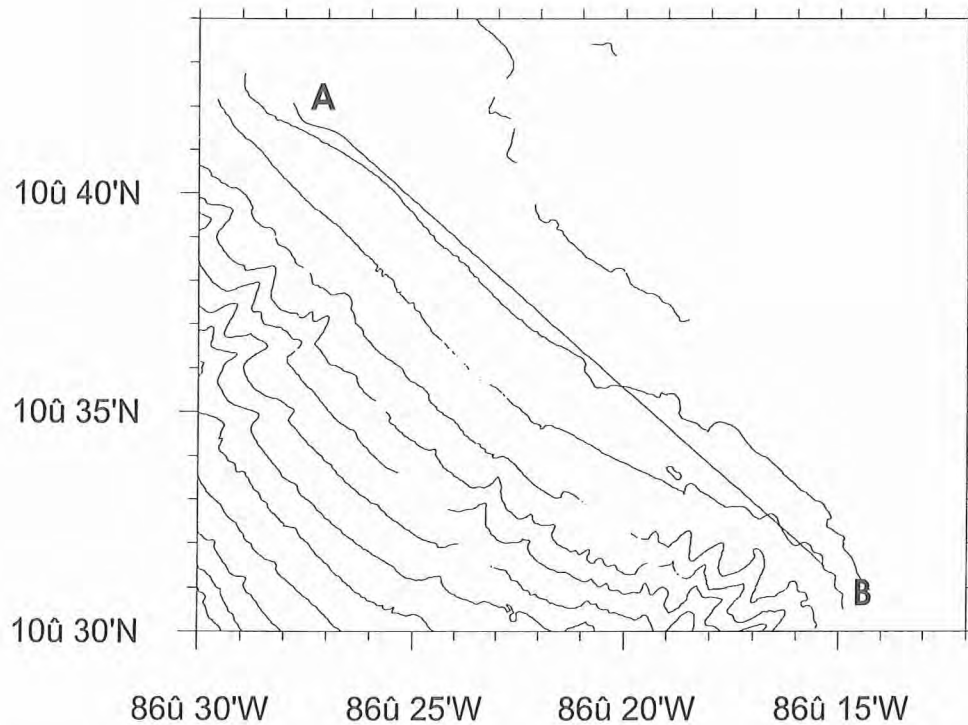
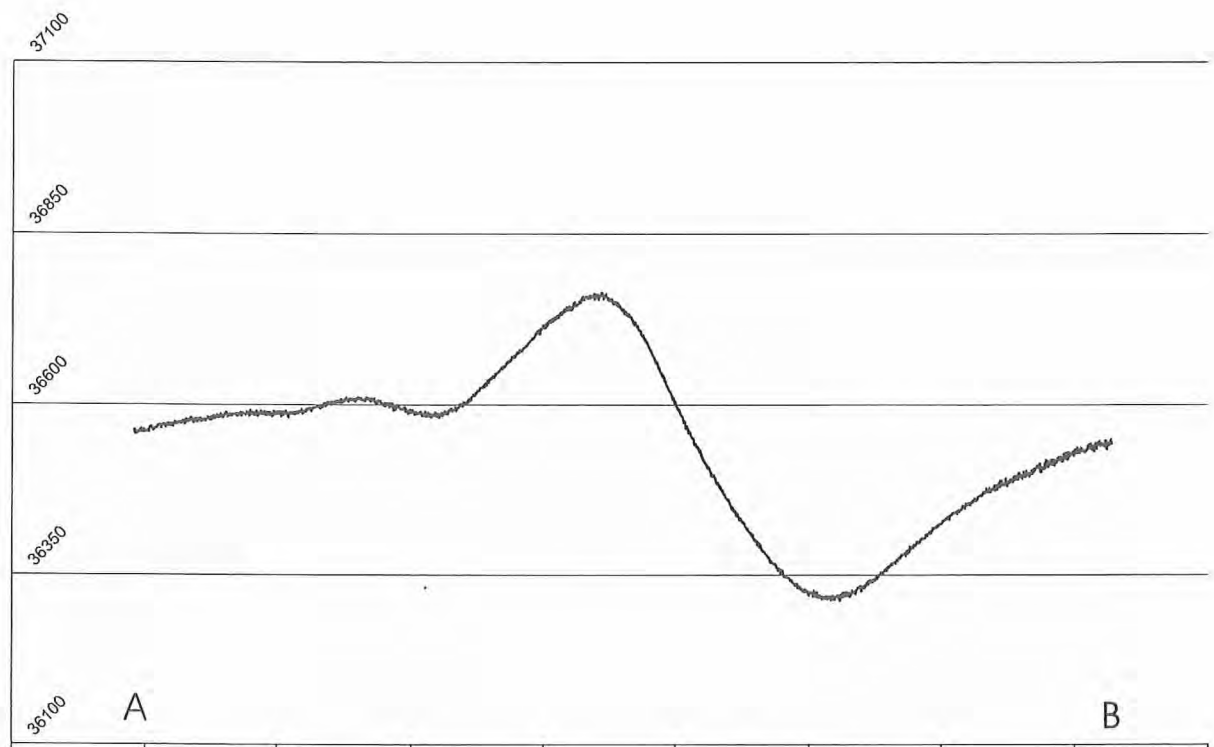
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Figure 6.5.5:
Magnetic field strength [nT] and track line of magnetic profile 05, SO 163-1



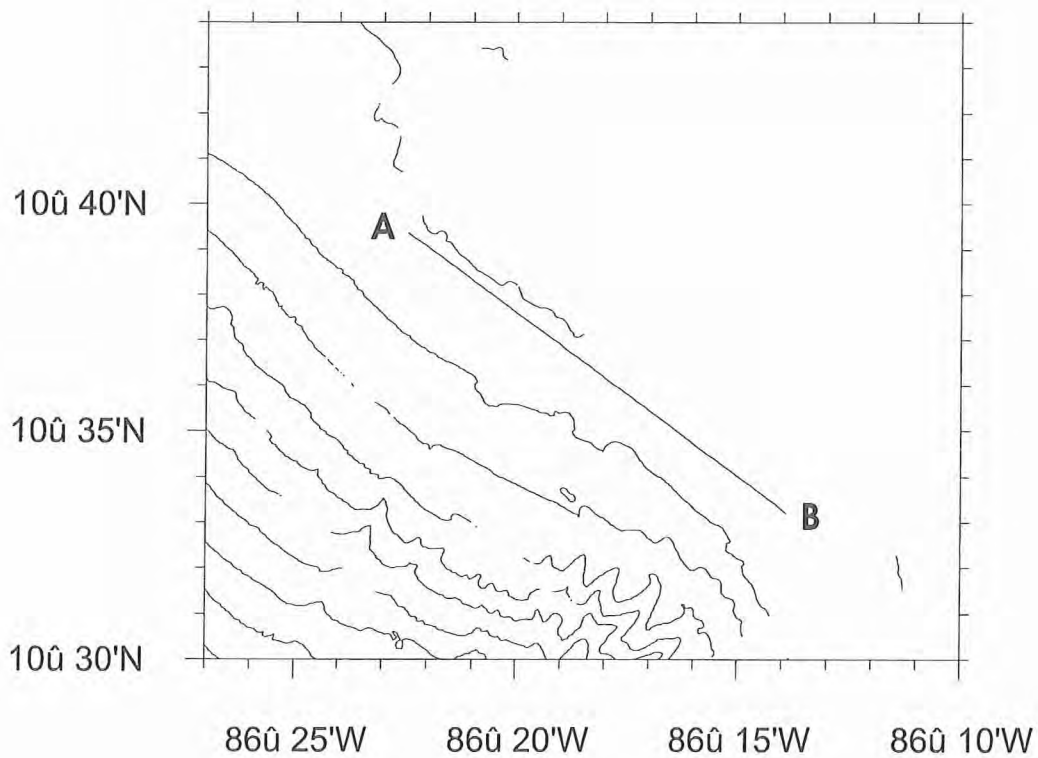
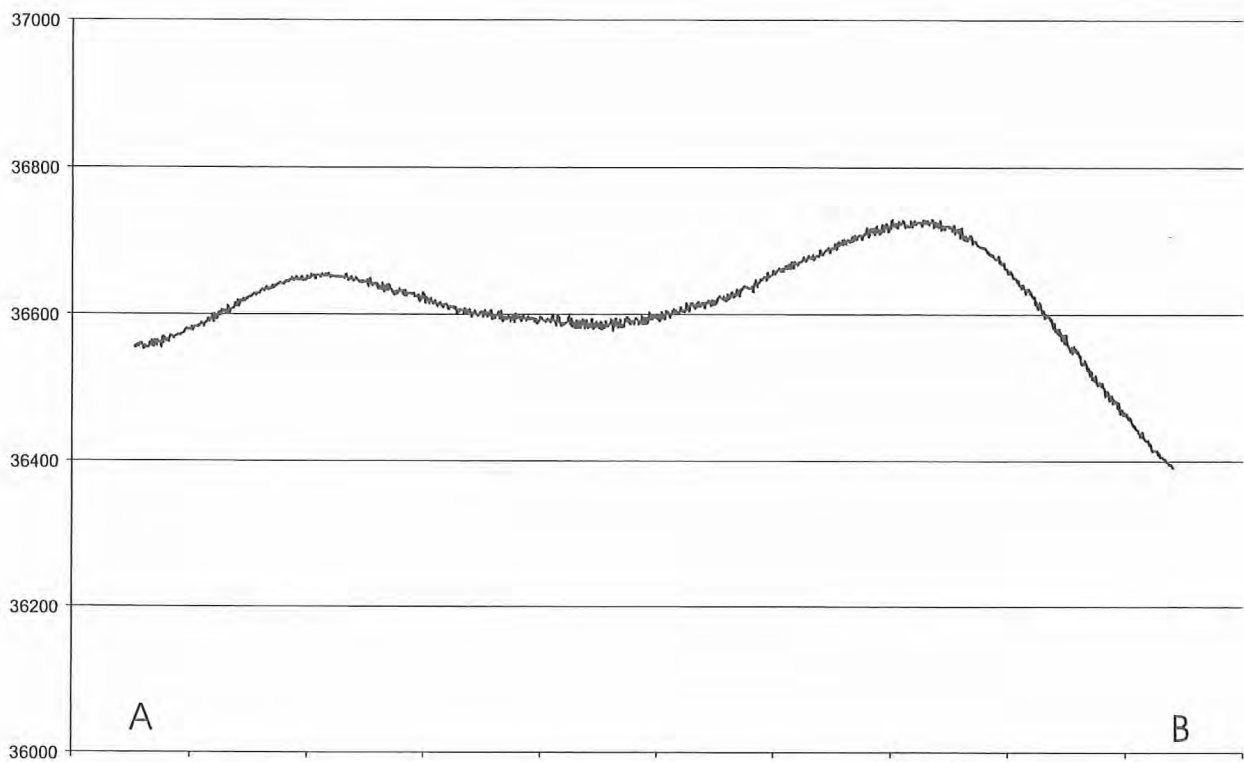
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Figure 6.5.6:
Magnetic field strength [nT] and track line of magnetic profile 06, SO 163-1



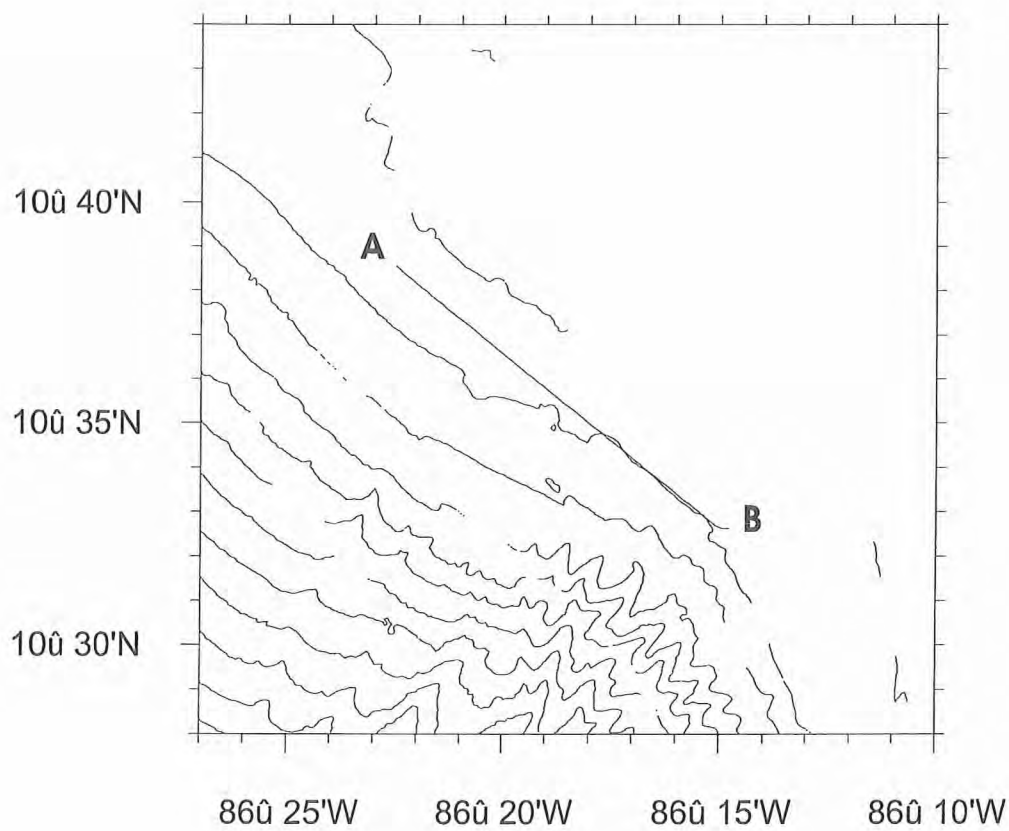
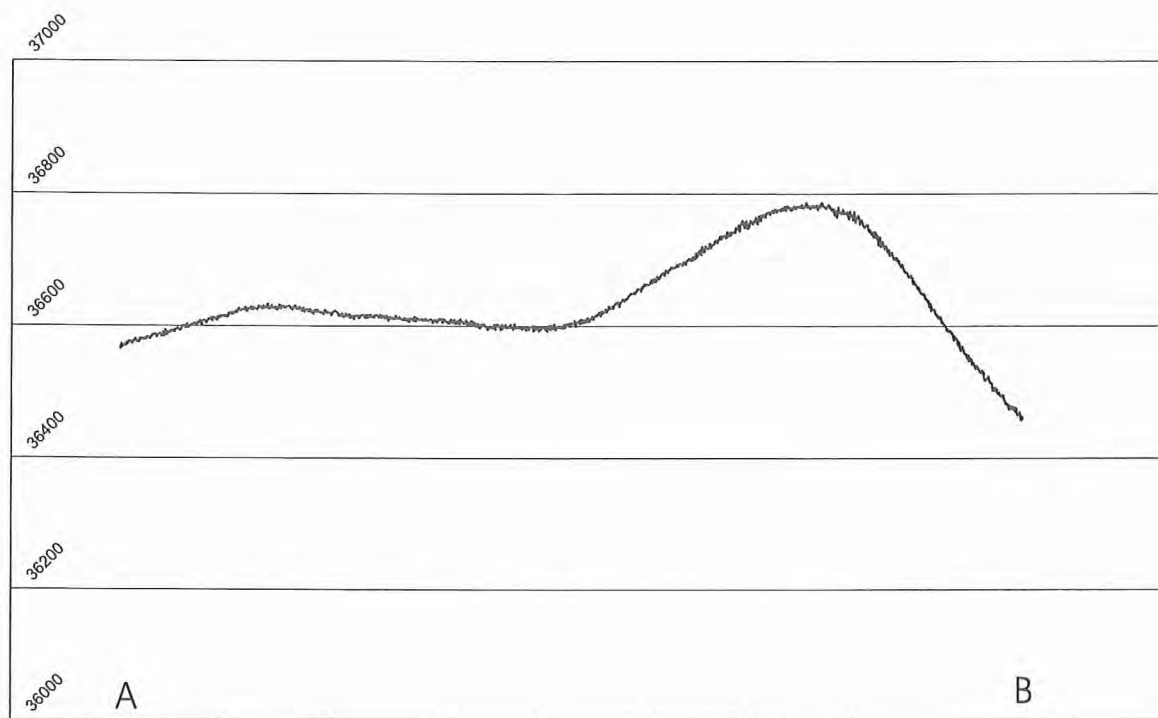
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Figure 6.5.7:
Magnetic field strength [nT] and track line of magnetic profile 07, SO 163-1



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Figure 6.5.8:
Magnetic field strength [nT] and track line of magnetic profile 08, SO 163-1
159



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Figure 6.5.9:
Magnetic field strength [nT] and track line of magnetic profile 09, SO 163-1

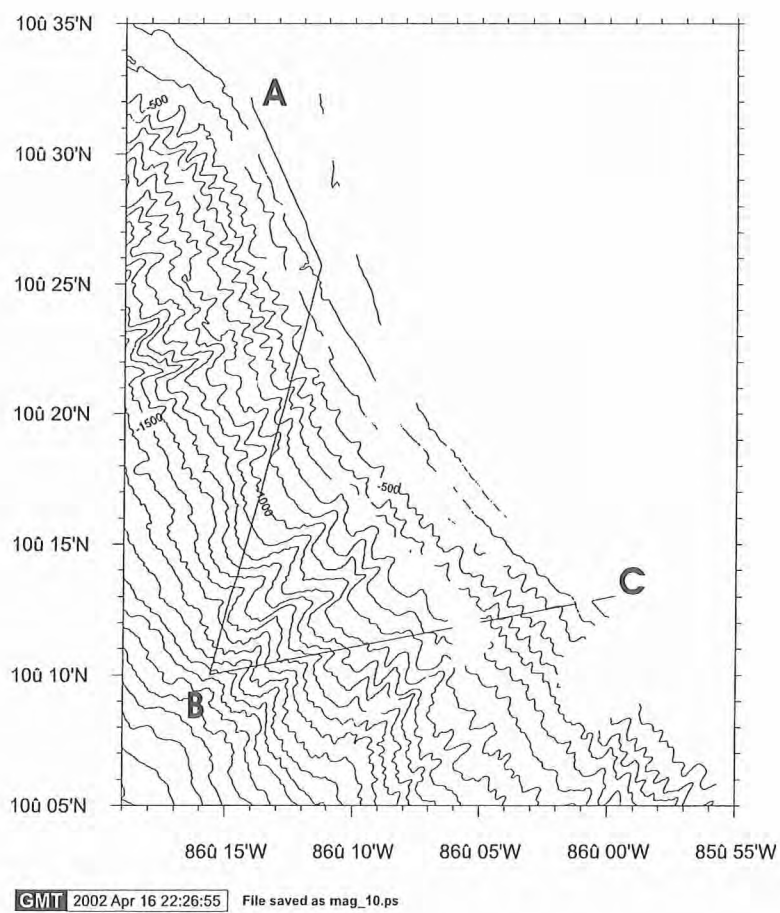
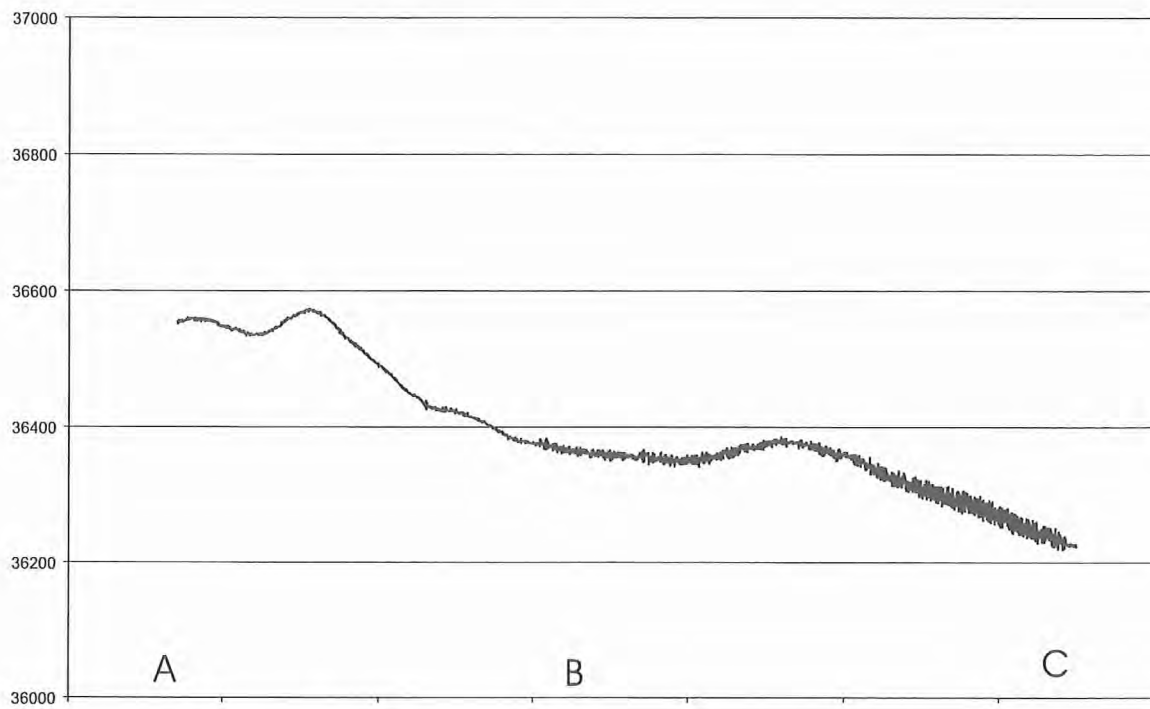


Figure 6.5.10:
Magnetic field strength [nT] and track line of magnetic profile 10, SO 163-1

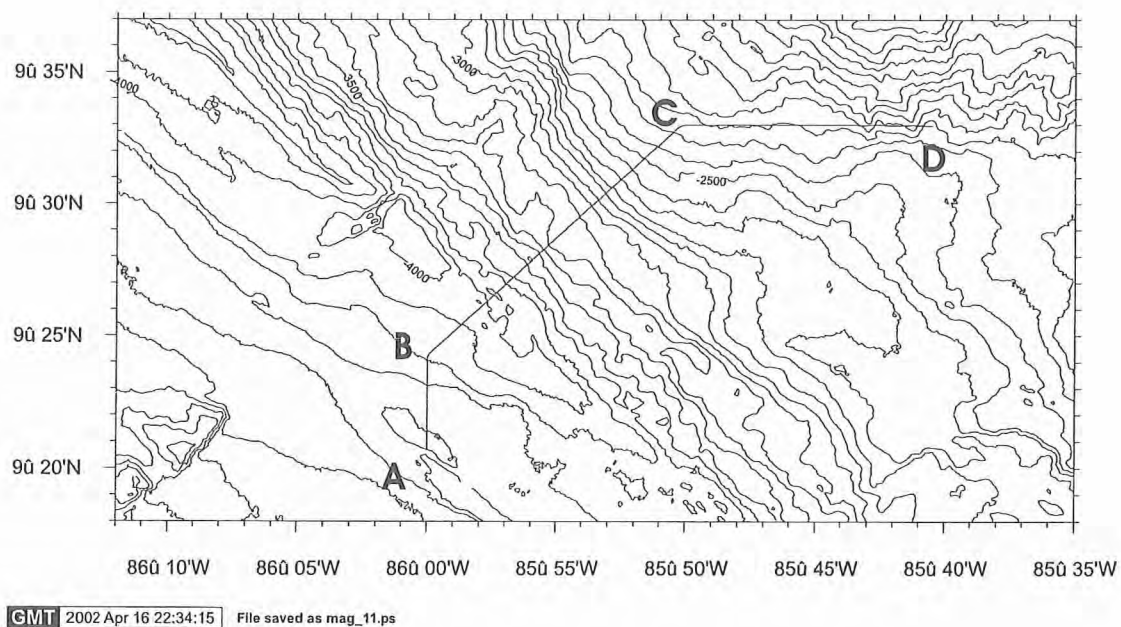
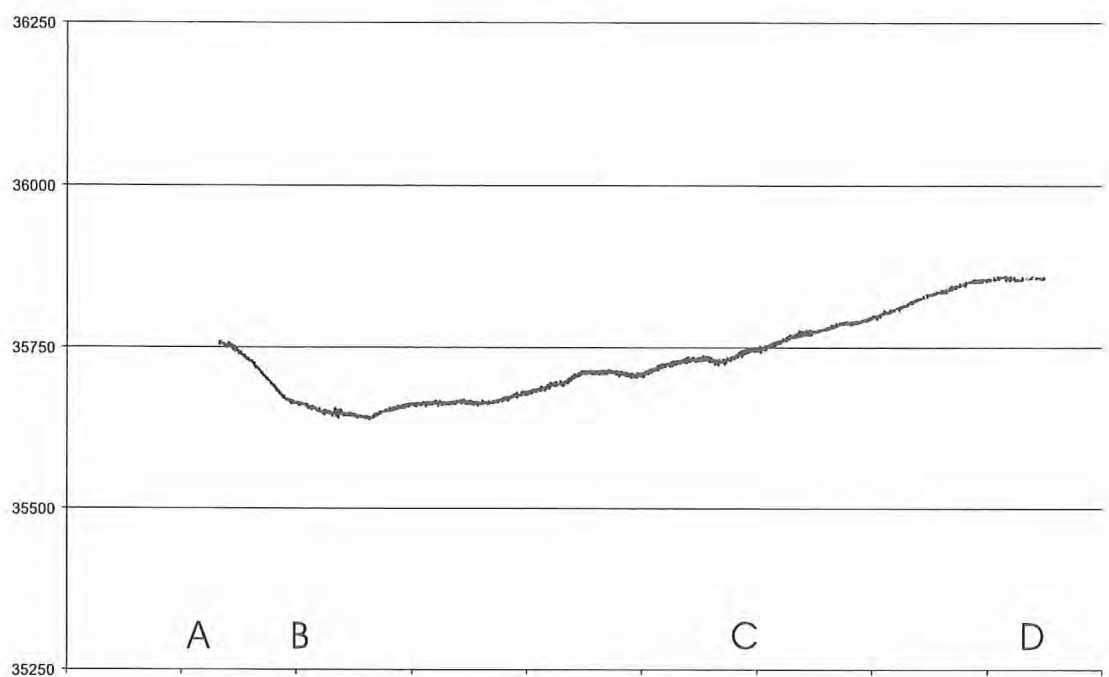
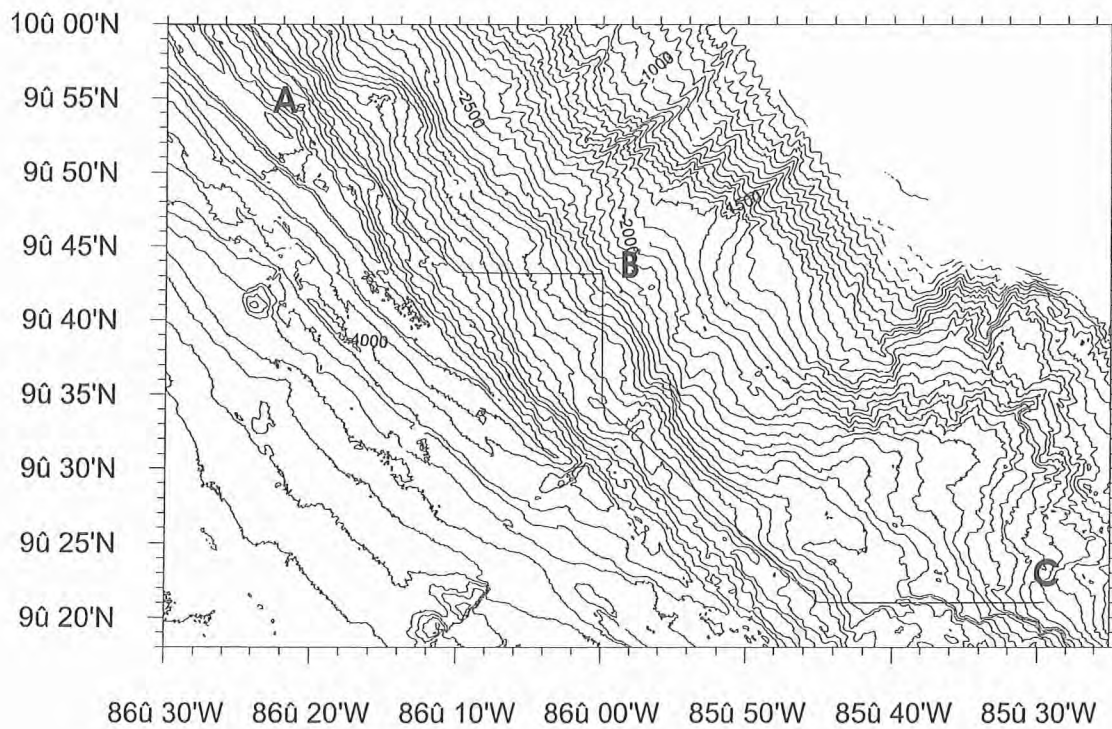
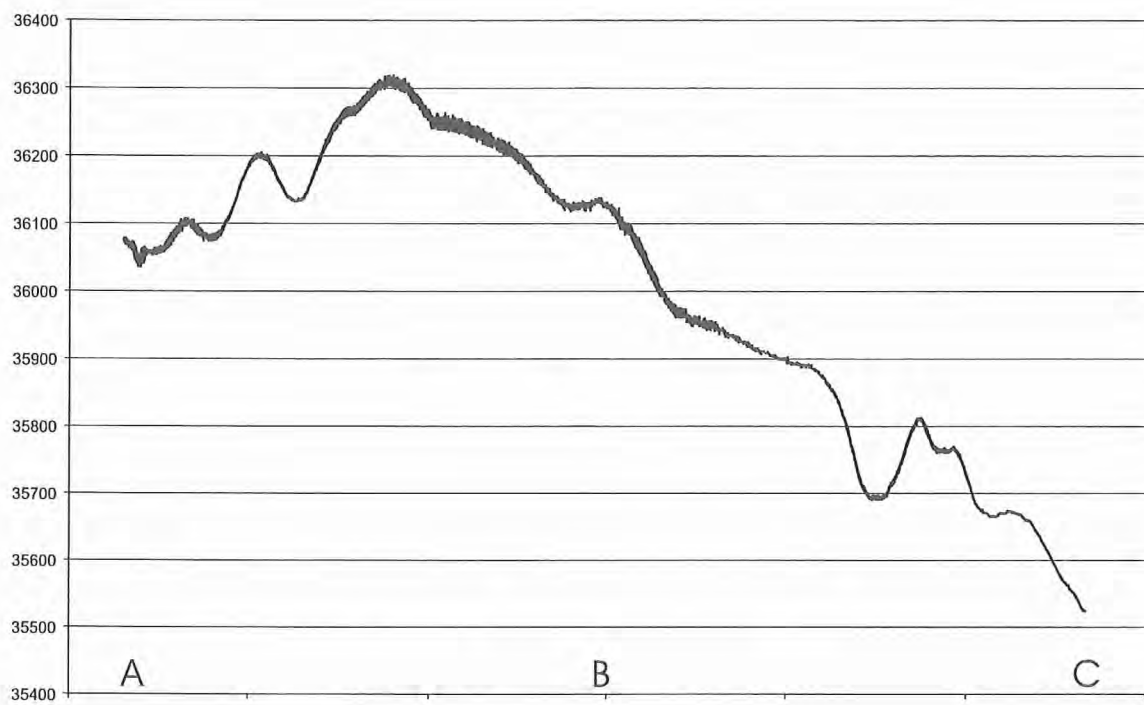


Figure 6.5.11:
Magnetic field strength [nT] and track line of magnetic profile 11, SO 163-1



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Figure 6.5.12:
Magnetic field strength [nT] and track line of magnetic profile 12, SO 163-1

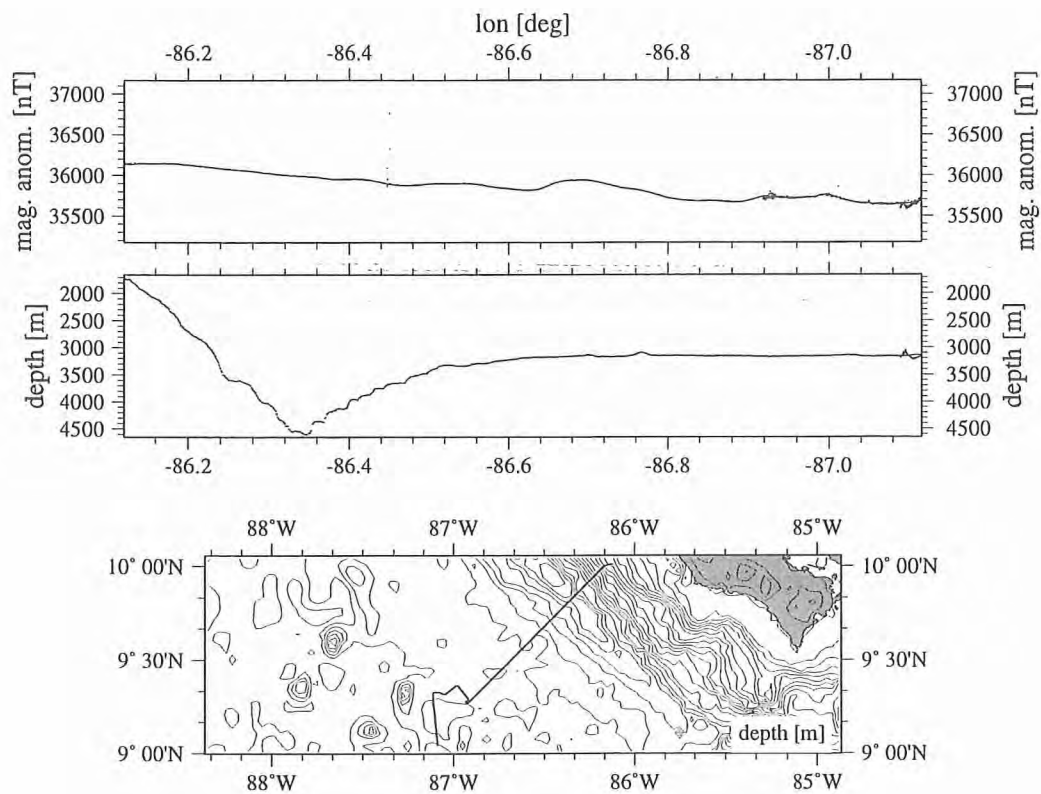


Figure 6.5.13: Magnetic field strength, bathymetry and track plot of magnetik profile SO 163, line 13

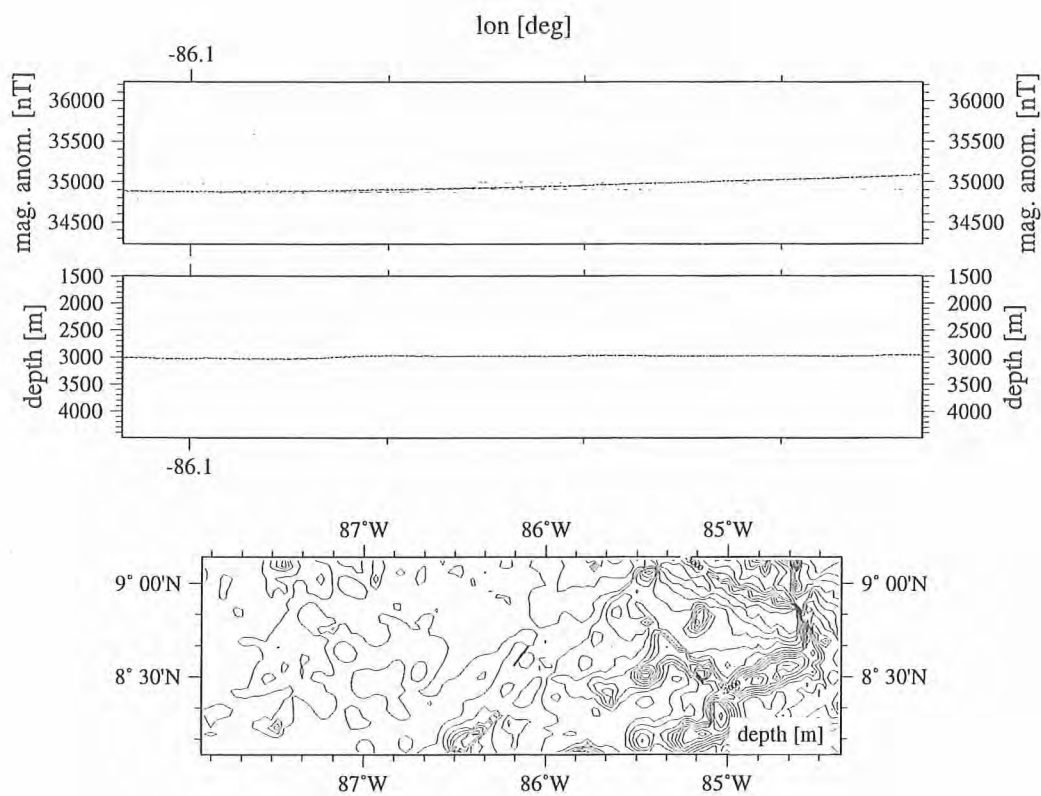


Figure 6.5.14: Magnetic field strength, bathymetry and track plot of magnetik profile SO 163, line 14

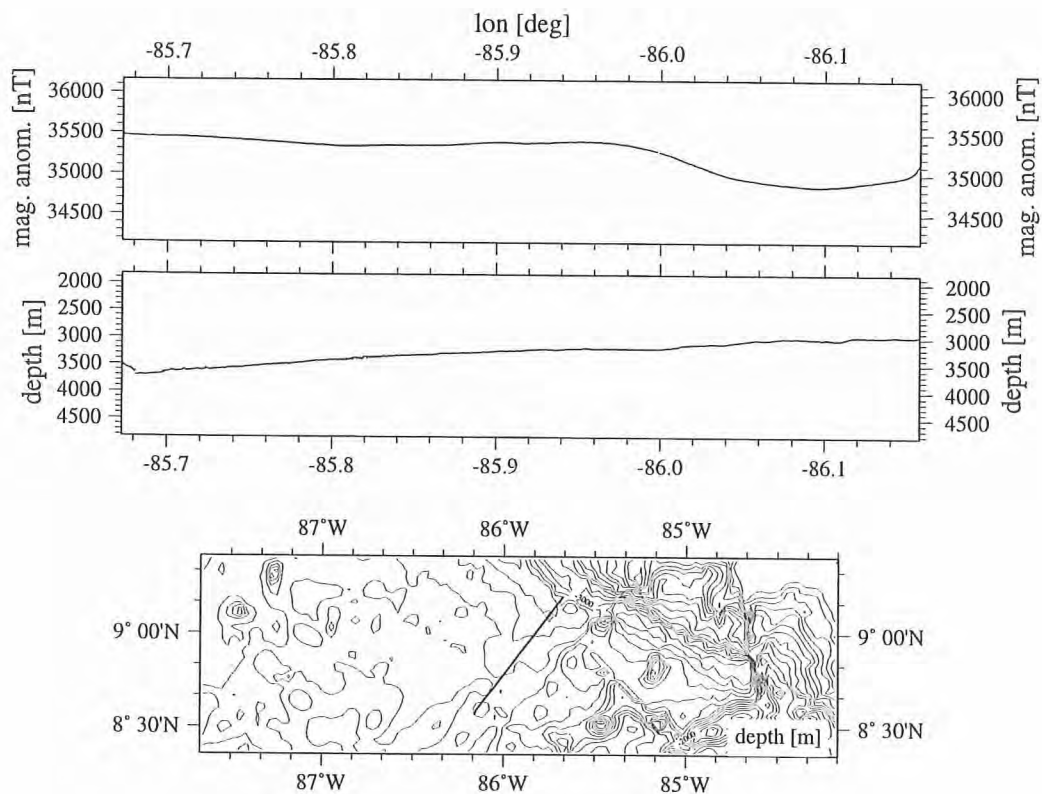


Figure 6.5.15: Magnetic field strength, bathymetry and track plot of magnetik profile SO 163, line 15

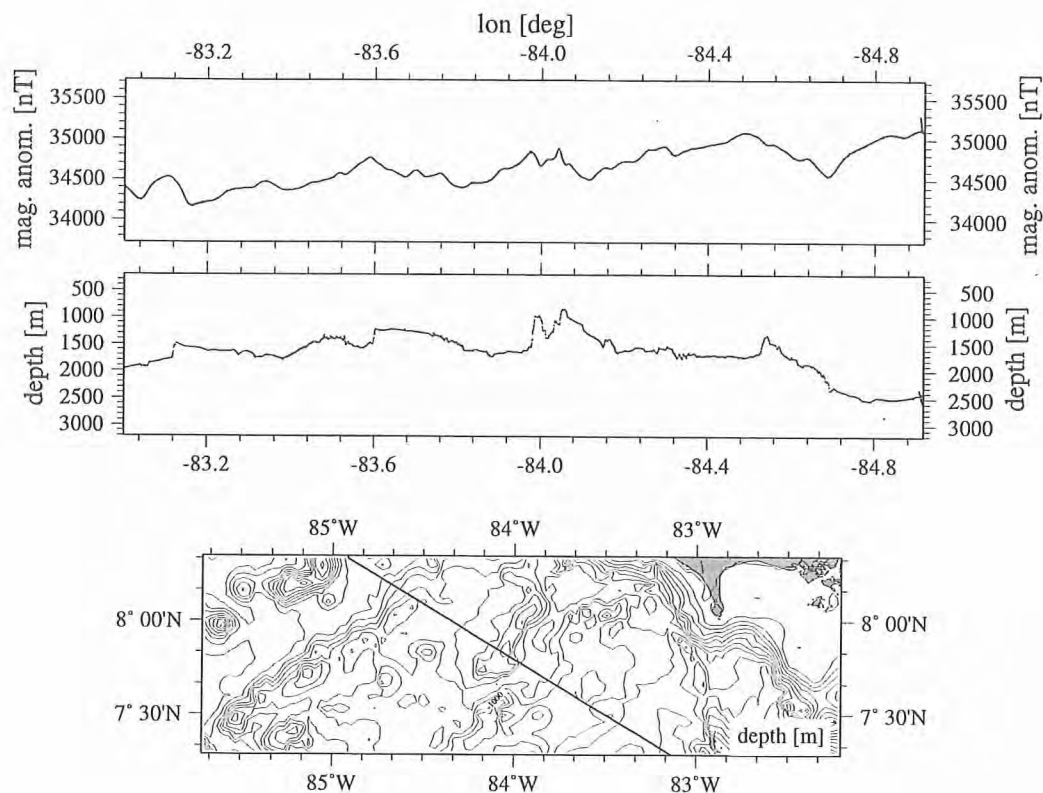


Figure 6.5.16: Magnetic field strength, bathymetry and track plot of magnetik profile SO 163, line 16

6.6 Seismic and seismological investigations

Introduction

During cruise SO163-2 numerous seismic profiles were collected, using ocean bottom recorders and a four channel streamer. Several different airguns were used. All instruments are described in chapter 5.3, details of the profiles can be found in Appendices 9.5 and 9.6. In addition, seismological investigations were made, including a local network above an active venting site, where 10 instruments were deployed during cruise SO163-1 and recorded for about one month (Chapter 6.6.1). A larger network was deployed centered on Jaco Scar, here a total of 28 instruments were deployed and shall be recovered during METEOR Leg M54/3B in early October 2002. For this network seven instruments were recovered and redeployed during the cruise (chapter 6.6.5).

The active seismic work focussed on three main topics. Several data set were collected to study the feasibility of modifying the receiver function method, that is well established in global seismology, for studies of shear wave velocities of gashydrates (chapter 6.6.2). The so called Megalense, a reflection identified on earlier MCS images, was a second target. Here, two dip and two strike lines were recorded by ocean bottom recorders and the streamer (chapter 6.6.3). The third target area was the decollement reflection, that in some places along the Costa Rican margin is especially strong and can be followed to the very tip of the upper plate. We choose the occurrence of the strongest reflection known from existing data and collected additional wide angle data that should provide the base for further lithological investigation using AVO and other attribute information (Chapter 6.6.4). Since during this cruise two G-Gun arrays were available (generously provided on loan from *Sodera Inc.*), several tests of the guns using different towing depths were made throughout the cruise, including the registration of the far field signal using hydrophones at large depth but away from the seafloor. The corresponding results are summarized in the chapters according to the area they were made.

Finally, a short profile was aquired across one of the carbonate mounds on the margin in order to detect the thickness of the carbonate layers (chapter 6.6.6). An overview map of all instruments and profiles is shown in Figure 4.2.2.

In the following, the data processing and archiving is described.

Seismic Processing: OBH/OBS Wide Angle Data and MCS Reflection Data

(A. Berhorst, N. Fekete, H. Kopp, L. Planert, M. Schnabel, and the GEOMAR wide-angle Working Group)

Data Conversion

The OBH/OBS data are recorded in continuous mode by the MLS and MBS units. Processing of the seismic and seismological data requires a prior conversion into pseudo SEG-Y format, which is conducted based on existing REFTEK routines modified for the instrument requirements and the available hardware platforms. Figure 6.6.1 illustrates the processing scheme applied to the raw data, of which a detailed description of the main tools is presented below.

send2pas

Data recording is performed on PC cards which store the compressed data file in the MLS and MBS recorders. After recovery of the instruments, the skew of the internal clock is determined and the flash cards are then transfered to a DOS/Windows based PC. The program send2pas decompresses the data stored on the cards. The data are converted to a PASSCAL data format in a 16 or 32 bit storage and written to the hard disk as well as copied to a SUN workstation via ftp as all other programs are UNIX based.

Send2pas will also detect any possible time slip errors which may have occurred during the recording on MLS units. The time slips are caused by a mismatch of the actual sampling rate of the MLS recorder compared to the desired sampling rate. This mismatch arises because the clock rate of the crystal oscillator in the MLS recorder is temperature dependent (Klaus Schleisiek, SEND GmbH, pers. comm.). The temperature dependence is known and corrected for in the determination of the

system time, but for performance reasons the sampling pulses are directly generated from the oscillator signal without any time correction. The send2pas routine detects and reports in the log-file when the accumulated inaccuracies of the sample rate cause an effective timing error of one sample. For the wide-angle data, the timing errors are considered by adding them to the skew of the internal clock, which is corrected within the dat2segy program (see below, note that the sign of skew and time slips are opposite, i.e., a negative sum of time slips corresponds to a positive skew). As each trace is at most a few tens of seconds long (vs. 1 hour and more between non-cancelling time slips), the corrected time is expected to be highly accurate with uncertainties well below one sample length.

ref2segy

The ref2segy program converts the output of send2pas to a pseudo SEG-Y trace consisting of one header and a continuous data trace containing all samples, as used by the PASSCAL suite of seismic utility programs. For each channel (consisting of pressure, vertical velocity, and velocity along two mutually perpendicular horizontal directions for OBS; pressure for OBH) one file is created with the name derived from the start time, the serial number of the Methusalem system, and the channel number. The file size of the pseudo-SEG-Y file is directly related to the recording time. For instance, a recording time of one hour sampled at 200 Hz (16 Bit) will produce a file size of 1.44 MB per channel. A record with two channels and a recording time of two days will produce a total data volume of 70 MB. After conversion to the pseudo SEG-Y trace, the seismic and seismological processing schemes diverge. The seismology data flow is described in the next section below.

pql

pql (Passcal Quick Look) is a simple display program for continuous seismic data. Its interactive zooming capability allows a rapid inspection of data quality.

segy2trig

The trigger signal, provided by the airgun control system, is recorded on an additional MBS unit during the shooting period. The trigger data are treated similarly to regular seismic data and downloaded to the hard disk via the send2pas and ref2segy programs. Then, the segy2trig program detects the shot times in the data stream by identifying the trigger signal through a given slope steepness, duration and threshold of the trigger pulse. The output is an ASCII table consisting of the shot number and the shot time. Accuracy of the shot time is one of the most crucial matters in seismic wide-angle work, and must be reproduced with a precision of a few ms. Due to this demand the shot times have to be corrected with the shift of the internal recorder clock. Additionally, the trigger file contains the profile number, the start/end time of the profile and the trigger recording, which will be merged with the ukooa file that links the information with the coordinates of the source and the hydrophones.

ukooa

The ukooa program is used to establish the geometric database by calculating the positions of sources at any given shot time and offset from the ship. The source is placed on the ship track using simple degree/meter conversions and then written to a file in UKOOA-P84/1 format. Corrections for offsets between antenna and airguns as well as consistency checks are included. This file will be used when creating a SEG-Y section via the dat2segy program. The program requires the trigger file to contain the shot times, the ship's navigation (see Chapter 5.1.1), and a parameter file containing information for the UKOOA file header.

dat2segy

The dat2segy program produces standard SEG-Y records either in a 16 or 32 bit integer format by cutting the single pseudo SEG-Y trace (the ref2segy output) into traces with a defined time length based on the geometry and shooting time information in the ukooa file. In addition, the user may set several parameters for controlling the output. These parameters contain information about the profile and the receiver station, number of shots to be used, trace length, time offset of the trace and reduction velocity (to determine the time of the first sample within a record). The clock drift of the recorder (skew) and possible time slip errors are also considered. The final SEG-Y format consists of the file header followed by the traces. Each trace is built up by a trace header followed by the data

samples. The output of the dat2segy program may be used as input for further processing with e.g. SEISMOS or Seismic Unix (SU).

relobs

Due to the drifting of the OBH and OBS instruments during deployment and possible errors in the ship's DGPS navigation system, the instrument positions may be mislocated by up to several 100 m. Since this error leads to asymmetry and incorrect travel time information in the record section, the instrument's position on the seafloor must be relocated, which is accomplished using the relobs program.

For input, the assumed OBH location, shot locations, and the picked travel times of the direct wave near to its apex are required. To simplify the picking procedure, a static correction with a hyperbolic equation is performed to flatten the direct wave. This yields a much more coherent direct arrival which would normally suffer from strong spatial aliasing in the uncorrected section making it difficult to track. By shifting the OBH position, relobs minimizes the deviation between computed and real travel times using a least mean square fitting algorithm (assuming a constant water velocity).

The distance from stern to the center of the airgun array was measured using a floating line. The source offset, i.e. the distance from the vessel's GPS antenna position was determined to be approximately 80 m, which is in good agreement with earlier calculations from previous cruises on RV SONNE.

syntrigger

In case of a corrupted trigger recording, it is necessary to generate a synthetic trigger signal to merge the shot times with the navigation coordinates in the ukooa file. This may be done using the syntrigger program which uses the shot interval as well as the time of the first and last shot of the profile to calculate a synthetic trigger slope. Due to the fact that the trigger has a small but linear drift, even non-integer trigger intervals (59.99983 s instead of 60 s) can be handled. The drift has to be calculated from neighbouring profiles.

• OBH/OBS-data analysis and processing

Raw data: As an example which was acquired with the Bolt air-gun, the raw OBH record section 28 for profile 09 is shown in Figure 6.6.2a. A broadband frequency filter is applied for display purpose. For the analysis, offset ranges between 0-13 km and 26-38 km west are presented in detail. For comparison, the seismic section of OBH 59 on profile 15, which was shot with the two G-guns, is presented in Figure 6.6.2b. The offset ranges from 0-9 km and 31-42 km east are used for the analysis. The parameter tests for the Prakla airgun-array, which was used for only 3 profiles, are not presented, but the applied processing sequence is displayed (see below).

Frequency filter analysis: To determine the frequencies of the seismic energy, filter panels with narrow frequency band passes for the near- and far offset ranges are shown in Figures 6.6.3 to 6.6.6. In the lower section of the figures the amplitude spectra of the corresponding filter panels are appended. The amplitude spectra of the applied Ormsby frequency filter operators are characterized by linear slopes. The filter applied, which is minimum delay, is described by four corner frequencies: Lower stop/pass band boundary and upper pass/stop band boundary. The frequencies on the filter panels correspond to the lower and upper pass frequencies. The main energy for the reflected phases in the near offset range is between 3-25 Hz for the Bolt-gun and 3-37 Hz for the G-guns while the direct waves show signals up to 110 Hz. The main energy of the far offset phases for both air-guns is between 3-16 Hz. As a broad frequency range is contained in the data, time and offset dependent filtering was applied (see below).

Deconvolution analysis: To improve the temporal resolution of the seismic data a deconvolution is applied to compress the basic seismic wavelet. The recorded wavelet has many components, including the source signature, recording filter, and hydrophone/geophone response. Ideally, deconvolution should compress the wavelet components and leave only the earth's reflectivity in the seismic trace. We applied Wiener deconvolution in successive trace segments, based on the following assumptions:

1. The earth's reflectivity is 'white'.
2. The wavelet shows the minimum-delay phase behavior.

As in this wide-angle data the amplitude spectra of the seismic traces vary with time and offset (e.g. reflected and refracted phases), the deconvolution must be able to follow these time and offset variations. Each trace is therefore divided into 3 s data gates with 1 s overlap, in which time invariant deconvolution operators are computed from the autocorrelation function of the data segment and applied to account for the non-stationarity of the seismic signals. The overall deconvolved trace results from a weighted merging of the independently deconvolved gates.

The raw data serves as input to the deconvolution process. As several recordings were influenced by a DC shift, a 0-3 Hz high-pass minimum delay Kaiser frequency filter with 60 dB attenuation between the pass and reject zone was applied prior to deconvolution in order to center the amplitudes around zero. The deconvolution test panels are shown in Figures 6.6.7 to 6.6.10 for the near and far offset ranges, respectively. In the lower section of the Figures the autocorrelation function is appended. Constant operator lengths of 200 and 300 ms (predictive length included) and a variation of the predictive length from 4 (spike) to 160 ms is displayed.

On the undeconvolved data of the Bolt-gun in Figures 6.6.7 and 6.6.8, strong energy is clearly visible in the autocorrelation function. The best resolution is obtained for a predictive length of 4 ms and an operator length of 300 ms, with a reduction of signal-to-noise ratio especially on the far offset traces. A predictive length of 100 ms and an operator length of 200 ms was chosen for this data set which is a compromise between temporal resolution and signal-to-noise ratio.

The autocorrelation function of the raw data from the G-guns in Figures 6.6.9 and 6.6.10 show much weaker energy behind zero lag than the Bolt-gun. Therefore the wavelet could be much better compressed. A predictive length of 60 ms and an operator length of 200 ms was finally selected.

After deconvolution an offset- and time-variant Ormsby filter with minimum delay characteristic was applied. As the seafloor depth changes along the seismic lines, each trace was statically corrected to a fixed seafloor travel time of 11 s based on the water depth before filtering. This information is available in the trace headers. After this filter was applied, the data were shifted back to their original travel times.

Processed data: Comparison of the data in Figure 6.6.11 to the unprocessed data in Figure 6.6.2a (Bolt-gun) and Figure 6.6.12 to Figure 6.6.2b for the G-guns respectively shows a clear reduction of the low and mono-frequency noise in the near and far offset traces and moderate (Bolt-gun) to good (G-gun) compression of the wavelet signal. For the picking of events and model building by raytracing both sections were used to keep all available seismic information.

• Final processing sequence

- Input: SEG-Y-data, 4 ms or 5 ms sampling rate with complete geometry information.
- Tapering the first 0.5 s to zero to reduce the response of the debias filter operator.
- Kaiser highpass (debias).
- Gated Wiener deconvolution: gate length 3 s, overlap 1 s, length of merge region 1 s,

Bolt-gun:	operator length 200 ms	prediction length 100 ms
G-gun:	operator length 200 ms	prediction length 60 ms
Prakla-array:	operator length 300 ms	prediction length 40 ms
- Static correction to a fixed seafloor traveltime of 11 s.
- Time and offset-dependent Ormsby frequency filter.

On time-shifted traces with a reduced time scale of 6 km/s the following filter parameters were used:

lower stop/pass	upper pass/stop [Hz]				
Bolt-gun					
G-gun					
Prakla-array		offset[m]	beginful[s]	endful[s]	
3/5	40/60				
3/5	57/77				
3/5	50/70	0	0	12.8	
		8000	0	12.6	
		48000	0	0	
3/5	30/40				
3/5	38/53				
3/5	30/50	0	13.7	14.3	
		8800	13.5	14.4	
		13200	13.0	13.9	
		52000	1.0	2.0	
		107000	0	0	
3/5	20/30				
3/5	22/35				
3/5	20/30	0	15.3	16.8	
		11700	15.1	16.6	
		19200	14.8	16.3	
		61700	7.0	10.1	
		114000	2.0	3.0	
		152000	0	0	
3/5	10/20				
3/5	14/23				
3/5	10/20	0	19.0	trace length	
		20000	18.4	trace length	
		130000	3.5	trace length	

• Data exchange

For the exchange of the OBH/OBS data, the SEG-Y-format on disk with a Sun tar-format was chosen. The raw segy data is in Integer2 format with trailer bytes between the record structure of SEG-Y. The processed data is in IBM-floating point without trailer bytes between the records.

For UTM transformation into Cartesian coordinates use: WGS84 spheroid, central meridian 85 0. 0. W, northern hemisphere.

This is the definition of the segy trace header for the GEOMAR OBS wide-angle reflection data. The extension of the standard SEG-Y header from 181 to 240 byte is a layout in order to process the data on the SEISMOS/SU software system. Reading bytes directly into this header will allow access to all of the fields.

• MCS processing

The airgun shots were also registered by a 4 channel mini-streamer. The group spacing is 50 m and the first offset was calculated from the direct wave to approximate 130 m. Due to the different shot intervals from 10 to 60 s (\cong 10 to 120 m) and the variable data quality, the processing sequence was adjusted for each profile after a check of the raw data. During the first 14 profiles all 4 channels could be used for processing, while only 2 channels recorded data of good quality for profile 15 to 28 after which we continued with only one channel. The bad traces are degraded by high amplitude noise and a low signal to noise ratio. Problems with streamer depth variations between first and last channel were nearly eliminated by attaching a buoy at the tail of the streamer.

A standard processing job (-) was created with some variable modules (*), depending on the profile parameters and the data quality.

- Input: SEG-Y-data, 2 ms or 4 ms sampling rate with complete geometry information.
- * Resampling from 2 ms to 4 ms.
- Kaiser highpass (debias).
- Static correction to a fixed seafloor travel time of 10 s.
- Time-variant gather-oriented predictive deconvolution: always 5 traces of a common receiver gather are used to calculate the operator for the centre trace, gate length 3 s,

- overlap 1 s, length of merge region 1 s.
- | | | |
|---------------|------------------------|----------------------------|
| Bolt-gun: | operator length 200 ms | prediction interval 100 ms |
| G-guns: | operator length 200 ms | prediction interval 40 ms |
| Prakla-array: | operator length 200 ms | prediction interval 32 ms |
- Time dependent Ormsby frequency filter.

	depth interval below seafloor [s]	lower stop/pass[Hz]	upper pass/stop [Hz]
Bolt-gun/	0 – 0.1	3/5	65/85
G-guns:	0.5 – 0.9	3/5	45/65
	1.3 – 1.9	3/5	30/45
	2.3 – 2.8	3/5	20/30
	3.6 – trace length	3/5	15/23
Prakla-array:	0 – 0.1	5/8	70/90
	0.5 – 0.9	5/8	50/70
	1.3 – 1.9	5/8	35/50
	2.3 – 2.8	5/8	25/40
	3.6 – trace length	5/8	15/23

- * Whole trace normalization for varying amplitude levels in the different channels.
- * Removing of high amplitude noise by comparing trace amplitudes with the average amplitude of the surrounding traces.
- Seafloor mute.
- Remove static correction.
- Binning of traces, depending on the desired output spacing.
- Dip dependent trace interpolation of empty bins.
- Finite difference time migration with 1500 m/s.
- Automatic gain control.
- Output: SEG-Y-data.

For profile 22 the raw plot of one channel is presented in Figure 6.6.13 with a broad bandpass filter (5 to 60 Hz) applied. In the processed section (two channels) in Figure 6.6.14 the reverberations of strong reflectors are removed and the resolution is remarkably improved. The same seismic line was also shot with the Bolt-gun (Figure 6.6.15) and the Prakla airgun-array (Figure 6.6.16). The bolt gun reveals two strong reverberations at around 200 ms and 400 ms, while the Prakla airgun-array shows a sequence of reverberations of up to more than 300 ms.

The MCS profiles are displayed in the individual chapters of the wide-angle profiles.

Processing of earthquake data (EQNET and JACO NET)

(I. Arroyo, J. Gossler, P.O. Thierer)

The initial processing of the data is identical to the processing sequence for wide angle data described in the previous section, i.e., reading of the flashcards, conversion into a REFTEK file, and conversion into a pseudo SEG-Y file. Following this step, the processing sequence of the earthquake data diverges from the one for wide angle data.

The single traces are cut into records of 25 hours length with one hour overlap between adjacent records, such that each record begins at 0:00:01 (except on the first day of recording, of course). The reason for this cutting is (a) to make file sizes smaller and hence the data more manageable, and (b) to be able to apply the time corrections on a daily basis. Routine for cutting: split_seg.y.pl

The timing of each 25 hour section of the data is corrected. Two timing errors need to be taken account of:

a) the occurrence of time slips (extra or missing samples) due to a mismatch of the desired and actual sample rates (see previous section for a more detailed discussion of time slips). The approach for correcting time slips differs from the approach used for the wide angle data because seismological analysis needs to refer back to the continuous data stream throughout. The occurrence time of all time

slips is read out from the send2pas log file. The time slips before the beginning of the record are added up and applied as static shift to the start time of the record. For time slips during the record a sample is added (positive time slips) or removed (negative time slips) at the appropriate time. This procedure is only approximate (a correct treatment would require resampling of the whole record, fraught with its own difficulties), and in general the apparent time of a sample can be off by up to half a sample length due to this approximation. Relative times within the same record can be off by up to one sample length if the time span straddles a time slip. Routine for time slip correction: `unslip`

b) the slow drift of the internal clock relative to GPS time. The system time is compared with GPS time at the beginning and end of the deployment, and a drift rate is inferred from the observed difference (skew). The time of each 25 hour record is corrected by applying the shift appropriate for the time 12.5 hours after the beginning of the record. The underlying assumption is that the system clock does indeed drift linearly, and that the drift over a 24 hour period is negligible, i.e. much less than one sample length, which was the case in this experiment. For some stations that ran out of power prematurely, no skew could be determined and therefore no drift correction carried out. Timing uncertainties for these stations accumulate at a rate 1.5-2 ms per day, starting with the first day of deployment. Routine for linear clock drift correction: `clock_cor.pl`

The data for each day are quality controlled and a table of station status prepared for each station (see Tables 6.6.1.1 and 6.6.5.1).

A short-term-average versus long-term-average (STA/LTA) trigger algorithm is then applied to the data to find out seismic events. Trigger parameters like length of the short term (s) and long term (l) time windows, the mean removal window length (m), the trigger (t) and dettrigger ratio (d), minimum number of stations (S) and at least the time window length (M) in which trigger signals contribute to the same event must be found individually for each data set. We take $s = 0.5$ sec, $l = 60$ sec, $m = 500$ sec, $t = 3.5$, $d = 0.8$, $S = 3$, $M = 10$ sec for the EQNET small scale network, where the data are very noisy and most event locations are very close to the network, so that the waveforms are coherent and have higher frequencies than the noise, so that is possible to use a relatively high trigger ratio. For the JACO network, where the noise is reduced drastically compared to EQNET we decreased the trigger ratio to 2.8, and increased M to 20 sec because of the larger spatial dimension of the network. Data has been checked by eye view also, and it seems that the selected trigger values delivered good results. The so found events are cut from the time-corrected data, and stored in sub-directories, one per event. The events have been quality checked again. Bad triggers have been sorted out. Events are also inspected for traces, which are unusable for a particular event, e.g. because they contain only noise. Routine for trigger and event cutting: `trig_all.sh`, which uses `refrig` program from the PASSCAL tool kit and `collate_ev.pl`, respectively.

The pseudo-SEG Y traces in the event directories are converted first into SAC, and then into SEISAN waveform format, which makes it possible to store all traces associated with an event in one file. After the conversion the data are registered into the SEISAN database (Havskov and Ottemöller, 2001). Routines for conversion: `segy2sac_all.csh`, `sacsei.sh`, `seisei.sh`

At last, P and where possible S phases are picked and events preliminarily located with the program HYP, which employ an iterative solution to the nonlinear localisation problem (Lienert, 1994). For the sake of simplicity of onboard processing, a one-dimensional velocity model is assumed based on offshore refraction profile line 3 of RV Sonne cruise SO76 west of Nicoya peninsula (Ye et al., 1996). The coda magnitude is also determined for each event.

Table 6.6.1: 1D-Velocity model for earthquake localisation

Vp(km/s)	Depth(km)
2.0	0.0
4.2	2.0
6.0	7.0
7.2	12.0
8.1	80.0

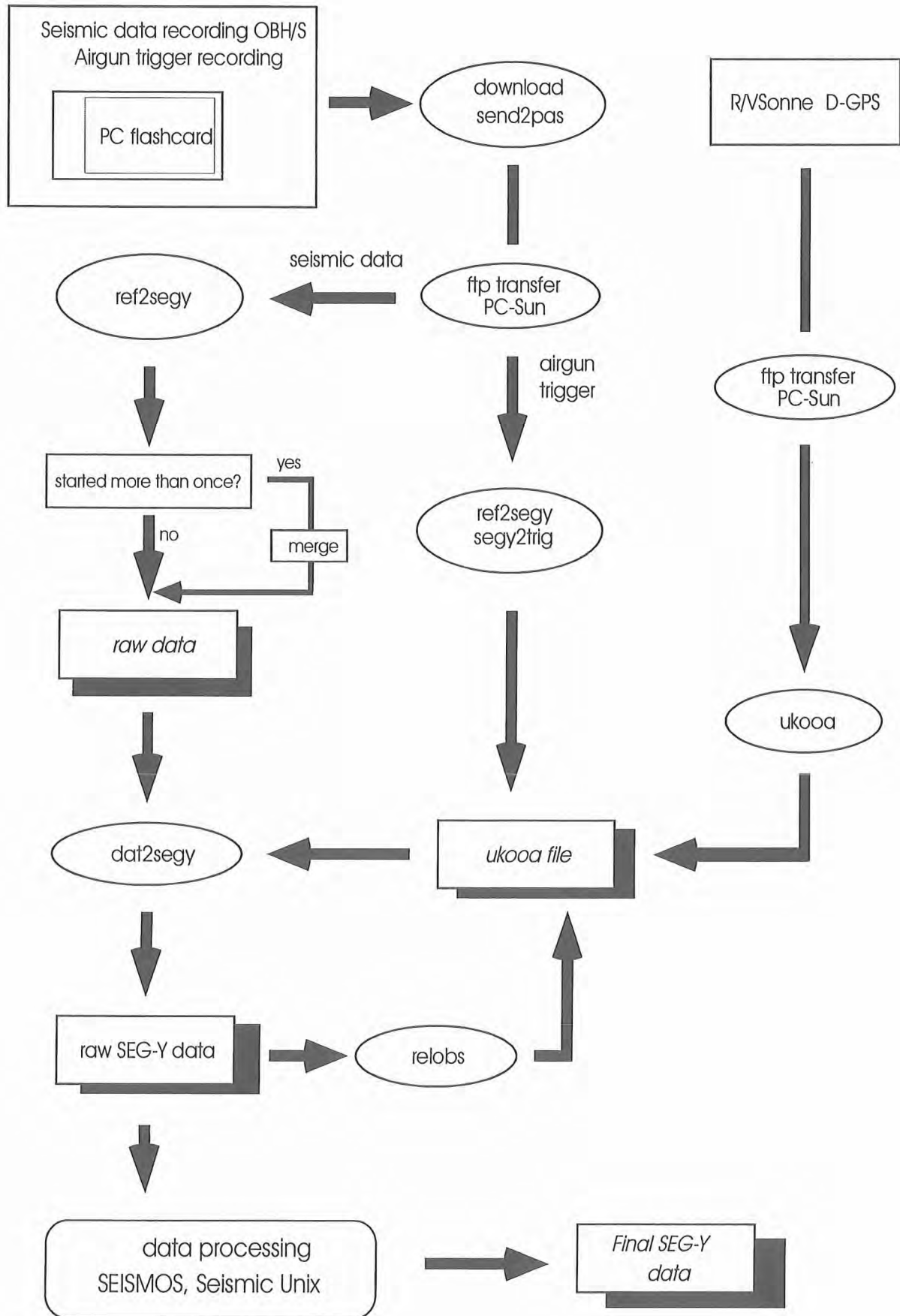


Figure 6.6.1: Processing flow of OBH/S data from raw data to SEG-Y records.

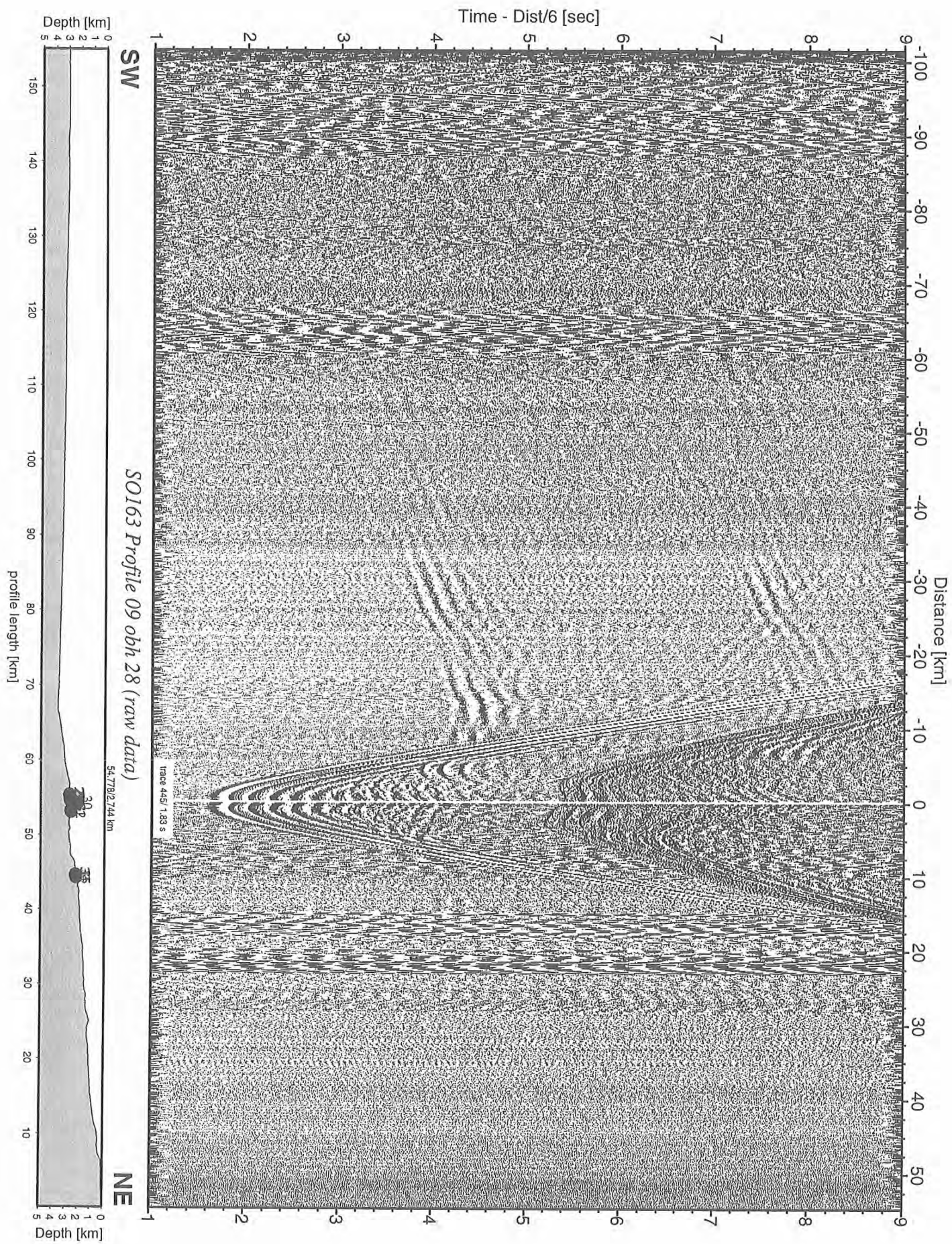


Figure 6.6.2a: Record section from obh 28 (raw data), Profile 09.

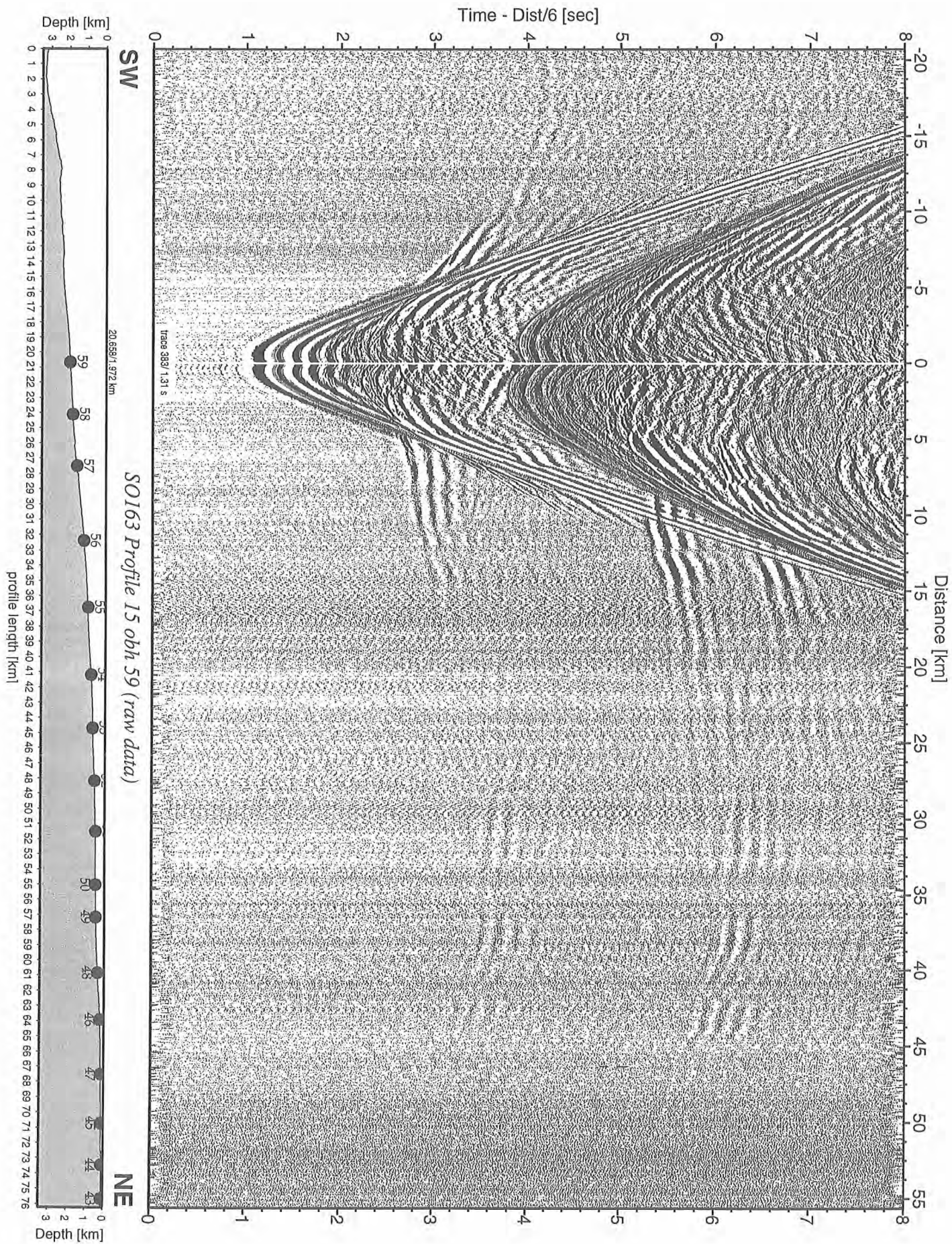


Figure 6.6.2b: Record section from obh 59 (raw data), Profile 15.

SO 163-2 OBH 28 Frequency Filters

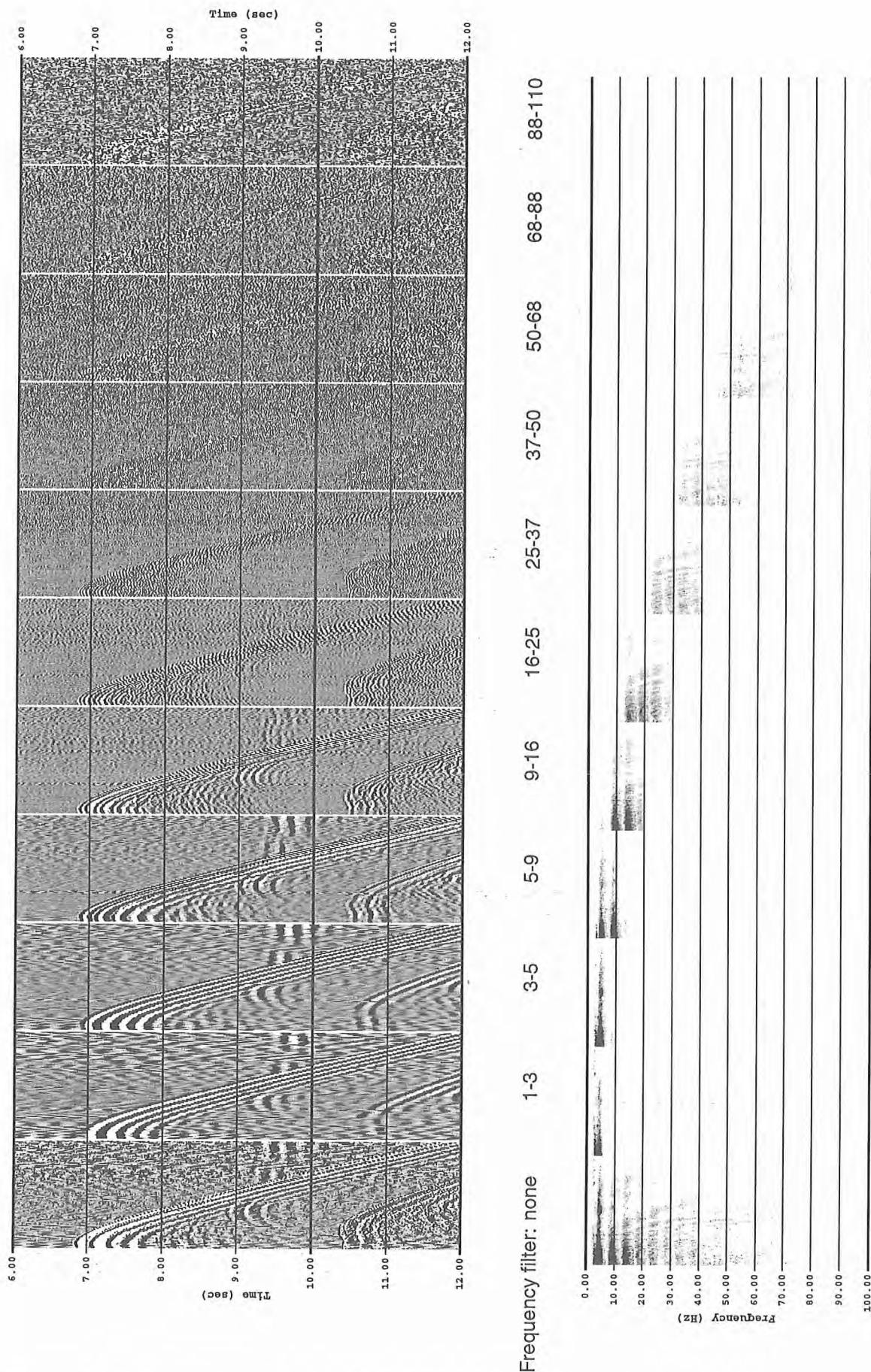


Figure 6.6.3: Frequency analysis of OBH28 (0-13 km offset, Bolt gun).

SO 163-2 OBH 59 Frequency Filters

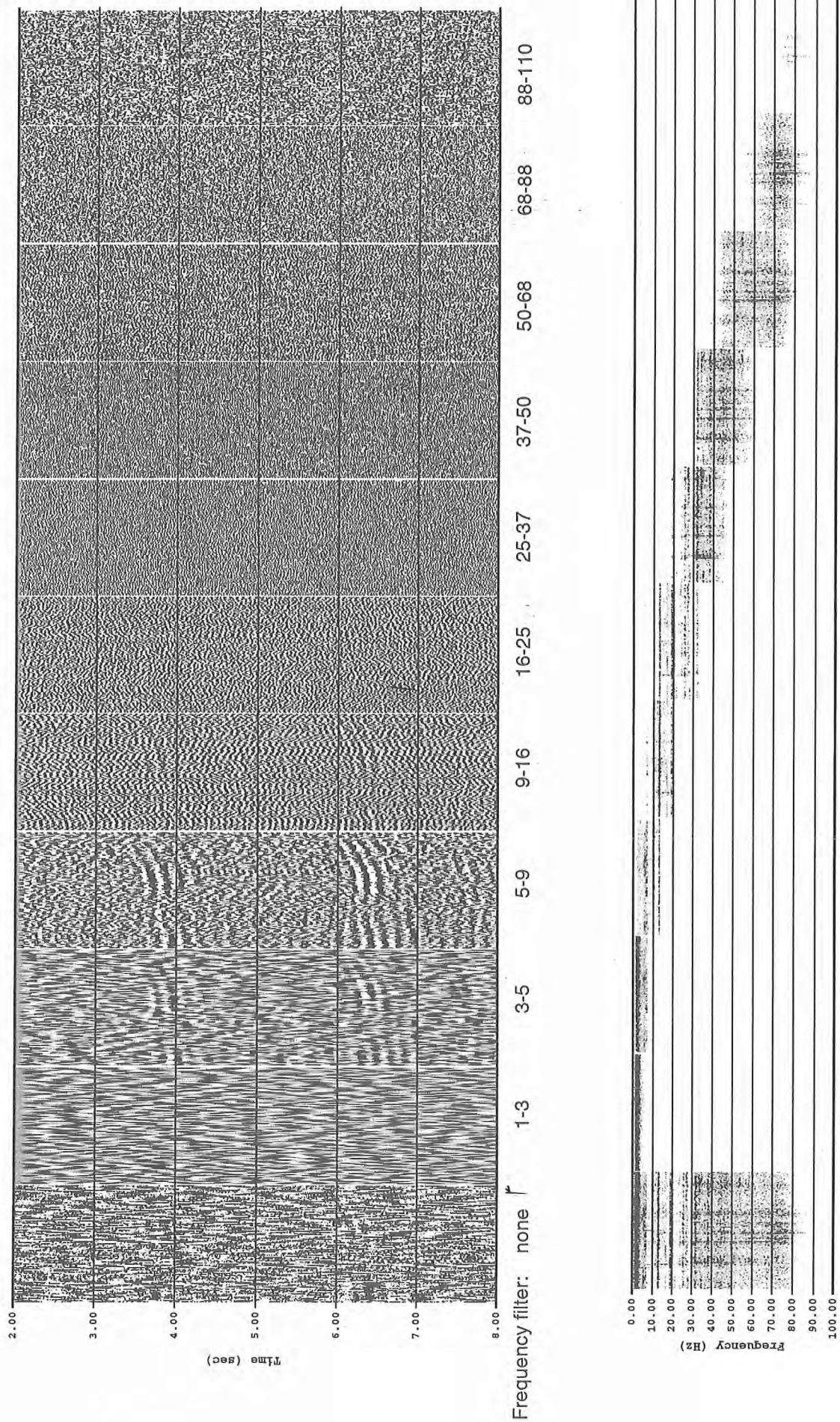


Figure 6.6.6: Frequency analysis of OBH59 (31-42 km offset, G-guns).

SO 163-2 OBH 28 Deconvolution Series

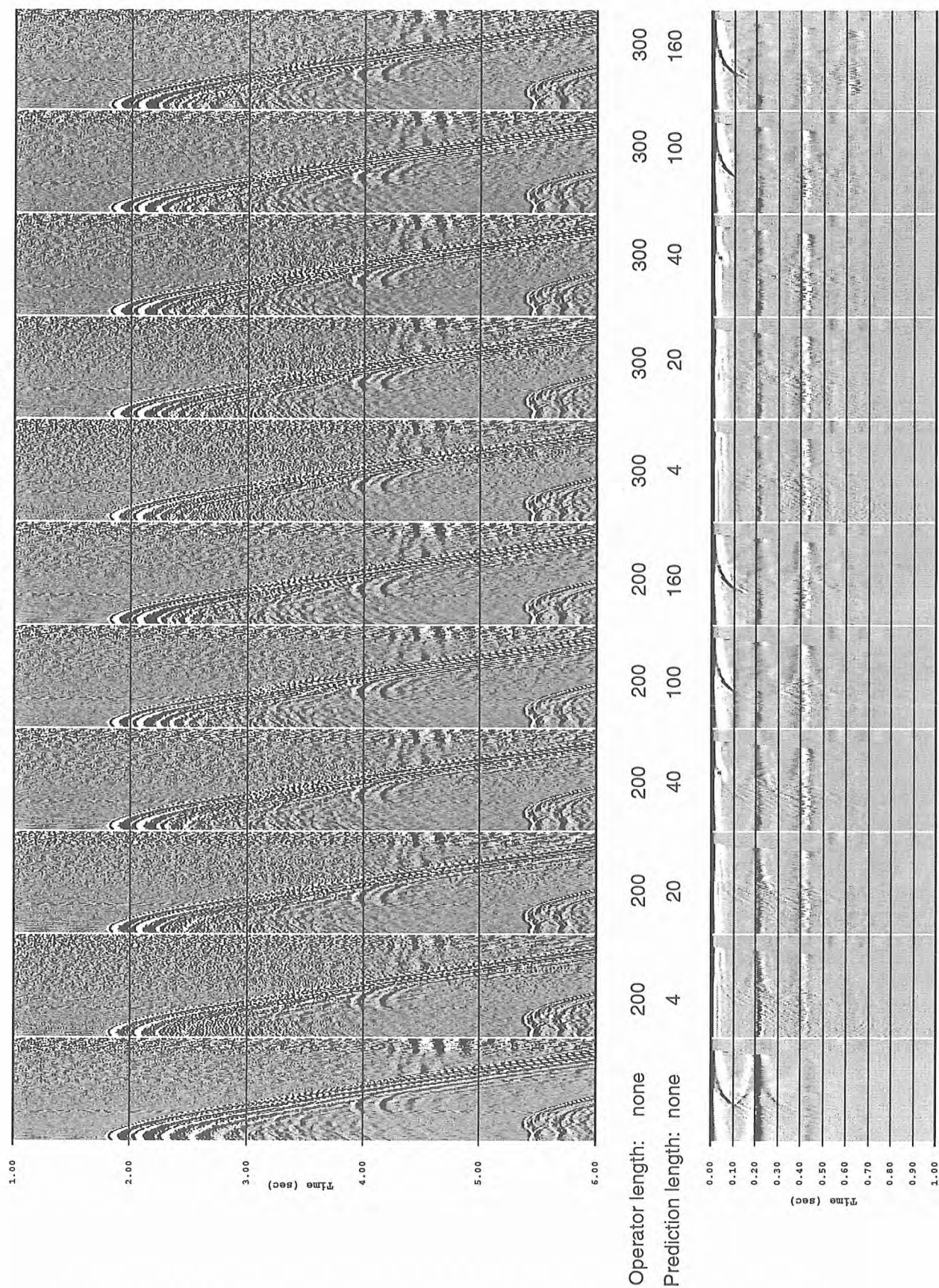


Figure 6.6.7: Deconvolution test of OBH 28 (0-13 km offset, Bolt gun).

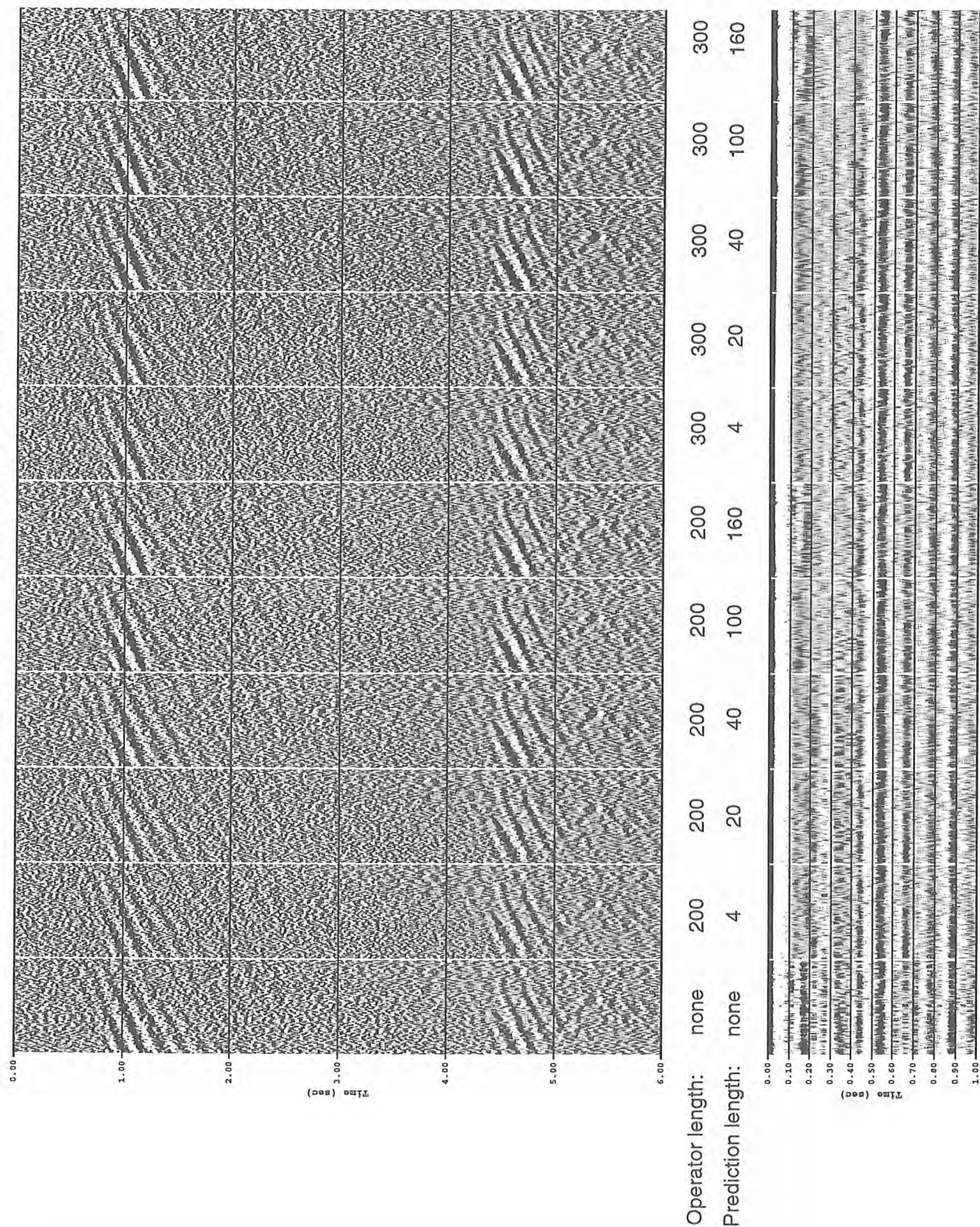


Figure 6.6.8: Deconvolution test of OBH 28 (26-38 km offset, Bolt gun).

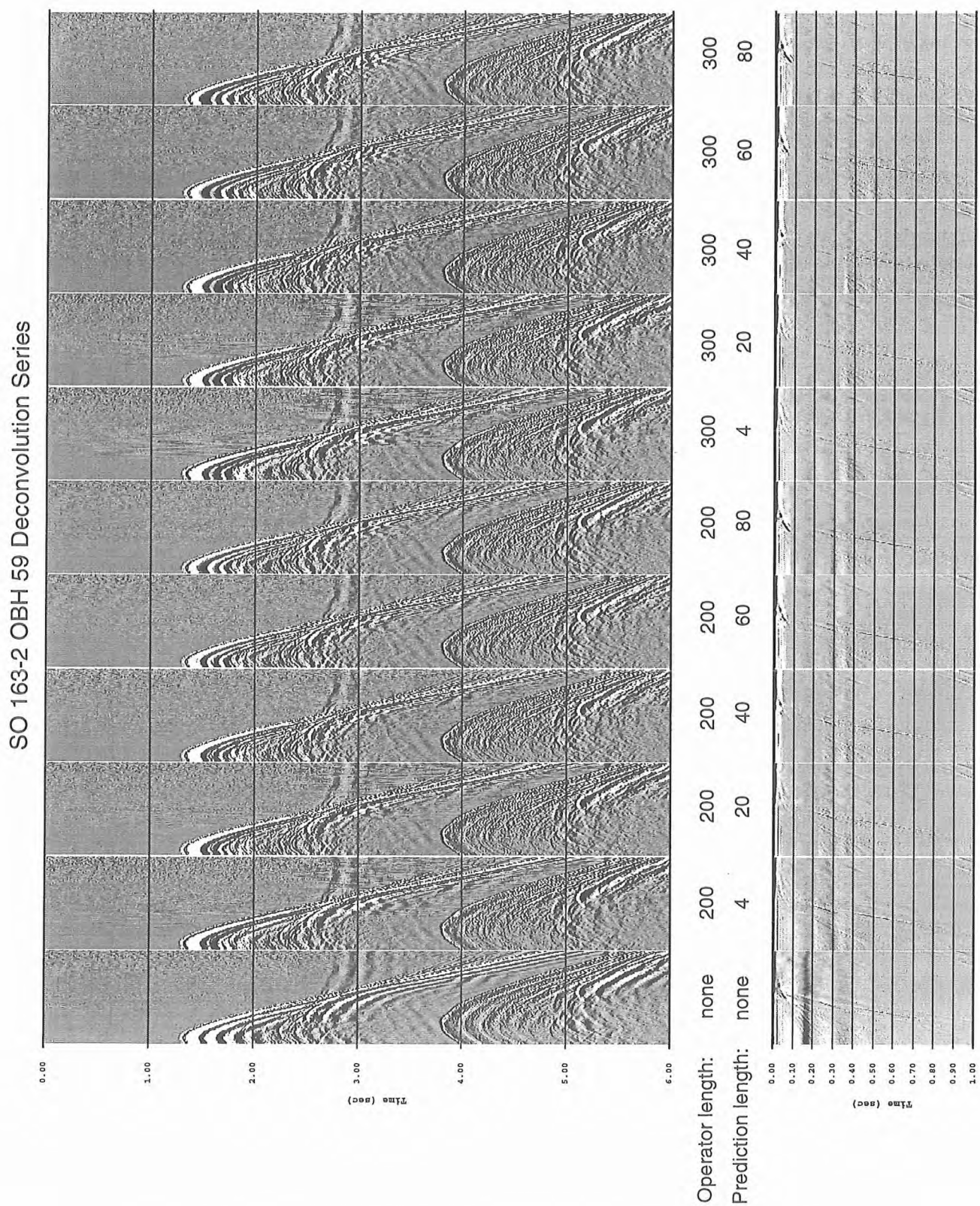


Figure 6.6.9: Deconvolution test of OBH 59 (0-9 km offset, G-guns).

SO 163-2 OBH 59 Deconvolution Series

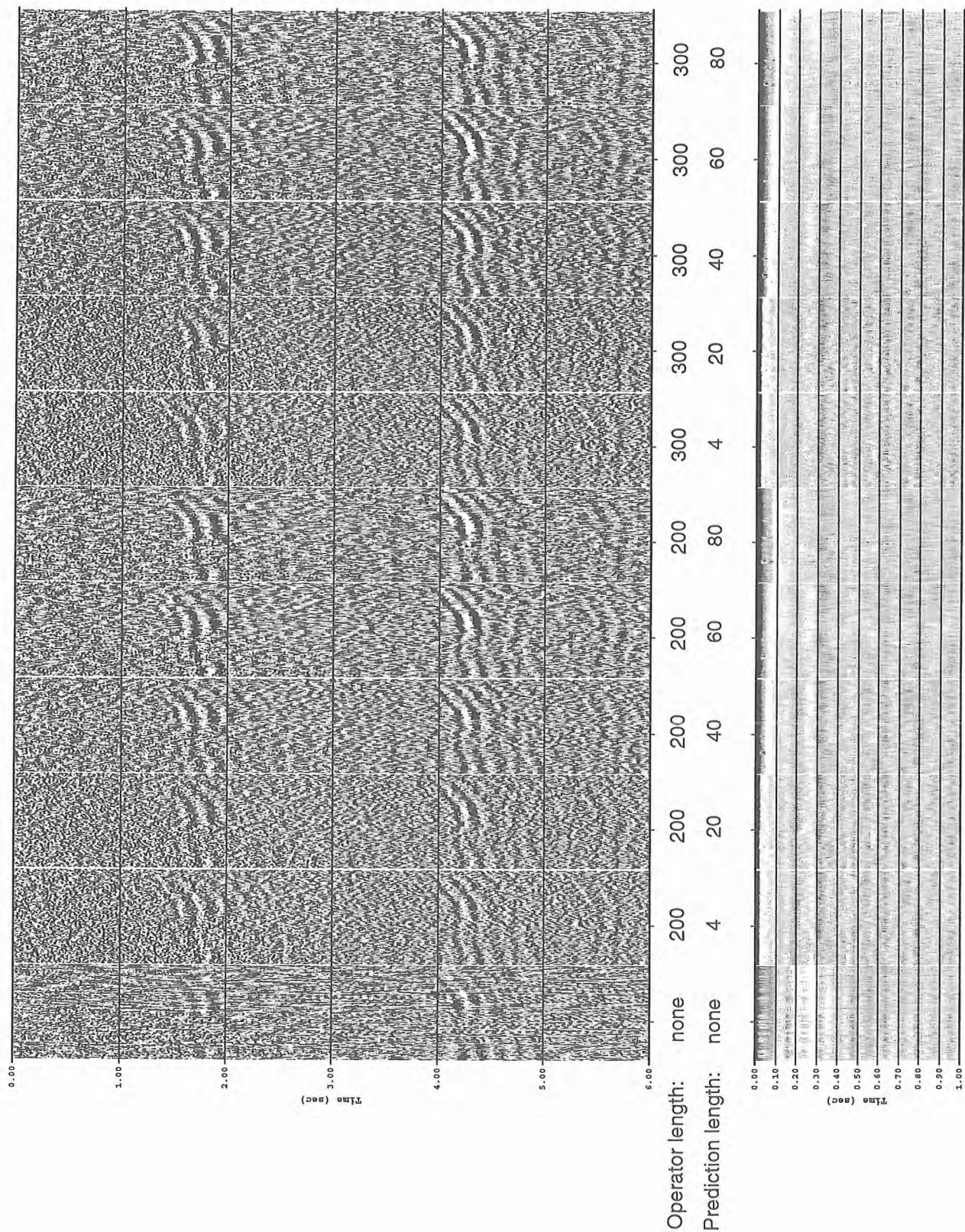


Figure 6.6.10: Deconvolution test of OBH 59 (31-42 km offset, G-guns).

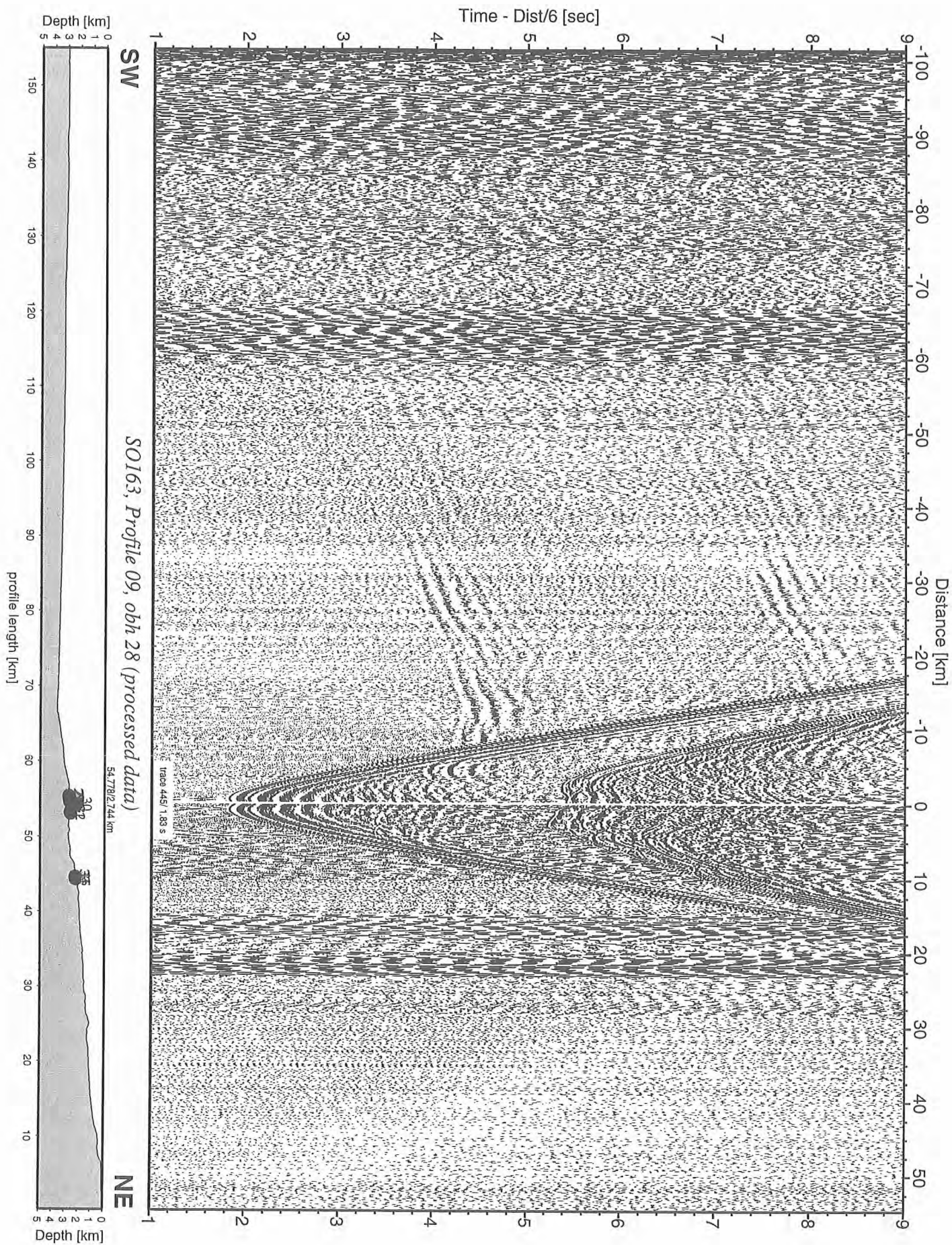


Figure 6.6.11: Record section from obh 28 (processed data), Profile 09.

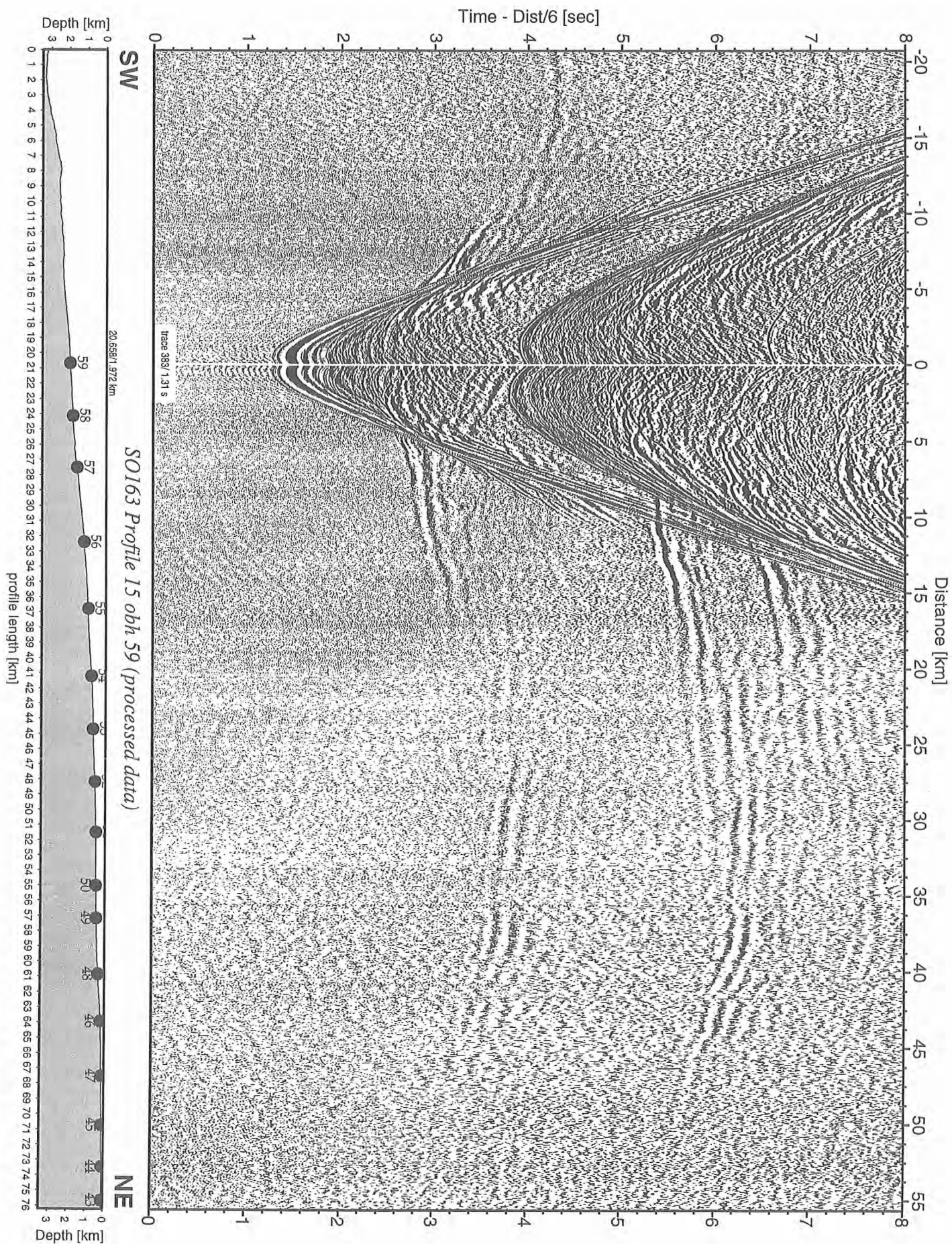


Figure 6.6.12: Record section from obh 59 (processed data), Profile 15.

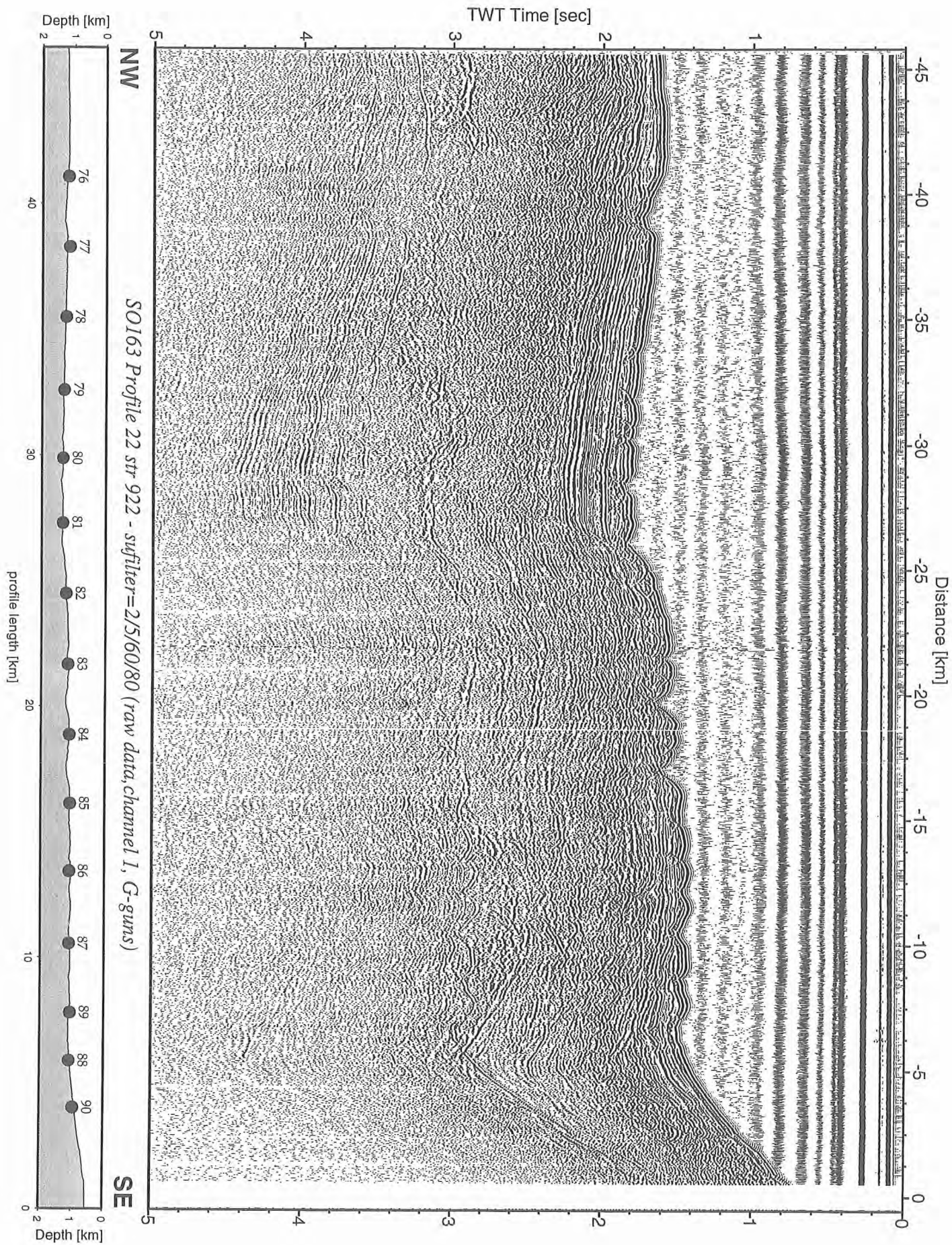


Figure 6.6.13: Record section from str 922 (raw data, channel 1, G-guns), Profile 22.

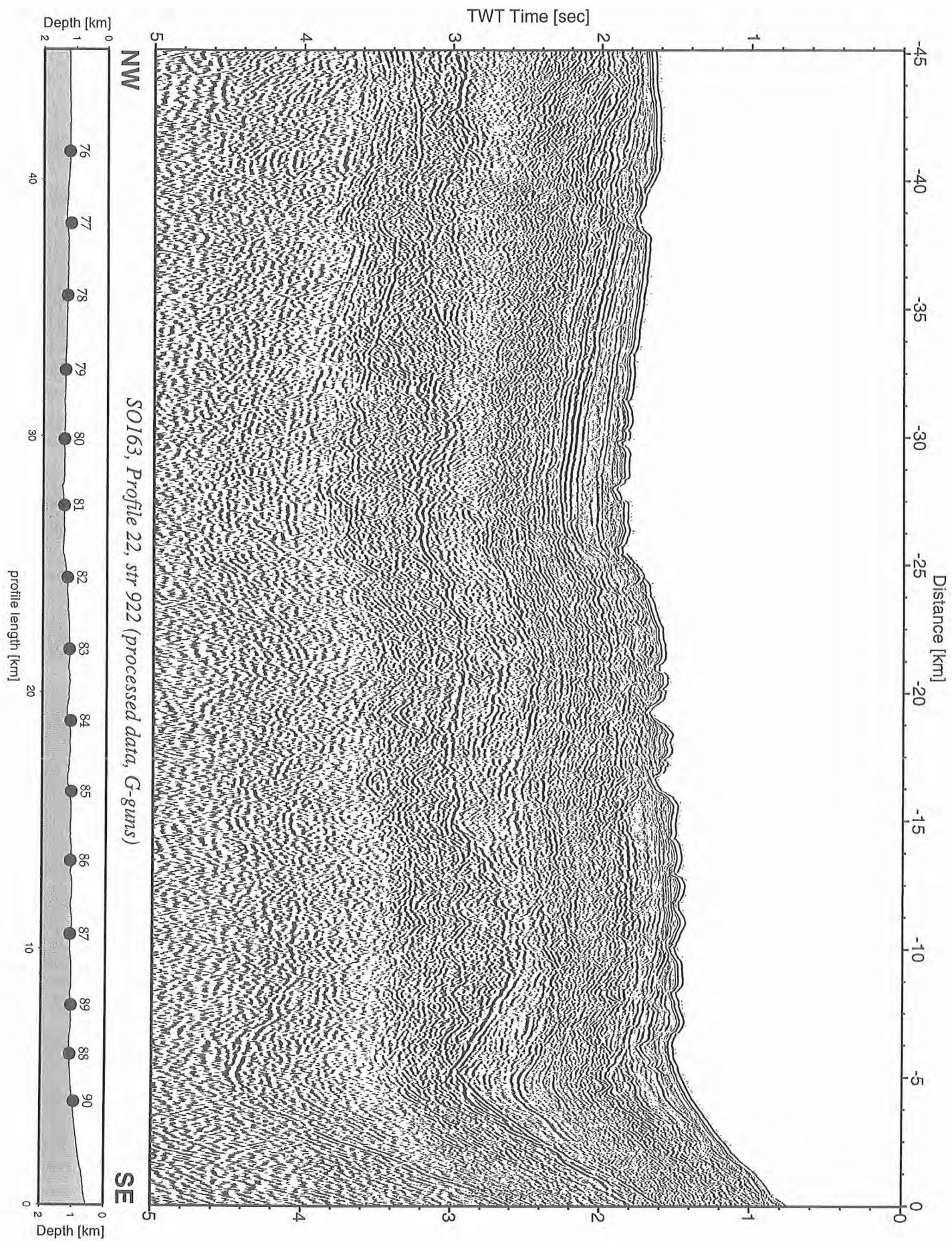


Figure 6.6.14: Record section from str 922 (processed data, G-guns), Profile 22.

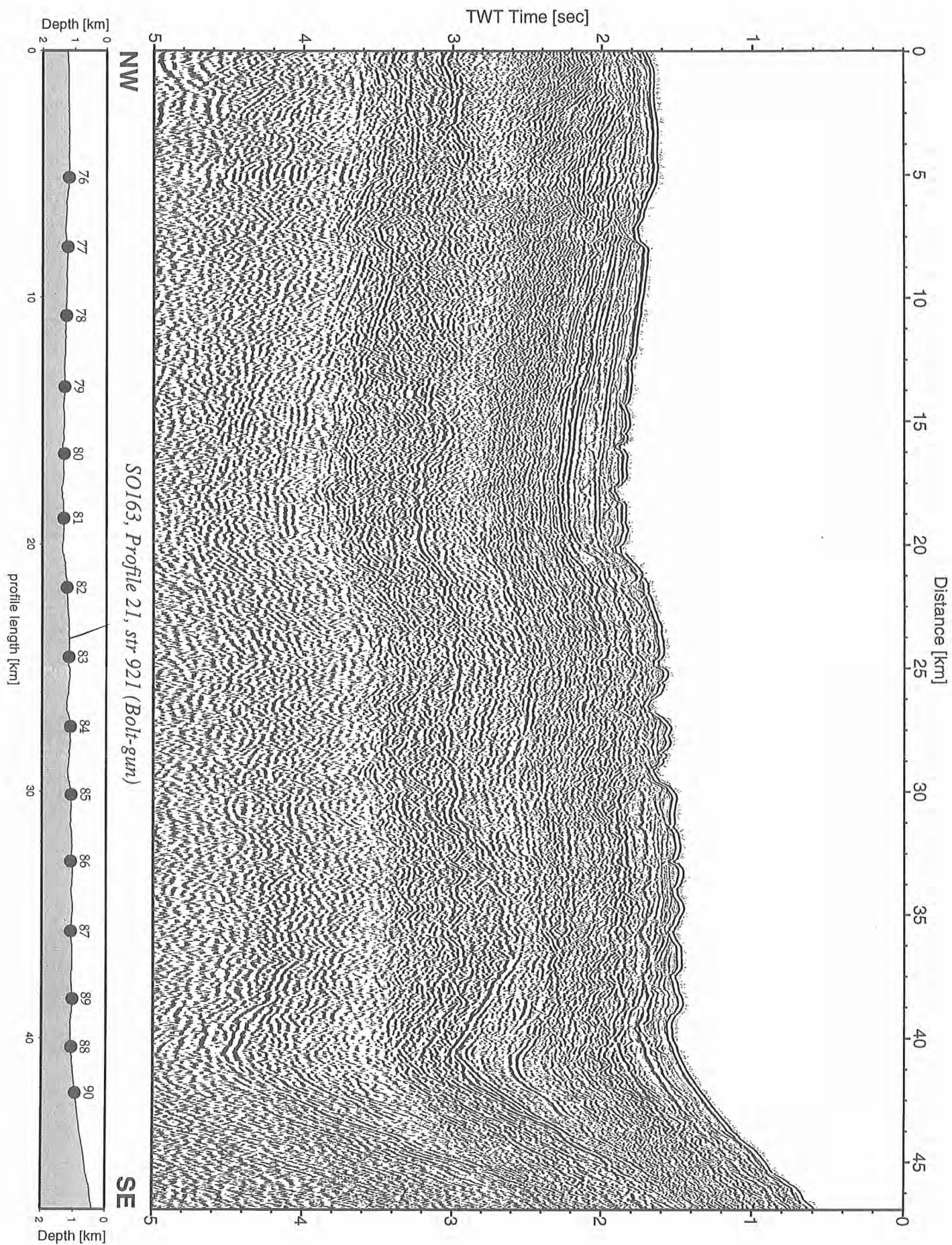


Figure 6.6.15: Record section from str 921 (Bolt-gun), Profile 21.

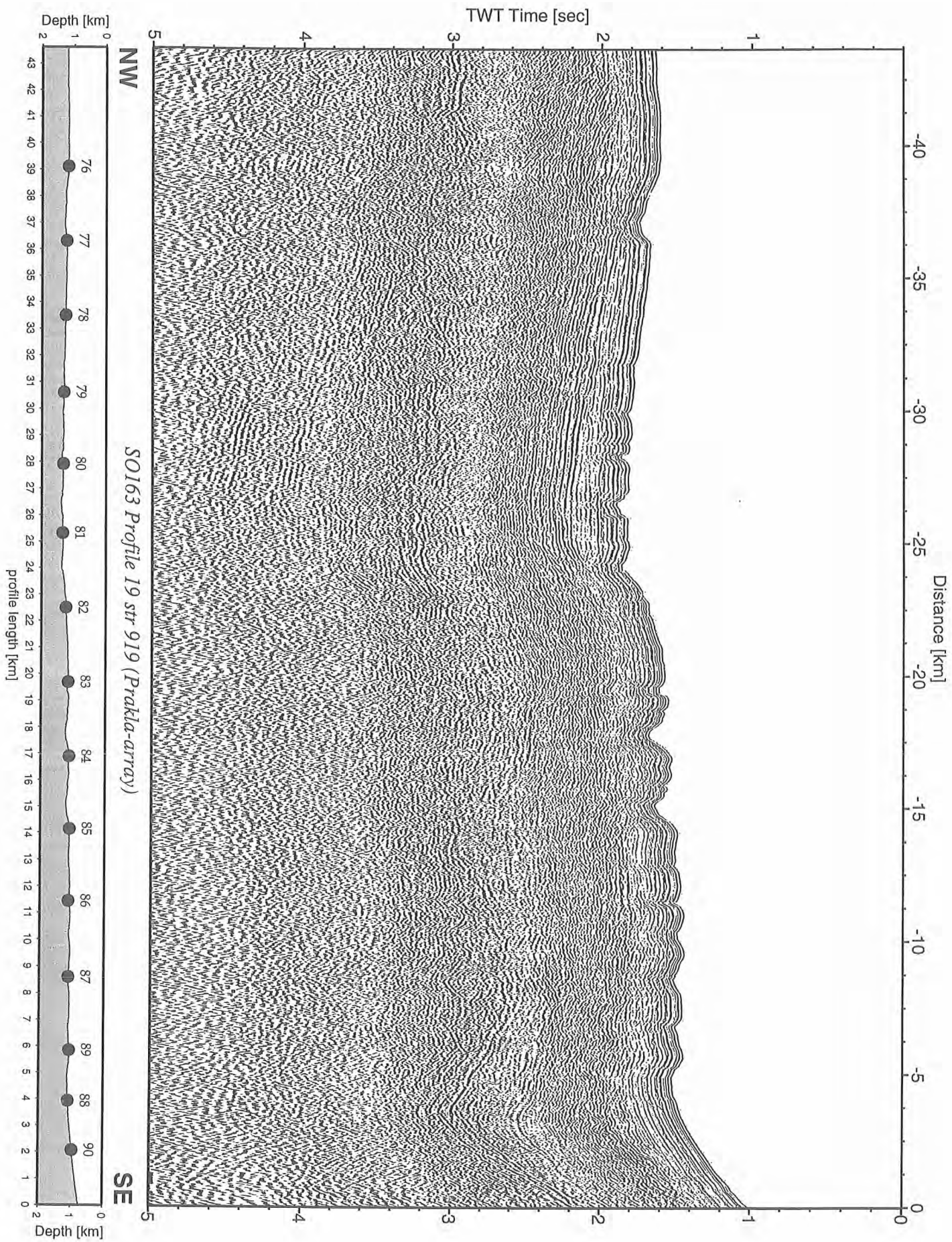


Figure 6.6.16: Record section from str 919 (Prakla-array), Profile 19.

6.6.1 The Local Earthquake Network and profiles P01-P06 (EQNET)

Determination of the optimum source depth of the Bolt-gun (P01 – P06)

(A. Berhorst)

In order to determine optimum depth of the Bolt-gun, profiles P01 to P06 across EQNET (Figure 6.6.1.1) were shot along the same transect (Figure 6.6.1.1.1), revealing a doming structure of 1 km length, 100 m high and a central depression. The cable length between the airgun and its floatations has been varied between 3.5 m and 14 m. Medium and far-field source signal recorded by the streamer and an OBH placed 500 m above seafloor are presented in Figures 6.6.1.9 f. and 6.6.1.10 f., respectively. Peak amplitudes recorded by the streamer were clipped, but as all sections are recorded with the same gain, the signals are comparable in time domain. In the frequency domain, the clipping causes unpredictable frequency distortions, making this records less reliable for an analysis. The registrations of the OBH are not clipped. Due to the 500 m long anchor, the direct wave is not mixed with the first reflections of the seafloor, thus presenting the far-field source signal. In Figure 6.6.1.12 the images for a source depth of 7.5 m and 9.5 m are exchanged. The number of reverberations is increasing with the source depth, while the time interval between first peak and bubbles is decreasing with increasing source depth. The increasing time interval between first peak and first bubble is well imaged in the first maximum of the frequency spectrum and its overtones. This first frequency maximum is shifted from 4.5 Hz for a source depth of 3.5 m to 6 Hz for a source depth of 14 m. The peak amplitudes of the first arrivals indicate a maximum for a cable length of 7.5 m.

The frequency analysis of the refracted arrivals (Figure 6.6.1.13) is in good agreement with these observations, thus indicating that the analyzed wavelet is our far-field signature convoluted with the subsurface reflectivity. In order to operate the Bolt-gun with maximum amplitude of the first peak, we used a cable length of 7.5m after these test profiles.

Velocity determination

(I.Arroyo)

The processed record of the OBH23 during the shooting of Profile 2 (Figure 6.6.1.3) was analyzed in order to obtain an 1-D velocity model of the upper sediment cover (~ 4 km). To achieve further studies in the far field, OBH23 was located 500 m above the sea floor. The pickings of the record are shown in Figure 6.6.1.14, together with the predicted arrivals modeled with the program R1D (Luetgert, 1992). The curvatures of the predicted arrivals nearly coincide with the observed phases. Four reflections were found, with P-velocities ranging from 1,5 to 2,0 km/s :

Depth (m)	P velocity (km/s)
0 – 1 430 (water)	1,480
1 430 – 1 973	1,468 – 1,766
1 973 – 2 504	1,791 – 1,905
2 504 – 3 199	1,922 – 1,959
3 199 – 3 593	1,953 – 2,027

SO163-2: EQNET1 and P01-P06 Profiles and Stations

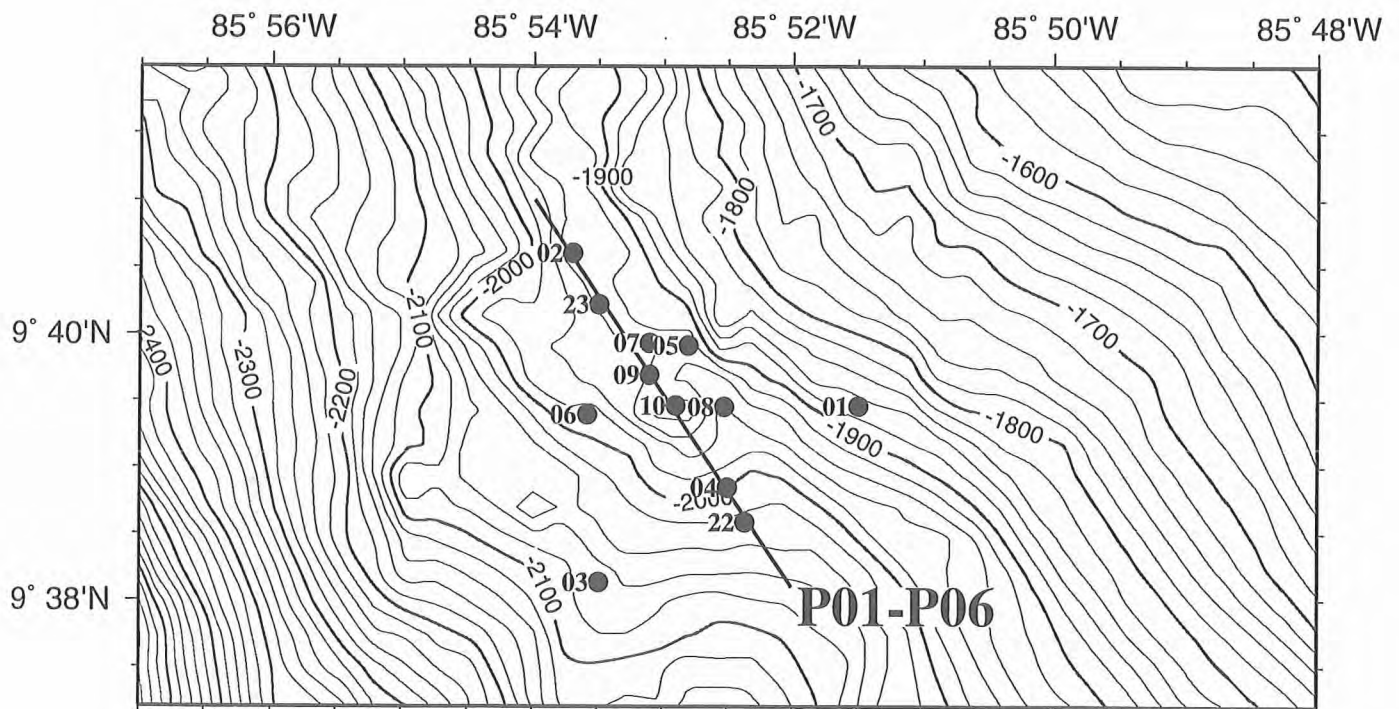


Figure 6.6.1.1: Location map of Eqnet1 and profiles 01-06 with 20 m isolines. Superimposed are the 10 OBH locations, covering the OFOS 10 mud volcano.

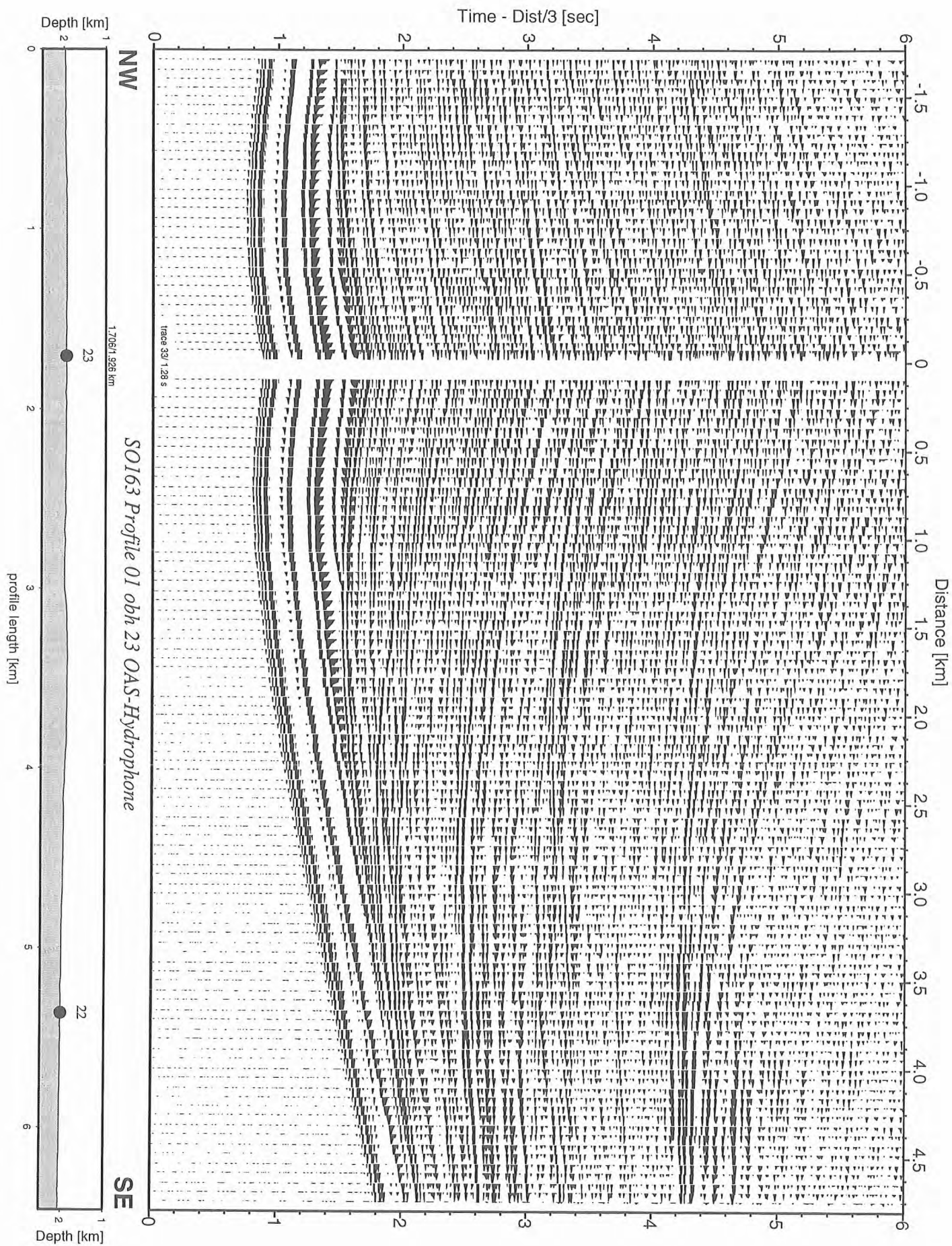


Figure 6.6.1.2: Record section from obh 23 OAS-Hydrophone, Profile 01.

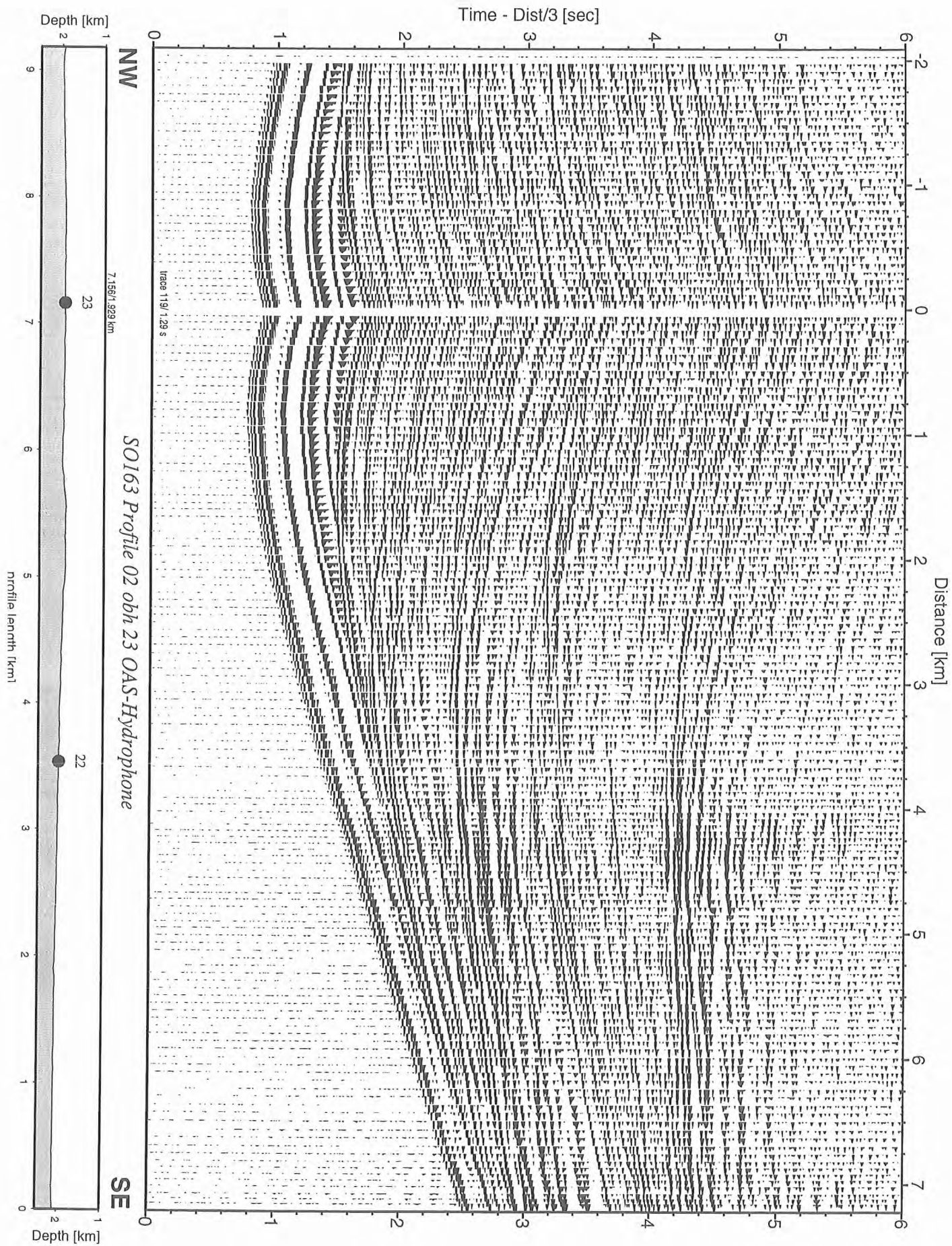


Figure 6.6.1.3: Record section from obh 23 OAS-Hydrophone, Profile 02.

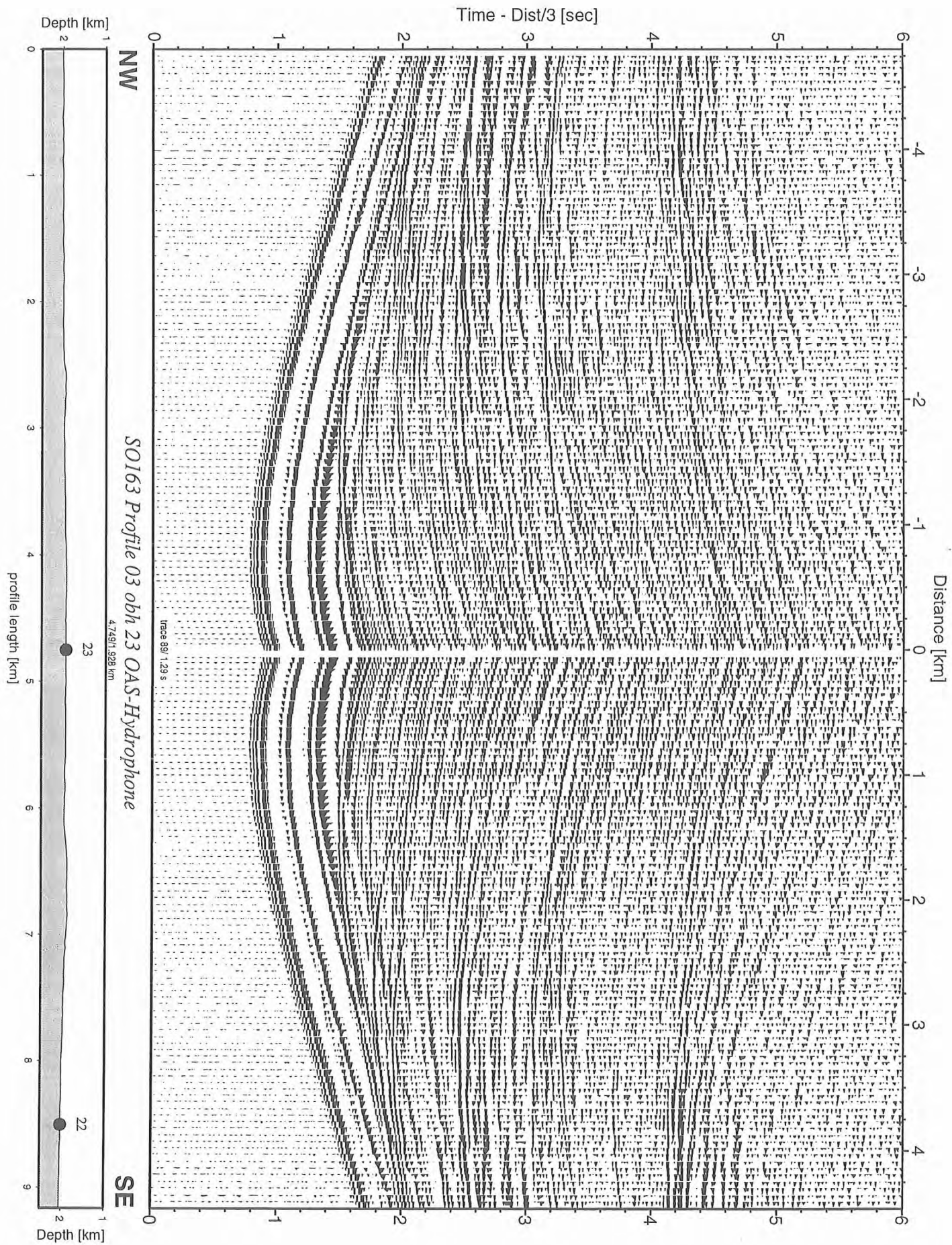


Figure 6.6.1.4: Record section from obh 23 OAS-Hydrophone, Profile 03.

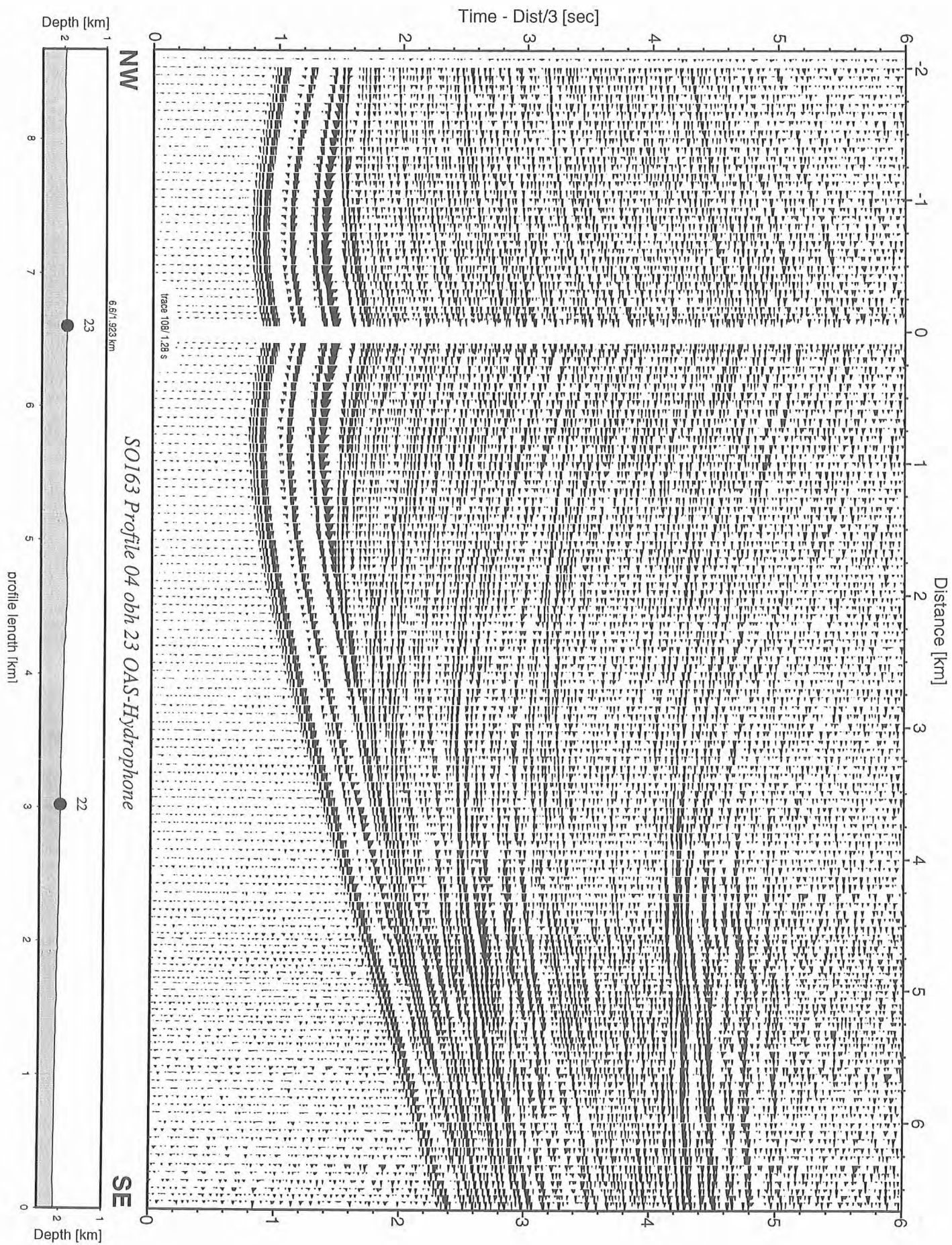


Figure 6.6.1.5: Record section from obh 23 OAS-Hydrophone, Profile 04.

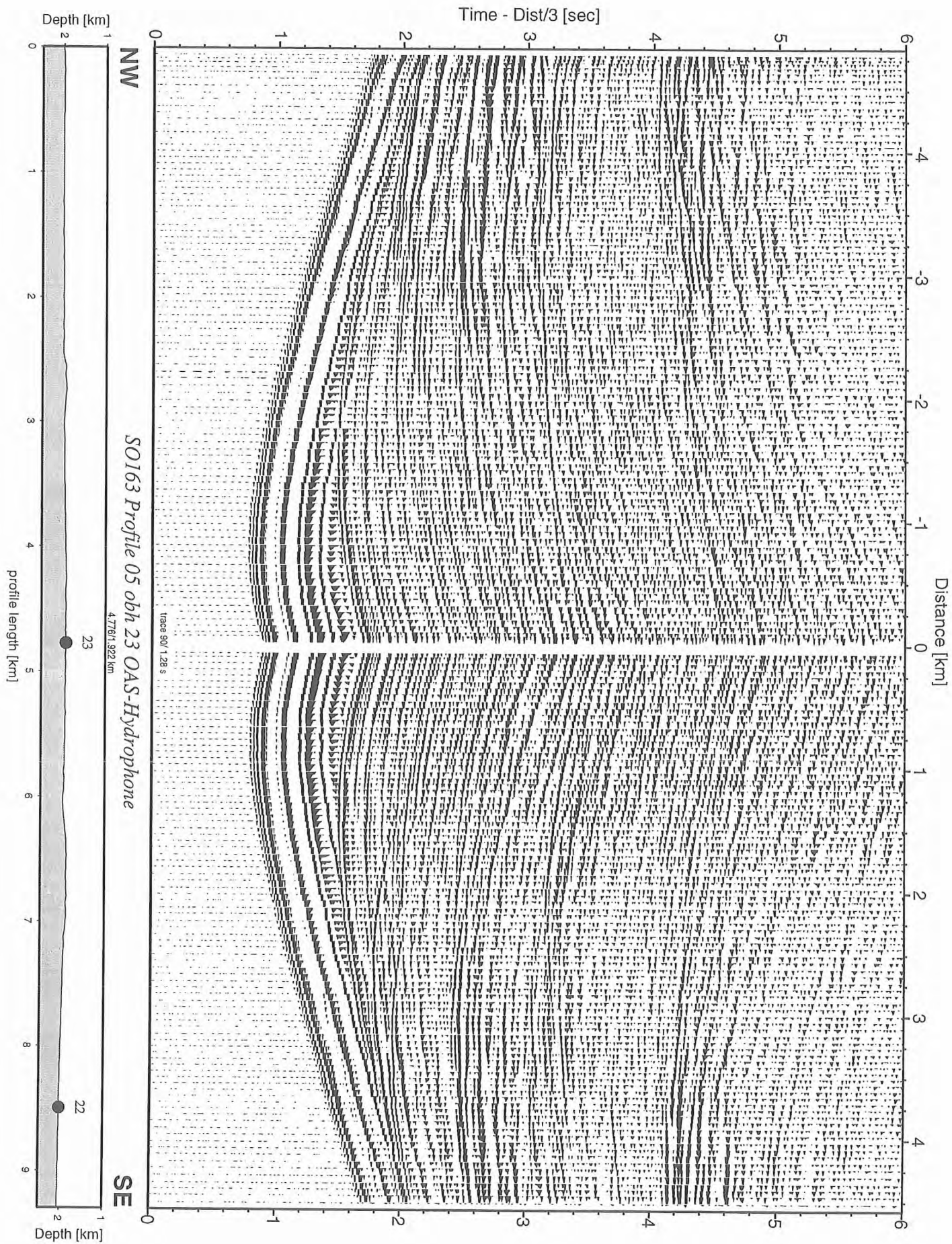


Figure 6.6.1.6: Record section from obh 23 OAS-Hydrophone, Profile 05.

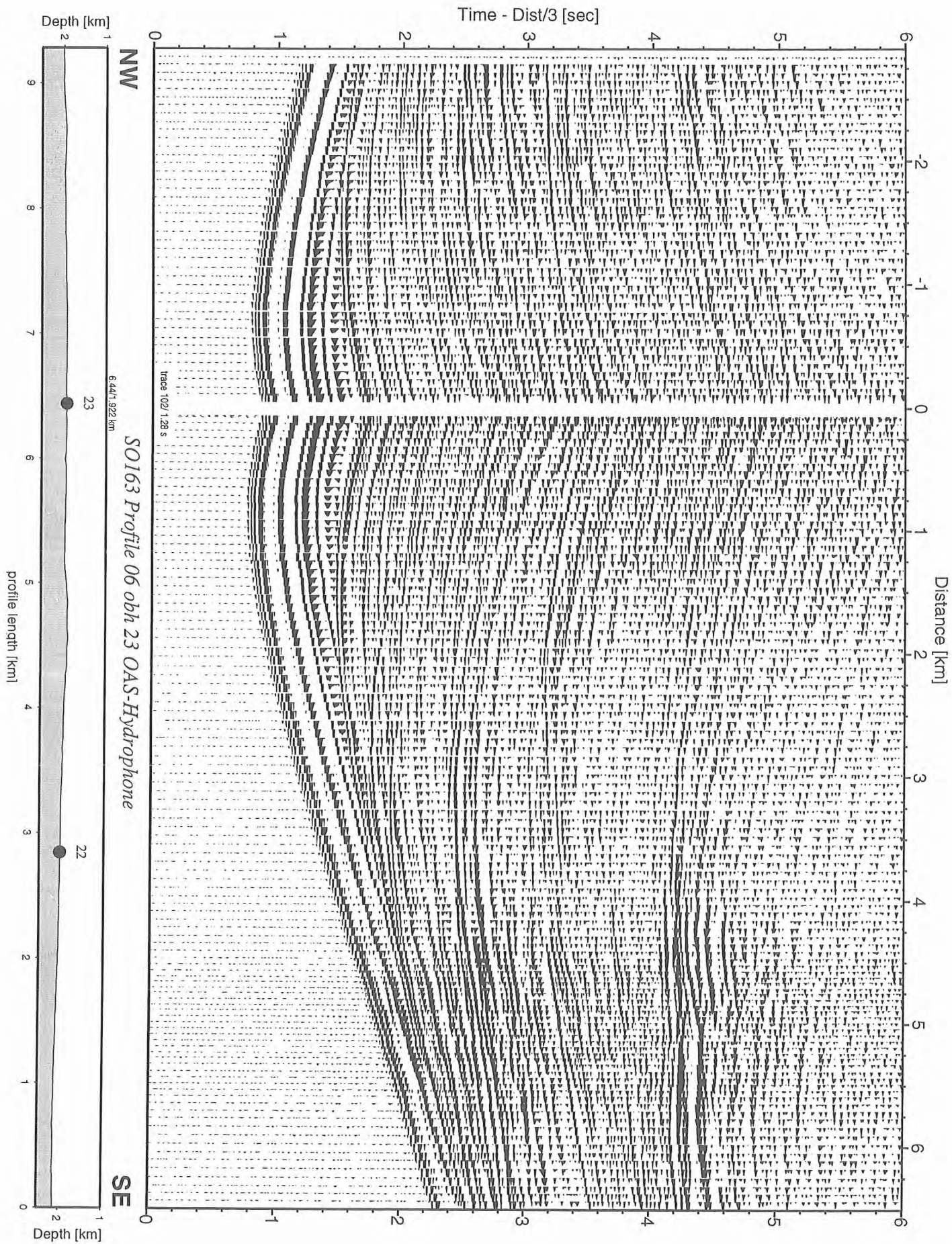


Figure 6.6.1.7: Record section from obh 23 OAS-Hydrophone, Profile 06.

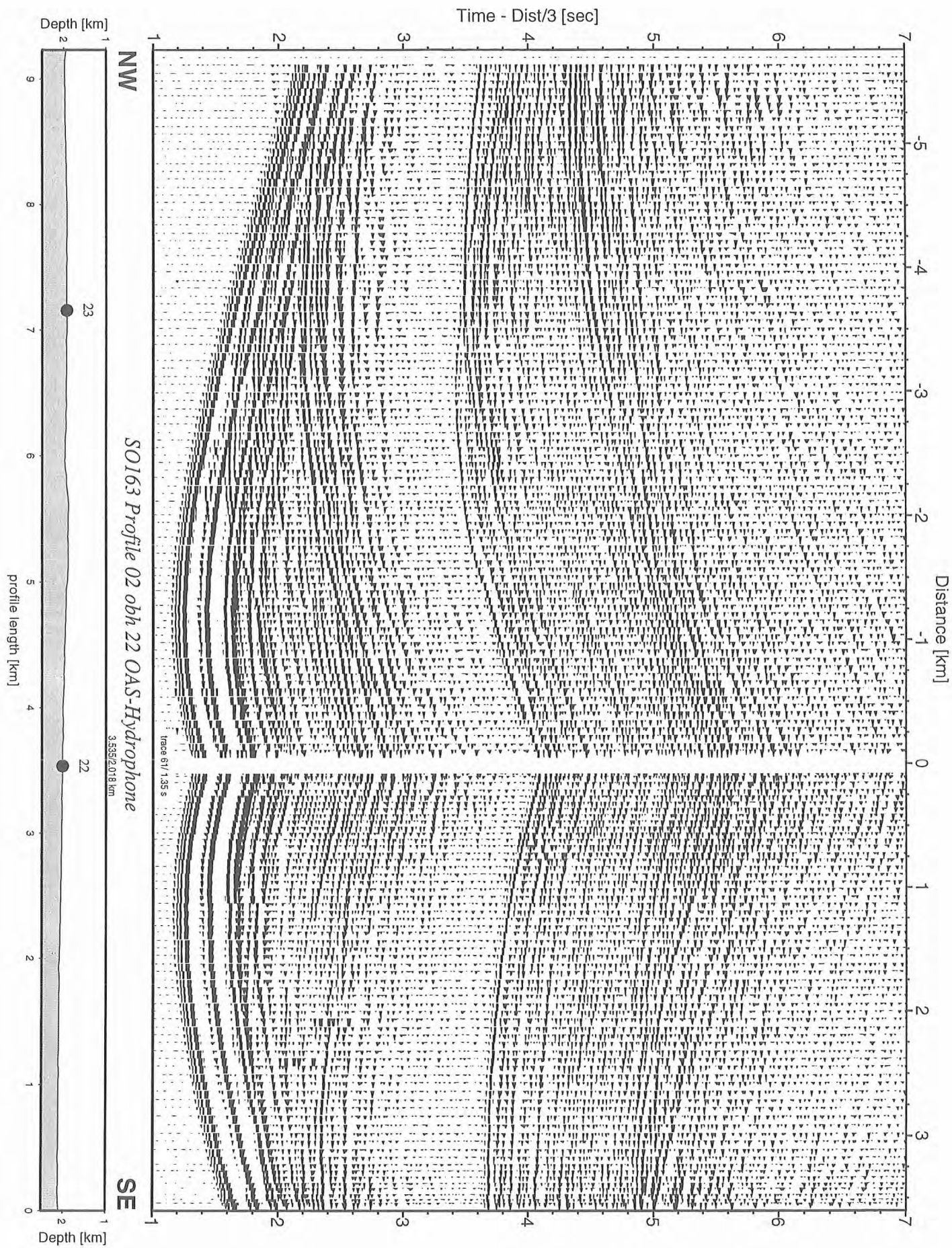


Figure 6.6.1.8: Record section from obh 22 OAS-Hydrophone, Profile 02.

SO163-2 Profiles P01 to P06 Frequency Analysis

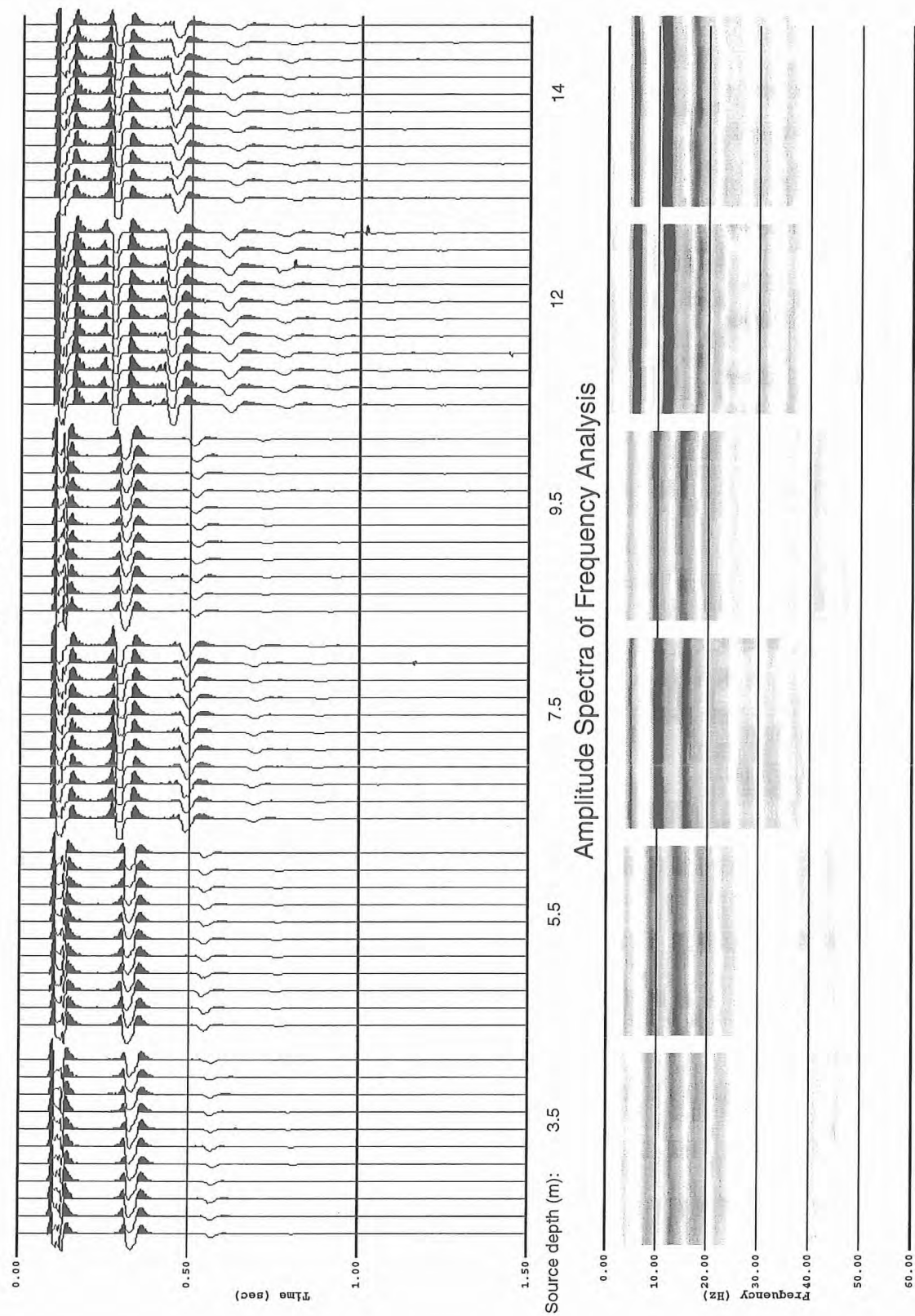


Figure 6.6.1.9: Frequency analysis of direct arrivals for different source depths (Bolt gun).

SO163-2 Profiles P01 to P06 Frequency Analysis

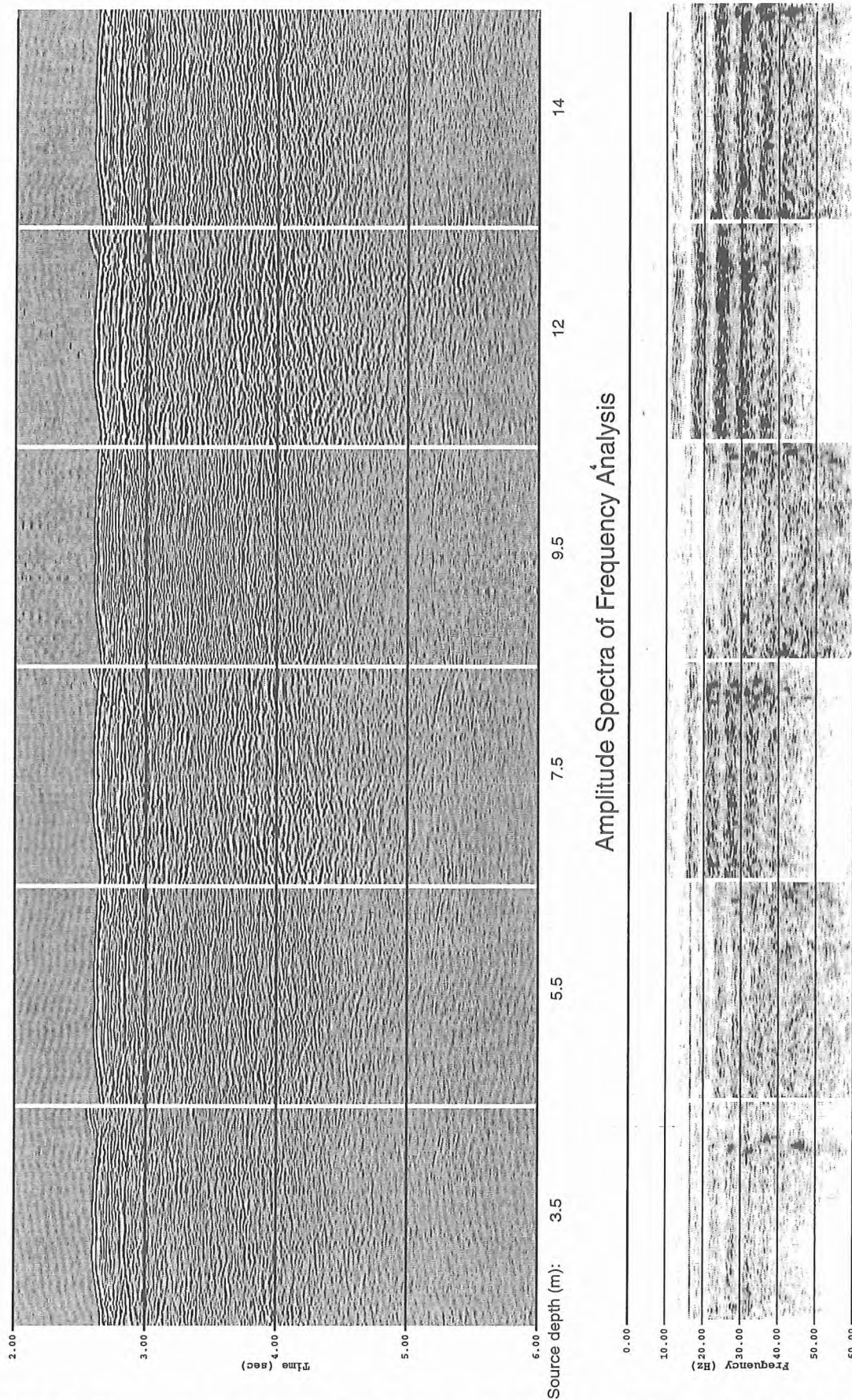


Figure 6.6.1.10: Frequency analysis of reflected waves for different source depths (Bolt gun).

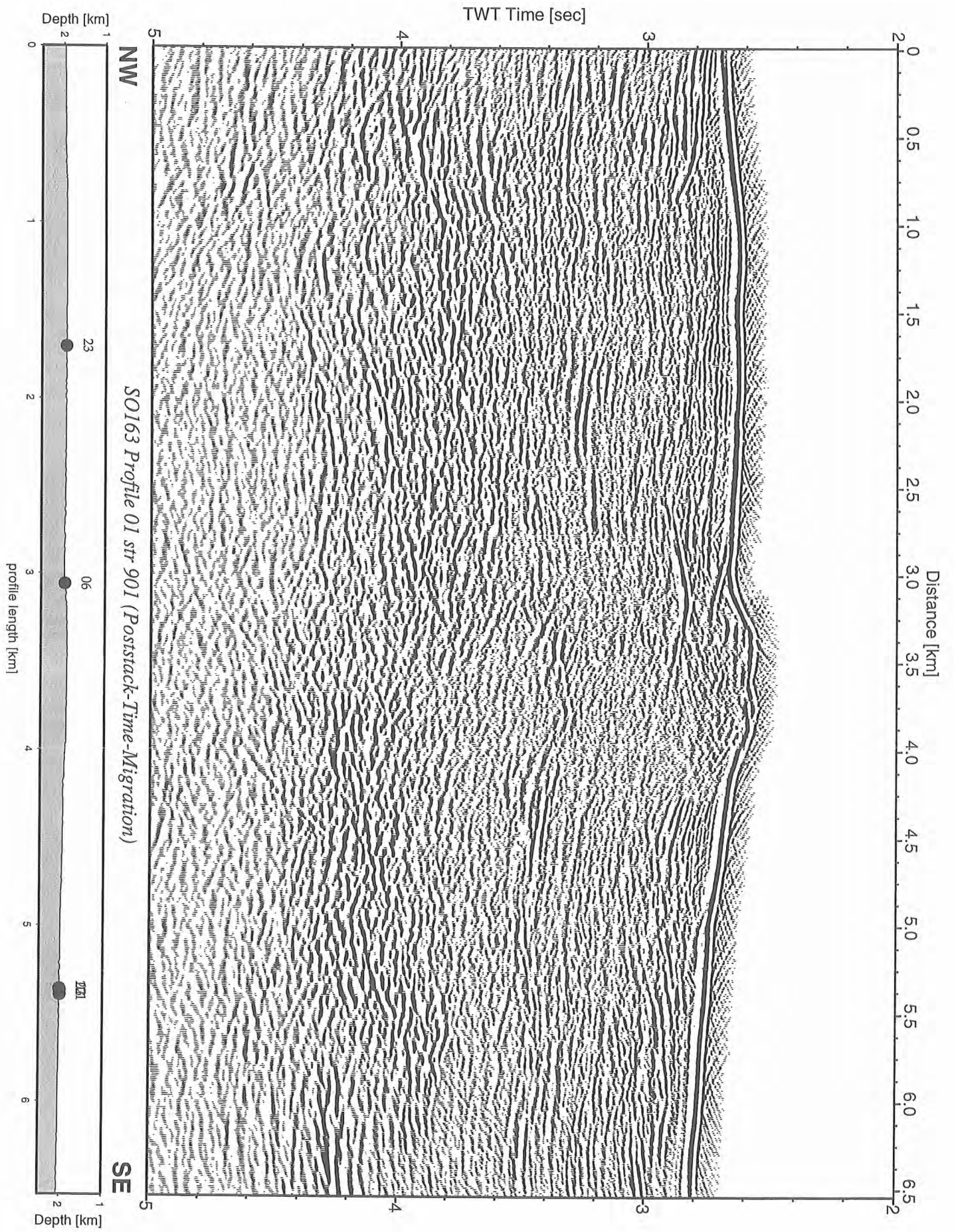


Figure 6.6.1.11: Record section from str 901 (Poststack-Time-Migration), Profile 01.

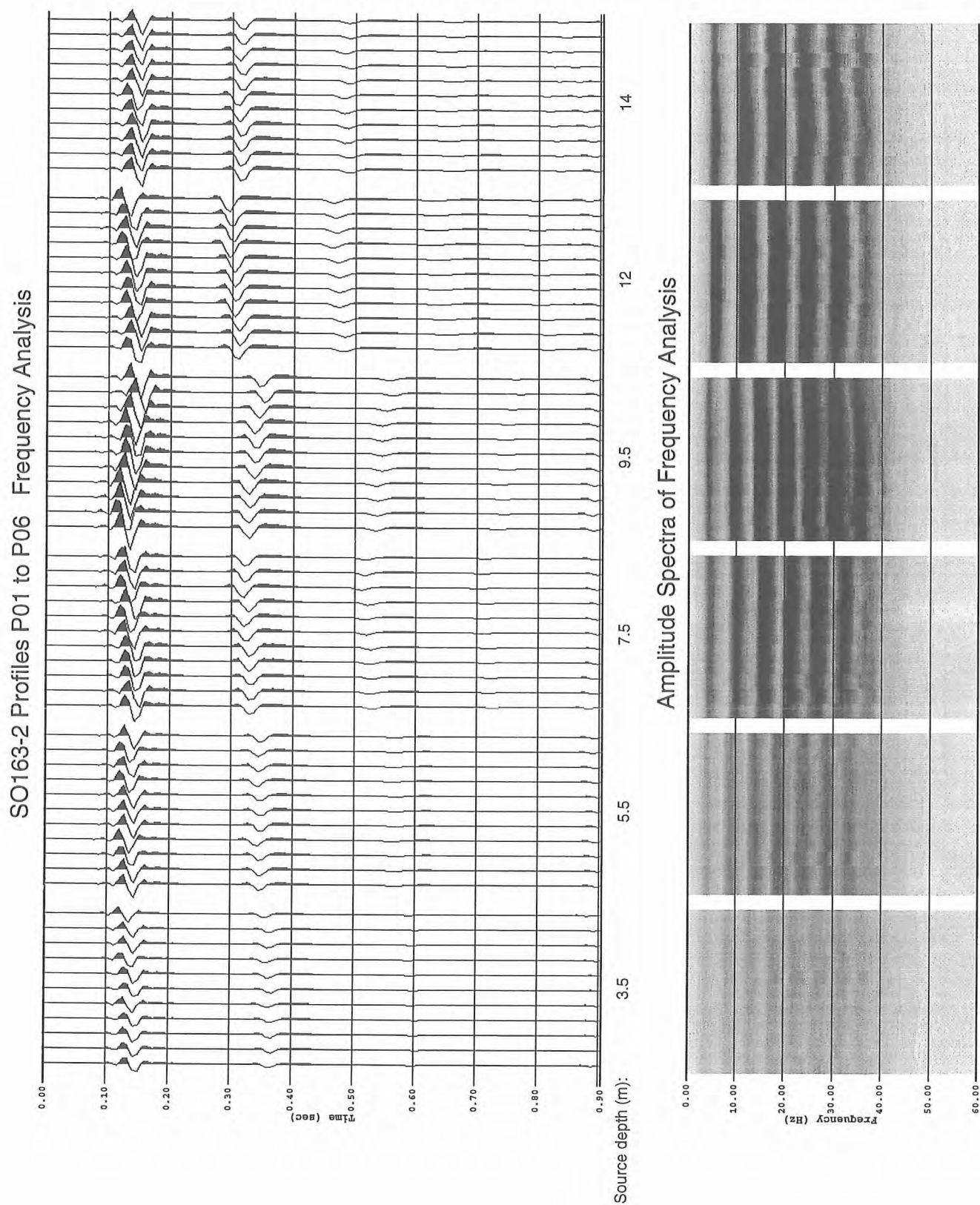


Figure 6.6.1.12: Frequency analysis of direct arrivals for different source depths (Bolt gun) recorded by station OBH 23 located 500 m above the seafloor.

SO163-2 Profiles P01 to P06 Frequency Analysis

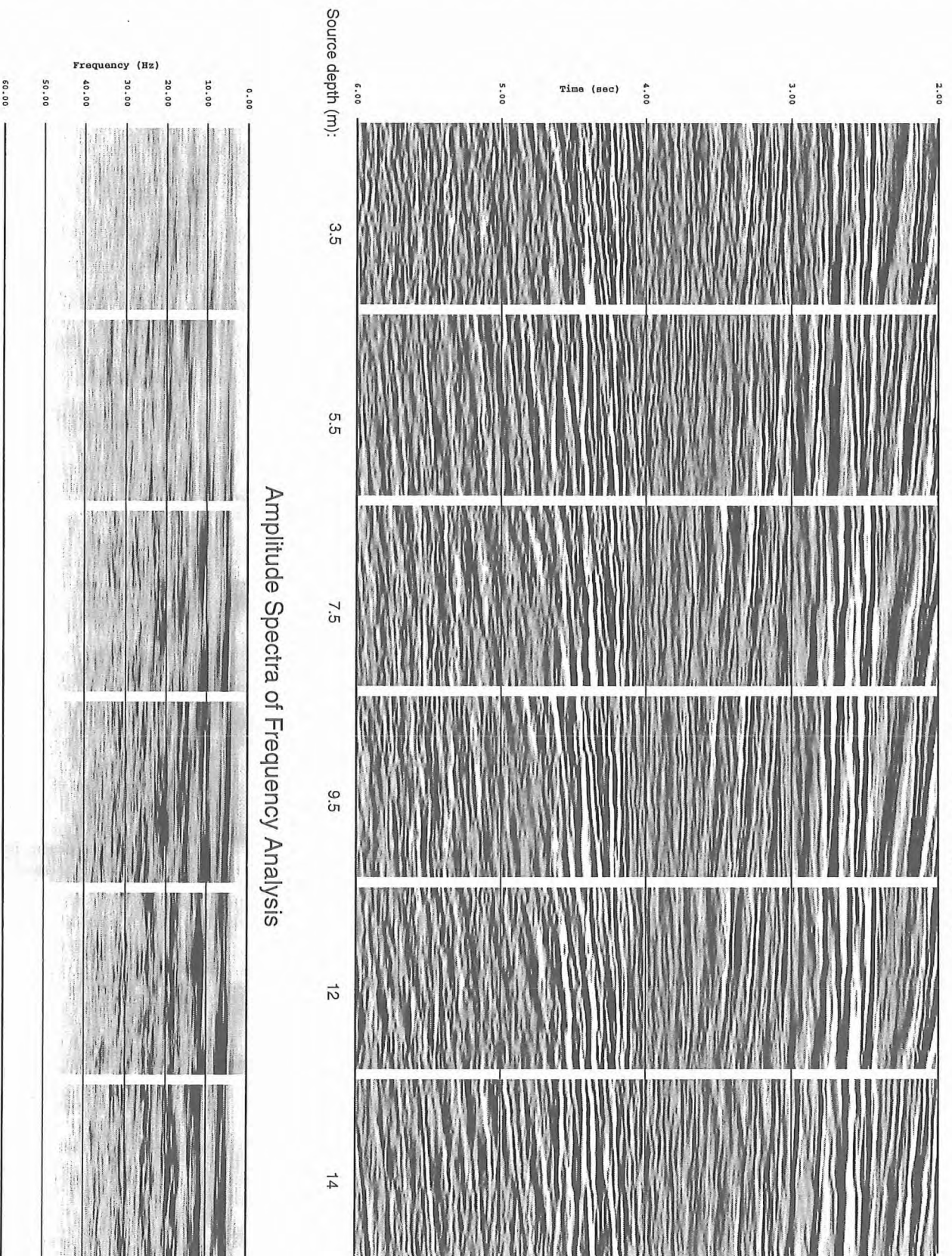


Figure 6.6.1.13: Frequency analysis of refracted arrivals for different source depths (Bolt gun) recorded by station OBH 23 located 500 m above the seafloor.

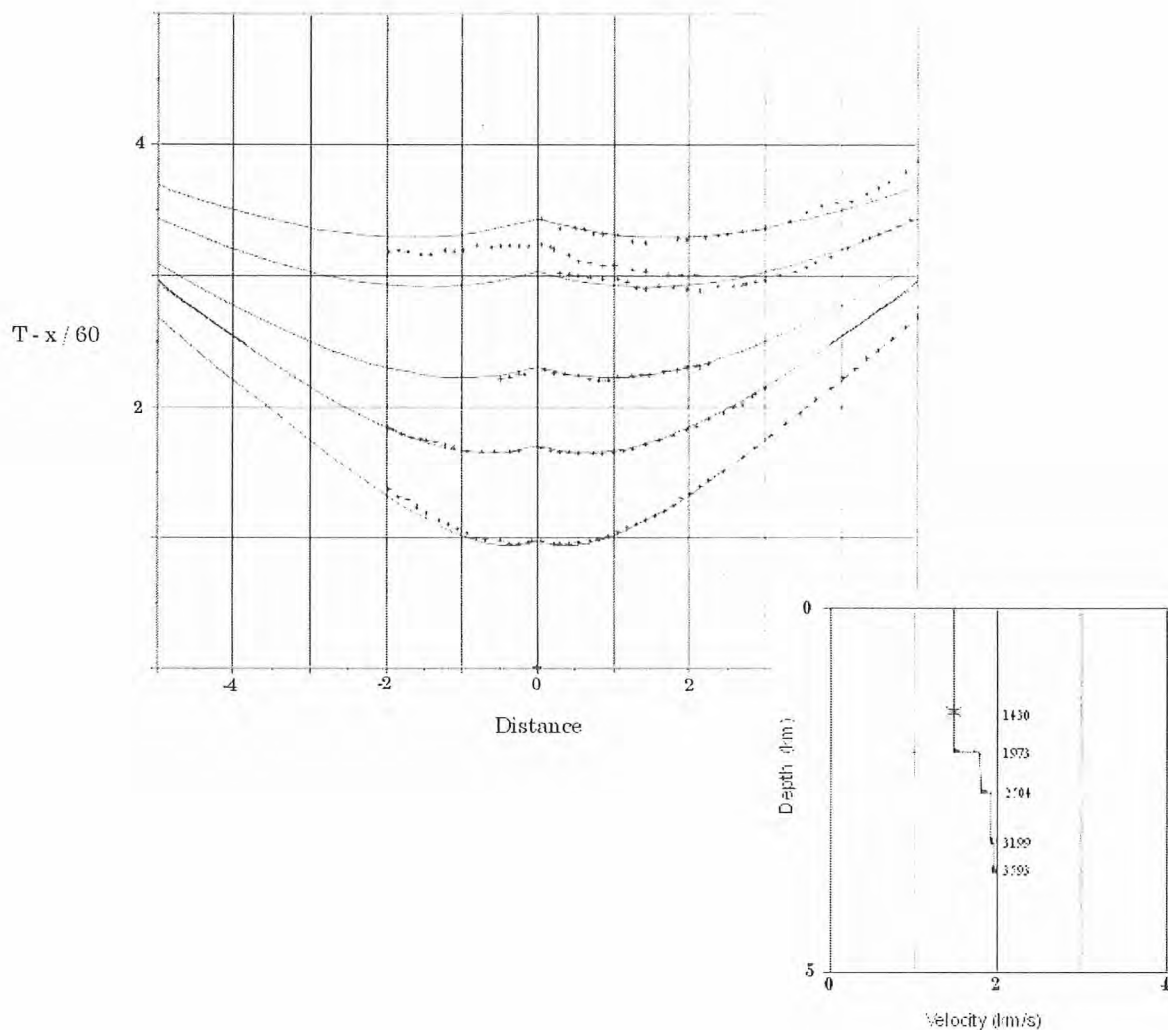


Figure 6.6.1.14: Pickings (crosses) from the record of the OBH23, Profile 2, and predicted arrivals according to the 1-D velocity model shown to the right.

The processed record of the OBH23 during the shooting of Profile 2 (Figure 6.6.1.3) was analyzed in order to obtain an 1-D velocity model of the upper sediment cover (~ 4 km). To achieve further studies in the far field, OBH23 was located 500 m above the sea floor. The pickings of the record are shown in Figure 6.6.1.14, together with the predicted arrivals modeled with the program R1D (Luetgert, 1992). The curvatures of the predicted arrivals nearly coincide with the observed phases. Four reflections were found, with P-velocities ranging from 1,5 to 2,0 km/s :

Depth (m)	P velocity (km/s)
0 – 1 430 (water)	1,480
1 430 – 1 973	1,468 – 1,766
1 973 – 2 504	1,791 – 1,905
2 504 – 3 199	1,922 – 1,959
3 199 – 3 593	1,953 – 2,027

The Local Earthquake Network (EQNET)

(I. Arroyo, J. Gossler, P. O. Thierer)

Description of the network

On March, 16, 2002, a total of 10 Ocean Bottom Hydrophones were deployed in the Pacific Ocean off Costa Rica in order to record local earthquakes and microseismicity data of the OFOS 10 mudvolcano (9° 39.450 N and 85° 52.900 W). Furthermore, active GI-Gun shooting along the array, following the base lines of the array should provide a dense quasi 3-D dataset of the volcano. The duration of the passive part of the experiment is about 6 weeks. Nine stations were installed in three triangles, each triangle is rotated of 60 degrees against their neighboured one and one auxiliary station is placed in the center of the array (see Figure 6.6.1.15).

The main objective of this passive seismological experiment is to detect and register microseismicity events which are effected by the activity of the OFOS 10 mudvolcano. Furthermore, local and regional earthquakes which affects from the subducting Cocos Plate underneath the Caribbean Plate are supposed to be registrated. The experiment might moreover clear up possible relations between both kinds of seismological events, the large scale one in the context to plate dynamics and the local ones, produced by the active mudvolcano.

Active GI-Gun shooting along the the array should provide a dense quasi 3-D seismic dataset which allows to developpe a precise velocity-depth model. Furthermore, when using the GI-Gun signals, a distinct relocation of the instruments could be done.

Two different types of hydrophones were installed; five short-period OAS hydrophones and five Differential Pressure Gauge hydrophones (DPG). Table 6.6.1.1 lists the station number, the hydrophone type used and the exact position of the ship when the station was deployed. These positions are not the final positions of the instruments at the seafloor, because while sinking to the seafloor, the instruments might drift with the current. The determination of the final position would be done by a full joint inversion of earthquake and station location and by using a minimum offset function of the GI-Gun shot points.

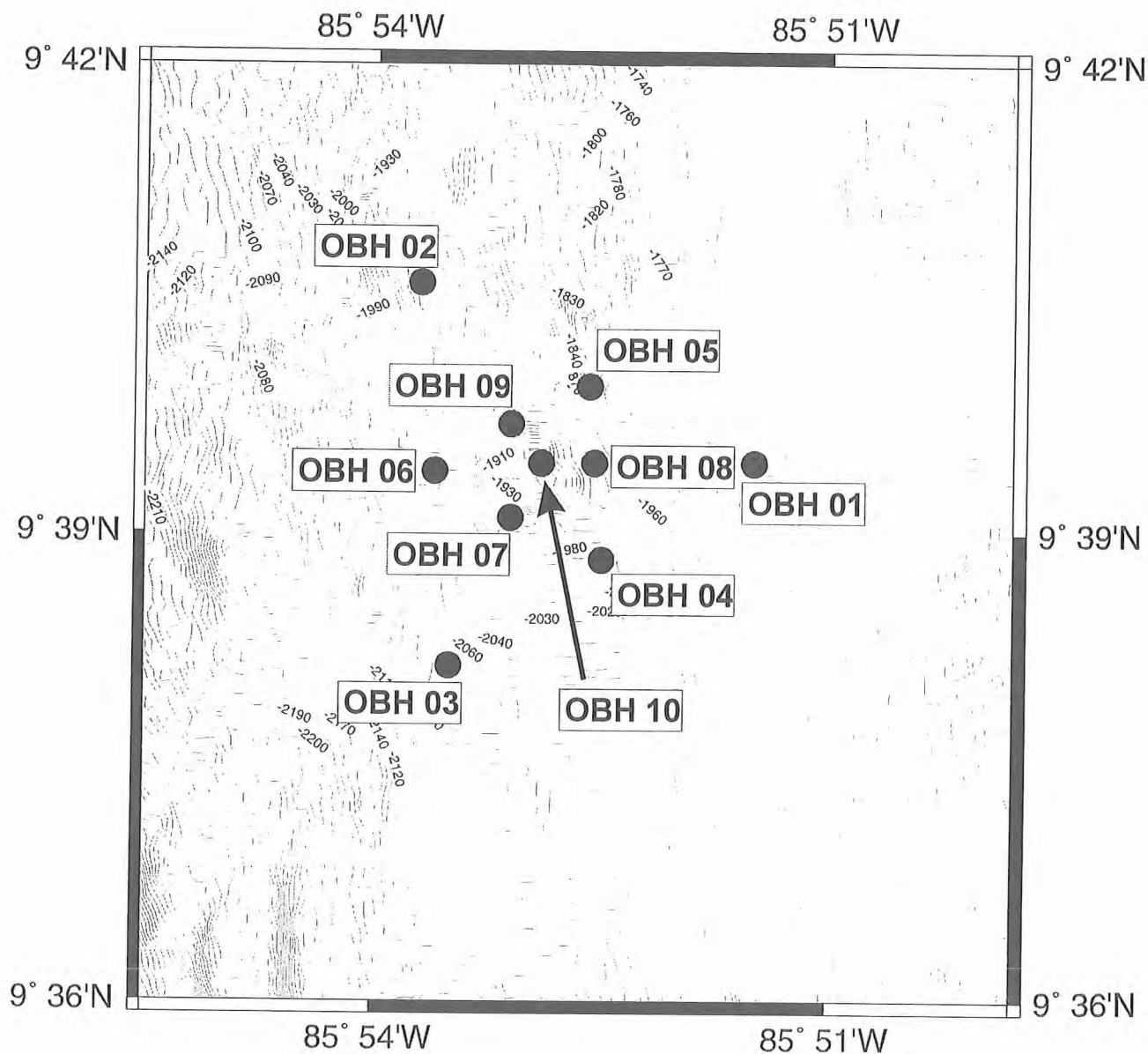
Station Number	Sensor	Latitude	Longitude
OBH 01	OAS	9° 39.456 N	85° 51.494 W
OBH 02	OAS	9° 40.605 N	85° 53.698 W
OBH 03	OAS	9° 38.135 N	85° 53.489 W
OBH 04	DPG	9° 38.849 N	85° 52.508 W
OBH 05	DPG	9° 39.911 N	85° 52.805 W
OBH 06	DPG	9° 39.394 N	85° 53.584 W
OBH 07	OAS	9° 39.930 N	85° 53.109 W
OBH 08	DPG	9° 39.450 N	85° 52.528 W
OBH 09	OAS	9° 39.691 N	85° 53.108 W
OBH 10	DPG	9° 39.461 N	85° 52.905 W

Table 6.6.1.1: Station number, type of sensor and ship position of deployment of the local earthquake network.

The ocean bottom array has been recovered on April 23 and 25, 2002 , at the beginning of RV SONNE cruise 163-2, the second Leg of the “SUBDUCTION” experiment.

Data Quality and Noise

As mentioned above, all stations of the network were recovered successfully. Most of the stations operated as intended and recorded a large number of local earthquakes as well as the airgun source



GMT 2002 Apr 8 22:56:29

Figure 6.6.1.15: Location map of the 10 OBH, covering the OFOS 10 mud volcano

from MCS profiles 01 to 06 shot during leg 2 of SO163 (Figure 6.6.1.1). However, some problems occur with the instruments.

The DPG pressure sensor of station obh05 shows up with an almost complete loss of useful data, instead it produced only spikes and noise. This kind of sensor seems to be very delicate, because it is known for problems. The hydrophones of station obh01 and obh09 sometimes ran out of range. Because this problem occurs more often during the cruise with different instruments, it should be subject of a more detailed error analysis.

Apart from these faults the data from the EQNET network are very noisy, overall, which must be connected to environmental influences in this area. The network status is shown in table 6.6.1.2.

[illegible]

Tabelle 6.6.1.2: Station status of EQNET seismic network

Seismic activity observed by EQNET

During the operating period of EQNET of about 40 days we could identify 184 events (Table 6.6.1.2). For 33 events a preliminary determination of the hypocenter was made now (Figure 6.6.1.16).

We found a very strange looking kind of event, that we called “Hongo”, which is Spanish and means mushroom, because of their shape (Figure 6.6.1.17-19). Those events, which have an almost dual-frequent spectrum (Figure 6.6.17), consist of a high frequent first part which lasts about 1 second with a frequency of about 36 Hz. The second part has smaller amplitudes, lasts about 2-3 seconds at a frequency of about 20 Hz.

We could identify 25 of those Hongo events and fairly locate most of them in the vicinity of the OFOS 10 mud volcano (Figure 6.6.1.16). Another two Hongos are located on the edge of the depletion where a big subducted seamount entered the continental plate. The others are located on the continental margin.

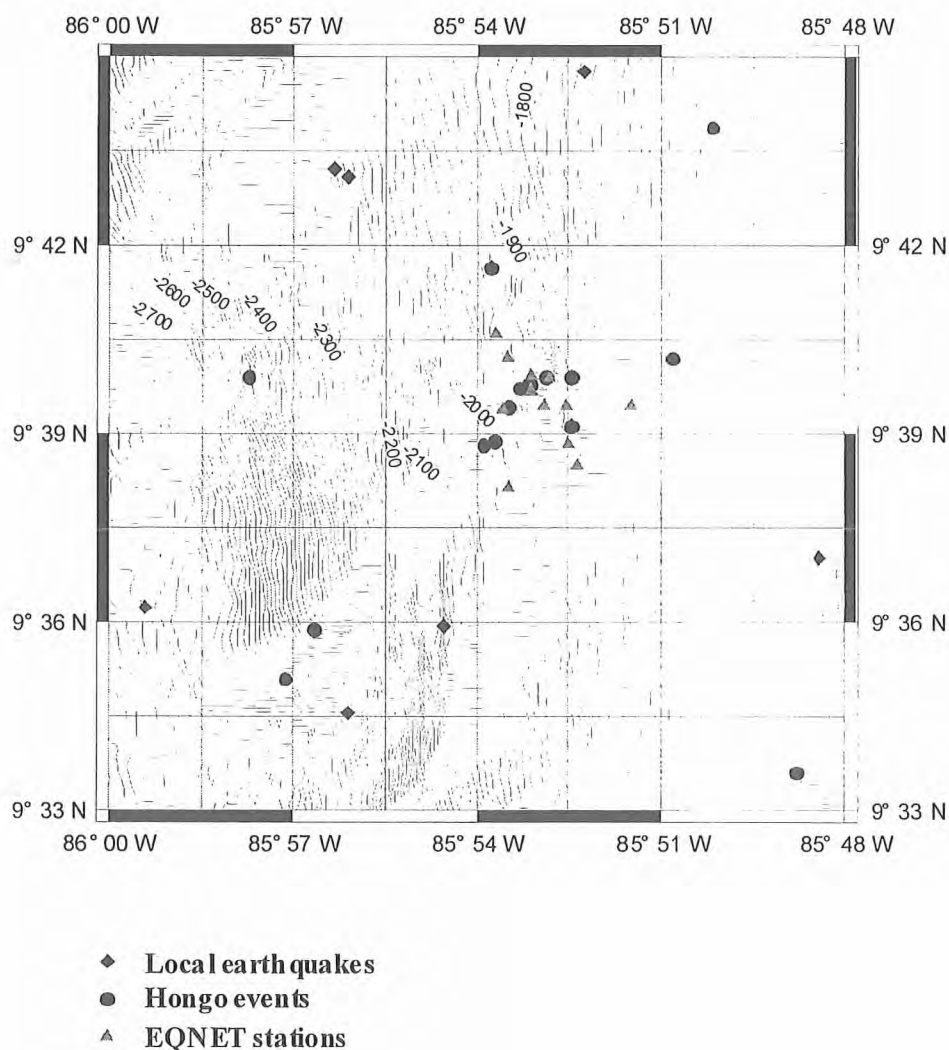


Figure 6.6.1.16: Epicenter locations of events located with the EQNET seafloor seismic network. The so-called Hongo events are located mainly very close to the OFOS 10 mud volcano in the center of the map. Some events occurred in the vicinity of the scarp of a big seamount. The EQNET array is marked due to better orientation.

Date			Time			Type	Latitude	Longitude	Depth (km)	Magnitude
2002	3	16	21	51	16.0	D				
2002	3	17	2	52	37.0	L	9.744	-85.759	43.0	3.2C
2002	3	17	4	29	0.0	L				
2002	3	17	8	44	9.8	R	11.177	-87.412	20.0	4.7C
2002	3	17	8	49	33.9	L	9.718	-85.935	14.3	4.6C
2002	3	17	9	1	4.5	L				
2002	3	17	12	40	3.5	L	9.776	-85.860	18.2	
2002	3	18	6	43	45.0	L				
2002	3	18	9	42	11.0	L				
2002	3	18	15	33	5.0	L				
2002	3	19	1	55	6.0	L				
2002	3	19	22	30	36.0	L				
2002	3	20	2	30	39.5	L				
2002	3	20	3	13	22.0	L				
2002	3	20	5	28	11.0	L				
2002	3	20	6	17	43.0	L				
2002	3	20	6	58	45.0	L				
2002	3	20	11	44	0.0	L				
2002	3	20	20	25	48.0	L	9.599	-85.909	13.9	2.8C
2002	3	21	3	1	26.3	L				
2002	3	21	12	18	5.0	L				
2002	3	21	13	28	3.0	L				
2002	3	21	16	17	9.7	L	9.576	-85.935	13.9	2.7C
2002	3	21	16	30	44.6	L	9.831	-85.716	43.7	3.1C
2002	3	21	17	29	42.6	H	9.585	-85.952	2.0	
2002	3	21	17	42	51.1	H	9.560	-85.813	2.0	
2002	3	21	17	57	13.8	H	9.598	-85.944	1.5	
2002	3	21	18	13	4.0	H				
2002	3	21	18	45	55.0	L				
2002	3	21	18	53	1.0	L				
2002	3	21	19	34	28.3	H	9.651	-85.744	1.0	
2002	3	21	20	16	31.0	H				
2002	3	22	3	27	22.2	L				3.3C
2002	3	22	3	48	26.0	L				
2002	3	22	4	37	29.0	L				
2002	3	22	10	29	20.7	H	9.752	-86.044	1.0	
2002	3	22	19	19	59.0	L				
2002	3	22	19	48	1.0	H				
2002	3	22	20	20	46.2	H	9.731	-85.836	0.7	
2002	3	22	20	28	38.4	H	9.665	-85.874	0.5	
2002	3	22	20	35	2.7	H	9.663	-85.886	1.7	
2002	3	22	20	43	26.9	H	9.662	-85.888	0.7	
2002	3	22	20	56	52.0	L				
2002	3	22	23	43	4.0	L				
2002	3	23	13	24	12.0	L				
2002	3	23	13	42	12.4	L	9.617	-85.807	5.0	
2002	3	23	17	31	37.0	H				
2002	3	23	17	40	46.0	H				
2002	3	24	1	27	36.0	L				
2002	3	24	3	29	16.0	L				
2002	3	24	6	32	37.5	L	10.112	-85.883	86.8	3.7C
2002	3	24	12	15	29.0	L				
2002	3	24	18	43	1.0	L				
2002	3	25	0	30	57.0	L				
2002	3	25	1	19	53.0	L				
2002	3	25	4	36	43.4	L				
2002	3	25	6	19	5.0	L				
2002	3	25	8	11	32.7	L	9.885	-86.117	69.9	4.4C
2002	3	25	10	25	40.0	L				
2002	3	25	11	16	15.9	L	9.746	-85.871	18.7	
2002	3	25	12	0	20.0	L				
2002	3	25	12	12	27.0	L				

Table 6.6.1.2: Seismic events registered by EQNET station network (L = local, H = hongo event)

Date	Time	Type	Latitude	Longitude	Depth (km)	Magnitude
2002 3 26	2 59 10.0	L				
2002 3 26	3 14 41.2	L				
2002 3 26	4 47 31.0	L				
2002 3 26	17 49 6.0	L				
2002 3 26	18 36 10.0	L				
2002 3 26	21 50 29.0	L				
2002 3 27	2 11 15.3	L	9.852	-85.941	43.8	
2002 3 27	5 54 30.0	L	9.720	-85.939	21.2	1.8C
2002 3 27	6 24 3.0	L				
2002 3 28	5 0 11.0	L				
2002 3 28	16 13 0.0	L				
2002 3 28	16 38 45.0	L				
2002 3 28	21 30 44.0	L				
2002 3 31	7 9 18.0	L				
2002 3 31	18 24 53.3	L	9.491	-85.749	1.7	
2002 3 31	23 8 54.0	L				
2002 3 31	23 16 50.0	L				
2002 4 1	1 10 9.0	L				
2002 4 1	20 4 15.0	L				
2002 4 2	1 33 12.0	L				
2002 4 2	8 24 36.0	L				
2002 4 3	2 11 47.0	L				
2002 4 3	8 37 36.0	L				
2002 4 3	8 45 21.0	L				
2002 4 3	10 35 26.0	L				
2002 4 3	16 26 9.0	L				
2002 4 3	21 53 25.0	L				
2002 4 4	4 20 28.1	L	9.731	-85.734	4.2	3.4C
2002 4 4	6 5 17.0	L				
2002 4 4	19 59 20.0	L				
2002 4 5	4 25 46.0	H				
2002 4 5	5 49 46.7	H	9.694	-85.896	9.5	
2002 4 5	5 57 59.3	H	9.665	-85.962	0.9	
2002 4 5	6 7 41.0	H	9.652	-85.874	2.5	
2002 4 5	6 13 45.2	H	9.648	-85.895	2.0	
2002 4 5	6 19 55.3	H	9.647	-85.898	1.5	
2002 4 5	6 25 58.8	H	9.657	-85.891	1.9	
2002 4 5	6 37 42.2	H	9.670	-85.847	1.4	
2002 4 5	6 46 24.6	H	9.663	-85.885	1.7	
2002 4 5	6 52 21.6	H	9.665	-85.881	1.0	
2002 4 5	14 42 58.0	L				
2002 4 5	15 31 2.0	L				
2002 4 5	21 0 53.0	L				
2002 4 6	3 26 42.0	L				
2002 4 6	6 24 49.0	L				
2002 4 6	9 23 38.0	L				
2002 4 6	9 31 35.0	L				
2002 4 6	16 38 40.7	H	9.604	-85.990	2.0	
2002 4 6	20 41 27.0	L				
2002 4 7	4 18 27.0	L				
2002 4 7	11 33 46.0	L				
2002 4 7	14 56 46.0	L				
2002 4 7	18 44 42.0	L				
2002 4 7	19 2 36.0	L				
2002 4 7	20 47 21.0	L				
2002 4 8	5 36 26.0	L				
2002 4 8	11 51 7.0	L				
2002 4 8	20 24 54.0	L				
2002 4 9	4 58 13.0	L				4.9C
2002 4 9	7 38 9.0	L				
2002 4 9	19 20 41.0	L				
2002 4 9	22 52 5.0	L				

Table 6.6.1.2: (continued)

Date	Time	Type	Latitude	Longitude	Depth (km)	Magnitude
2002 4 9	23 30 49.0	L				
2002 4 10	0 32 52.0	L				
2002 4 10	0 41 4.0	L				
2002 4 10	1 11 40.0	L				
2002 4 10	20 16 53.0	L				
2002 4 11	0 0 43.0	L				
2002 4 11	0 27 34.0	L				
2002 4 11	9 50 30.0	L				
2002 4 11	10 13 37.0	L				
2002 4 11	13 23 42.0	L				
2002 4 11	20 43 35.0	L				
2002 4 11	21 39 15.0	L				
2002 4 12	0 8 57.0	L				
2002 4 12	15 6 20.0	L				
2002 4 12	18 15 4.0	L				
2002 4 13	19 28 34.0	L				
2002 4 14	5 37 26.0	L				
2002 4 15	9 17 44.0	L				
2002 4 15	9 24 31.0	L				
2002 4 15	10 7 54.0	L				
2002 4 15	10 11 33.0	L				
2002 4 15	10 28 31.0	L				
2002 4 15	10 32 7.0	L				
2002 4 15	10 53 36.0	L				
2002 4 15	19 4 48.0	L				
2002 4 16	16 20 21.0	L				
2002 4 16	21 31 32.0	L				
2002 4 17	4 53 23.0	L				
2002 4 17	5 43 3.0	L				
2002 4 17	12 35 40.0	L				
2002 4 17	16 1 53.0	L				
2002 4 17	18 51 47.0	L				
2002 4 17	21 30 54.0	L				
2002 4 17	21 34 22.0	L				
2002 4 17	22 0 34.0	L				
2002 4 18	5 3 45.0	L				
2002 4 18	7 41 53.0	L				
2002 4 18	8 46 21.0	L				
2002 4 18	12 15 28.0	L				
2002 4 18	12 18 43.0	L				
2002 4 18	14 12 35.0	L				
2002 4 18	16 13 7.0	L				
2002 4 18	17 58 40.0	L				
2002 4 18	18 15 21.0	L				
2002 4 18	18 18 40.0	L				
2002 4 19	1 42 59.0	L				
2002 4 19	2 14 39.0	L				
2002 4 19	2 45 56.0	L				
2002 4 19	3 16 21.0	L				
2002 4 19	11 26 39.0	L				
2002 4 20	9 21 29.0	L				
2002 4 21	12 19 25.0	L				
2002 4 21	19 9 41.0	L				
2002 4 21	19 57 11.0	L				
2002 4 21	20 21 24.0	L				
2002 4 22	3 13 42.0	L				
2002 4 22	11 27 10.0	L				
2002 4 22	14 2 34.0	L				
2002 4 23	5 16 6.0	L				
2002 4 23	7 48 31.0	L				

Table 6.6.1.2: (continued)

We can imagine, that these Hongo events are caused by fluid currents which are pressed through a tube-like system, or something similar which lasts for about one second, e.g. the first part of the event, followed by a depression where those fluids may expand or flow back what can be observed as the second part of the event which lower frequency and amplitudes.

Further investigations should be done due to these strange events.

Further on we observed 53 emergent events which have not been located yet, due to the impossibility of identifying phase onsets (Figure 6.6.1.20). The origin of those events is not clear yet, although we know them from volcanoes. There, emergent events are tectonic events that run through a highly scattering media (U. Wegler, pers. com.).

Two tectonic events have been located in the vicinity where the subducted seamount entered the continental lithosphere (Figure 6.6.1.16). Other earthquakes are spread over the continental margin.

Only one regional earthquake from Nicaragua has been registered during the operating period (Figure 6.6.1.21)

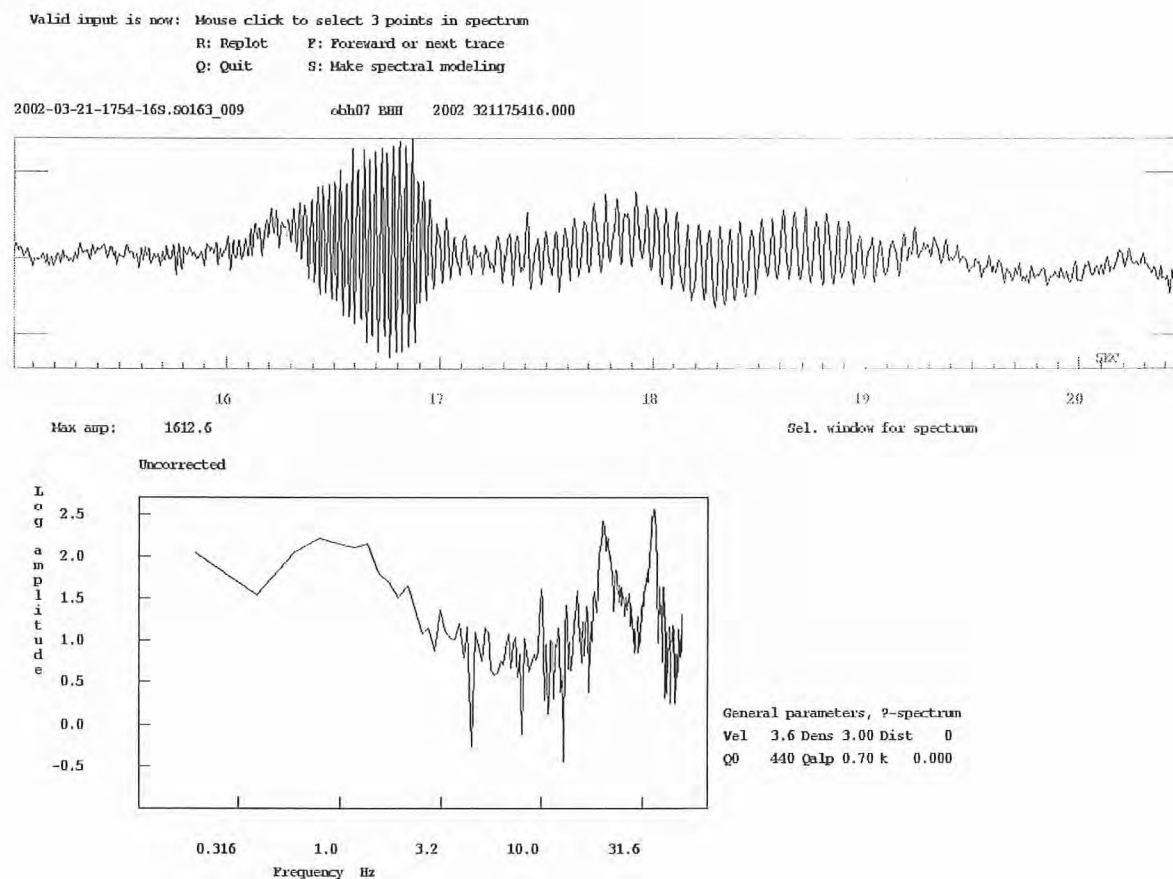


Figure 6.6.1.17: Hongo event from March 21, 2002 17:57 UTC registered at OBH07. The event is divided into two parts: the first one takes one second with a frequency of about 36 Hz (second peak in the spectrum below), while the second part has smaller amplitudes and stretches over 2.5 seconds. Ist frequency is at about 20 Hz (first prominent peak in the spectrum). The event may be connected to fluid processes of the OFOS 10 object.

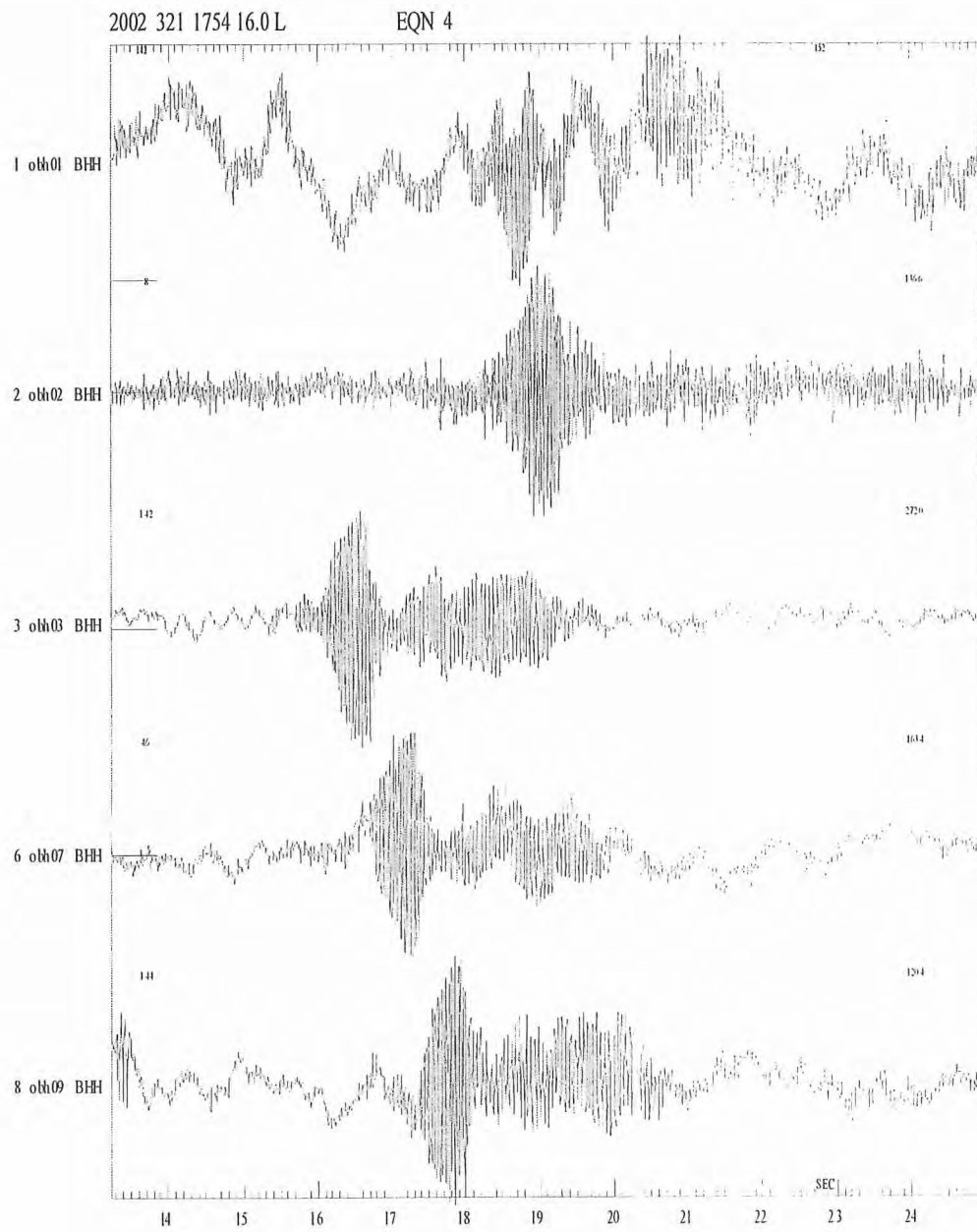


Figure 6.6.1.18: Seismograms of the Hongo event from previous figure, but for all stations of EQNET seafloor network.

SO163 2002-03-22-2017-42S.SO163_009
Plot start time: 2002 3 22 20:20 46.937

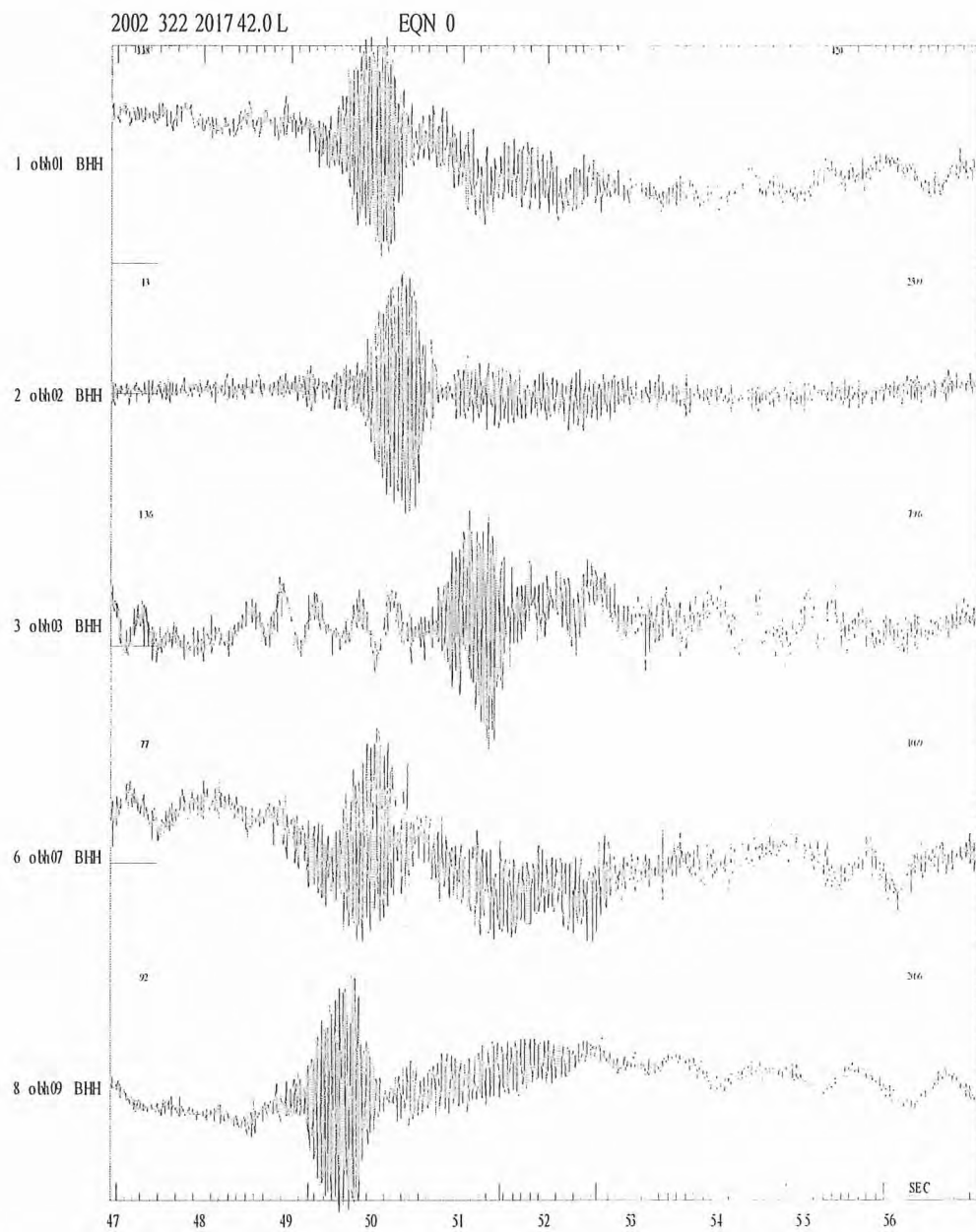


Figure 6.6.1.19: Another Hongo event from March 22, 2002 20:20 UTC observed at EQNET.

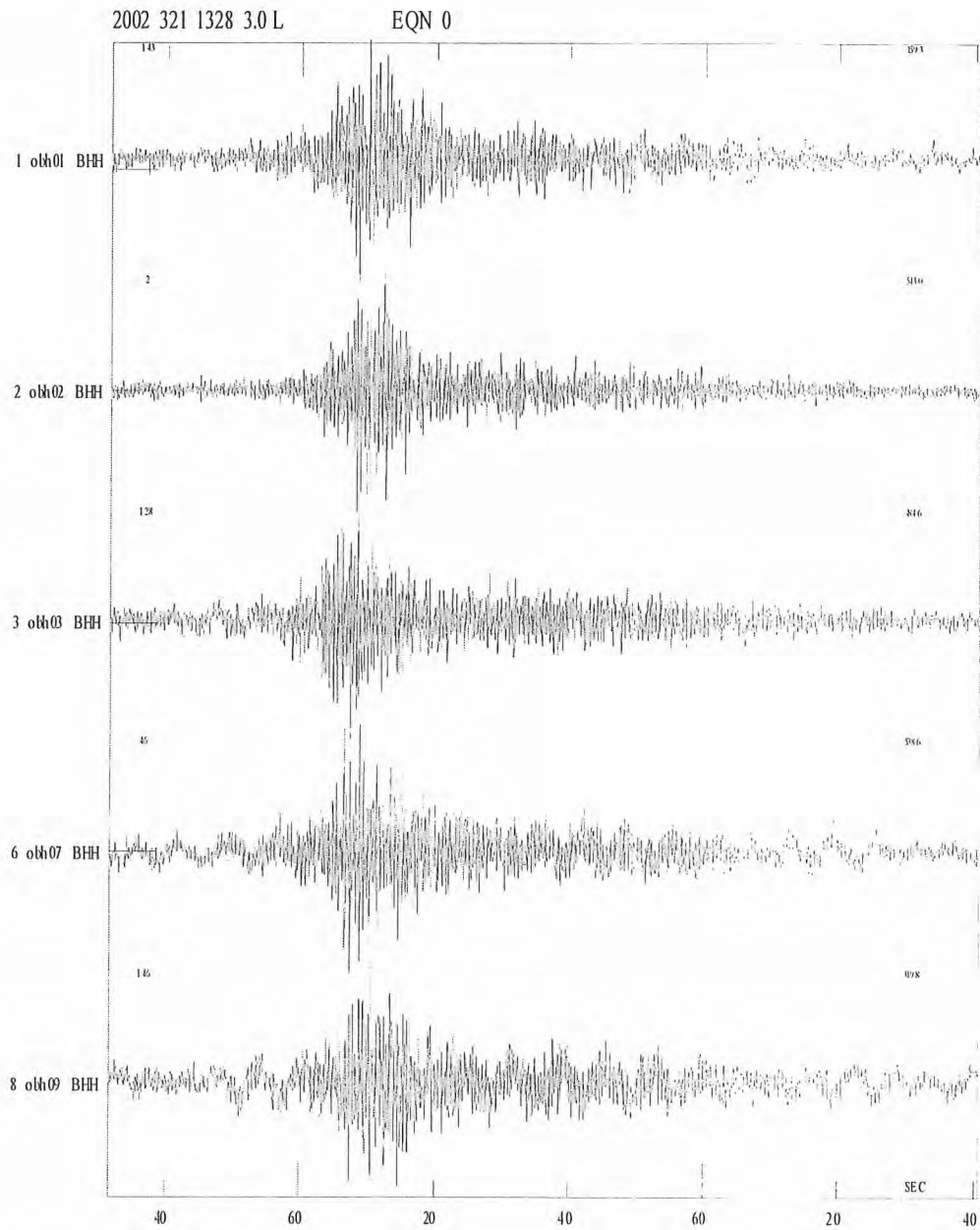


Figure 6.6.1.20: Emergent event observed on March 21, 2002 13:30 UTC at EQNET. The event cannot be easily located due to missing phase onsets. Those events are specific to volcanoes.

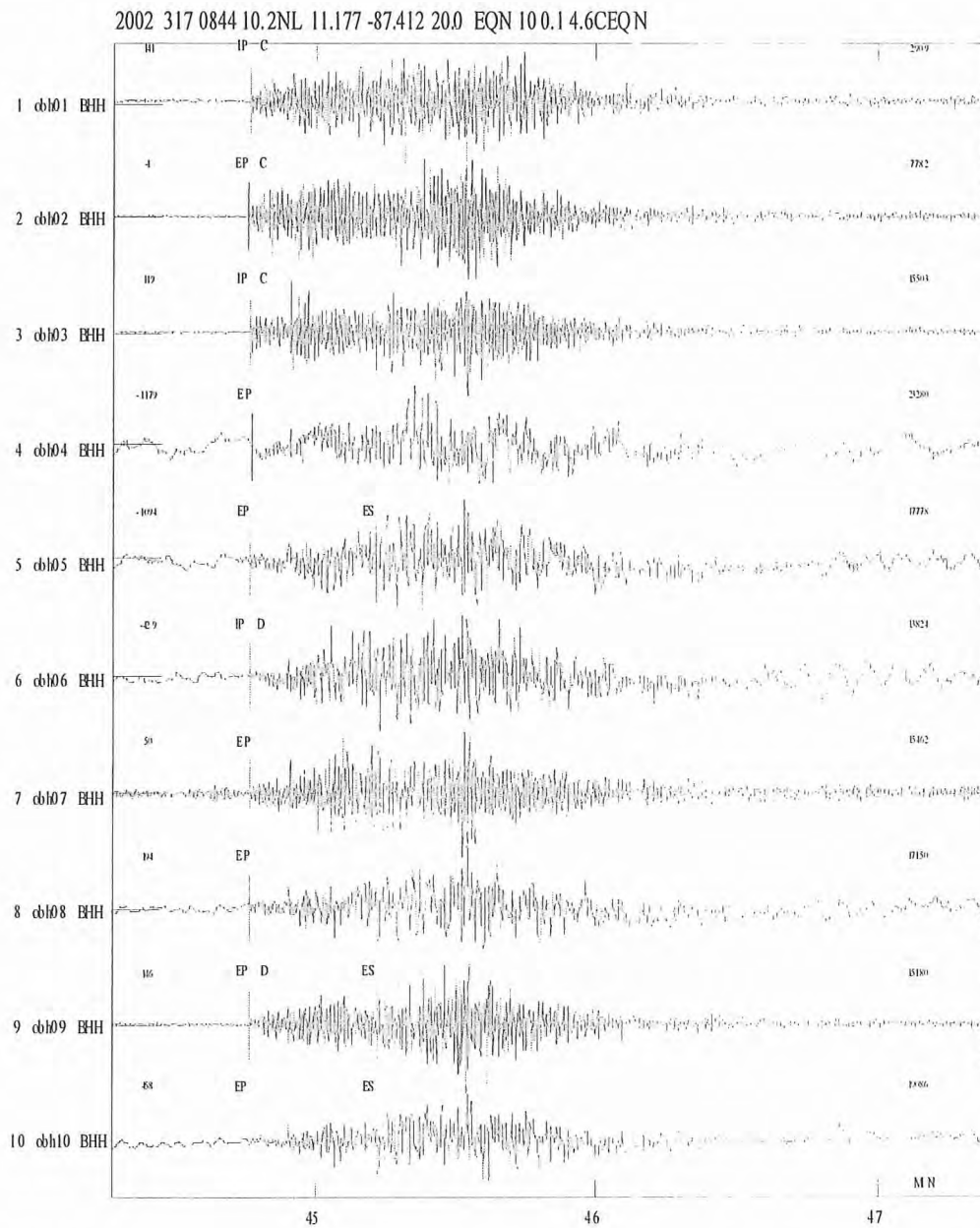


Figure 6.6.1.21: A magnitude 4.6 (Coda-magnitude) regional earthquake from Nicaragua on March 17, 2002 08:44 UTC. Localization was made with EQNET station array using mostly P onset. For distant events the earthquake waves are steep, so that their S waves oscillate parallel to the seafloor and do not produce pressure signals that can be registered from the OBH.

6.6.2 Gas-hydrates on Nicoya Slide

(D. De Nil, J. Gößler, A. Berhorst)

One objective of project B1 of the Sonderforschungsbereich 574 is to develop and to apply a new technique for determining shear wave velocity of marine gas-hydrates in order to discriminate different ways, in which gas-hydrate may be incorporated in sediment, as well as to estimate its amount. Gas-hydrates may be part of the pore fluid, part of the sediment matrix, or act as a cement for sediment grains in different ways. Especially, if the hydrate acts as a cement as described for cores from site 1041 of the Ocean Drilling Program (Kimura et al., 1997), shear wave velocity is increased significantly, even for small amounts of hydrate, because of the significantly risen shear strength. Discrimination of the different lithologies is especially important for understanding the role of gas-hydrates in slope stabilization and slope failures. Large slope failures may cause tsunamis and result in emitting large amounts of climate-relevant methane gas. Free gas that has been trapped beneath gas-hydrates may escape through faults and gas-hydrate in the sediment may dissociate due to reduced overburden load. Near ocean-bottom hydrate may also ascent directly due to its buoyancy.

An indicator for the presence of gas-hydrates are bottom-simulating reflectors (BSR), which are more or less distinct reflectors with a usually negative polarity due to an inverse velocity contrast caused by the bottom of hydrate-bearing sediment and possibly free gas beneath. They approximately follow the ocean-bottom reflection in a distance according to the pressure and temperature of the stability zone of gas-hydrates. The effect of the presence of the hydrate can be estimated by determining shear wave velocity of gas-hydrate saturated sediment above the BSR and the sediment without gas-hydrates below the BSR. The absence of a BSR does not indicate the absence of gas-hydrates. At the above mentioned drilling site off-shore Nicoya, gas-hydrates have been found included into sediment layers in different ways, but a BSR was not observed. However, for SONNE cruise SO-163-2 locations with relative strong BSRs have been chosen, because of the high dilatational wave velocity contrast there, which implies a possibly high shear wave velocity contrast. Furthermore, dilatational wave velocity and subsurface structure are already well determined at the chosen locations, so that shear wave velocity need not be determined absolutely, but can be determined by relative travel times as well.

According to the results of Pecher et al. (1998), BSRs observed offshore Costa Rica are best explained by a thin layer containing free gas beneath hydrate-bearing sediments. Poststack time-migrated streamer sections acquired and processed during SONNE cruise SO-163-2 show, that regions with strong BSRs coincide with low amplitudes of later phases (compare e.g. Figure 6.6.2.10), which can be attributed to attenuation by free gas beneath the gas-hydrate stability zone. This observation matches well with the result of Pecher et al. (1998). Since free gas can escape along fractures, bottom simulating reflectors are mostly absent in areas affected by slumping. A prominent example where a BSR has been observed in a structure interpreted as slump mass, is Nicoya slide. One possible explanation is, that the sediment slumped as a block and stratification was not disrupted at this location so that the gas cannot escape. Another possible explanation is, that the BSR re-formed after slumping. Different origins might result in a different fine structure causing different seismic velocities.

Profiles P09 and P10 (Figure 6.6.2.1a) follow the seismic reflection line SO-81-21 and refraction line SO-81-300 (Hinz et al. (1992), Pecher et al. (1998) and Ye et al. (1996)) across Nicoya Slide. The data has been recorded by the four-channel-streamer and ocean bottom seismometers and hydrophones. The new data complements SO-81-21 and SO-81-300 especially by three-component registration of reflected and refracted waves. While profile P09 was shot by the 321-Bolt-gun (with 7.5 m cable to floatations, about 34 m behind and about 3 m portside the vessel), profile P10 was shot by the four-G-gun-cluster (with 4 m cable to floatations, portside guns about 32 m and starboard side guns about 22.5 m behind the vessel). A group of four OBSs and five OBHs, including one OBH placed 500 m above seafloor for recording the far-field source signal, was deployed onto Nicoya Slide, a smaller group consisting of one OBS and two OBHs was deployed north-east of the headwall landward of this slide (Figure 6.6.2.1b, respectively Appendix 9.5). The smaller group has been recovered after

shooting of line P10 was completed, while the larger group was complemented by two test-instruments (OBH36 and OBH37). Sampling rate was 200 or 250 Hz, depending on the capabilities of the instruments. The following profiles P11, P12, P13 and P14 have been shot mainly for testing different source configurations (varied cable lengths of the G-gun-clusters of 4m and 6m, varied distances behind the vessel and lateral symmetries) for later measurements and for enabling a three-dimensional localization of the larger group of OBSs and OBHs by cross-lines. These profiles also provide additional information about lateral variations near to the ocean bottom. After profile 14 was finished, all instruments were recovered successfully. Except for OBS 24 and one approximately horizontal channel of OBS 25, all ocean bottom instruments recorded useful data. Standard processing as described in Section 6.6.1 has been applied. Selected sections are presented in Figures 6.6.2.2 ff.

The individual locations where the instruments were deployed onto the slump mass were mainly chosen because of the seismically almost transparent region below the BSR on the slide observed in migrated MSC sections provided by C.R. Ranero (personal communication). The stations landward of the headwall serve as similar but non-slided reference. Additionally, waves which entered the sediment through the headwall can be observed at this location, but all in all, subsurface structures appeared relatively simple, compared to other potential locations. According to preliminary two-dimensional re-localization of the ocean-bottom-instruments, their drift was in the order of the smallest planned station distances, but rather similar for all instruments. Deploying the instruments in groups, enables estimating very local effects, e.g. caused by small scaled structures and different coupling due to local variations of the ocean bottom material. Furthermore, the redundant information enables stacking of corresponding components in order to enlarge signal-to-noise ratio. The not completely known effective transfer functions of the rather different kinds of receivers makes comparisons difficult, however, it appears reasonable to assume that receivers of one kind have similar transfer functions. The difference in amplitudes between OBS 26 and OBS 31 might be caused by softer near-surface material at OBS 26. The transfer functions of the receivers shall be examined in more detail in future work, so that receiver effects may be eliminated from the data and differences due to local variations in subsurface and coupling at the surface may be estimated more precisely from different kinds of instruments, too. Unfortunately, during these measurements there were only the five deployed OBSs available. However, at intermediate and large angles of incidence, shear waves have a vertical displacement component and can be recorded by hydrophones as well.

The area around Nicoya Slide will probably be studied during Meteor cruise 54-1 in July and August 2002 by high frequency 3D reflection seismic and further ocean bottom seismic measurements, again. The data acquired during SONNE cruise SO-163-2 serve as a basis for later studies on gas-hydrates at locations with weak and perhaps lacking BSR and more complex subsurface structures. In addition to studies performed on Nicoya Slide during SONNE cruise SO-163-2, we also recorded two profiles offshore north-west Nicoya corresponding the MSC line BRG-99-39 and three profiles at a location on line SO-81-09 which has been studied in details by Pecher et al. (1998) and McGee (2000). This way, we are getting a rough idea of the variety of shear wave characteristics at locations with distinct BSRs offshore Nicoya. The subsurface structure at these locations is relatively simple, too. We chose two-dimensional layouts (Figures 6.6.2.1a, 6.6.2.1c and 6.6.2.1d and station-list in Appendix 9.5), so that we may identify phases arriving from outside the profile plane caused by lateral heterogeneities by array or network techniques. The OBSs were complemented by a larger number of OBHs in order to cover a larger region. Offshore north-west Nicoya, we deployed the stations at 2 km distances covering almost the whole slope, while we had chosen distances of 200 m at the location on SO-81-09 in order to study very local effects.

The source wavelet of the G-gun array is effectively shorter, but especially at large offsets the higher signal amplitudes of the Bolt-gun is advantageous, although the Bolt-gun has a rather powerful bubble interfering with phases to be evaluated. On recordings of some instruments of profiles P09 and P10, diving wave phases from the mantle may be observed even at very large distances up to 100 km, however, mostly the reasonably interpretable data probably ends at about 50 to 75 km due to a small signal-to-noise ratio and/or an insufficient dynamic and pre-amplification, respectively. At very large offsets, distortions from the direct water wave, its multiples and other late phases of the preceding shot are significant at a shot-interval of 60 s, which was chosen as a compromise with respect to spatial resolution. For the sake of higher source energy, profile P32 (corresponding to multi-channel

seismic line BGR-99-39) was shot by the Bolt-gun (with 7.5m cable to floatations, about 34 m behind and 3 m portside the vessel) and portside and starboard side GI-guns (with 4m cables to floatations, about 35 m behind the vessel) combined, while the cross-profile P32 (with 30 s shot interval) and profiles P35 to P37 offshore Nicoya were shot with 60 s shot-interval by the GI-gun cluster, only.

All ocean-bottom instruments deployed for profiles P32 (OBH142-OBH161) and P33 (OBH148-OBH165) and P35 to P37 (OBH173-OBH192) were recovered successfully. OBH146 and the seismometer channels of OBS159 did not record any useful data. The quality of the data of the other stations especially of line P32 and P33 is rather inhomogeneous and some of the stations recorded useful data only temporarily. Especially the data of OBH148, OBH173 and seismometer channels of OBS174 were severely contaminated by electronic noise and several stations e.g. OBH149 and the seismometer-channels of OBS184 recorded rather weak amplitudes. At other hydrophones the direct wave was already clipped. OBH147, OBH163 and OBH173 additionally served as hydrophone tester. Probably, only some of the recordings can be evaluated reasonably, depending on the processing technique. Comparability of different shots of line P33 might be affected by problems with the floatations of the starboard side G-guns resulting in a different source depth. Examples of processed OBH- and OBS-sections are presented in Figures 6.6.2.22 ff. Lines P35-P37 were shot only a few days before the end of the cruise have not been processed, yet. Data examples are presented in Figures 6.6.2.38 ff.

Further processing and analysis of the acquired data to be performed at Kiel University after the expedition includes evaluation of the three-component data by receiver functions (e.g. Yuan et al., 1997) focussing on determining shear wave velocity in the vicinity of the BSR. Receiver functions result of a deconvolution of a dilatational wave signal from a corresponding converted shear wave signal and describe the conversion operator. They are comparable to deconvolution of the source signal from an acoustic signal yielding reflectivity or transmissivity and the corresponding impulse response, respectively. Since dilatational and shear arrivals cannot be separated by component rotation completely in general, and since exact deconvolution of sampled and noisy data is instable (Oldenburg, 1981), the procedure is not completely straight forward. However, several sophisticated approaches to stabilise receiver functions have been developed and successfully applied in seismology, where receiver functions are a well established technique. Refer to Endrun (2001) for an overview and comparison of different approaches to receiver functions with seismological examples. Since active seismic data is not extremely different from earthquake recordings beside a difference in typical main frequencies of the signals and the different scale of the structures we are interested in, it is reasonable to assume that the principle ideas of receiver functions are transferable to active seismics. Possibly, receiver functions are a valuable complementary technique for determining shear wave velocity directly by picking shear arrivals and amplitude versus offset analysis.

Relative high amplitude horizontal motion of reflected and refracted waves has been recorded by several stations. Especially in the recordings of stations on SONNE lines SO-81-21 and SO-21-09, arrivals that are probably P to S-converted shear waves from the BSR can be identified even in unprocessed data (see Figure 6.6.2.47), although e.g. the waveform inversion performed by Pecher (1995) at the location on profile SO-81-09 works well without assuming a shear wave velocity contrast at the BSR which is necessary for the conversion of waves. However, waveform inversion of streamer data is not very sensitive to shear wave velocity and especially according to the models of Carcione and Tinivella (2000) and the results of Tinivella and Accaino (2000), we expect at least a small change in shear wave velocity causing at least a weak shear wave signal that might not necessarily be visible in unprocessed data. For a reliable, detailed interpretation, further processing is necessary, but could not be conducted on board due to a lack of time.

SO-163-2: Profiles P09-P14 & P35-P37 (Nicoya-Slide)

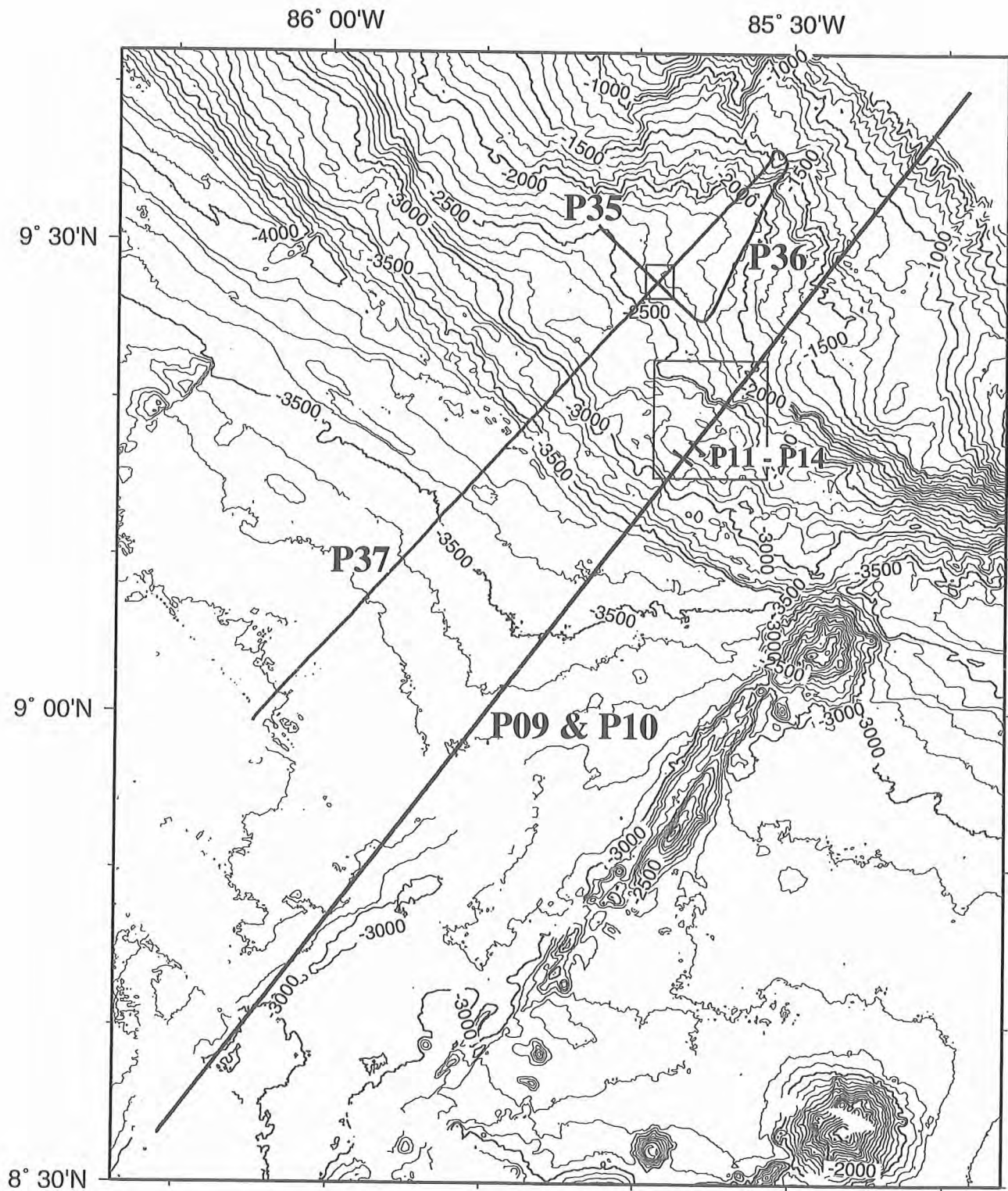


Figure 6.6.2.1a: Location map of profiles P09-P14 and P35-P37 (overview) with 100m isolines, compare figures 6.6.2.1b and 6.6.2.1c.

SO-163-2: Profiles P09-P14 and Stations OBS25-OBH37

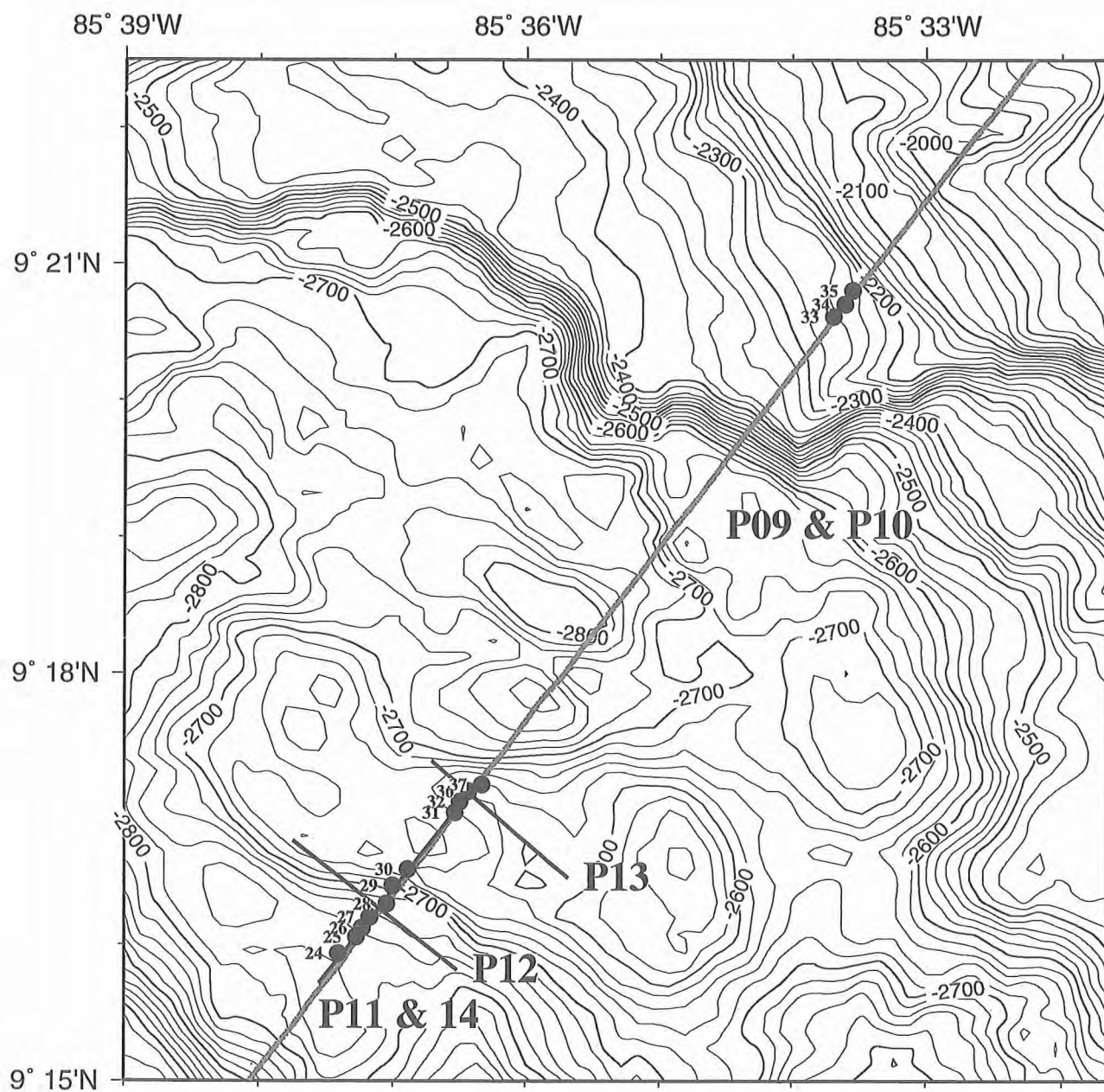


Figure 6.6.2.1b: Location map of profiles P09-P14 and stations OBS25-OBH37 (relocated) with 20m isolines, compare figure 6.6.2.1a.

SO-163-2: Stations OBH173 - OBH 192 of lines P35 - P37

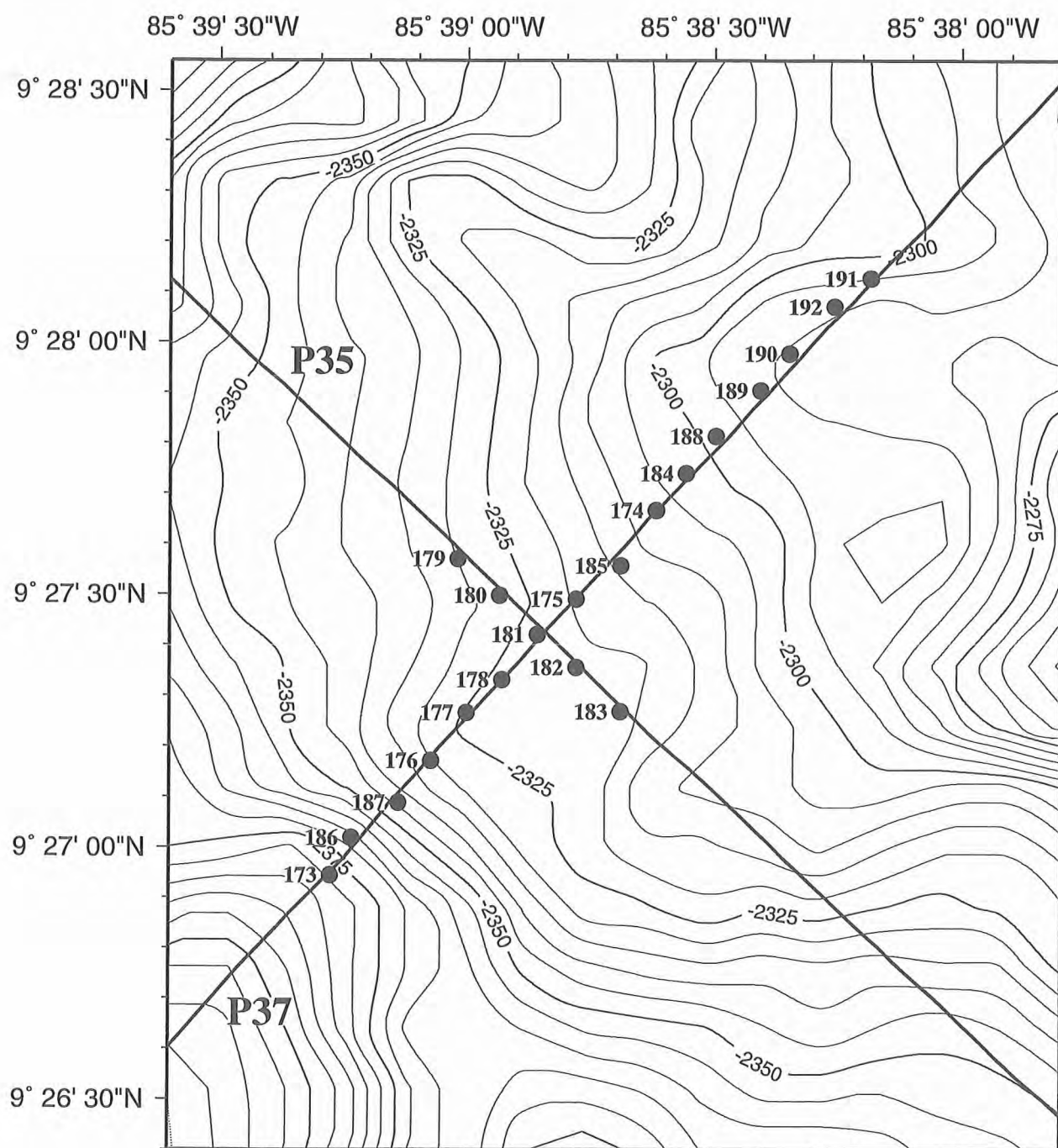


Figure 6.6.2.1c: Location map of OBH173-OBH192 (deployment co-ordinates) on profiles P35 and P37 with 5m isolines, compare figure 6.6.2.1a.

SO-163-2: Profiles P32 & P33 (Gas-hydrates offshore north-west Nicoya)

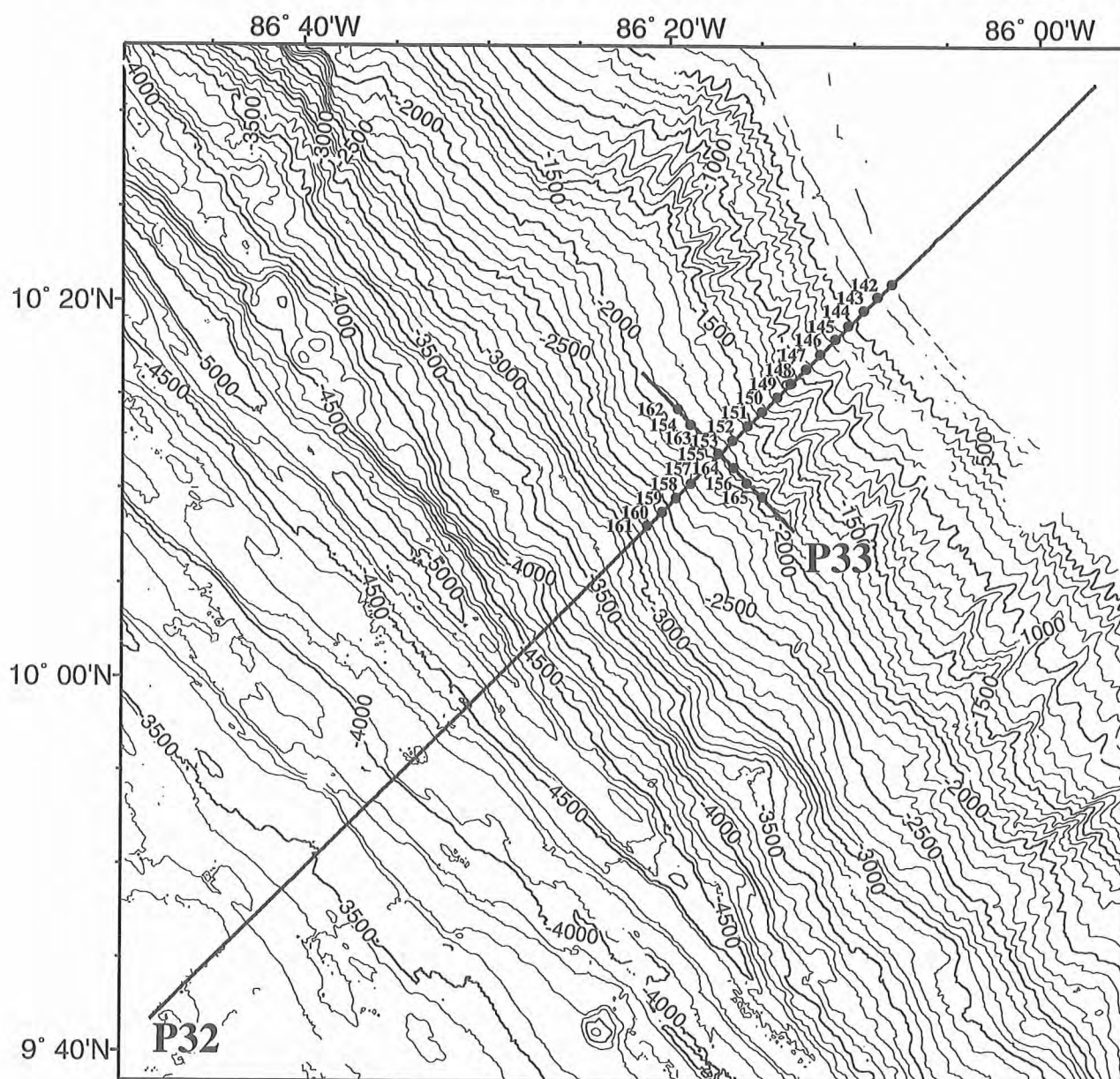


Figure 6.6.2.1d: Location map of profiles P32 and P33 with 100m isolines. Stations OBH142-OBH165 are superimposed.

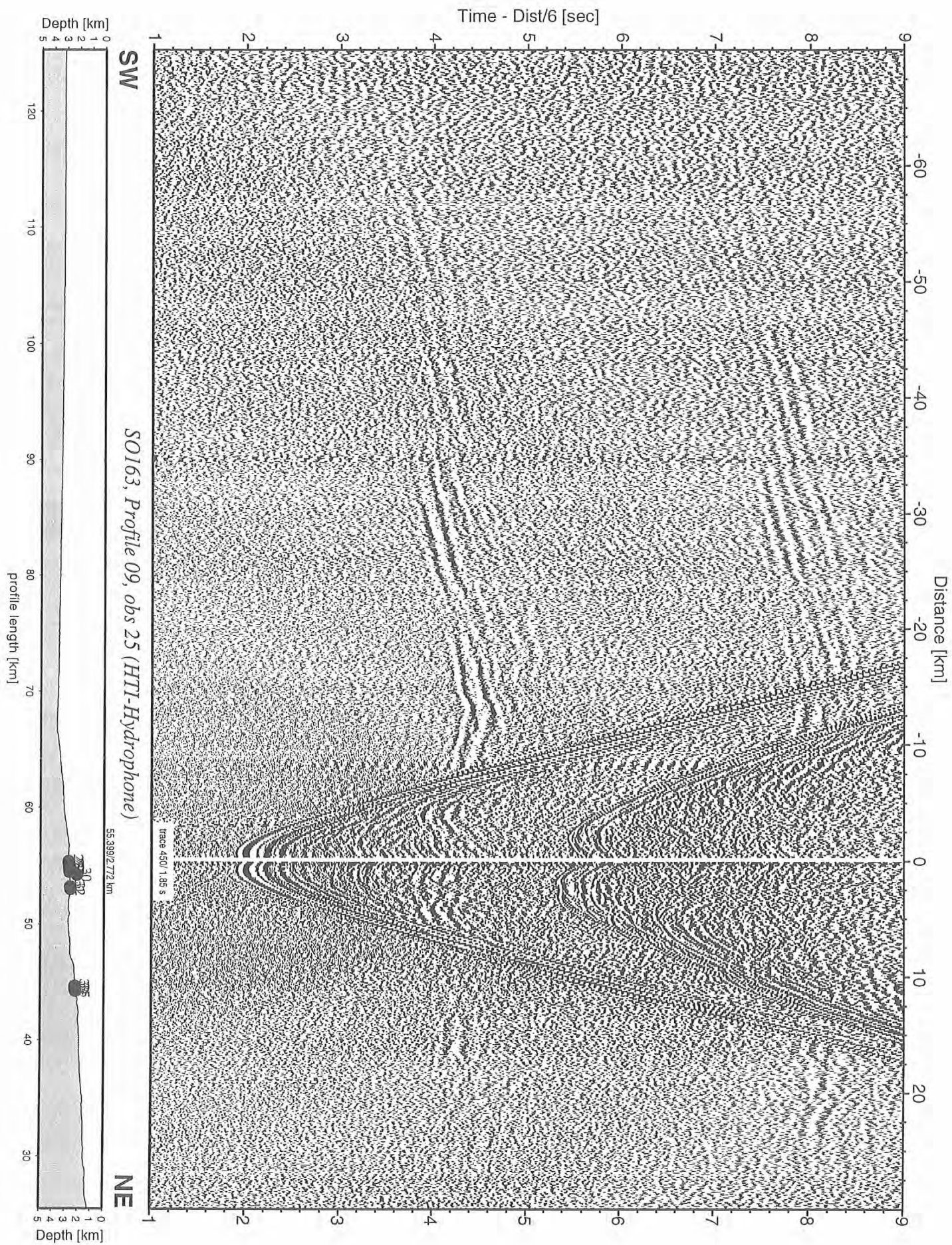


Figure 6.6.2.2: Record section from obs 25 (HTI-Hydrophone), Profile 09.

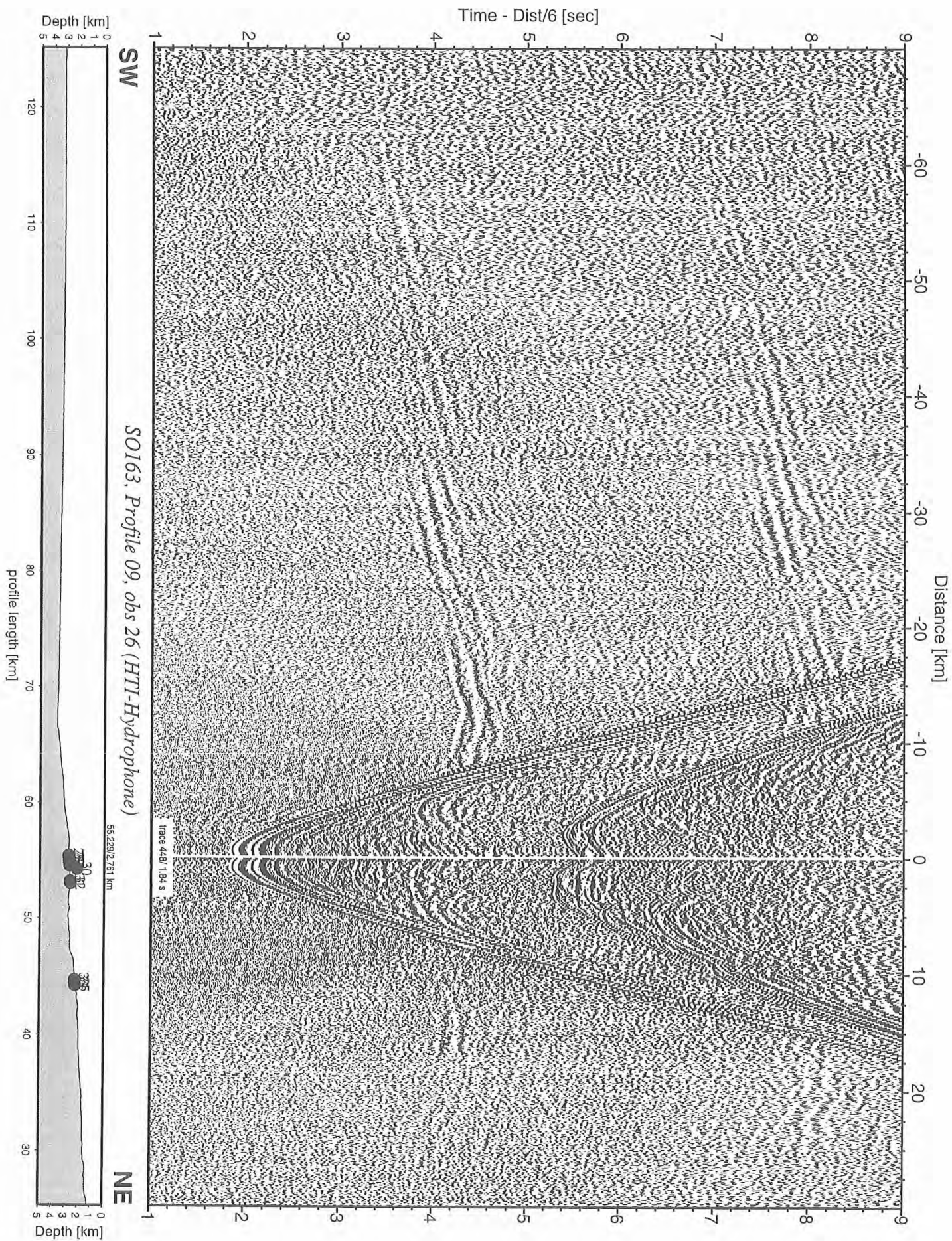


Figure 6.6.2.3: Record section from obs 26 (HTI-Hydrophone), Profile 09.

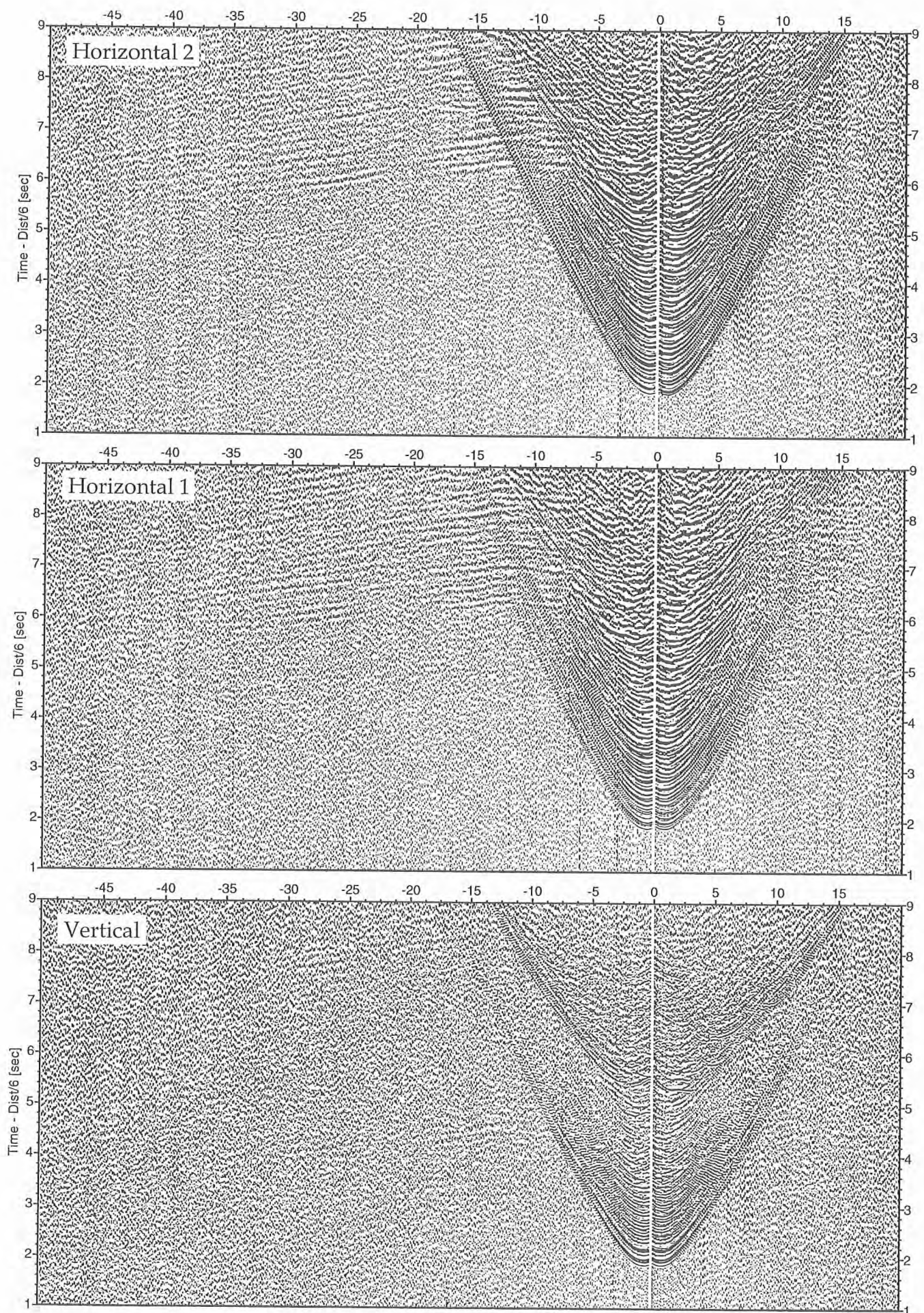


Figure 6.6.2.4: Record sections from obs 26 (Owen-4.5Hz), SO163, Profile 09.

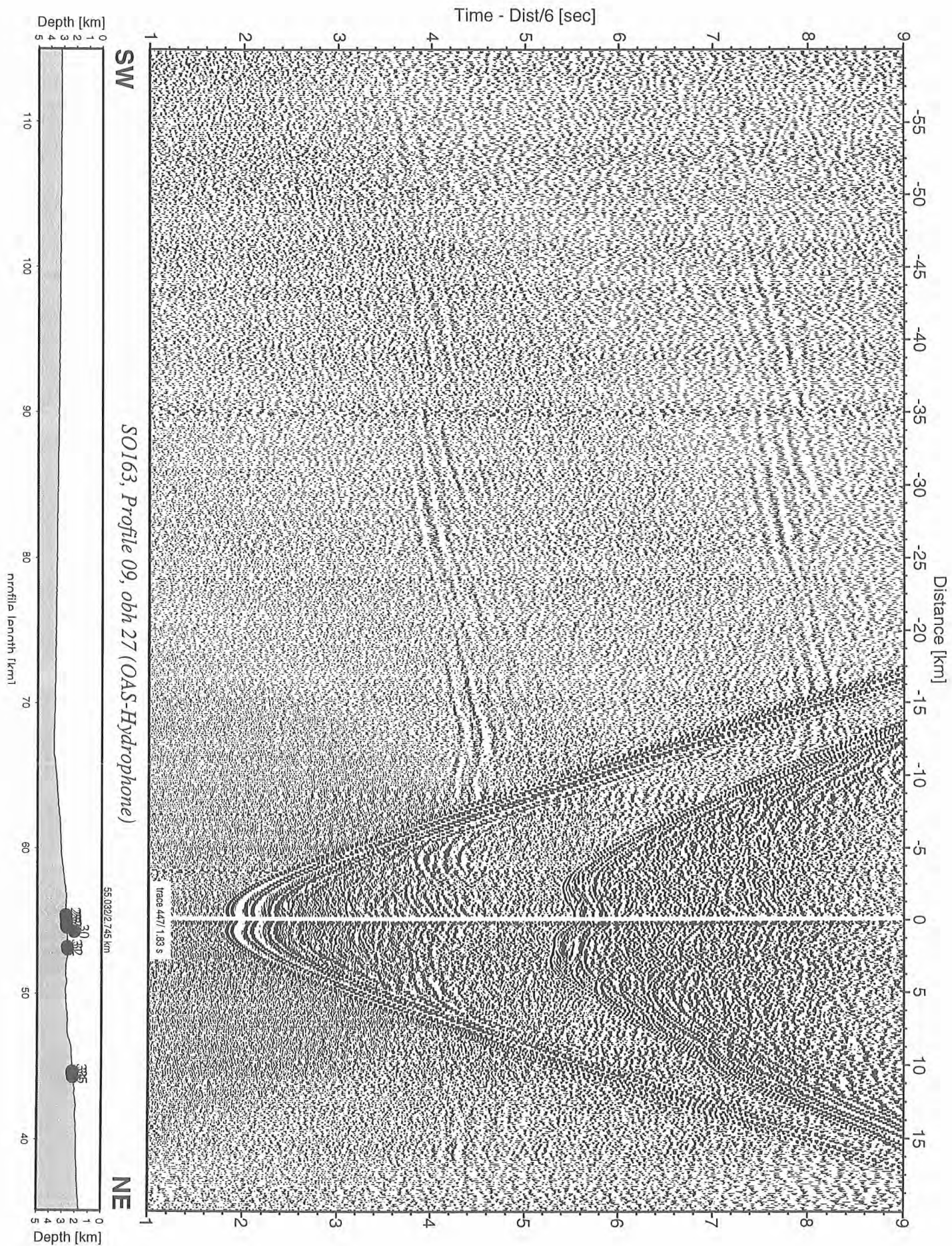


Figure 6.6.2.5: Record section from obh 27 (OAS-Hydrophone), Profile 09.

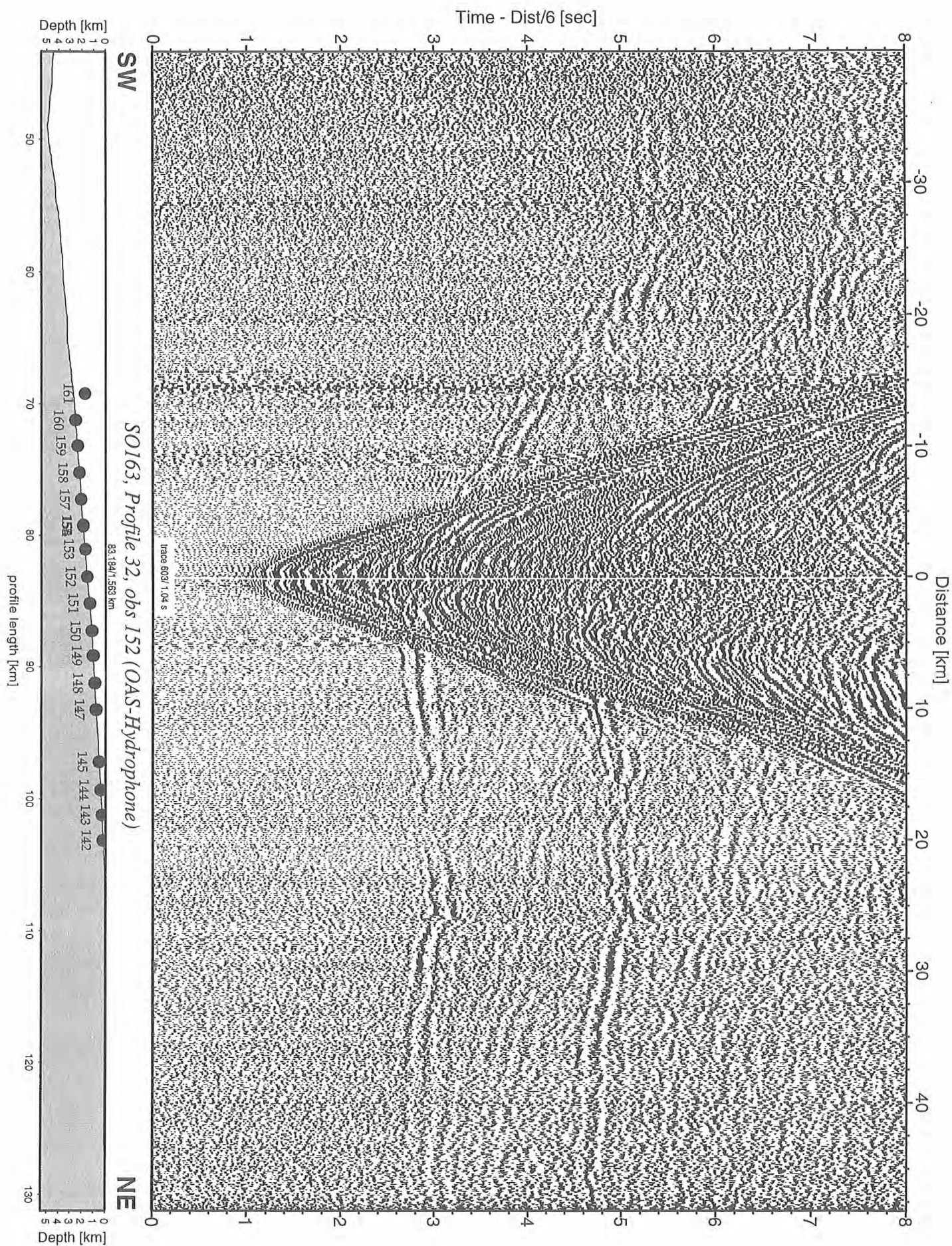


Figure 6.6.2.28 Record section from obs 152 (OAS-Hydrophone), Profile 32.

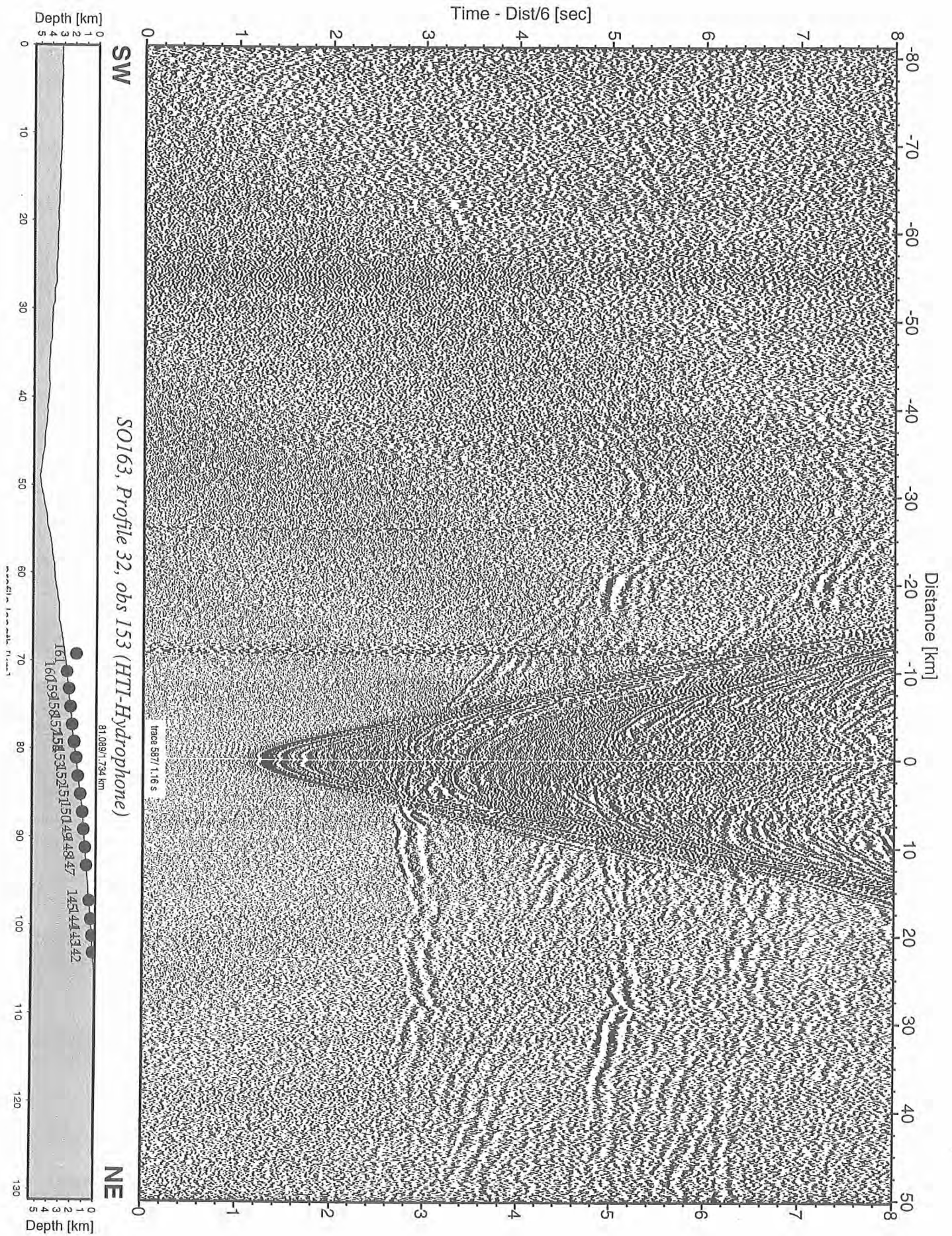


Figure 6.6.2.29 Record section from obs 153 (HTI-Hydrophone), Profile 32.

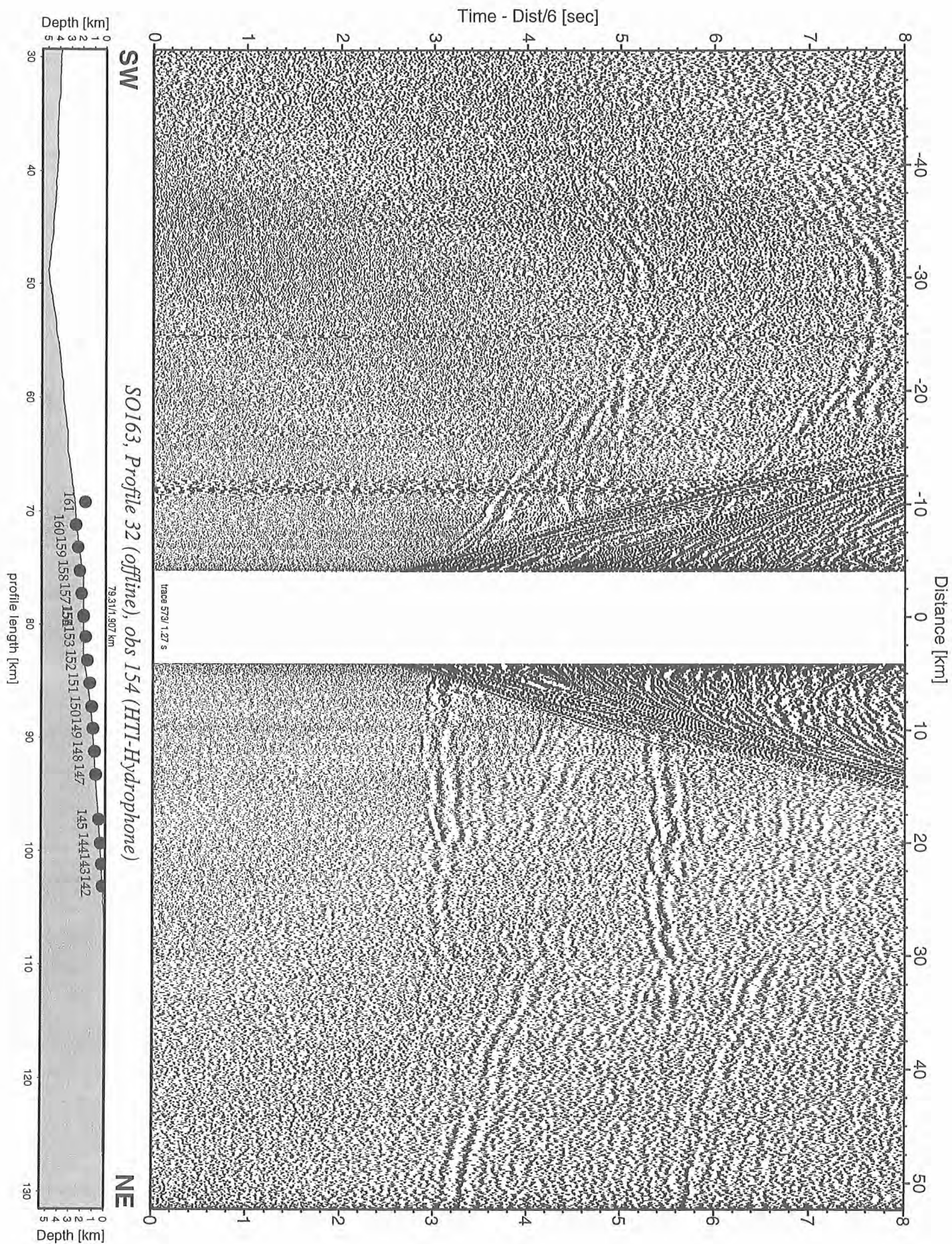


Figure 6.6.2.30 Record section from obs 154 (HTI-Hydrophone), Profile 32.

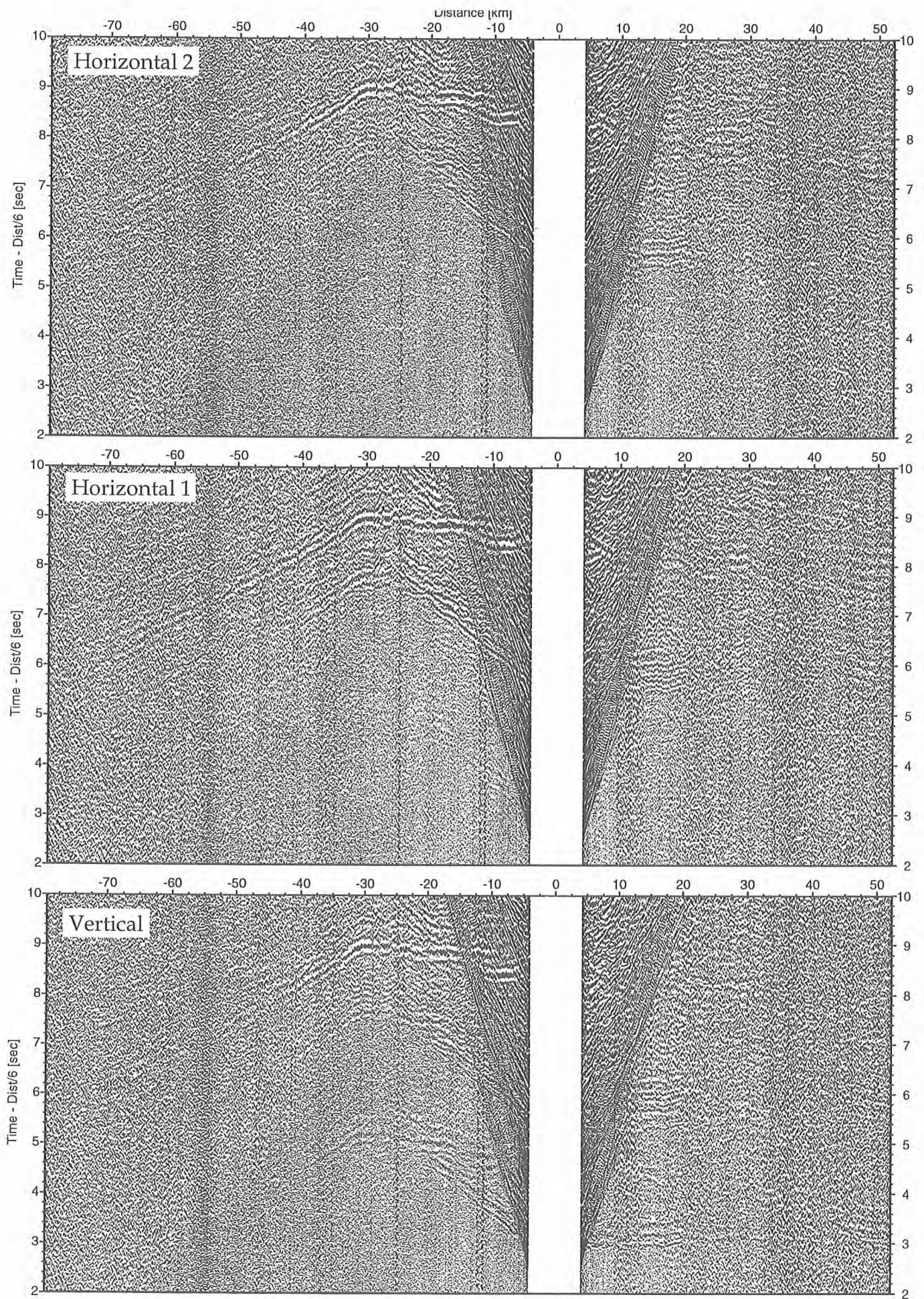


Figure 6.6.2.31 Record sections from obs 154 (Owen-4.5Hz), SO163, Profile 32 (offline).

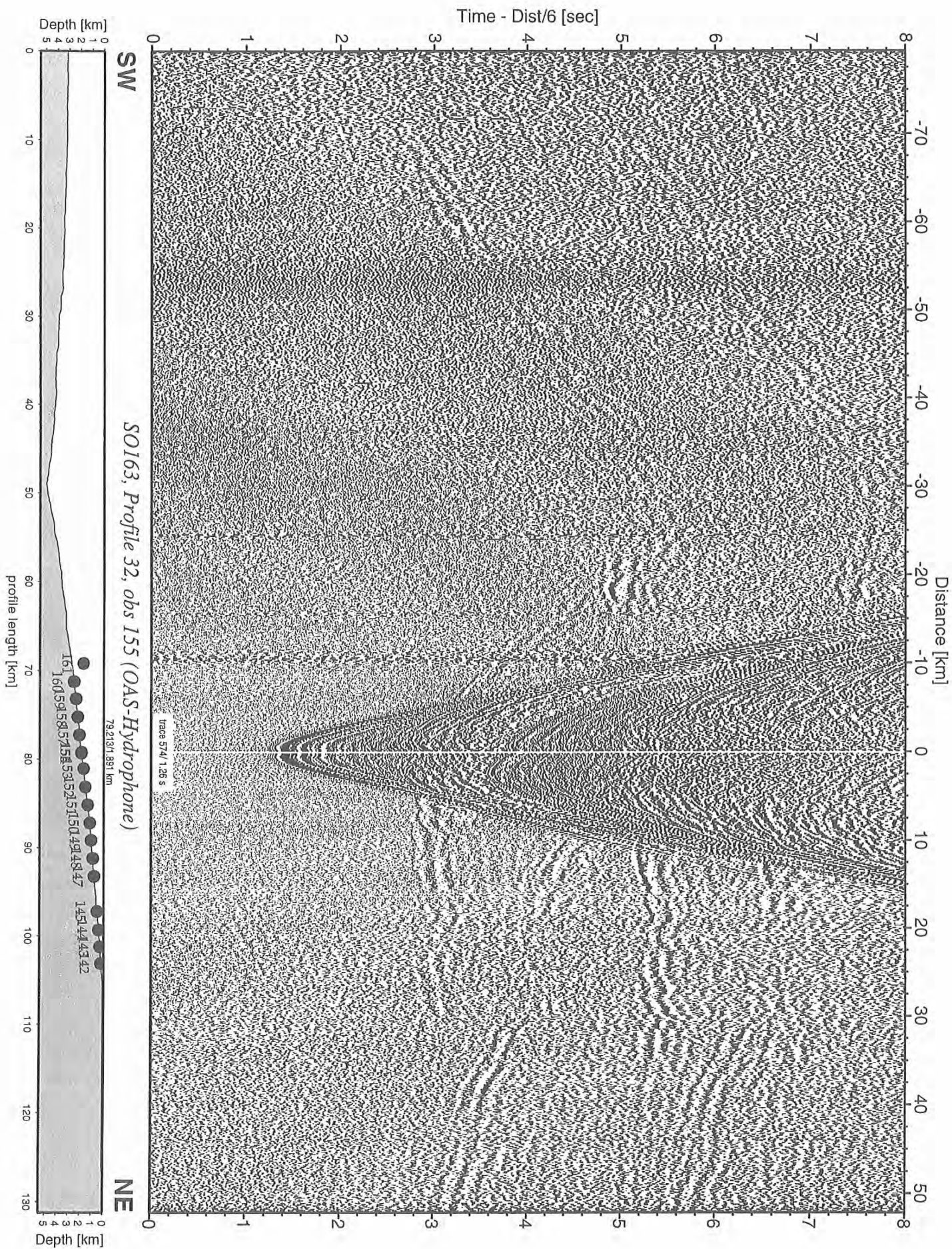


Figure 6.6.2.32 Record section from obs 155 (OAS-Hydrophone), Profile 32.

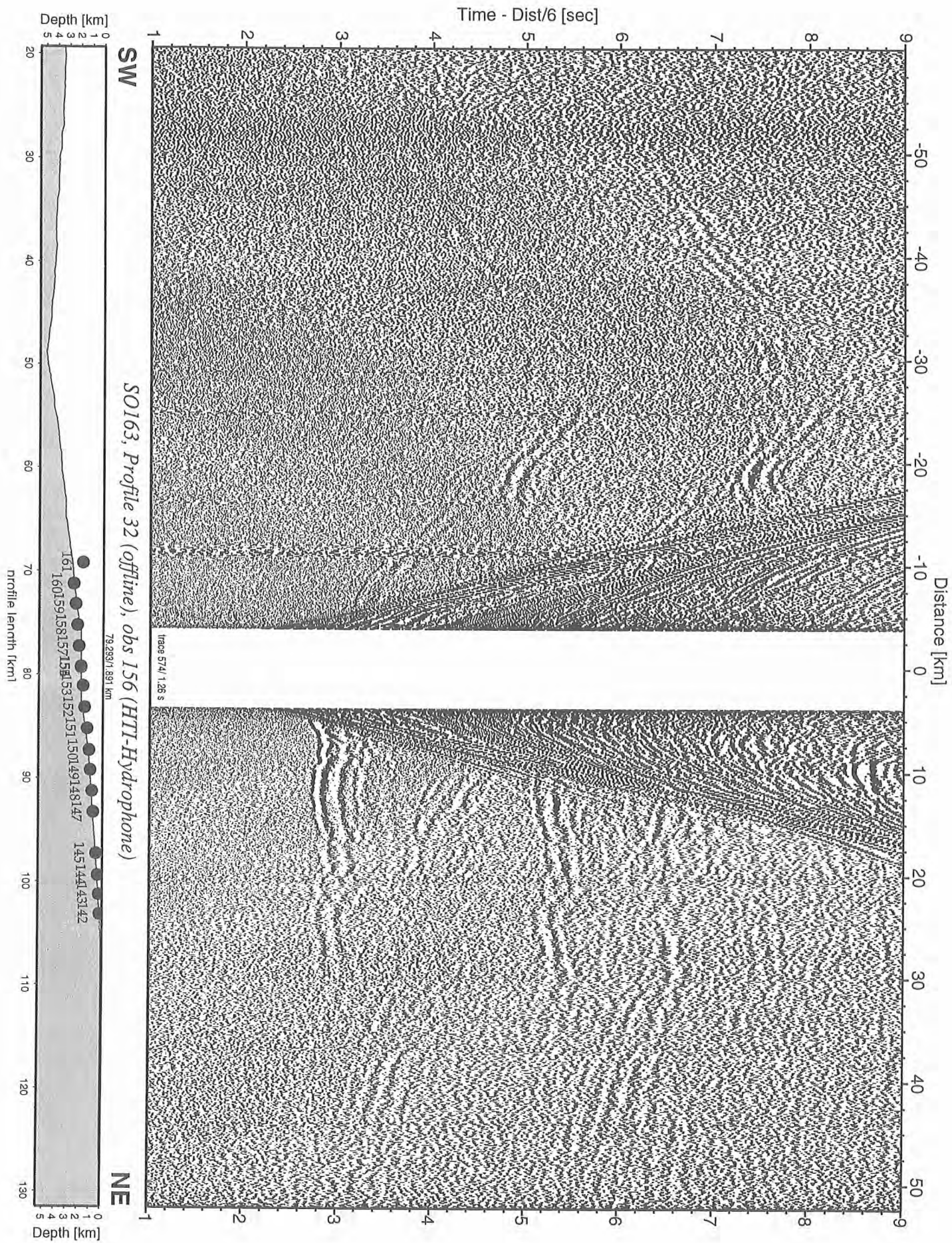


Figure 6.6.2.33 Record section from obs 156 (HTI-Hydrophone), Profile 32.

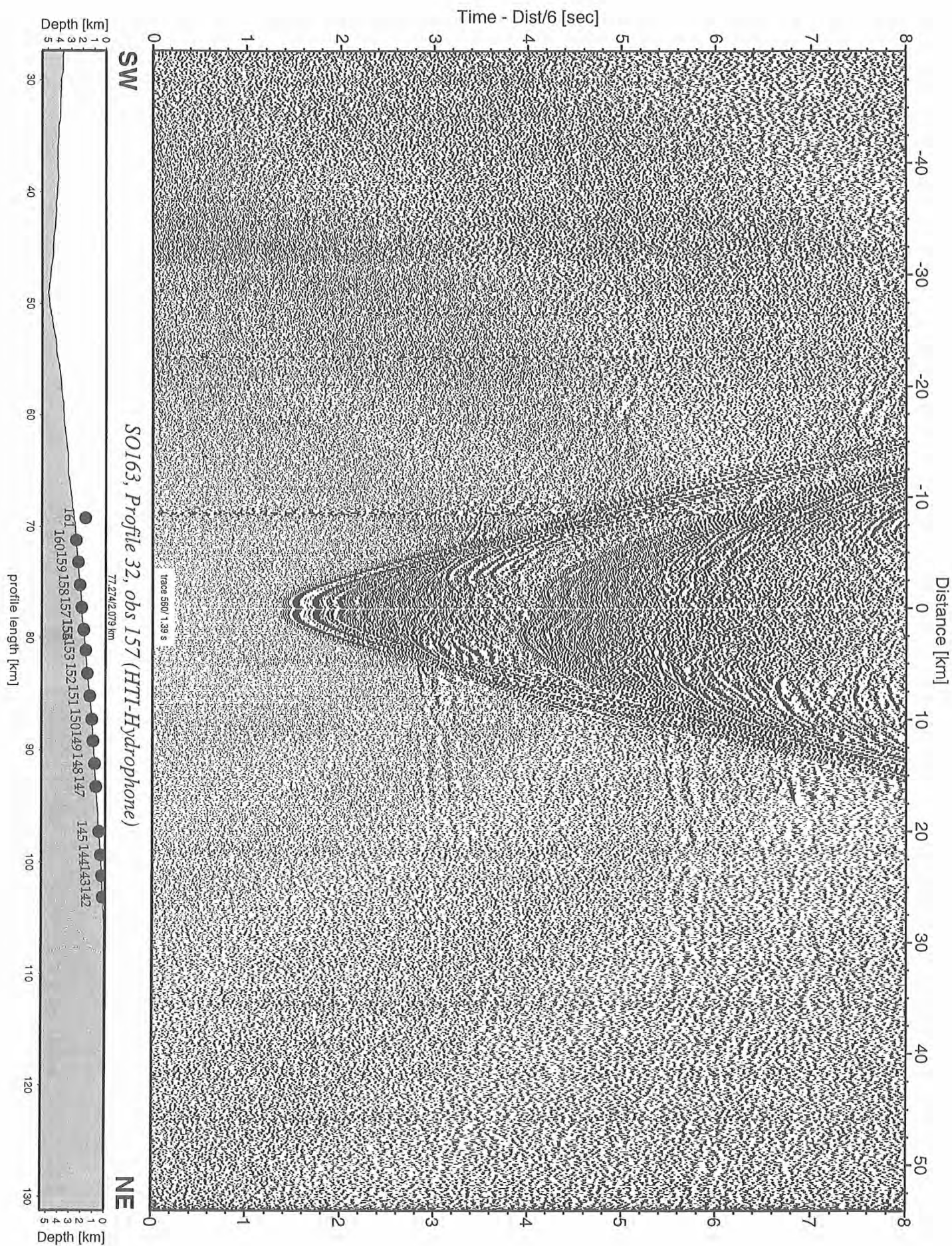


Figure 6.6.2.34 Record section from obs 157 (HTI-Hydrophone), Profile 32.

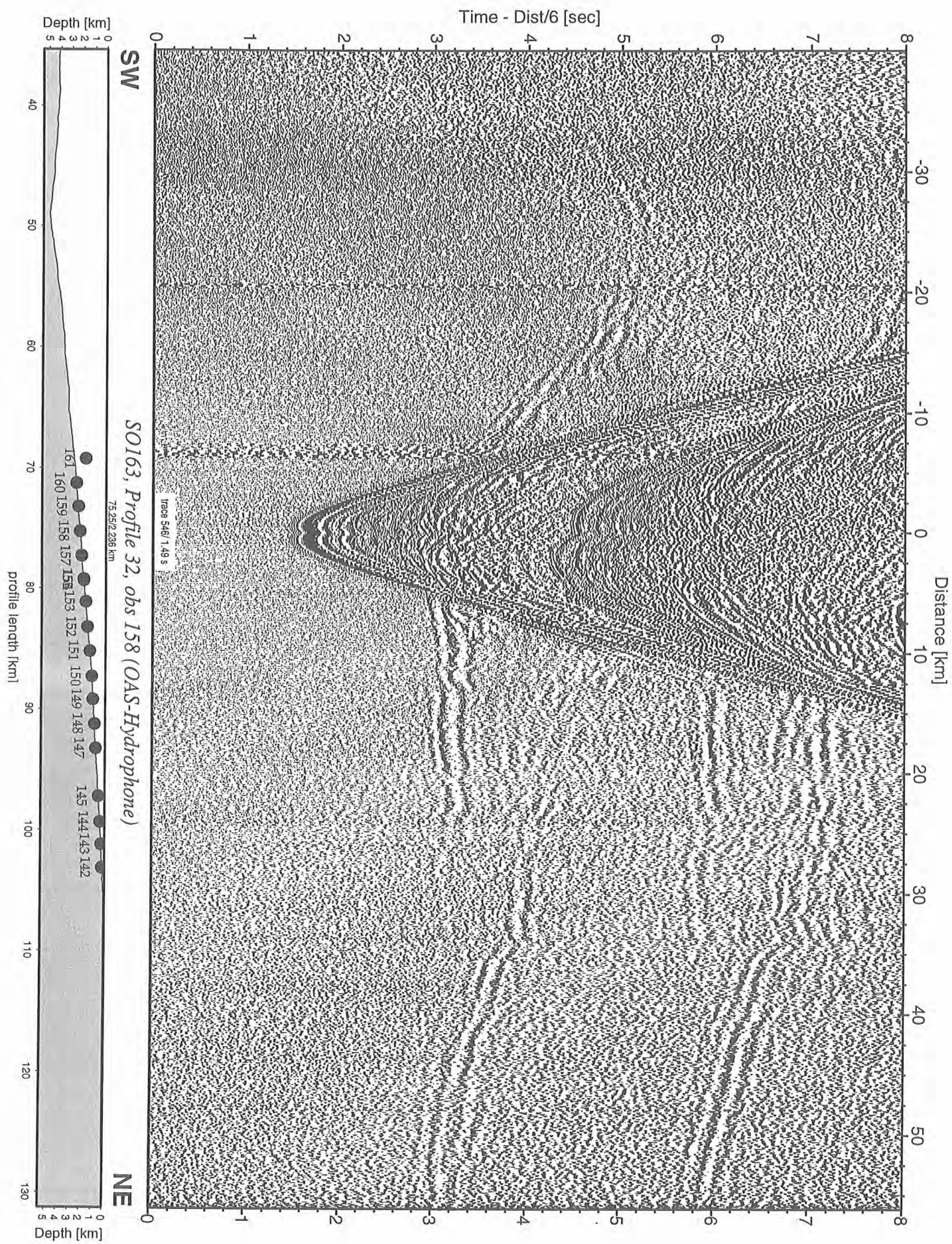


Figure 6.6.2.35 Record section from obs 158 (OAS-Hydrophone), Profile 32.

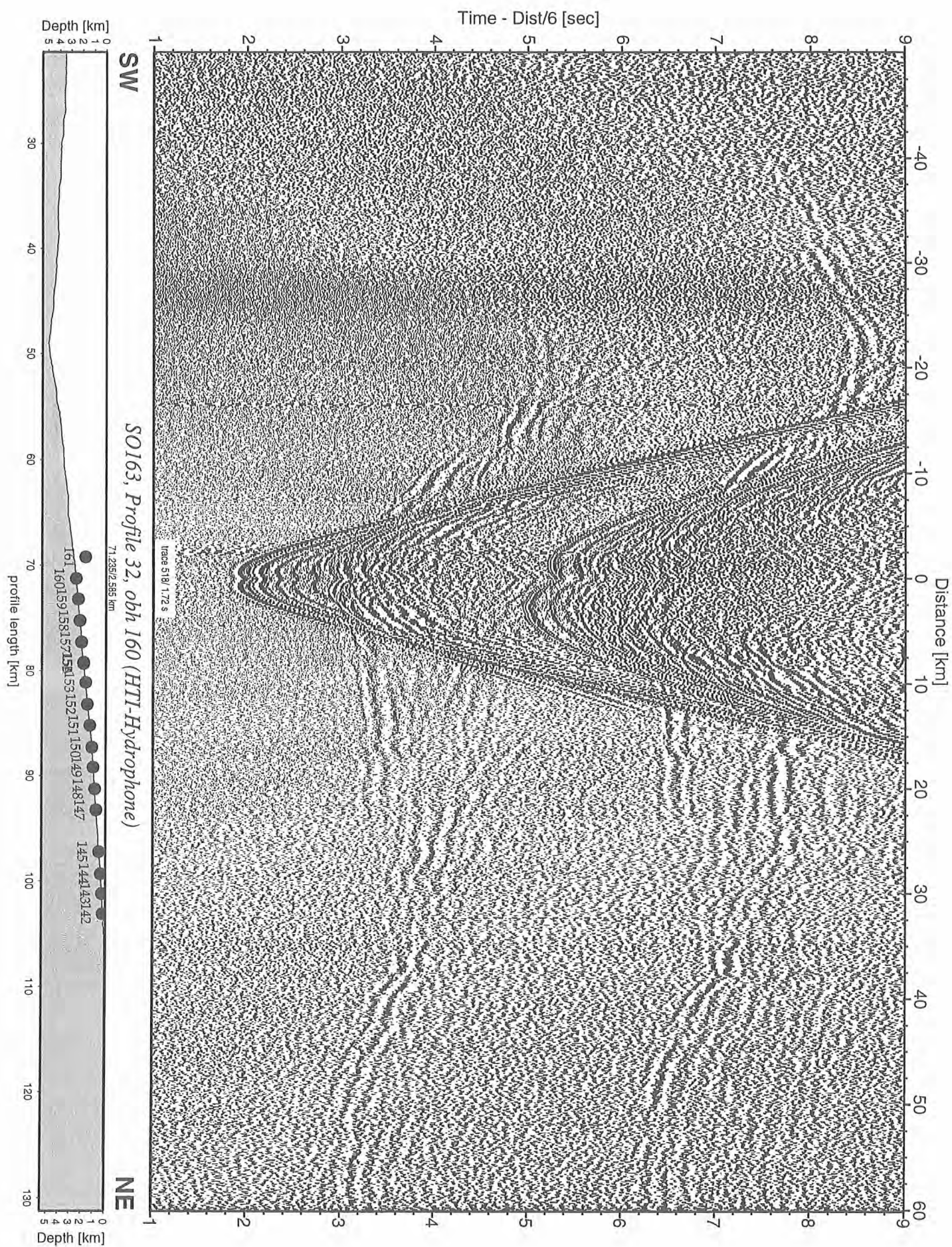


Figure 6.6.2.36 Record section from obh 160 (HTI-Hydrophone), Profile 32.

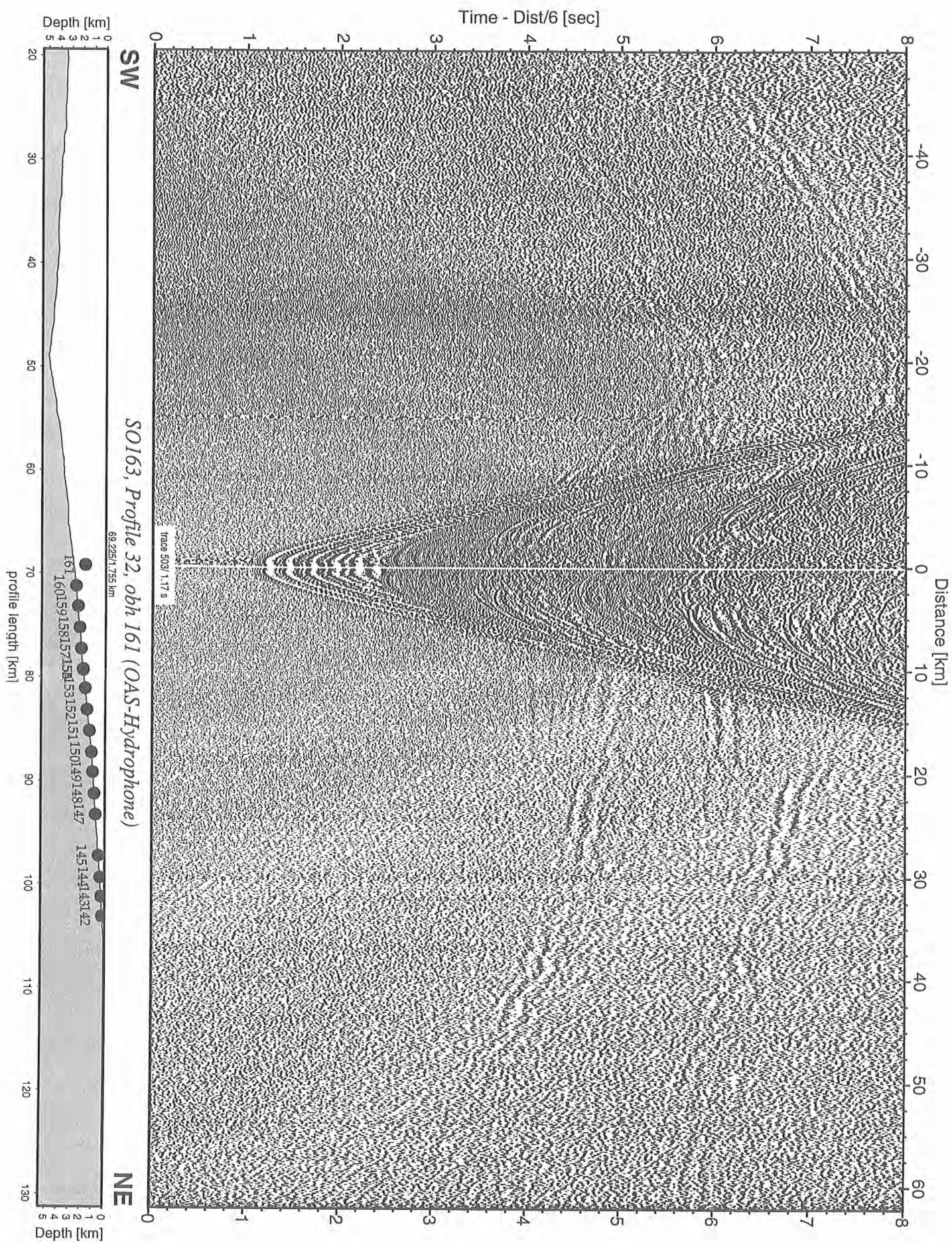


Figure 6.6.2.37 Record section from obh 161 (OAS-Hydrophone), Profile 32.

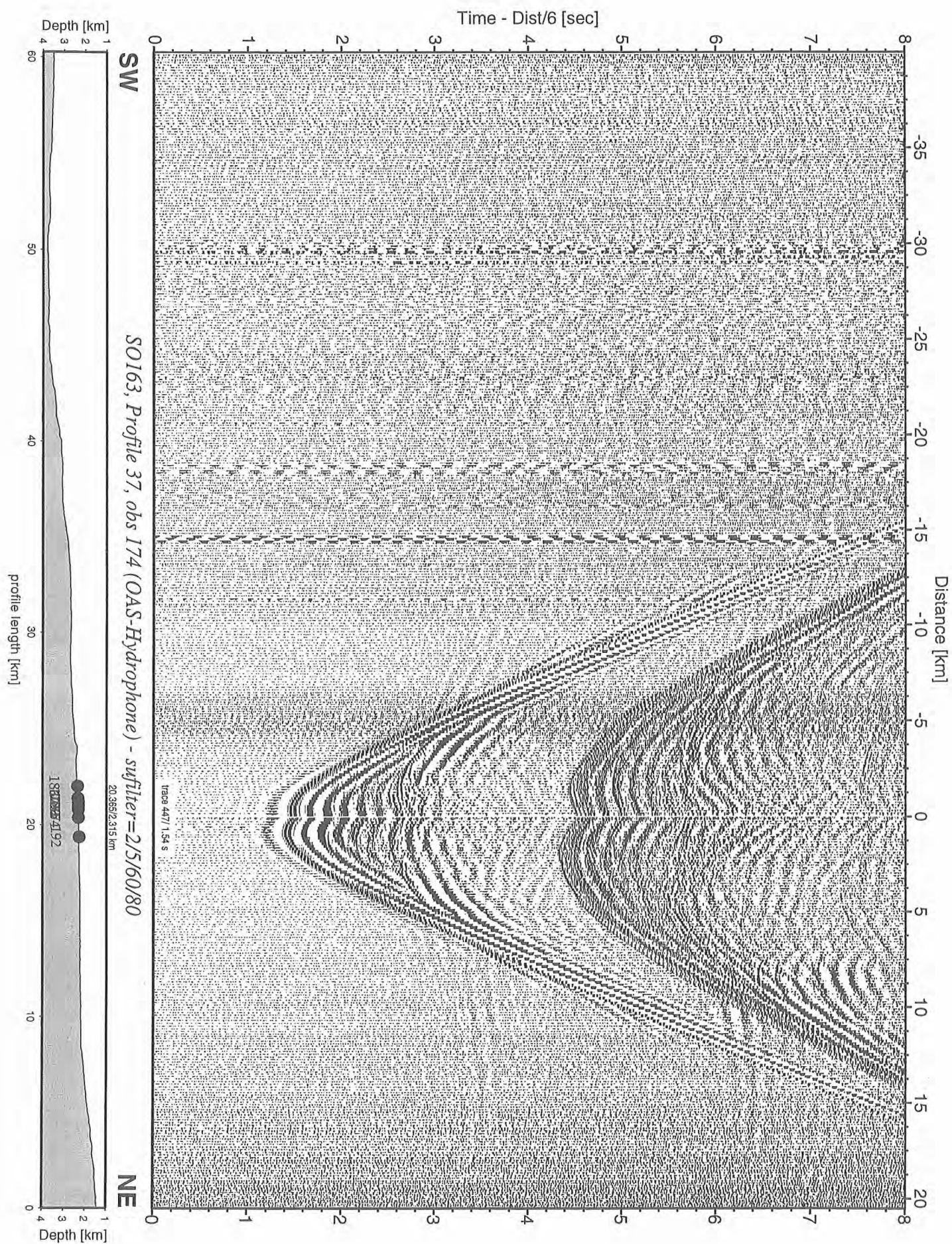


Figure 6.6.2.38: Record section from obs 174 (OAS-Hydrophone), Profile 37.

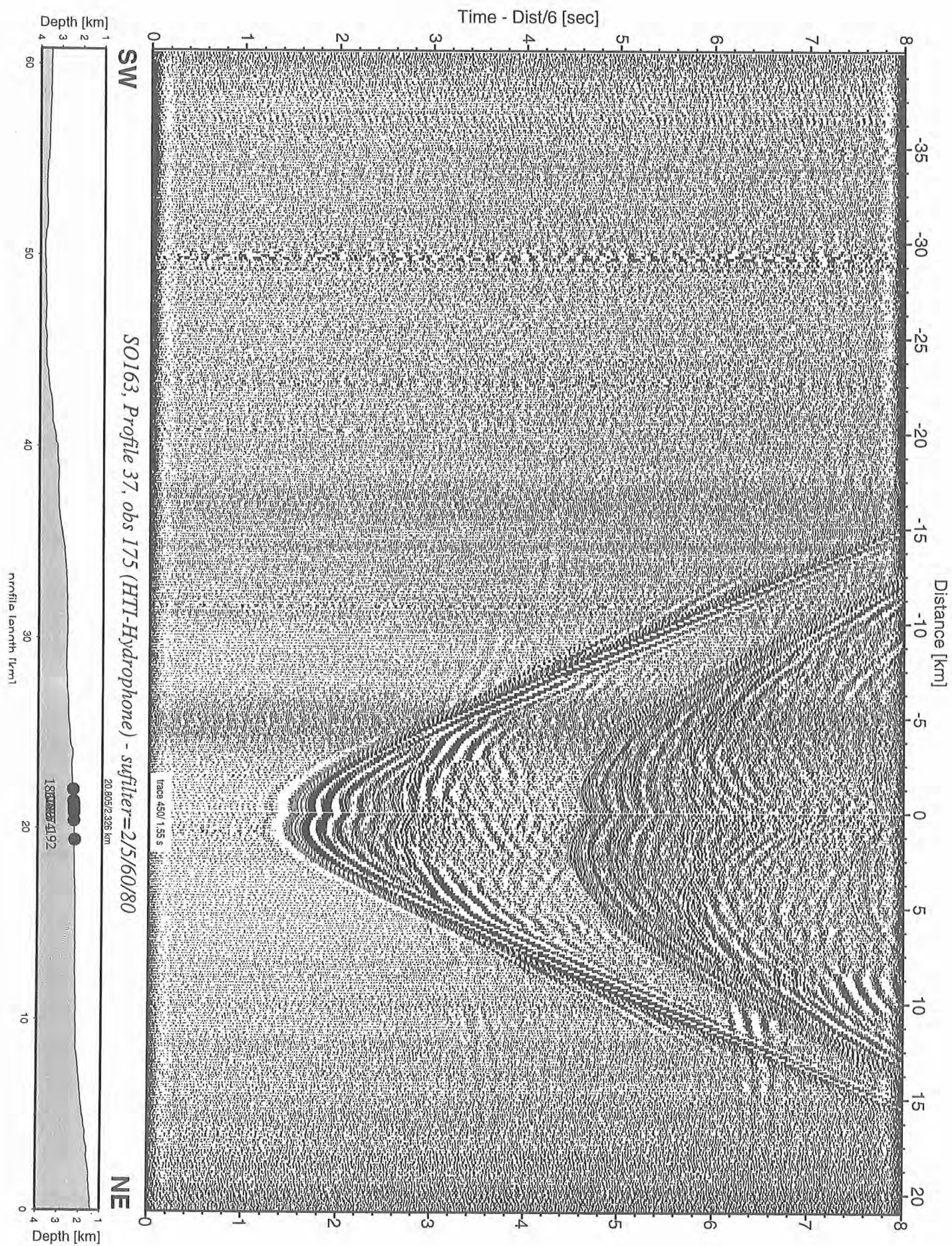


Figure 6.6.2.39: Record section from obs 175 (HTI-Hydrophone), Profile 37.

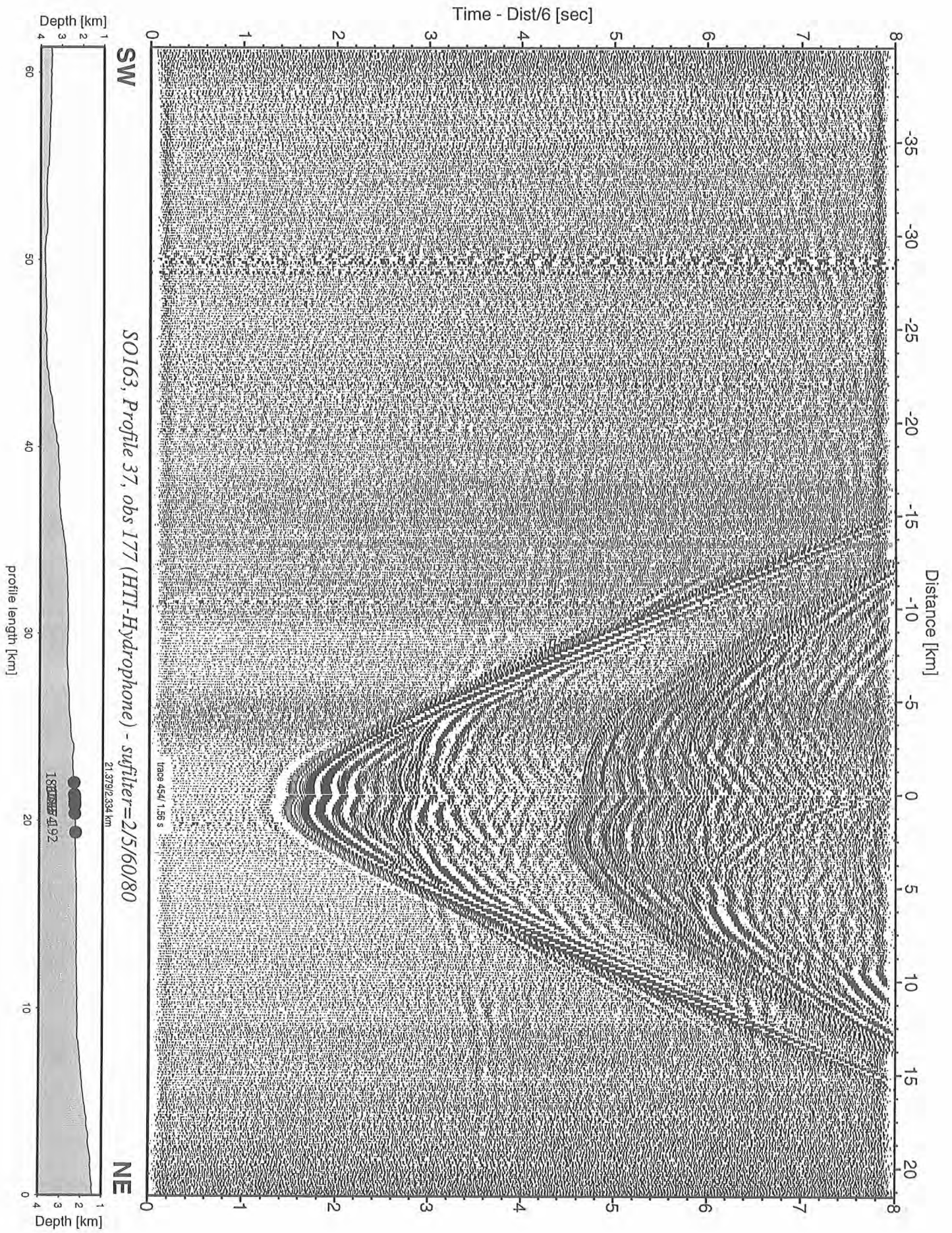


Figure 6.6.2.40: Record section from obs 177 (HTI-Hydrophone), Profile 37.

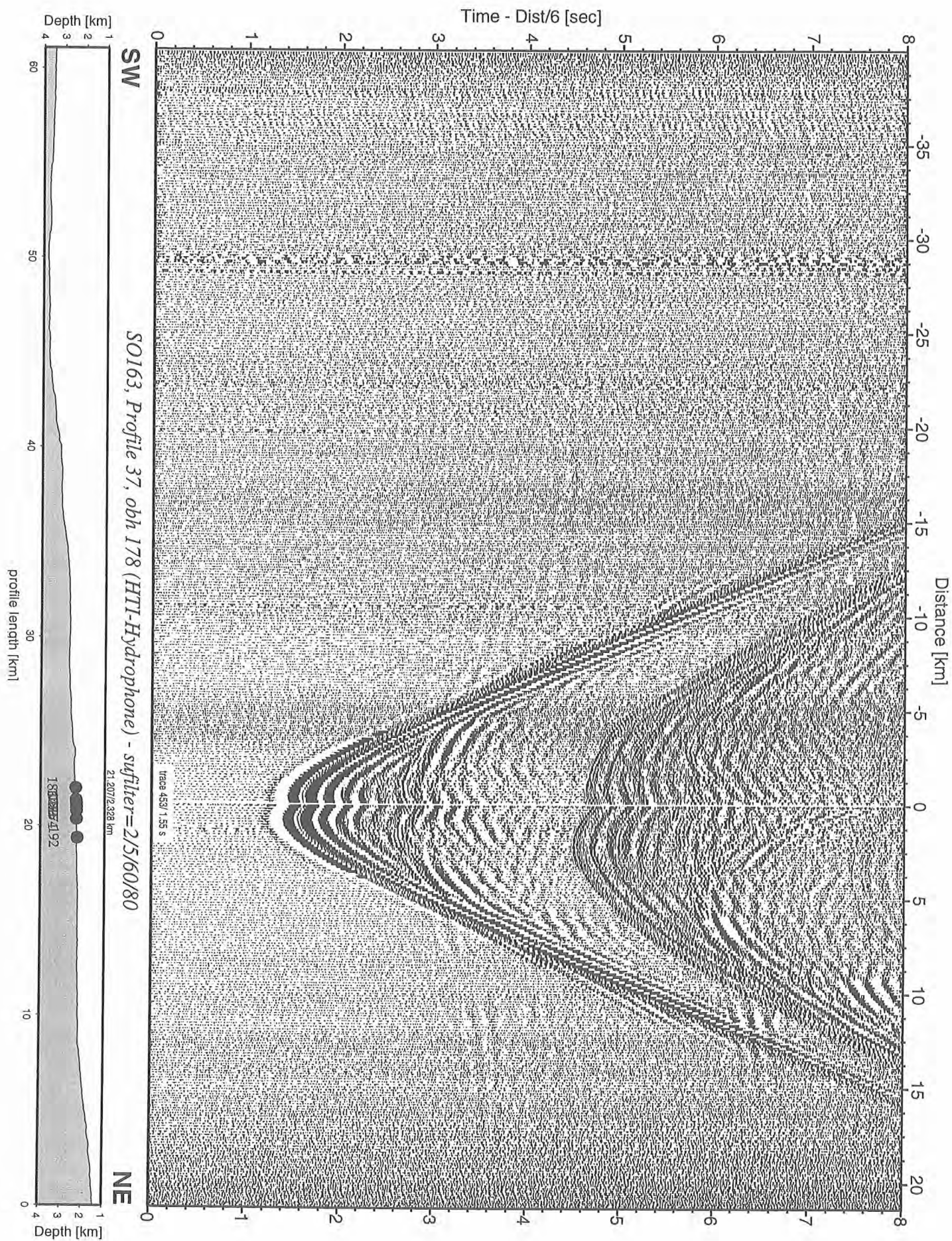


Figure 6.6.2.41: Record section from obh 178 (HTI-Hydrophone), Profile 37.

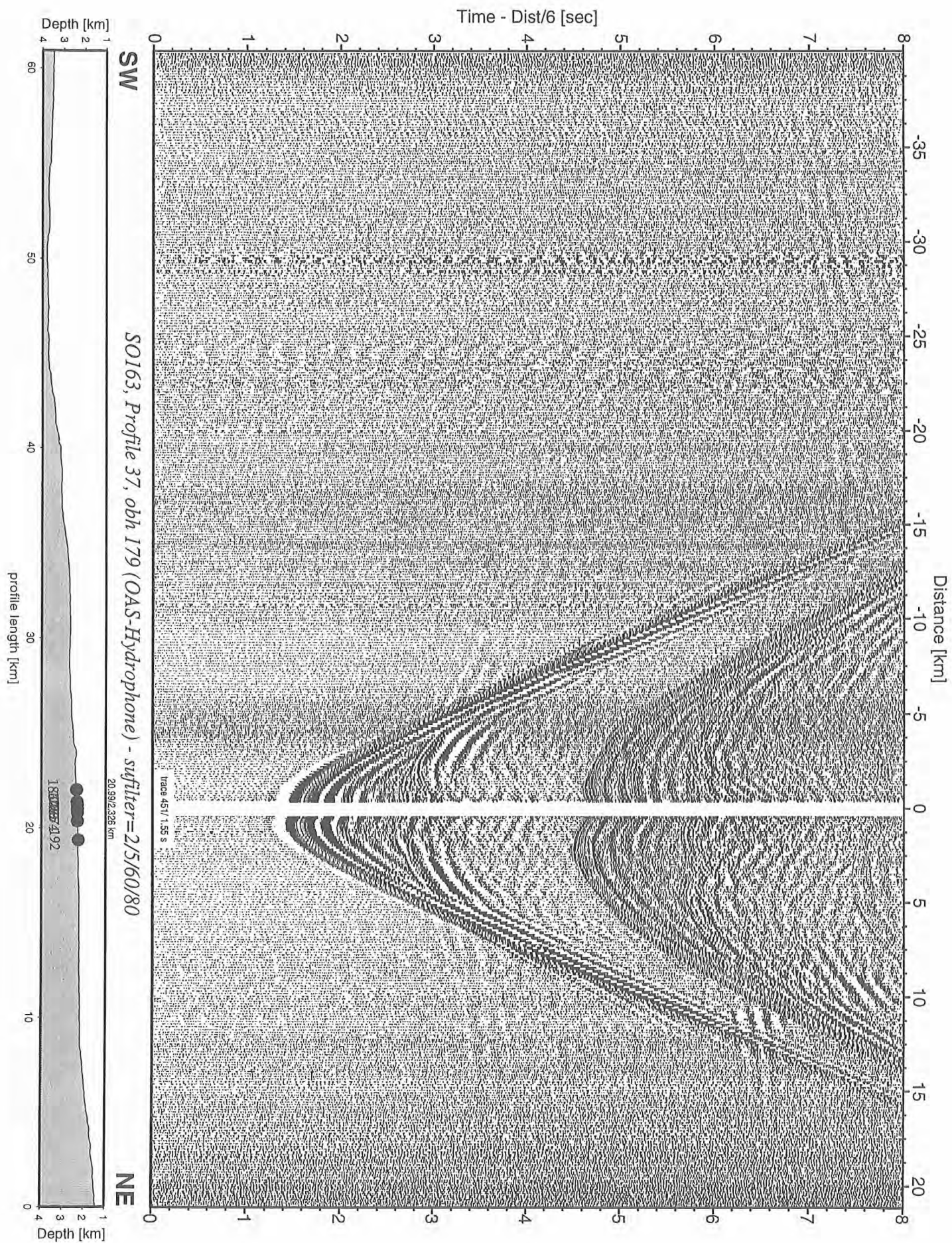


Figure 6.6.2.42: Record section from obh 179 (OAS-Hydrophone), Profile 37.

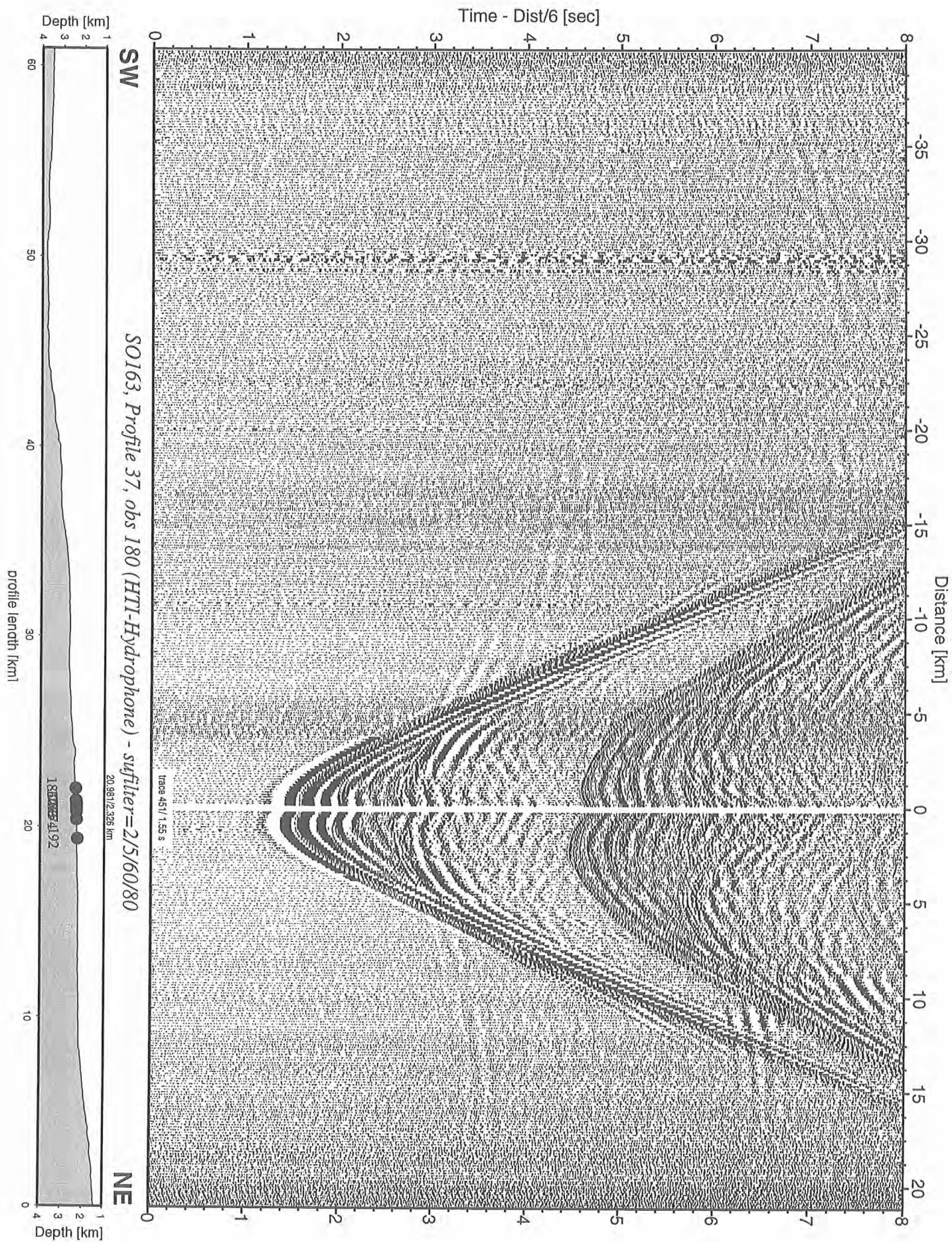


Figure 6.6.2.43: Record section from obs 180 (HTI-Hydrophone), Profile 37.

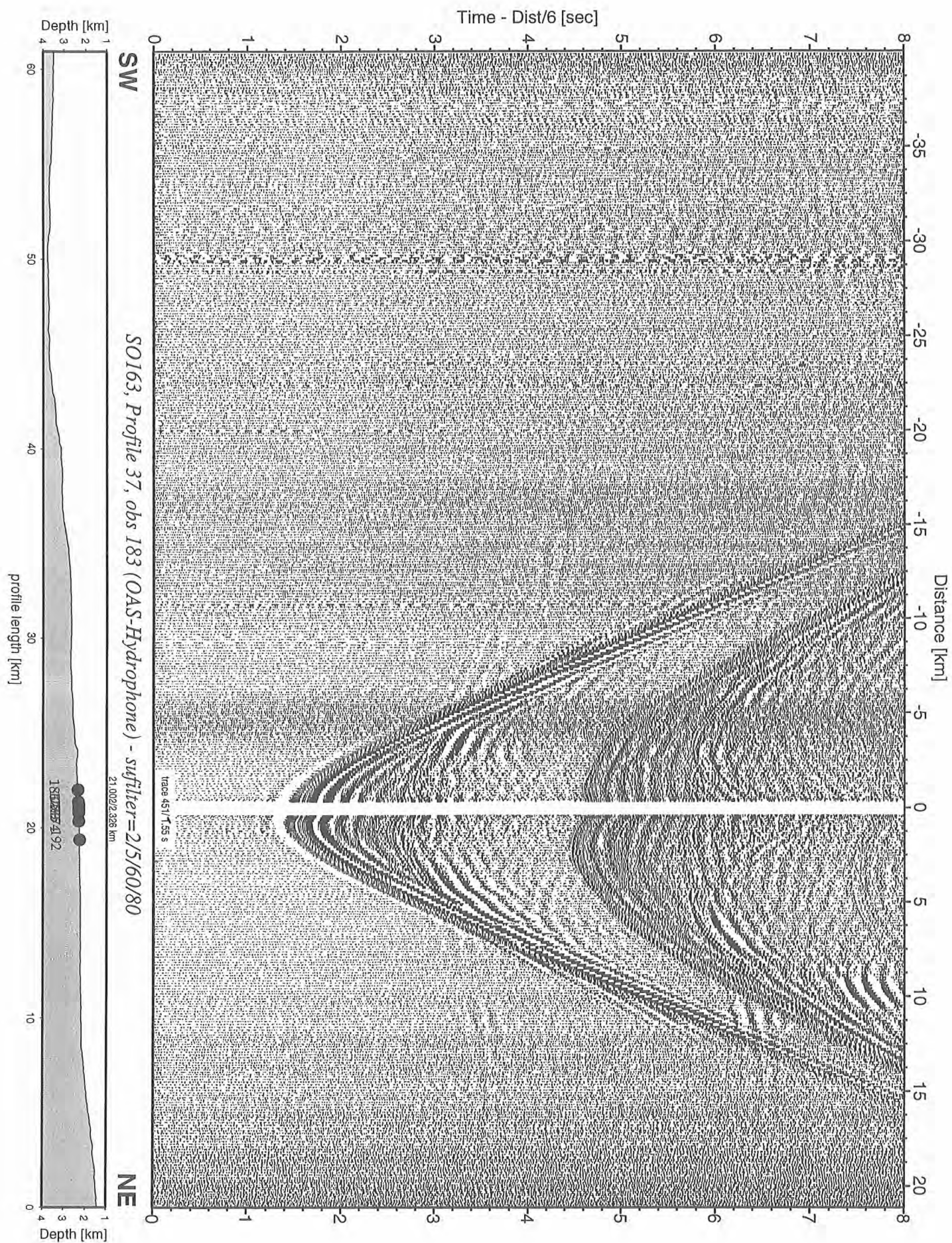


Figure 6.6.2.44: Record section from obs 183 (OAS-Hydrophone), Profile 37.

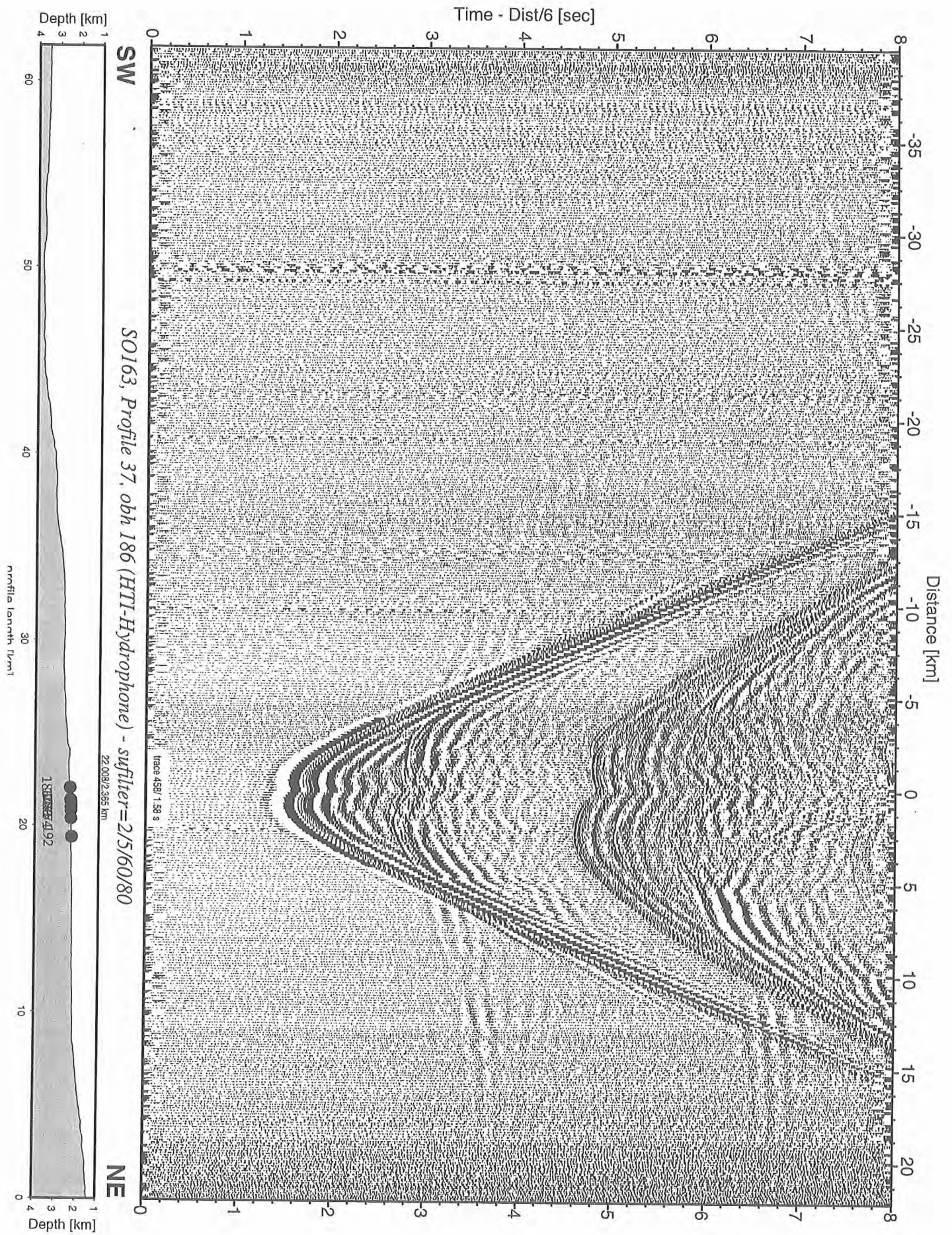


Figure 6.6.2.45: Record section from obh 186 (HTI-Hydrophone), Profile 37.

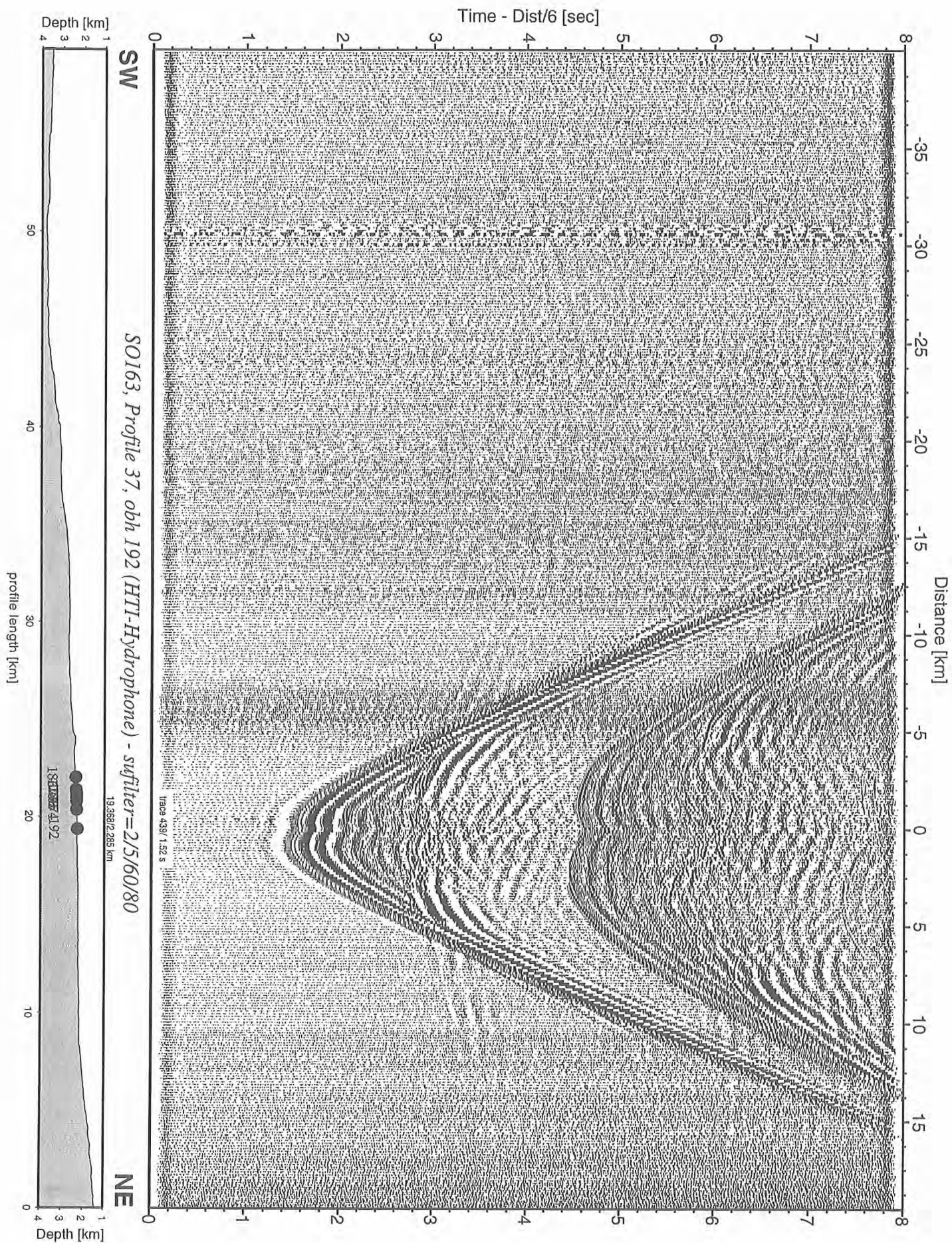


Figure 6.6.2.46: Record section from obh 192 (HTI-Hydrophone), Profile 37.

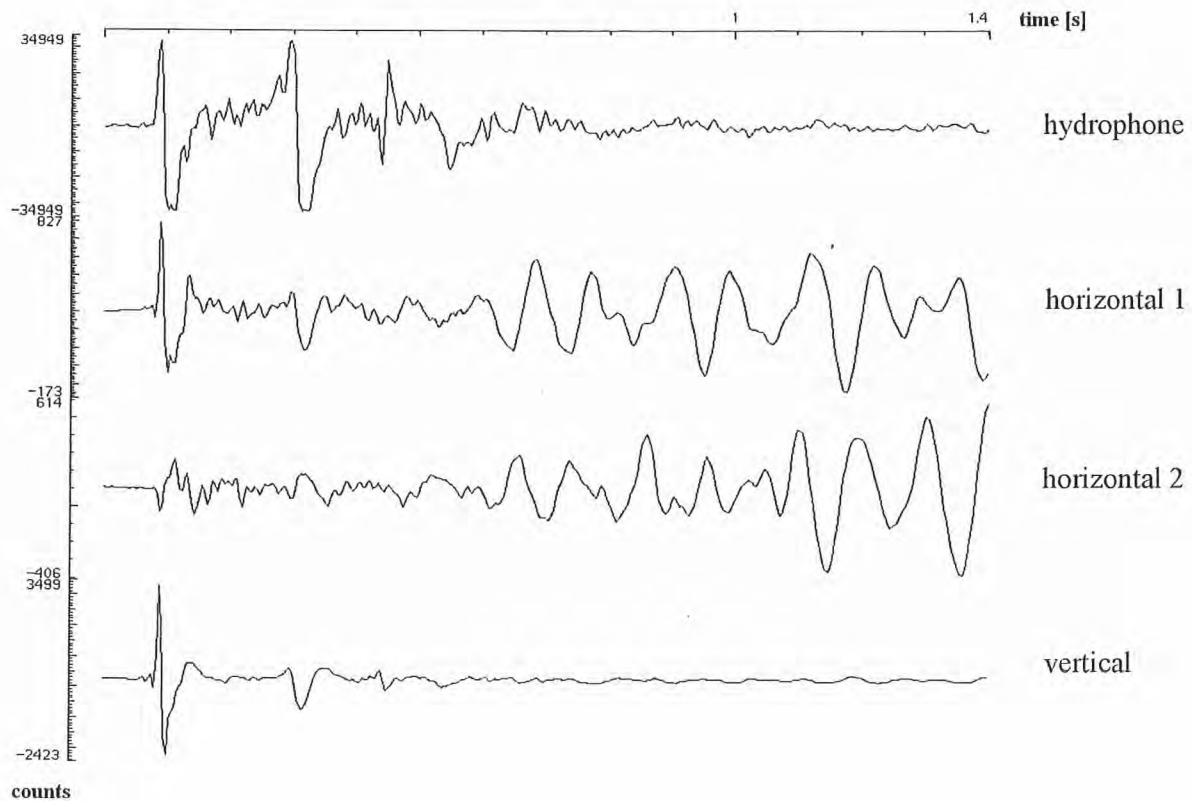


Figure 6.6.2.47: Example of unprocessed, trace-normalized components of a short offset recording of OBS31 on line P09. Hydrophone and approximately vertical trace are dominated by the direct wave consisting of three discrete oscillations and the reflected dilatational wave from the BSR with clear negative polarity on the hydrophone trace between second and third oscillation of the source signal. On the hydrophone trace peaks of the direct wave are clipped. On the approximately horizontal components, after the less powerful signature of the source signal, a phase with approximately horizontal polarization (probably a P to S converted phase from the BSR) follows. One might even identify a second shear phase in the coda (which might be a shear wave reflection from the BSR which has been converted into a shear wave already at the seafloor).

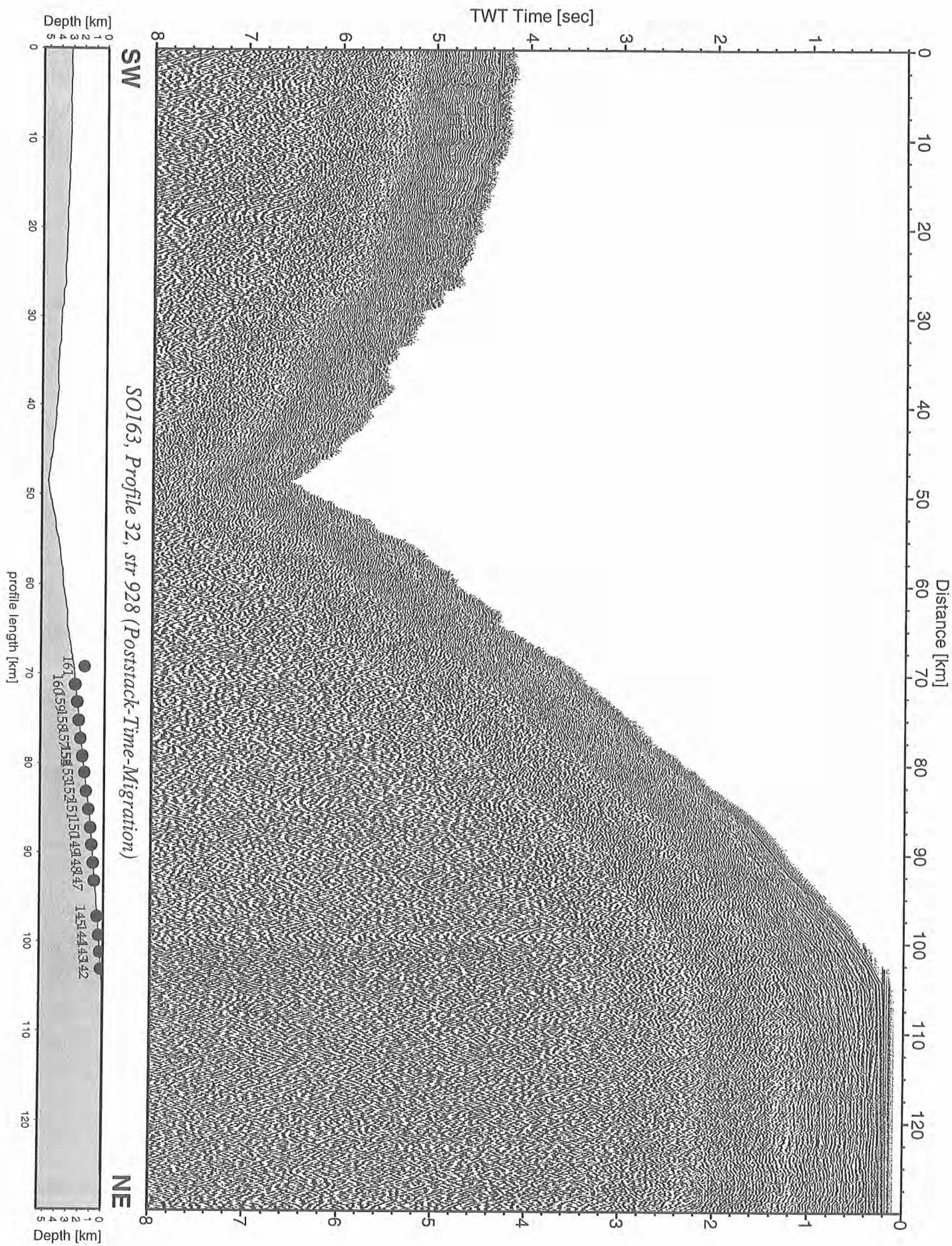


Figure 6.6.2.48: Record section from str 928 (Poststack-Time-Migration), Profile 32.

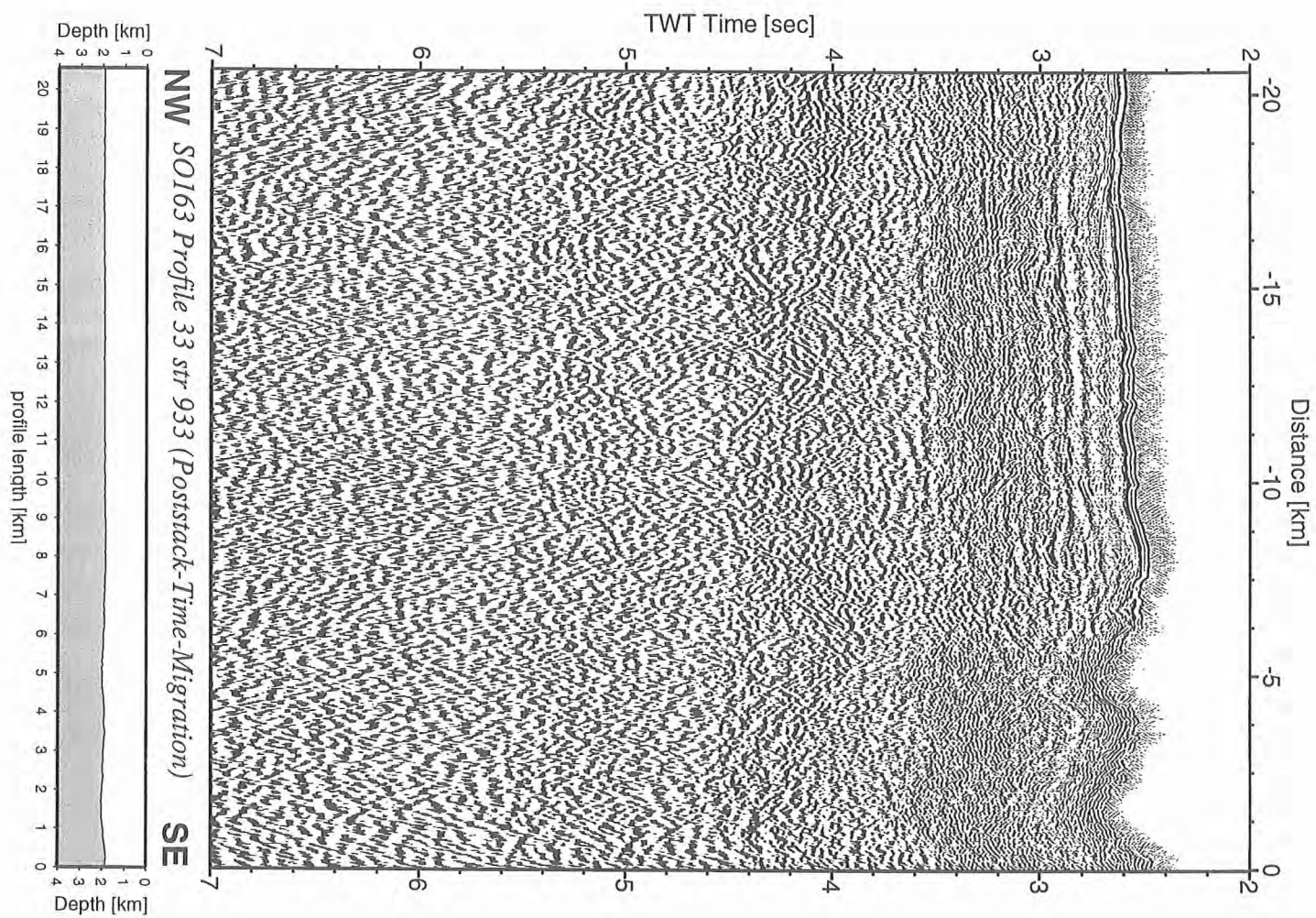


Figure 6.6.2.49: Record section from str 933 (Poststack-Time-Migration), Profile 33.

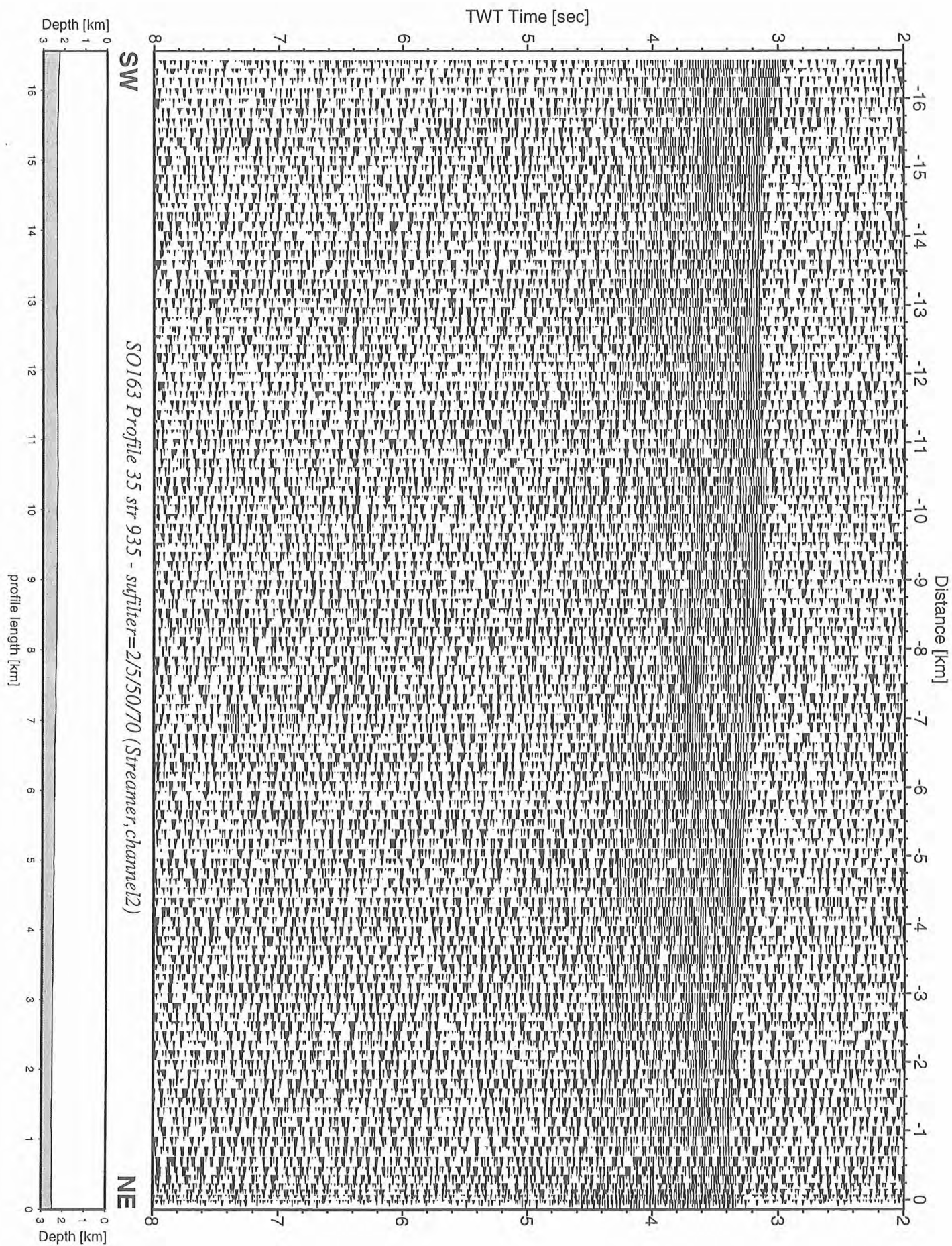


Figure 6.6.2.50: Record section from str 935 (Streamer,channel2), Profile 35.

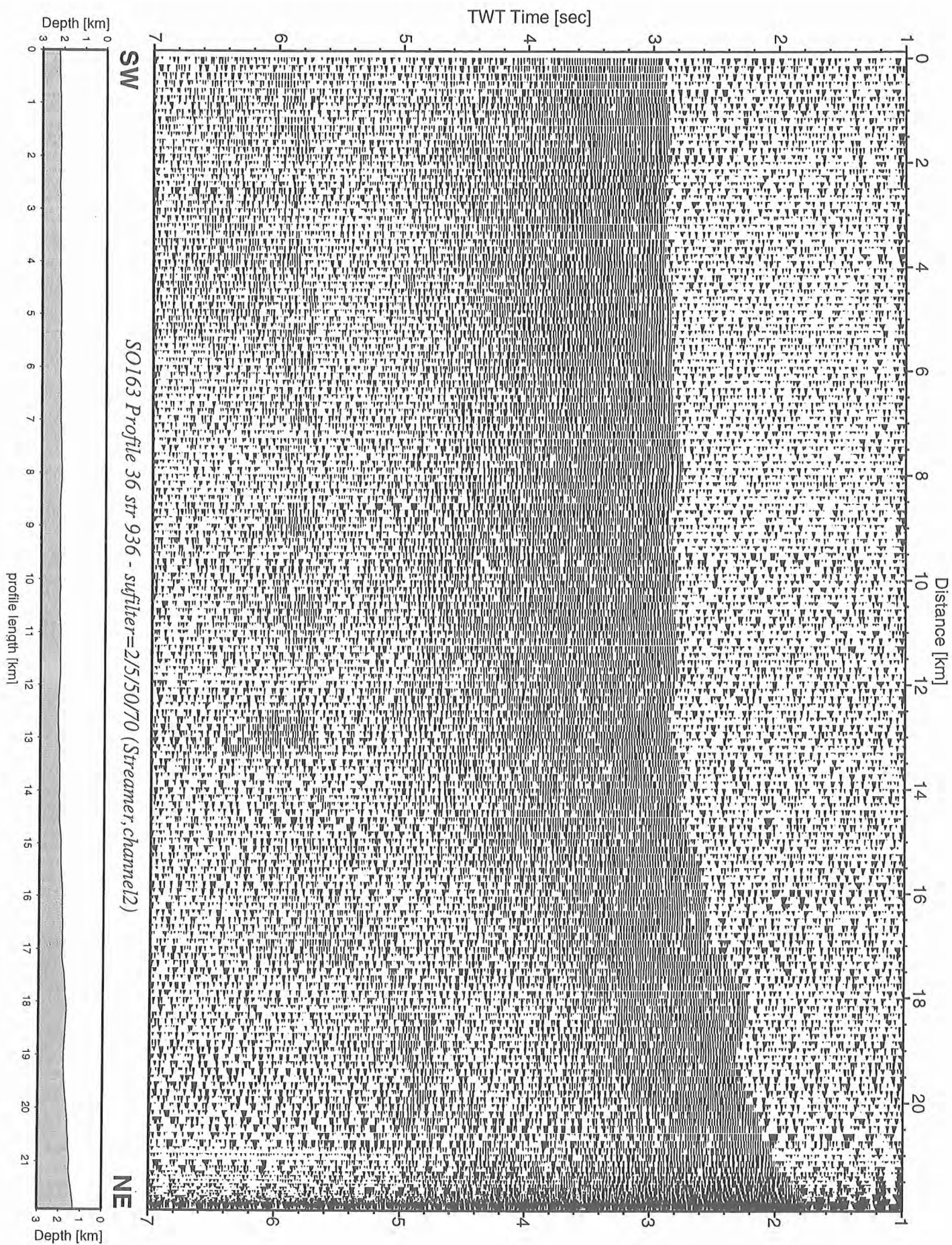


Figure 6.6.2.51: Record section from str 936 (Streamer, channel2), Profile 36.

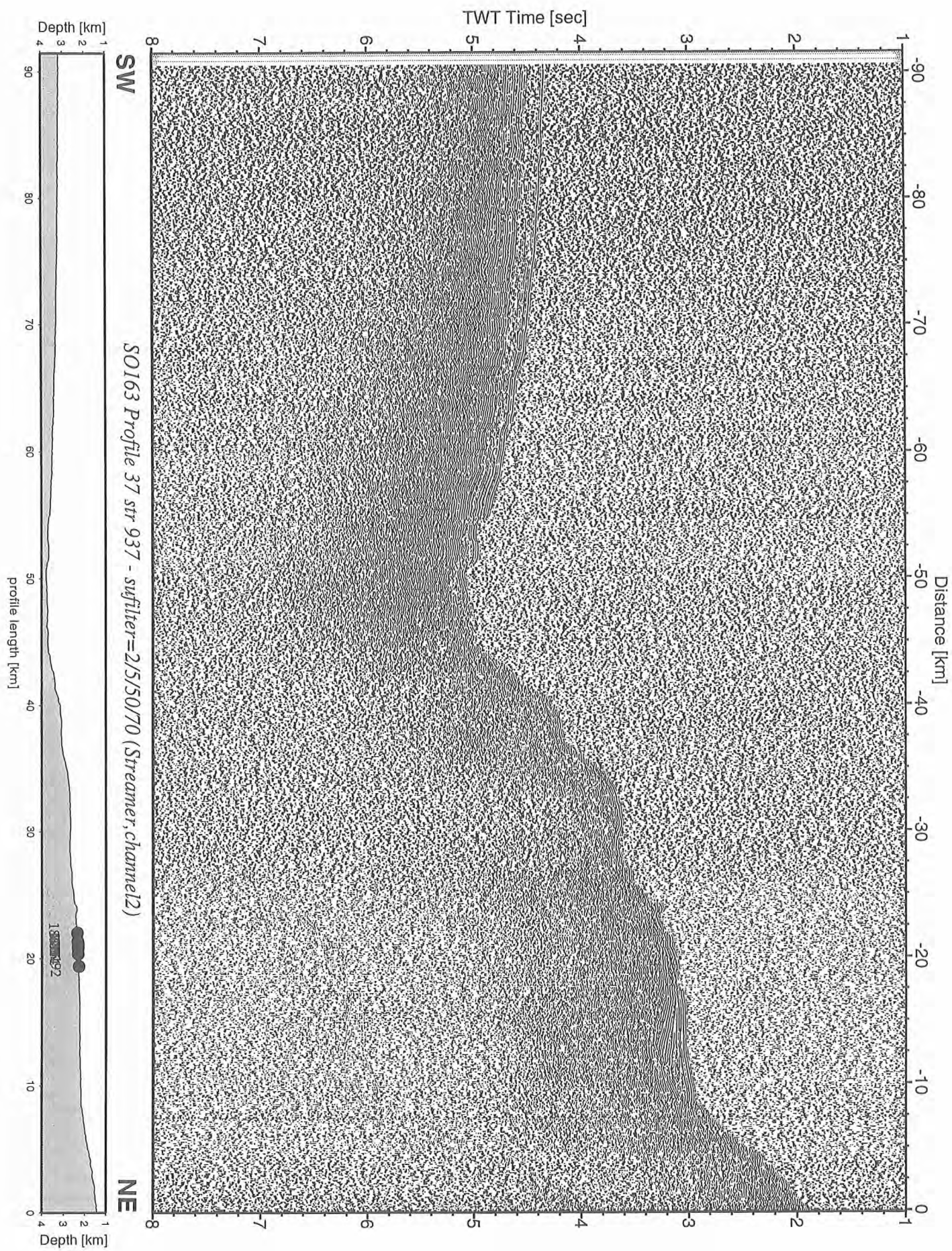


Figure 6.6.2.52: Record section from str 937 (Streamer,channel2), Profile 37.

6.6.3 The megalens

(M. Schnabel, P.-O. Thierier)

For the first period of the SFB 574, the largely erosive subduction zone off Central America was chosen as the target area for investigation. Concerning the erosion, Ranero & v. Huene (2000) identified two mechanisms: erosion by tunneling and basal removal of rock lenses from the overriding plate. Those megalenses are 10-15 km long and 1.5-2 km high, and they are surrounded by plate-boundary reflections. Seismics and bathymetry show a relative stable slope above these lenses, which is not consistent with a subducting seamount. The megalenses occur at depths at which temperatures indicate that the smectite-illite transition has occurred, and the plate interface is entering the zone of stick-slip behaviour. Thus, the lenses could be associated with active erosion in an environment of increasing friction. One such seismic reflection image is shown in Figure 6.6.3.1.

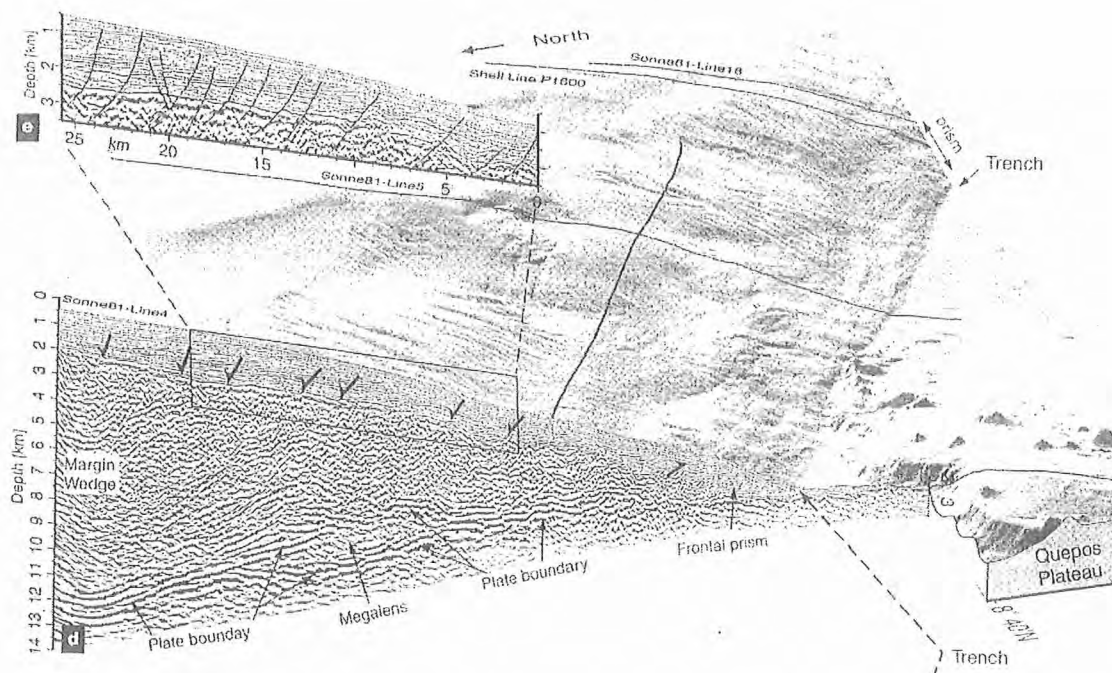


Figure 6.6.3.1: Prestack depth migration of SO81 line 4 projected on bathymetry perspectives. Plate-boundary reflections bifurcate at ~6km depth and surround a rock megalens (Ranero & v. Huene, 2000).

One of the aims of the subproject A2 within SFB 574 is to map the lateral extend of this megalense and its continuity. The lithological properties causing the surrounding reflections should be investigated. With the new data collected on SO163-2, we will be able to perform AVO-analysis to describe these reflections in greater detail.

Based on the results of Ranero & v. Huene (2000), we decided to collect seismic wide-angle data in the surrounding of profile 5 of Sonne-Cruise SO81. Four profiles were layed out, with a minimum of at least 9 OBH and 5 OBS. Along each line, several airgun configurations were used (see App. 9.5.4 to 9.5.6 and 9.6 for details on instrumentation and shots). In Figure 6.6.3.2 the location of the profiles and instruments are given. Selected record sections are shown in Figures 6.6.3.3 to 6.6.3.82.

The data quality along all profiles was quiet good. Only a few stations situated on the uppermost slope near the shore showed only short offsets. All other stations recorded shots over large distances (up to 50 km). Onboard forward modelling with the ray tracing application „MacRay“ (Luetgert 1992) resulted in the velocity-depth models shown in Figure 6.6.3.84 and Figure 6.6.3.87.

The principal geometry as deduced from the ray tracing show good agreement. The topmost layers are interpreted as sediments. The upper one shows P-wave velocities ranging from 1.6 to 2.0 km/s and has a thickness up to about 1000m. The velocity in the second layer is significant higher (2.5 to 4.0 km/s), and could therefore be interpreted as compacted and cemented sediments.

Refractions from the margin wedge are seen in all of the records. The velocity in the outer wedge is about 3.5km/s, increasing up to 4.2 near the coast. The velocity at the lower end of the wedge ranges from 4.0 km/s near the trench up to 4.8km/s.

For a few stations, reflections from the Mohorovičić-discontinuity could be identified (Figure 6.6.3.83a). The velocity of the oceanic layer 3 ranges from 6.0 to 6.9 km/s and the Moho is situated at a depth of 12.5 km at 30 km landward behind the trench.

Profile P22, SO 163-2

(I. Arroyo, H. Kopp, E. Flueh)

Profiles 19 and 21-23 were carried out as part of the detailed research focusing around the megalense and the Quepos mound. The location map with bathymetric lines every 100 m is shown in Figure 6.6.3.2. These profiles consisted of 15 OBHs (from 76 to 90), with an average offset between stations of 1,5 nm (2,7 km) for a total length of almost 40 km with an orientation NW-SE, parallel to the trench axis. Details and coordinates of the stations could be found in Appendix 9.5.4.

The shooting of the profiles were performed by using different airguns arrays: Prakla array with a trigger interval of 10 s (profile 19), 32 l airgun (profiles 21 and 23) and G-Gun (profile 22) with trigger interval of 30 s. The records were stored and processed as described in Section 6.6. The clearest records were obtained with profile 22 (Figure 6.6.3.49 to 6.6.3.63), which is the one analyzed on board.

The profiles were located over the upper continental slope, with OBHs at depths ranging from 960 to 1350 m. This sector is part of the broad embayment in the upper and middle slope that opposite the subduction of the Quepos Plateau. Here, the margin slope is characterized also by deep canyons that end at the upper part of the frontal prism of the margin (von Huene et al., 2000).

Profil 22 shows high quality data. Due to the relative similarity of the records and the limited time on board, only the phase arrivals of three stations (OBH 76, 80 and 86) were picked in order to achieve 1-D velocity models. OBH 80 shows one of the most interesting and rich records (see Figure 6.6.3.51). The pickings of the record are shown in Figure 6.6.3.86, together with the predicted arrivals modeled with the program R1D (Luetgert, 1992). A very good correlation between the observed and the predicted arrivals was achieved (Figure 6.6.3.86). From this 1-D modeling, the following layers were detected (Figure 6.6.3.86):

Depth (m)	P velocity (km/s)
0 – 1 350 (water)	1,500
1 350 – 1 664	1,550
1664 – 2 401	1,633 – 1,867
2 401 – 3 156	3,467 – 4,067
3 156 – 6415	4,300 – 5,400
6 415 – 6 621	4,833
6 621 – 10 669	5,467 – 5,533
10 669 – 12 487	6,333 – 6,533

There is a top cover of about 300 m of slope sediments (1,6 km/s), underlayed by another sedimentary section of about 700 m with velocities increasing from 1,6 to 1,9 km/s. Below this sedimentary layers, velocities raise up to 3,5 km/s at a depth of 3 km and increase to 5,5 km/s at 10 km. These velocities may correspond to the presence of igneous rocks of the Nicoya Complex. There is a 200-m thick low-velocity layer at a depth of around 6,5 km, included to explain a strong reflection at about 3 s reduced travel time. In this sector of the continental margin, the plate boundary is at 11-12 km depth, according to the seismic line 4 (Figure 6.6.3.1) of the leg Sonne 81 (Ranero & von Huene, 2000). The model presented here shows another pronounced discontinuity at this depth.

The streamer record of this profile is also in concordance with the described model. Figure 6.6.3.81 shows an upper well-layered sedimentary cover, with some disturbances to the SE, approaching the erosional scarp that represents the edge of the embayment. This layer fits the described velocities from 1,6 to 1,9 km/s. A strong reflection comes from a depth of about 3 s that coincides with the strong change in velocities.

These results are in agreement with the continental margin configuration proposed by von Huene & Flueh (1994) and Hinz et al. (1996) for Costa Rica, where the subduction erosion is the dominant process.

Profile 18, SO 163-2

(Fekete, N.)

Profile P18, with a total length of 53.058 km, was acquired using the bolt gun in a NW-SE direction, parallel to the trench axis as well as to profiles P21-23 and crossing the profile P24 approximately at its middle. The data of 13 instruments (OBH62 – 69 and OBH 71 – 75) were of good quality, showing offsets of up to 50 km. The distance between the instruments was approximately 3.5 km.

Onboard forward modelling with the ray tracing application „MacRay“ (Luetgert 1992) resulted in the velocity-depth model shown in Figure 6.6.3.85. The uppermost two layers are identified as sedimentary, and have velocities of 1.6 – 2.1 km/s and 3.1-4.3 km/s and thicknesses of about 1 and 1.5 km, respectively. The sediments show a topographic stepdown in the middle of the profile, accompanied by an increase in velocity. The next layer is interpreted as the wedge, with seismic velocities ranging from 4.4 – 4.9 km/s and a thickness of 3 km. Beneath that, a 4-km thick oceanic upper crust is found, characterised by a velocity between 5.0 – 5.2 km/s. The topography of these latter two layers do not reflect that of the sediments. At a depth of about 9.5 km, a layer of 5.5 - 6.2 km/s can be found. This is interpreted as the oceanic basement, and reaches a thickness of as much as 7 km, down to the seismic boundary at a depth of 15.5 km that is thought to be the Mohorovičić discontinuity.

SO-163-2: Profiles P15-25 (Megalens & Quepos Mound)

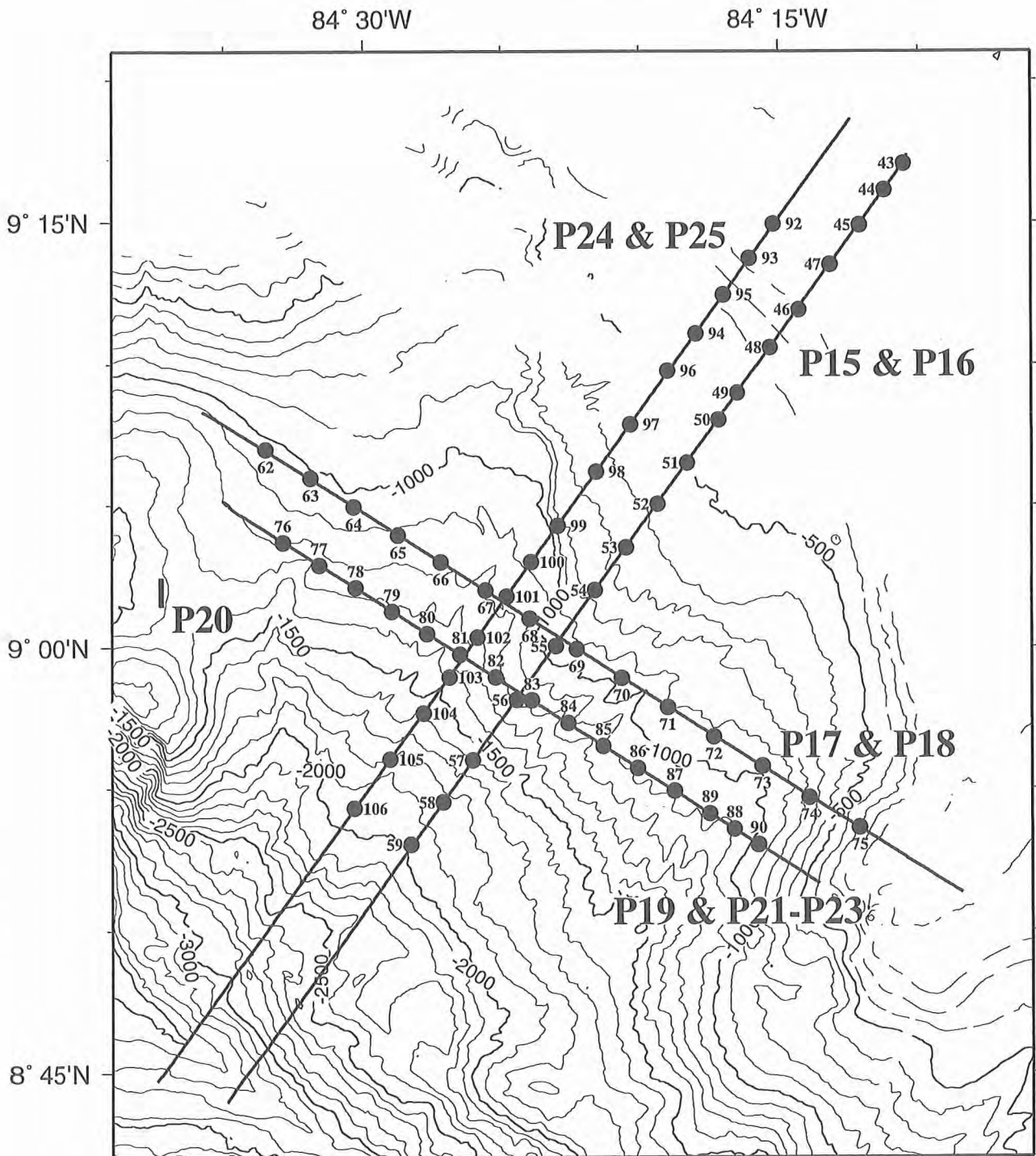


Figure 6.6.3.2: Location map of profiles P15-P25 with 100 m isolines. Stations OBH43-OBH57 (profiles P15 and P16), OBH62-OBH75 (profiles P17 and P18), OBH76-OBH90 (profiles P19-P23) and OBH92-OBH106 (profiles P24 and P25) are superimposed.

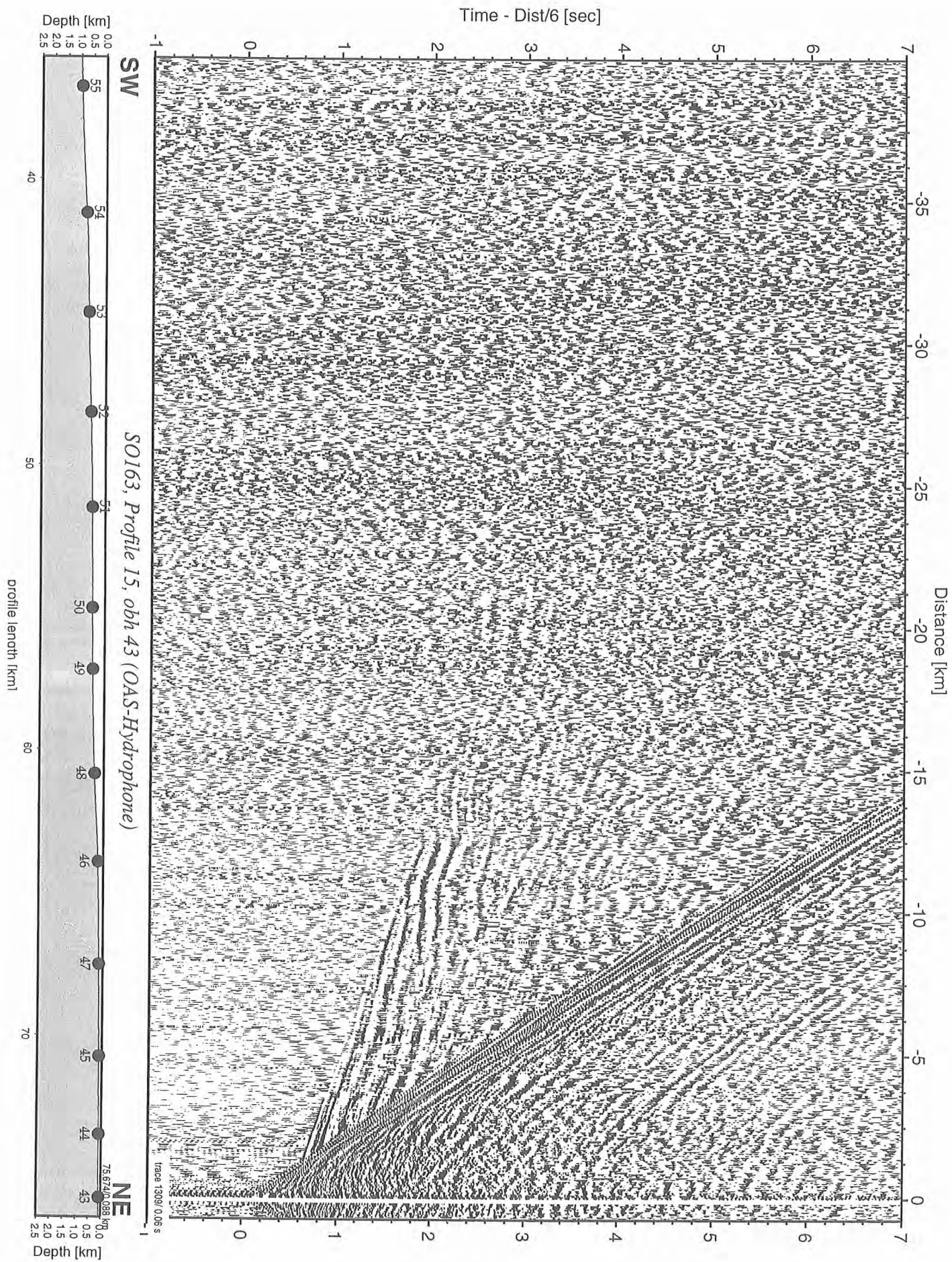


Figure 6.6.3.3: Record section from obh 43 (OAS-Hydrophone), Profile 15.

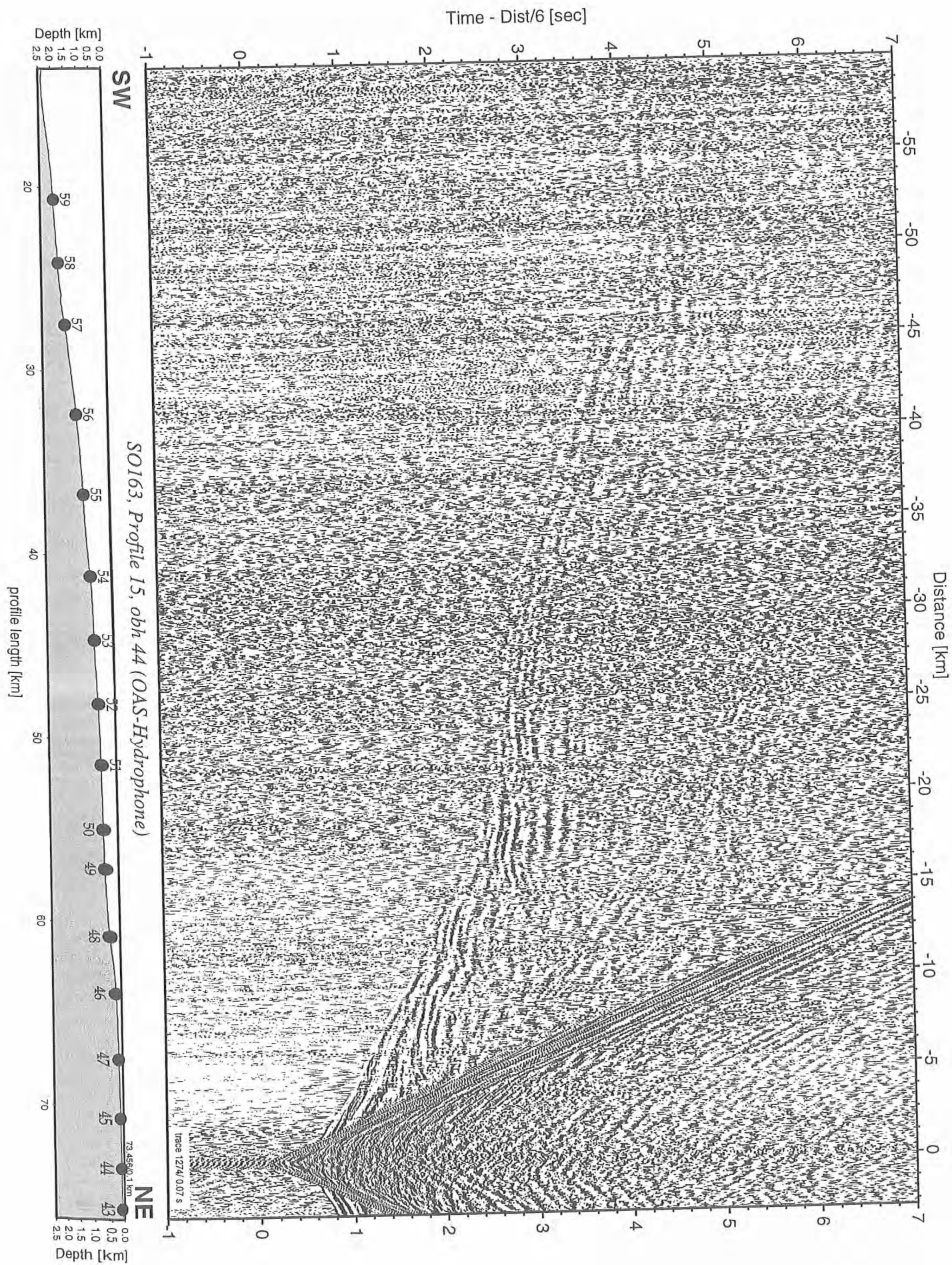


Figure 6.6.3.4: Record section from obh 44 (OAS-Hydrophone), Profile 15.

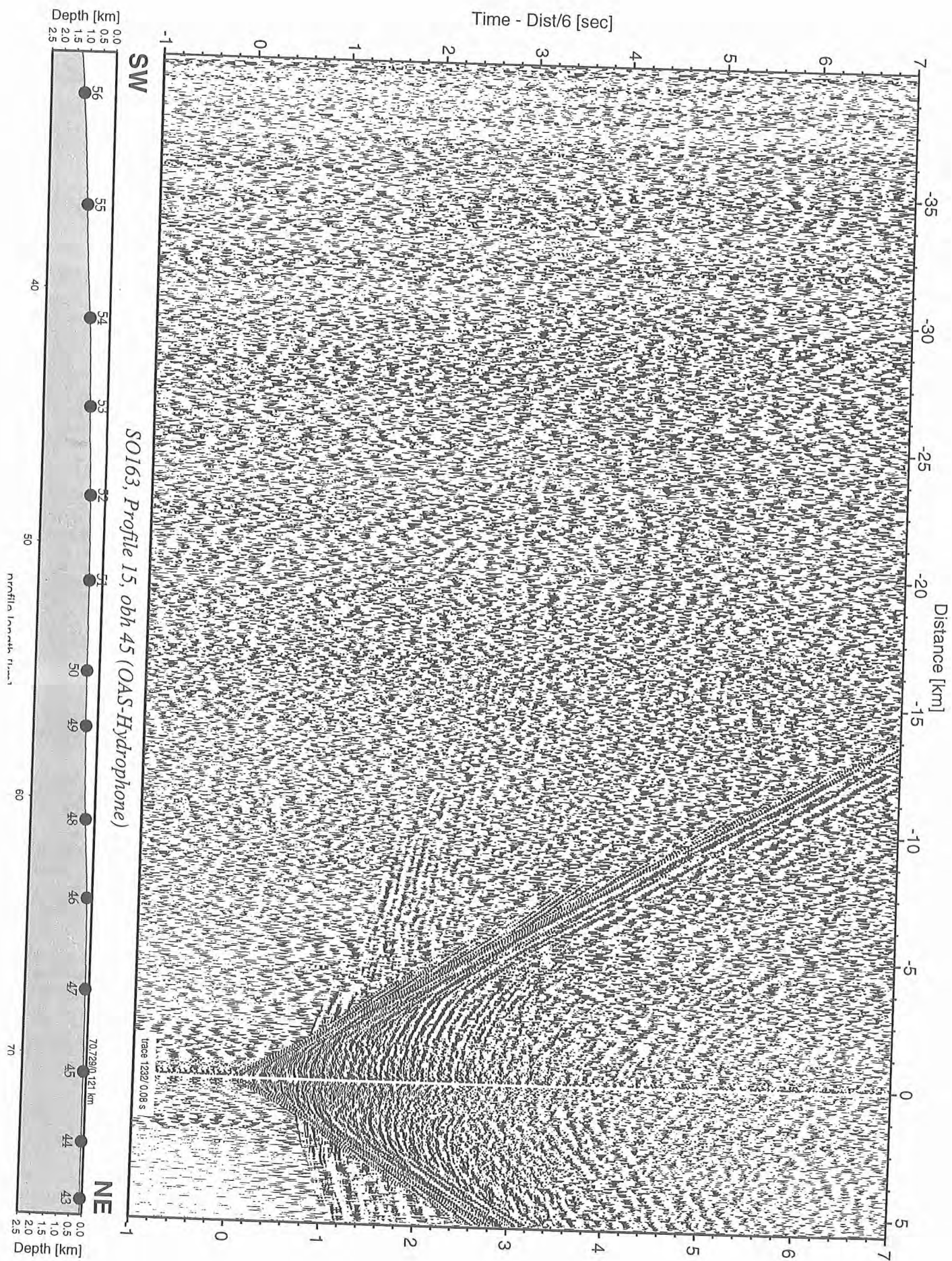


Figure 6.6.3.5: Record section from obh 45 (OAS-Hydrophone), Profile 15.

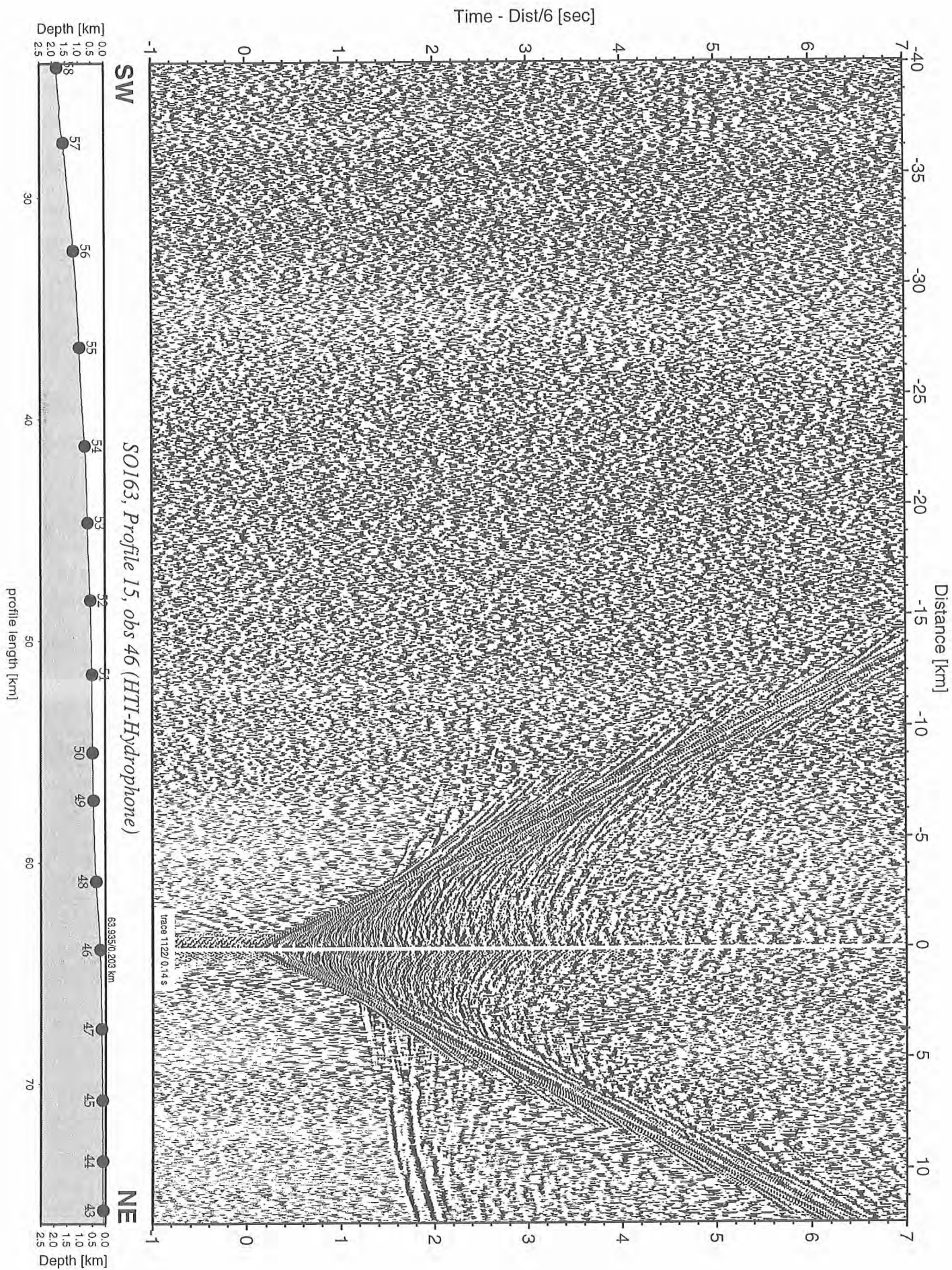


Figure 6.6.3.6: Record section from obs 46 (HTI-Hydrophone), Profile 15.

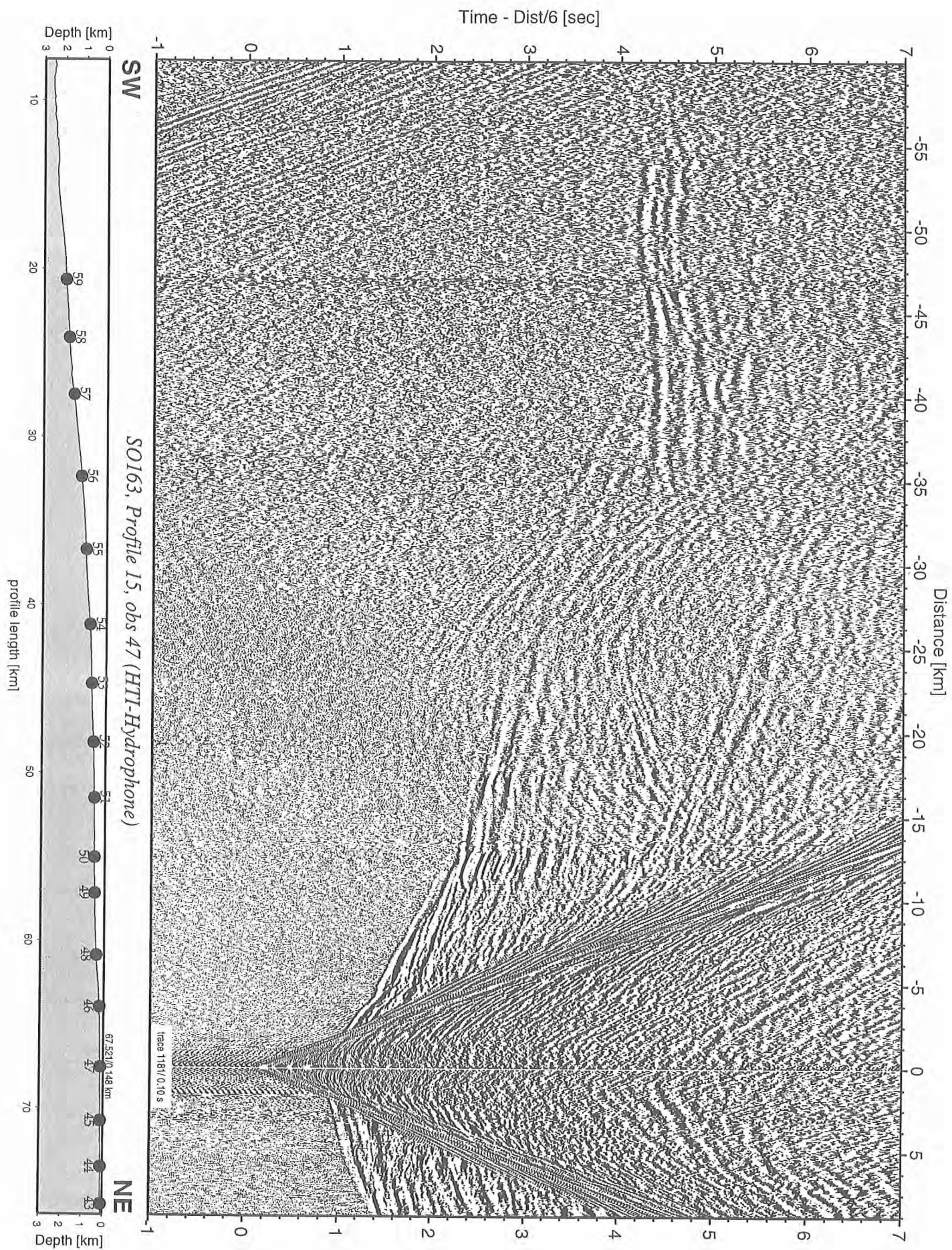


Figure 6.6.3.7: Record section from obs 47 (HTI-Hydrophone), Profile 15.

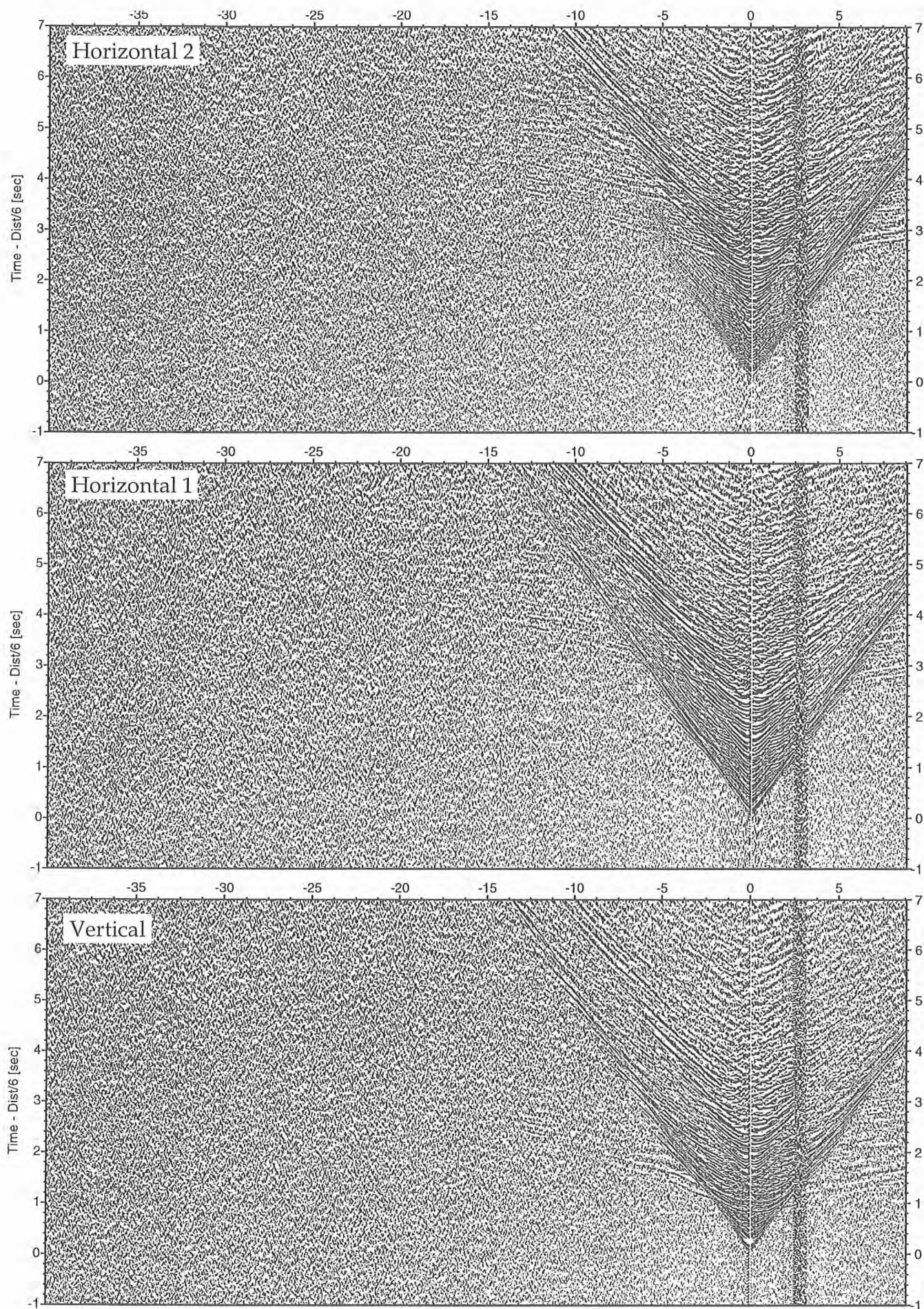


Figure 6.6.3.8: Record sections from obs 47 (Owen-30Hz), SO163, Profile 15.

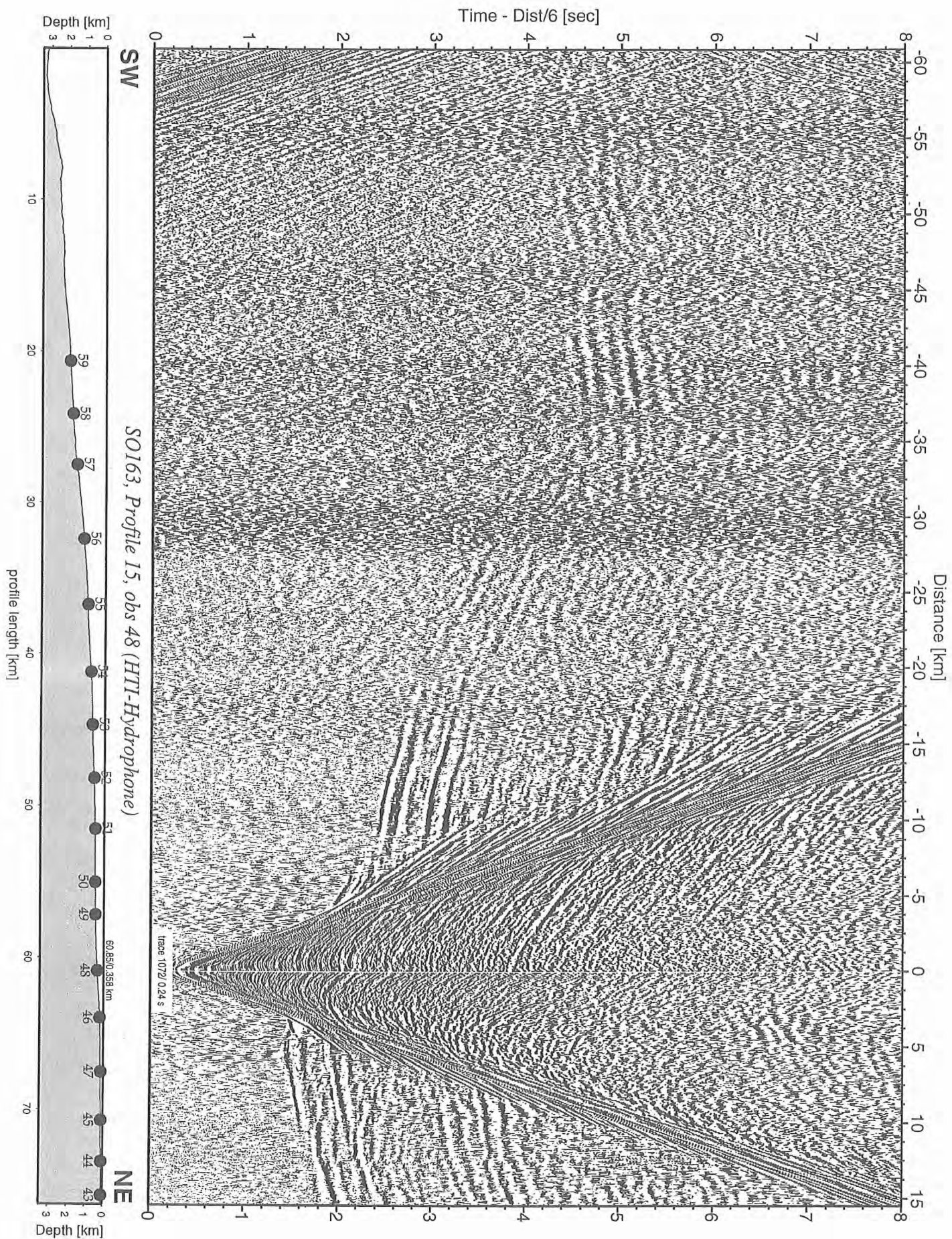


Figure 6.6.3.9: Record section from obs 48 (HTI-Hydrophone), Profile 15.

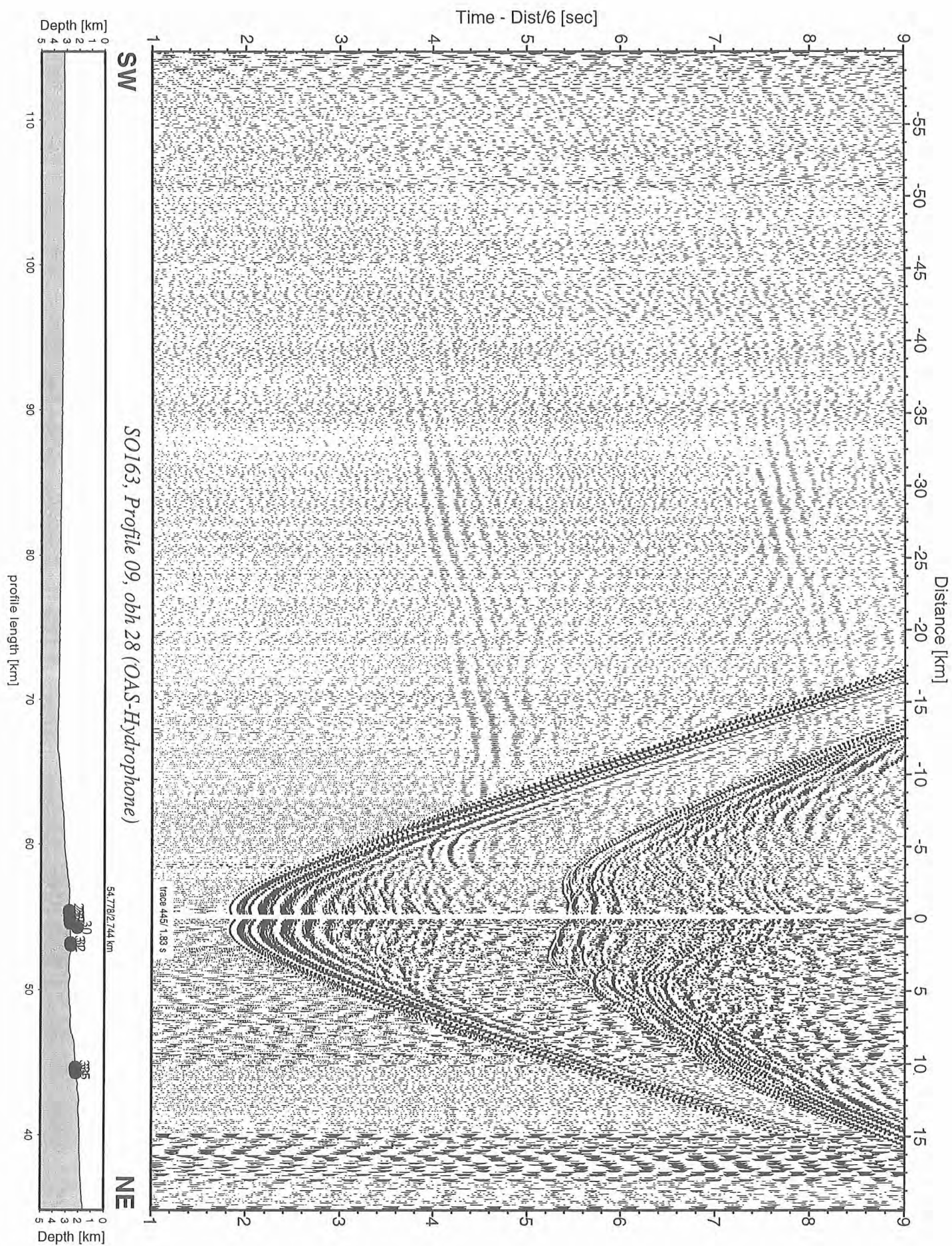


Figure 6.6.2.6: Record section from obh 28 (OAS-Hydrophone), Profile 09.

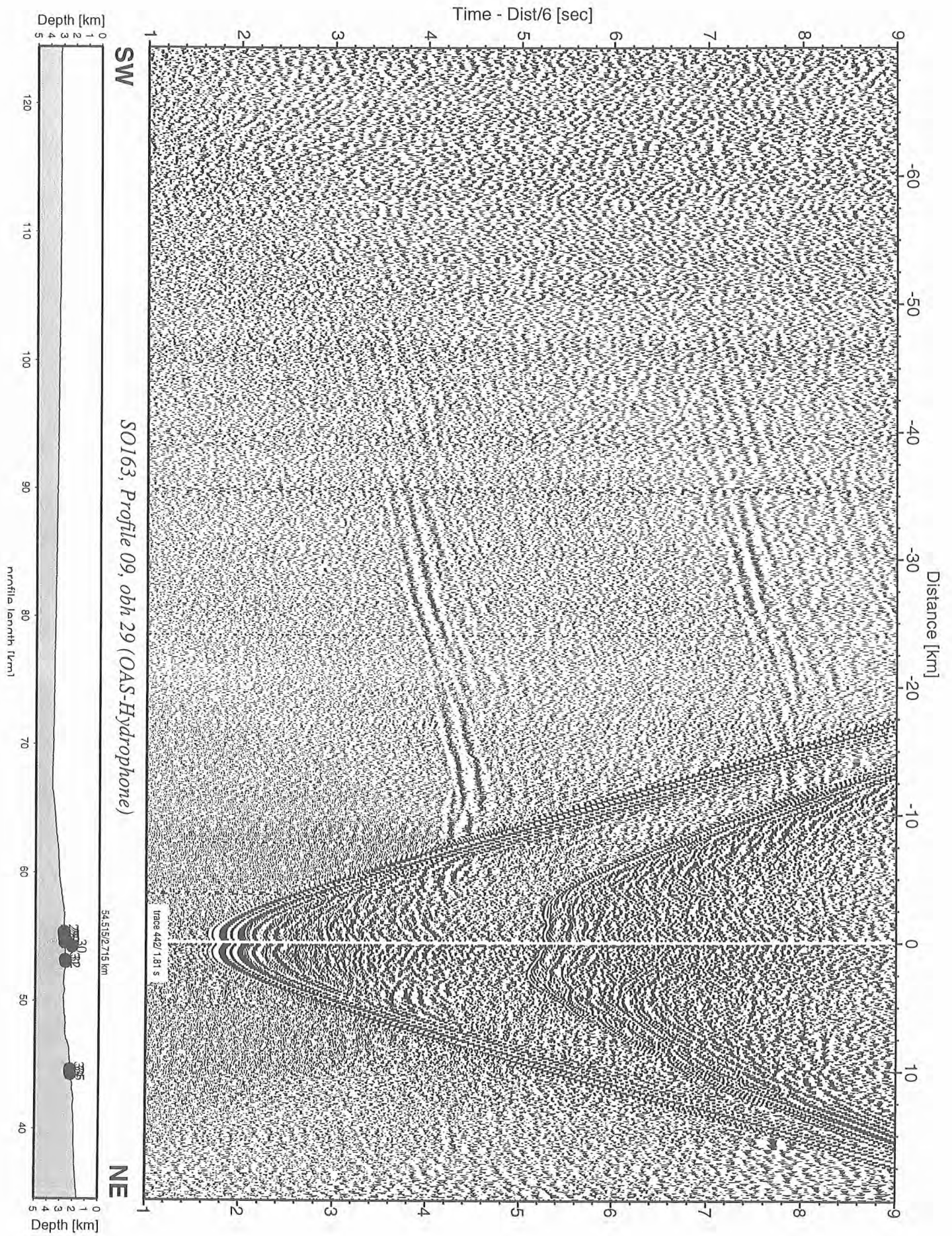


Figure 6.6.2.7: Record section from obh 29 (OAS-Hydrophone), Profile 09.

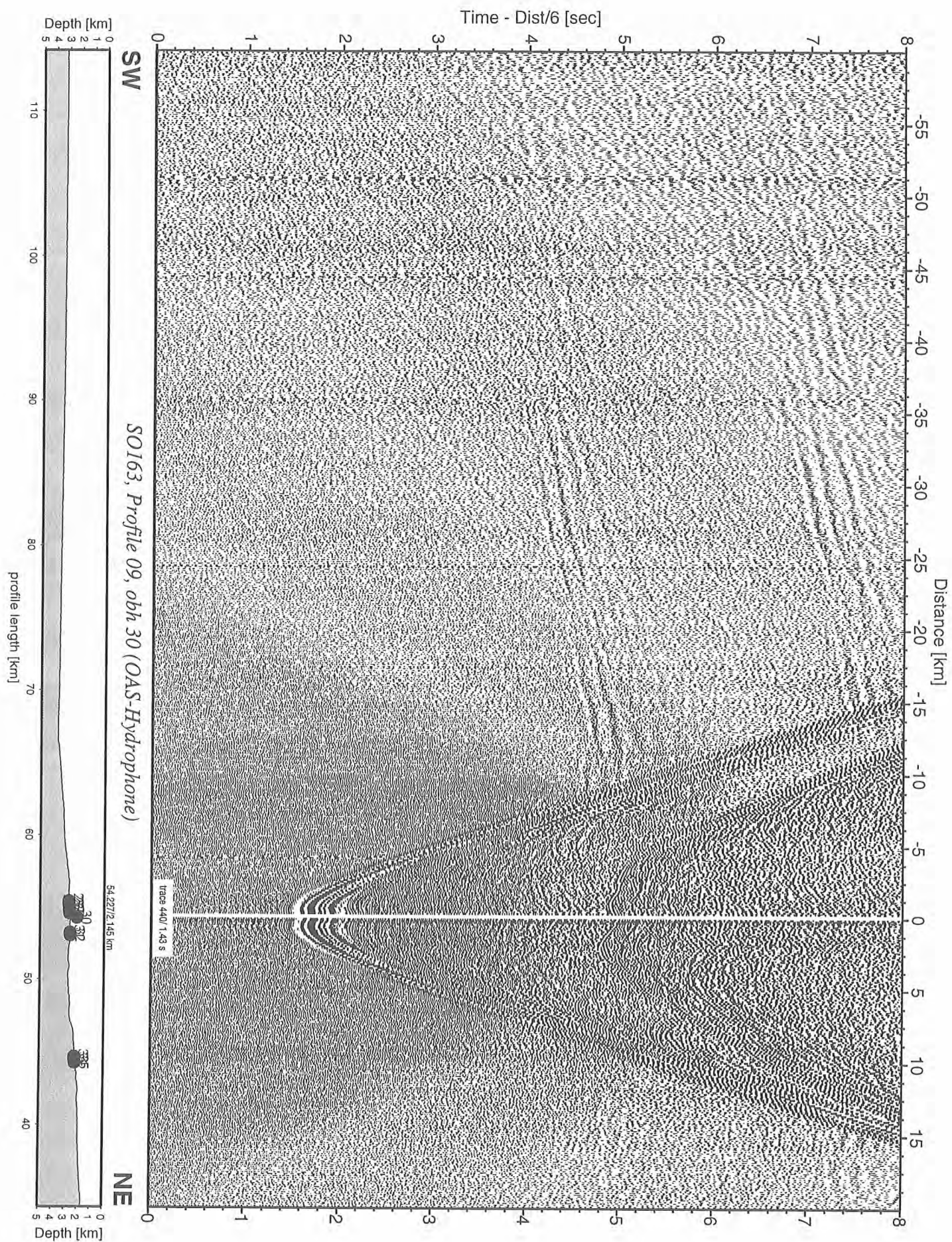


Figure 6.6.2.8: Record section from obh 30 (OAS-Hydrophone), Profile 09.

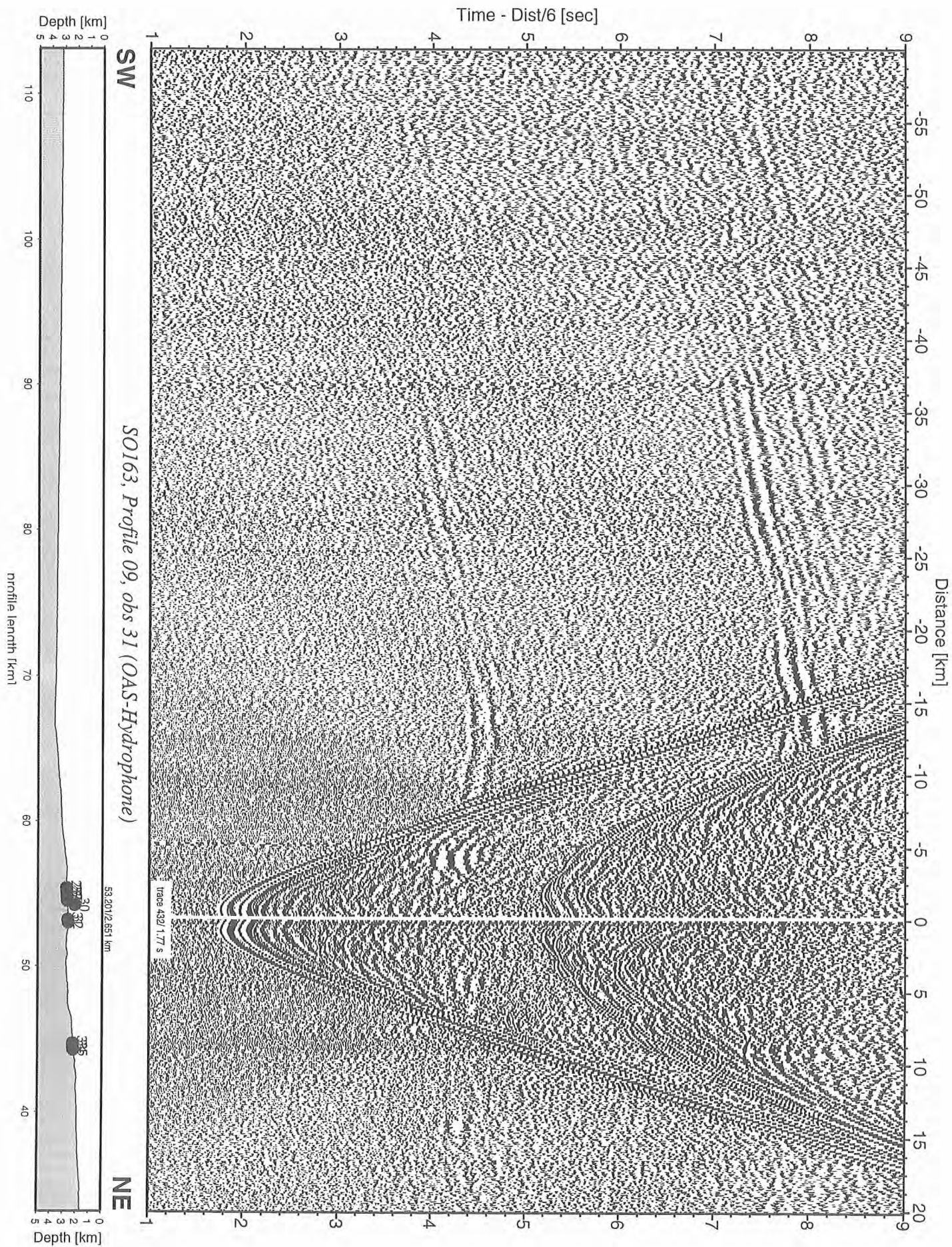


Figure 6.6.2.9: Record section from obs 31 (OAS-Hydrophone), Profile 09.

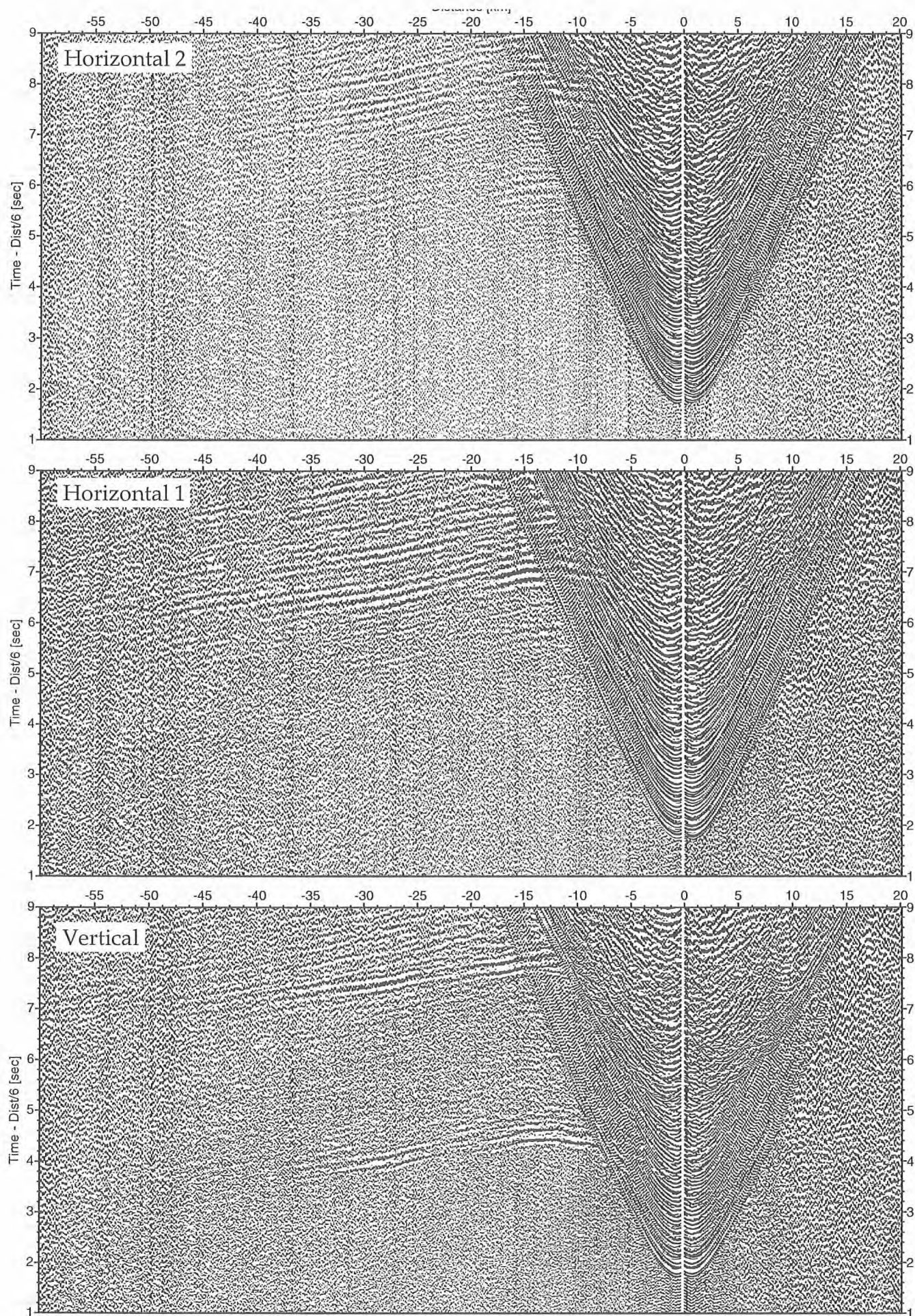


Figure 6.6.2.10: Record sections from obs 31 (Owen-4.5Hz), SO163, Profile 09.

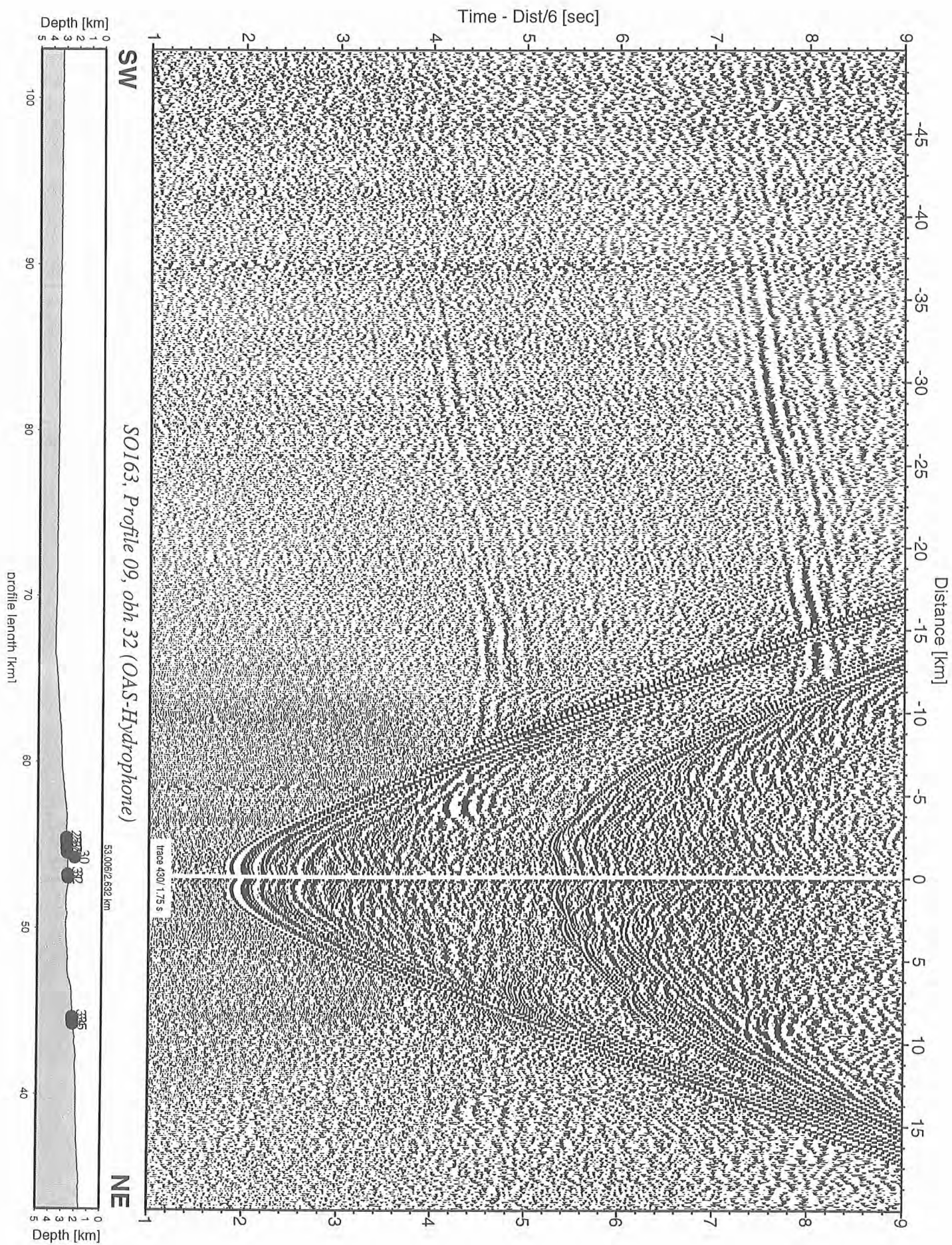


Figure 6.6.2.11: Record section from obh 32 (OAS-Hydrophone), Profile 09.

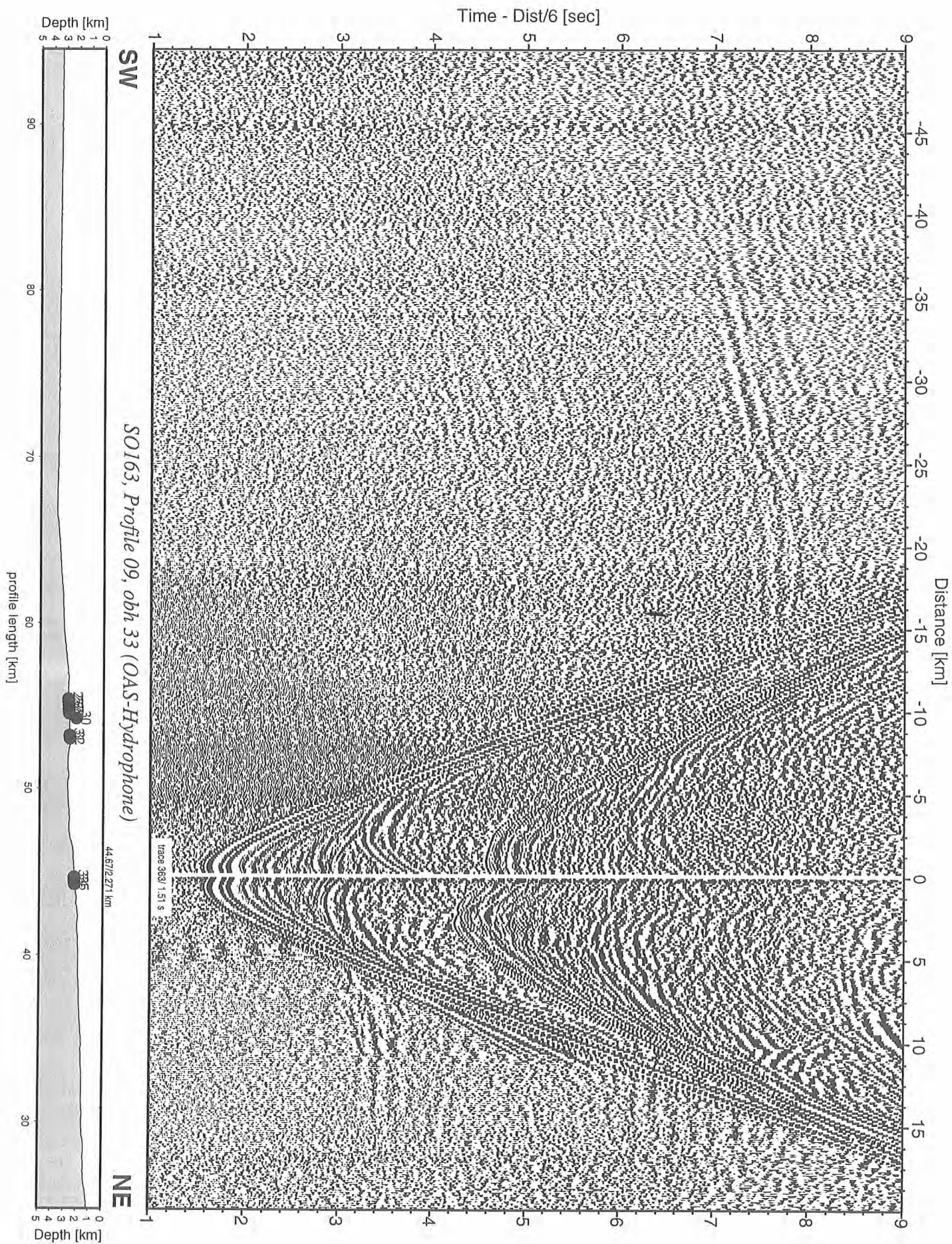


Figure 6.6.2.12: Record section from obh 33 (OAS-Hydrophone), Profile 09.

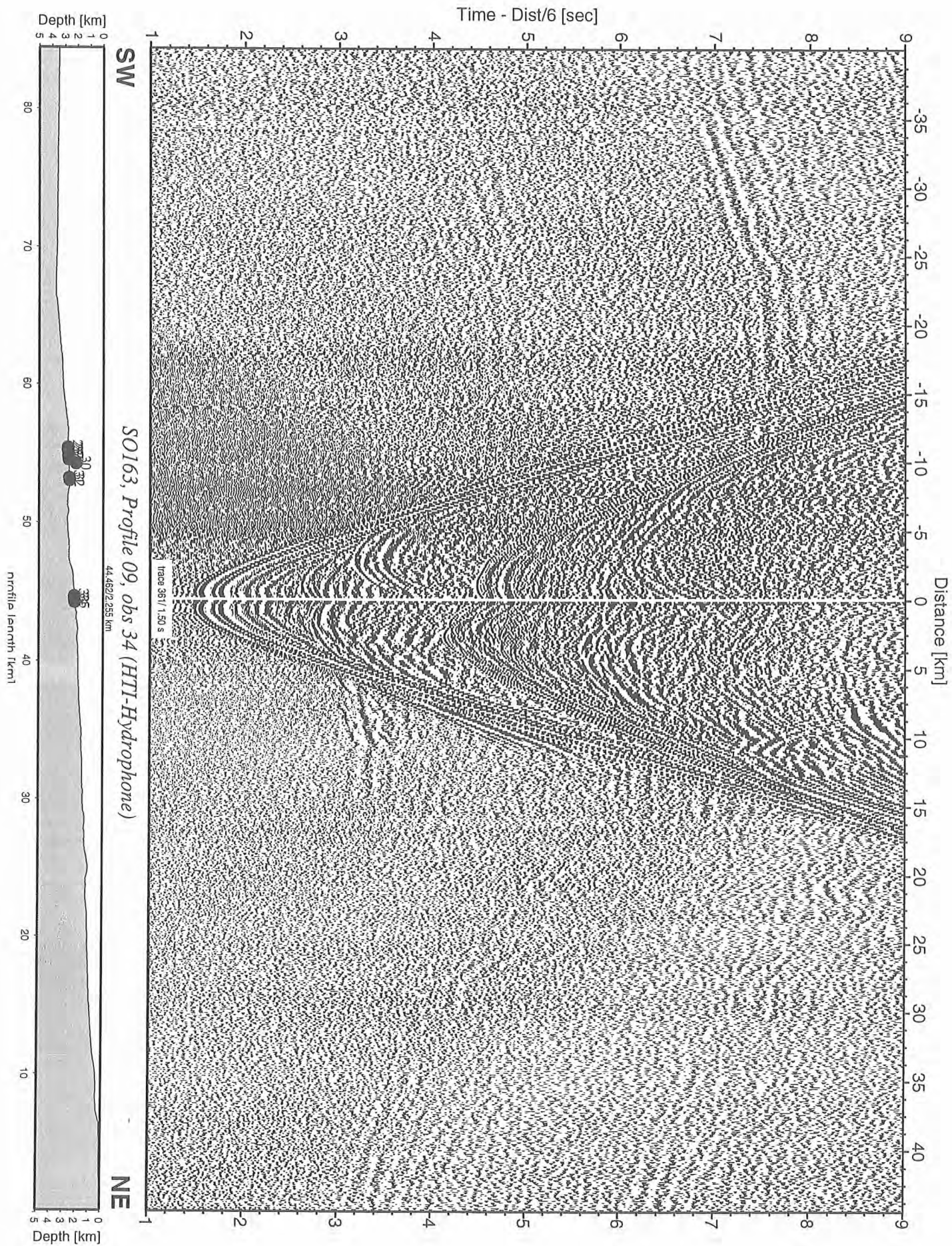


Figure 6.6.2.13: Record section from obs 34 (HTI-Hydrophone), Profile 09.

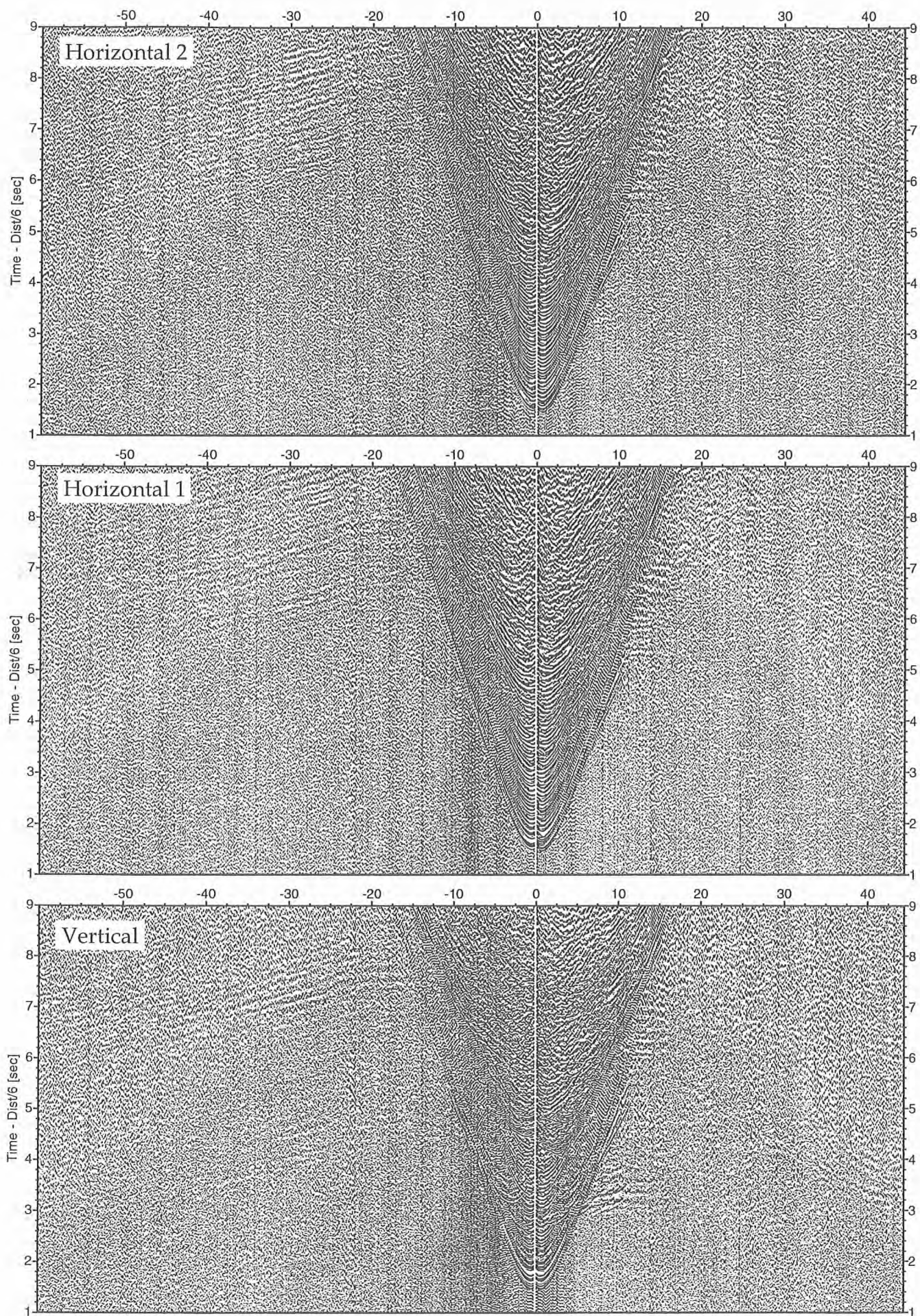


Figure 6.6.2.14: Record sections from obs 34 (Owen-4.5Hz), SO163, Profile 09.

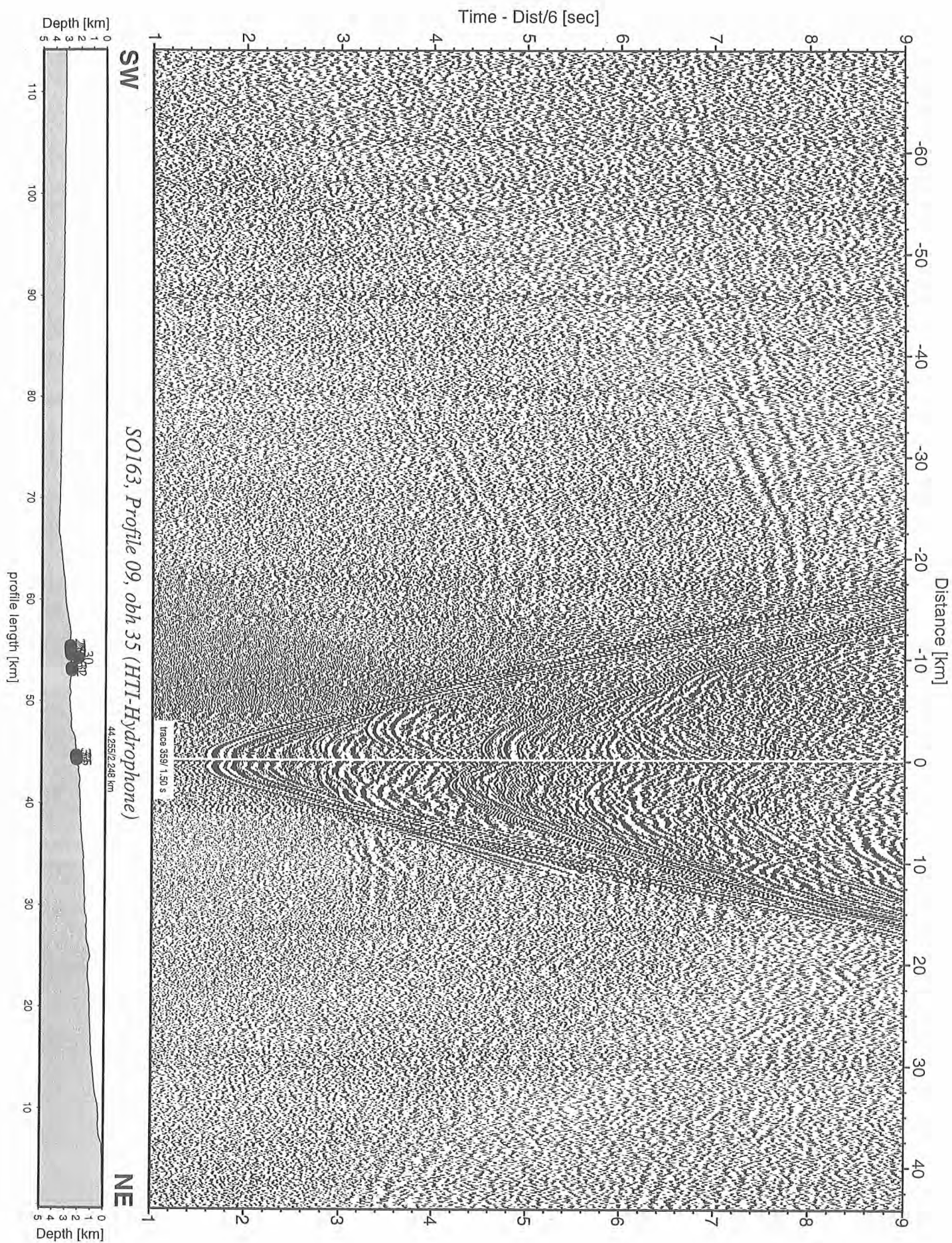


Figure 6.6.2.15: Record section from obh 35 (HTI-Hydrophone), Profile 09.

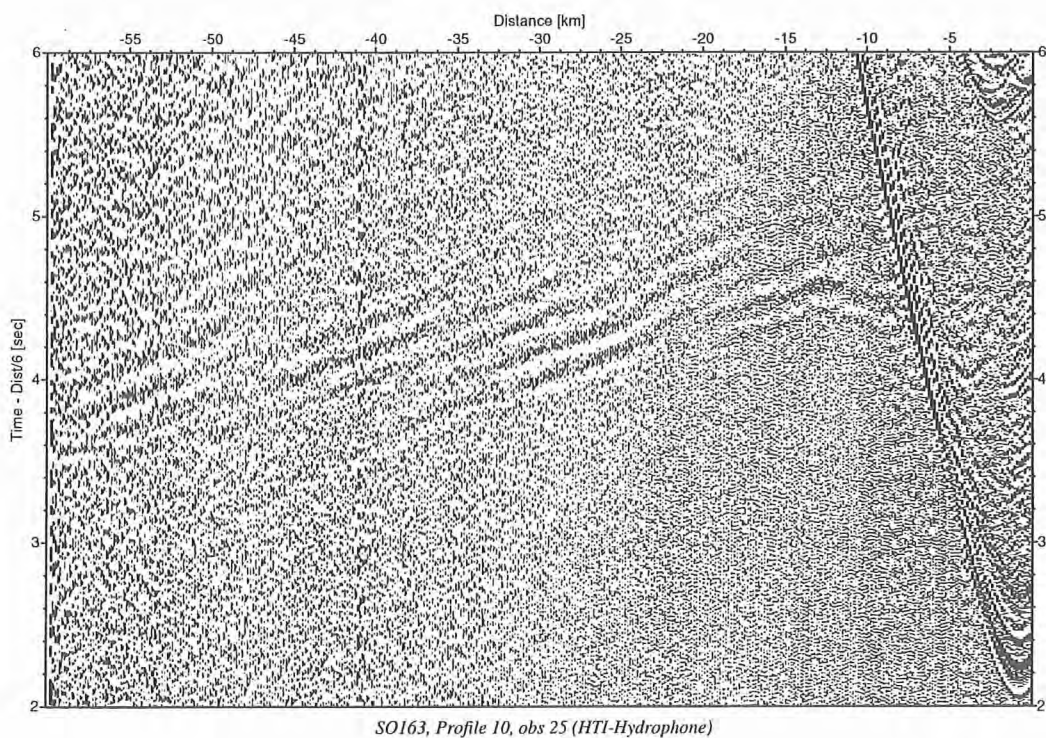
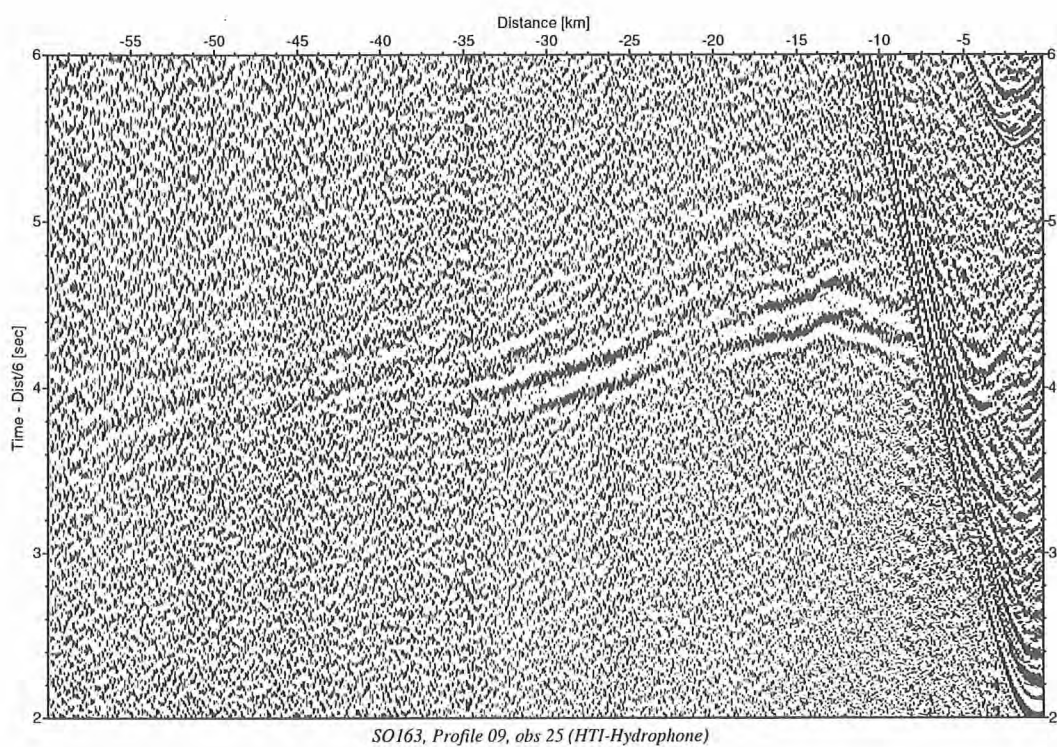


Figure 6.6.2.16.: Comparison between profiles P09 (Bolt gun) and P10 (G-gun) for station OBS 25.

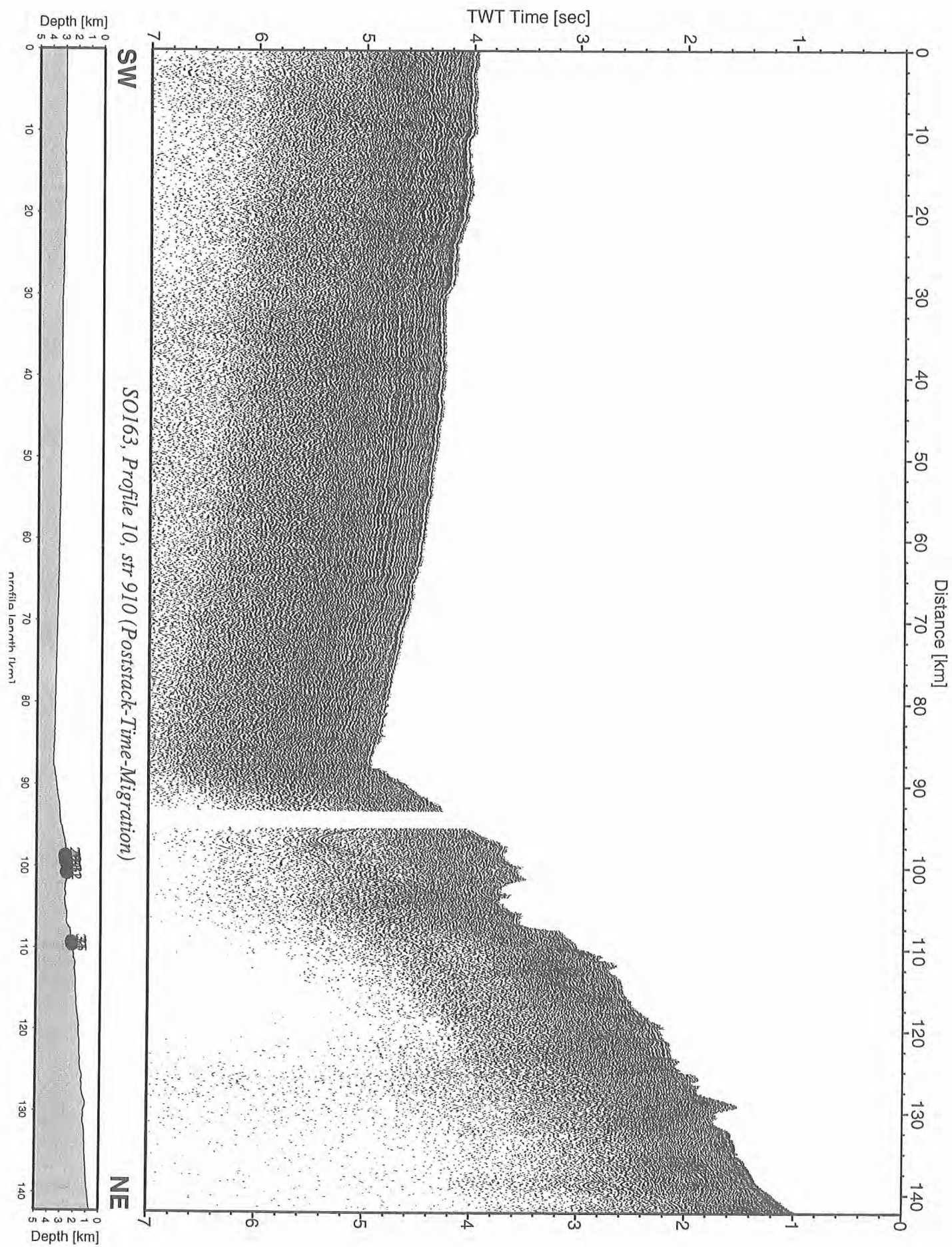


Figure 6.6.2.17: Record section from str 910 (Poststack-Time-Migration), Profile 10.

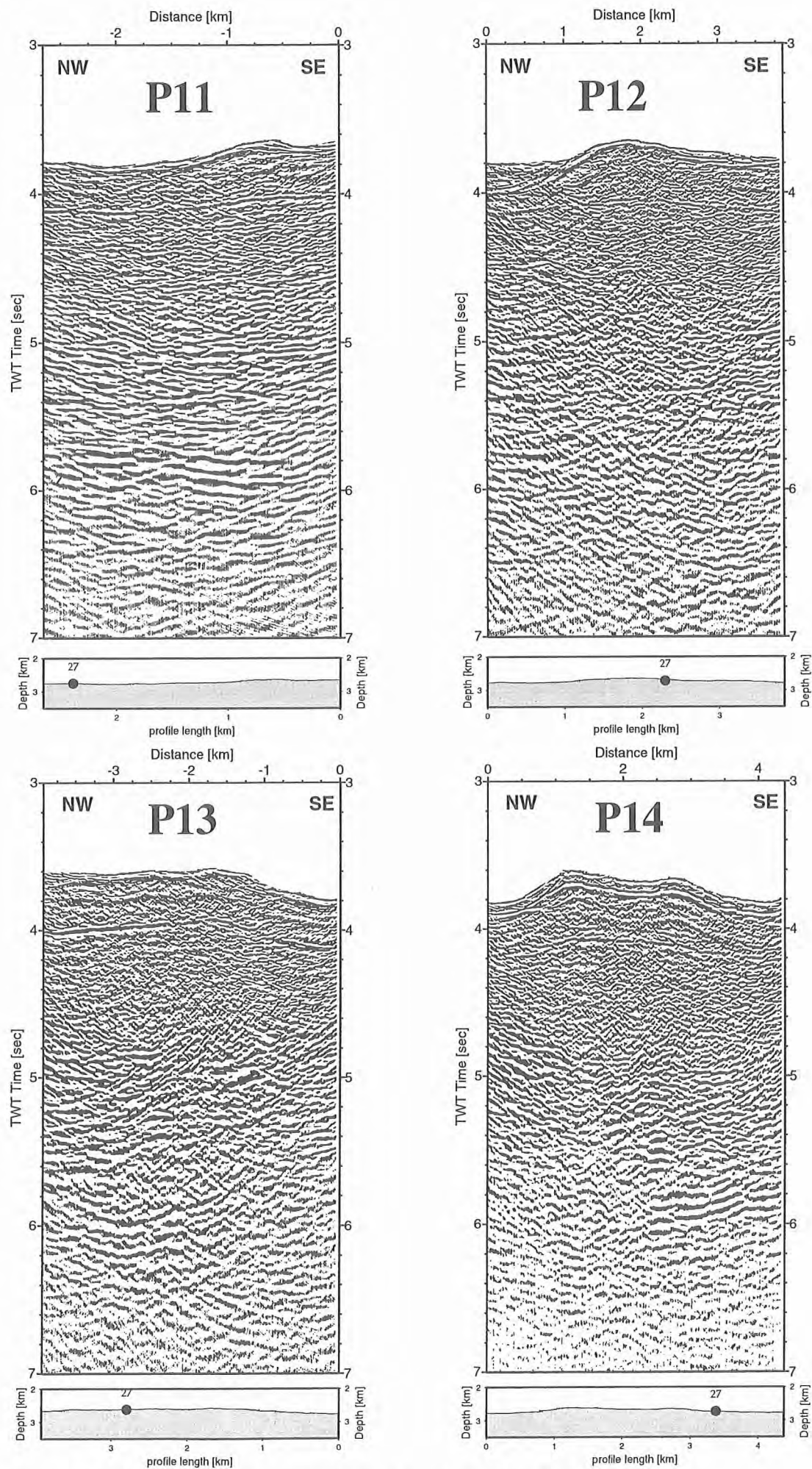


Figure 6.6.2.18: Post-stack time migrated sections for profiles P11-P14.

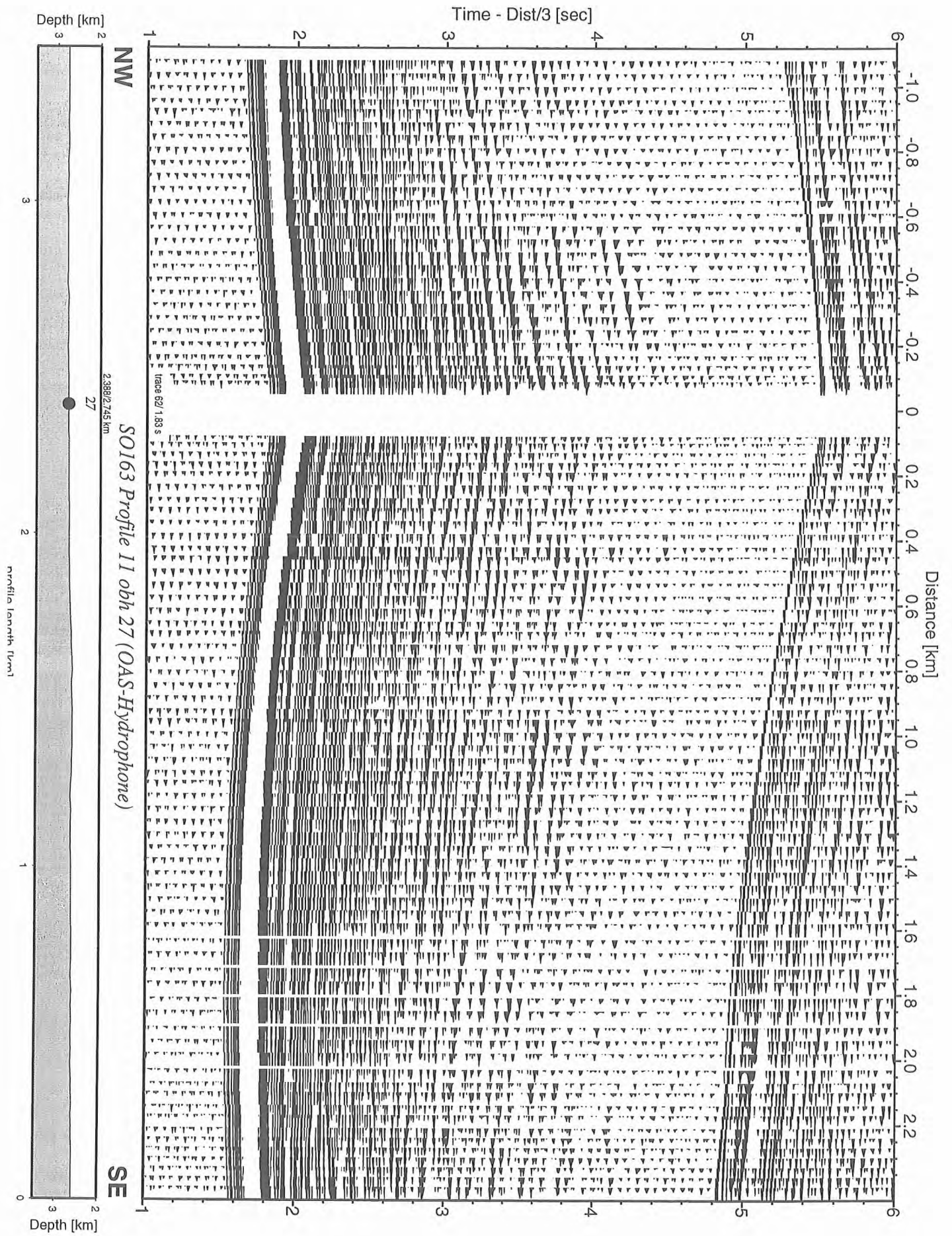


Figure 6.6.2.19: Record section from obh 27 (OAS-Hydrophone), Profile 11.

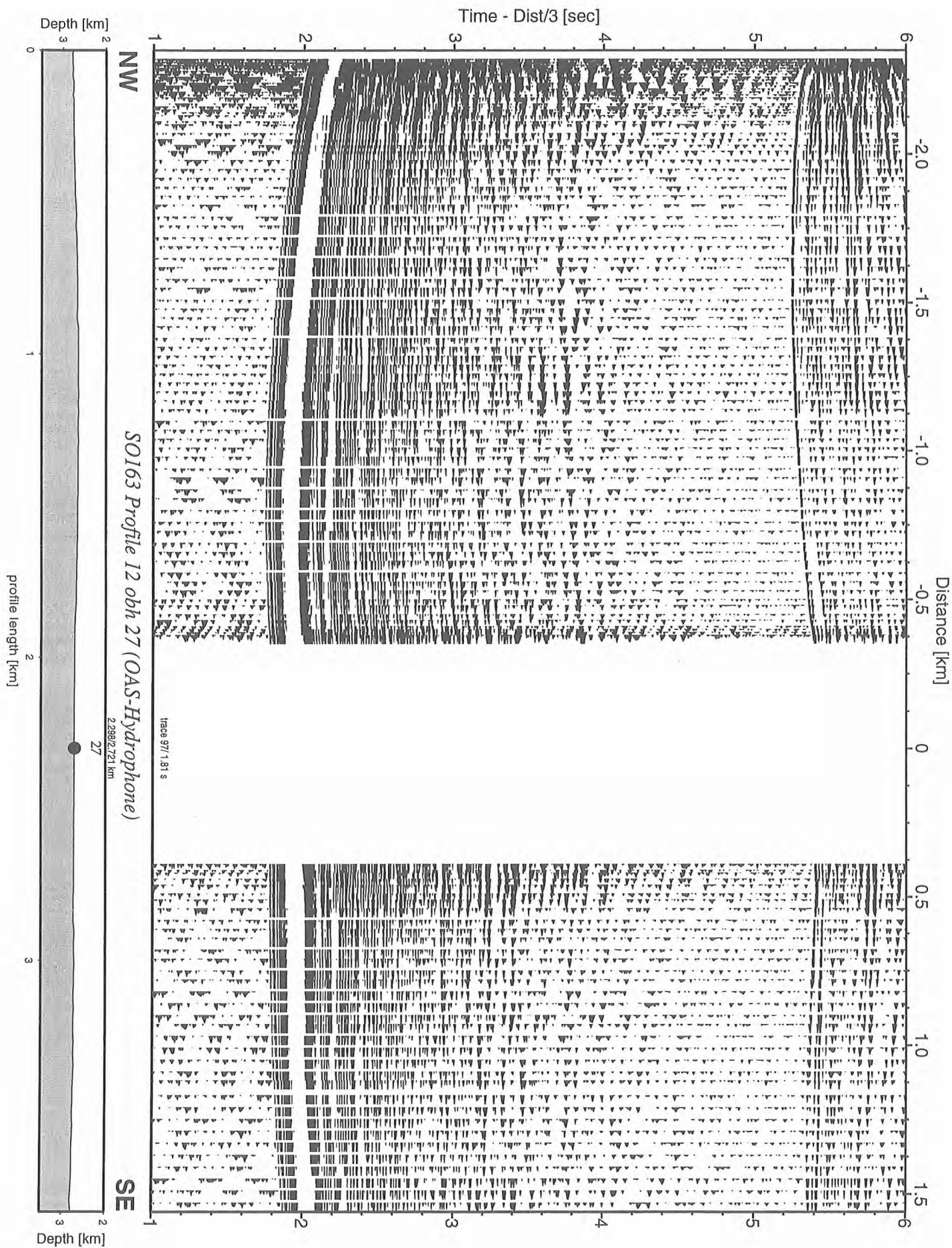


Figure 6.6.2.20: Record section from obh 27 (OAS-Hydrophone), Profile 12.

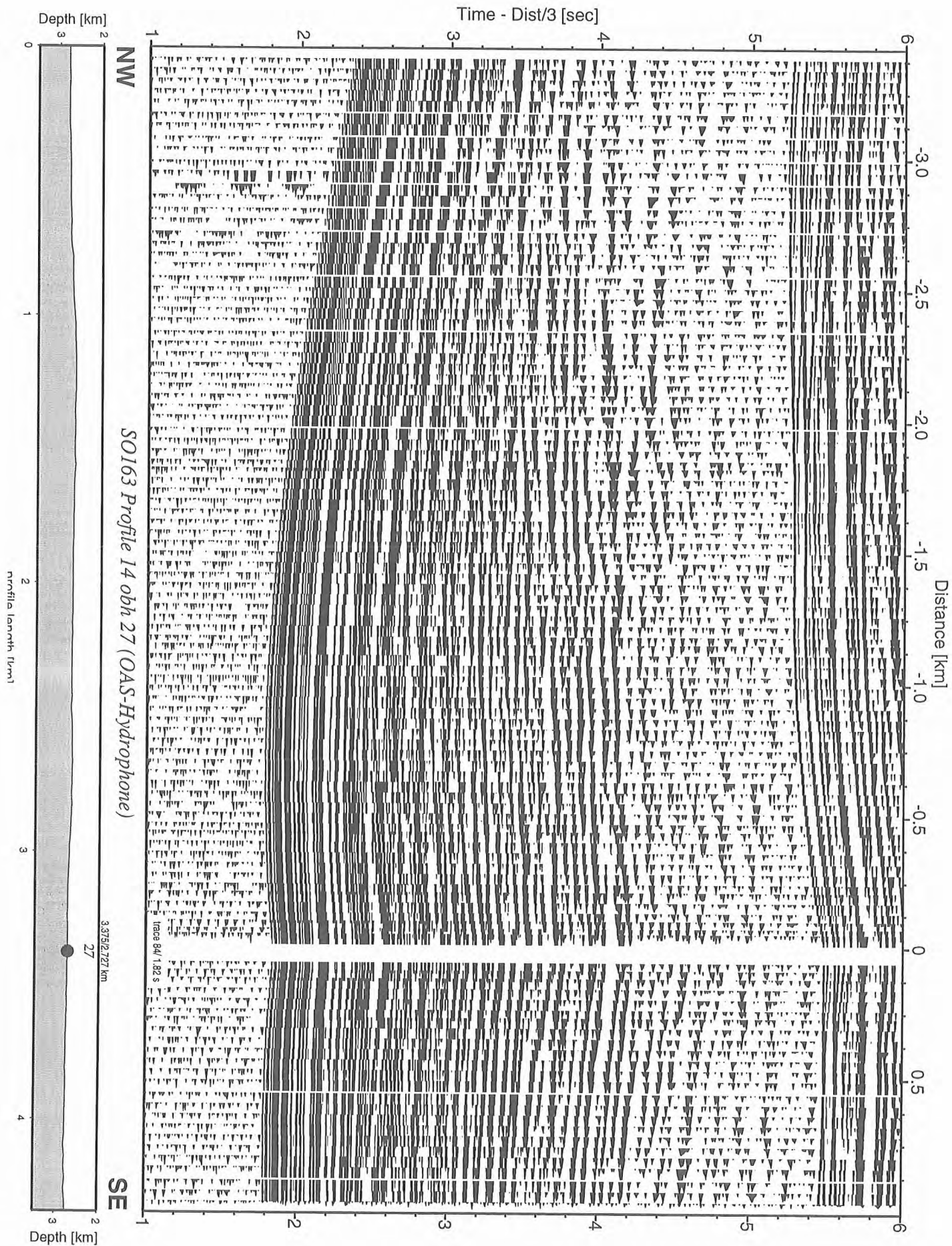


Figure 6.6.21: Record section from obh 27 (OAS-Hydrophone), Profile 14.

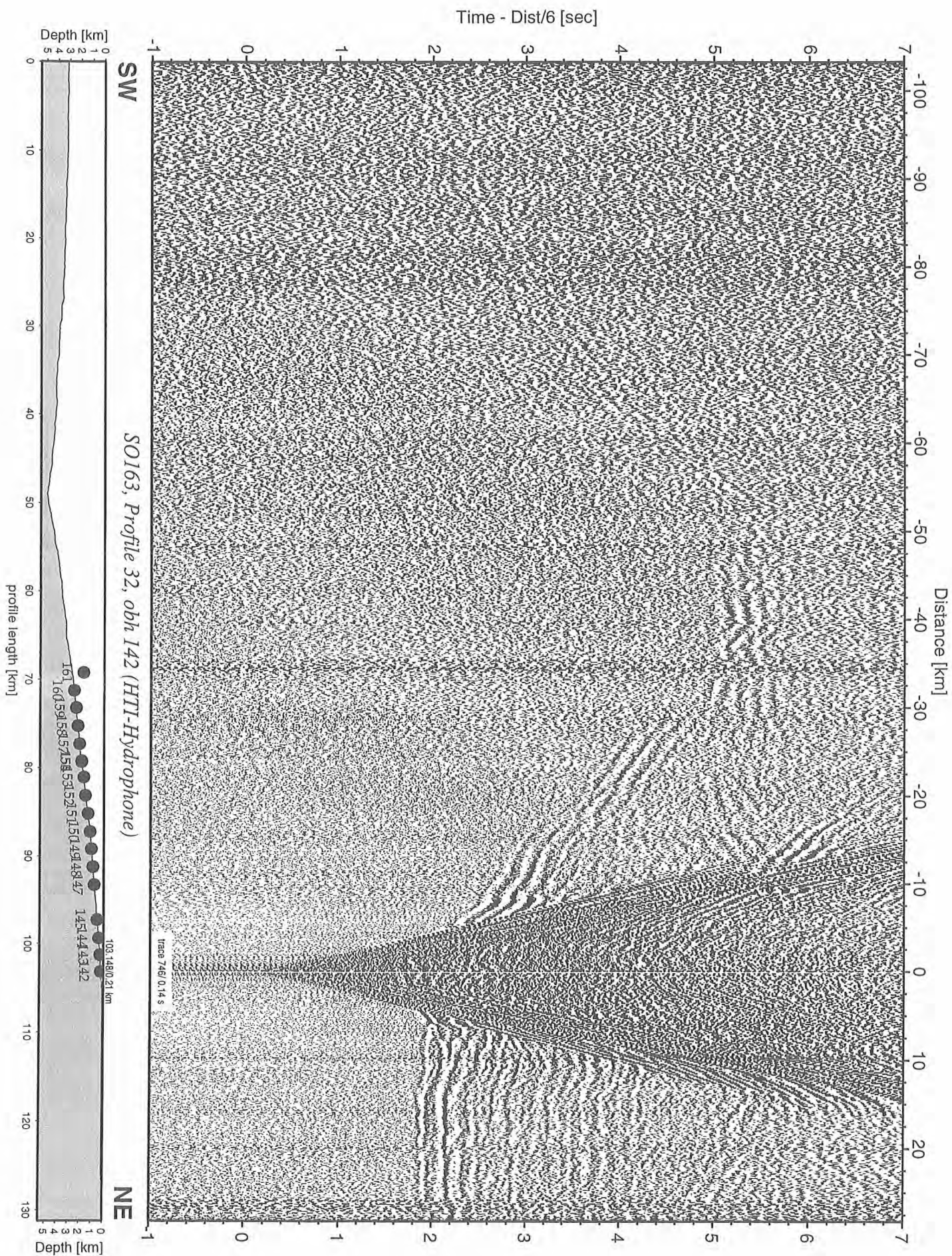


Figure 6.6.2.22 Record section from obh 142 (HTI-Hydrophone), Profile 32.

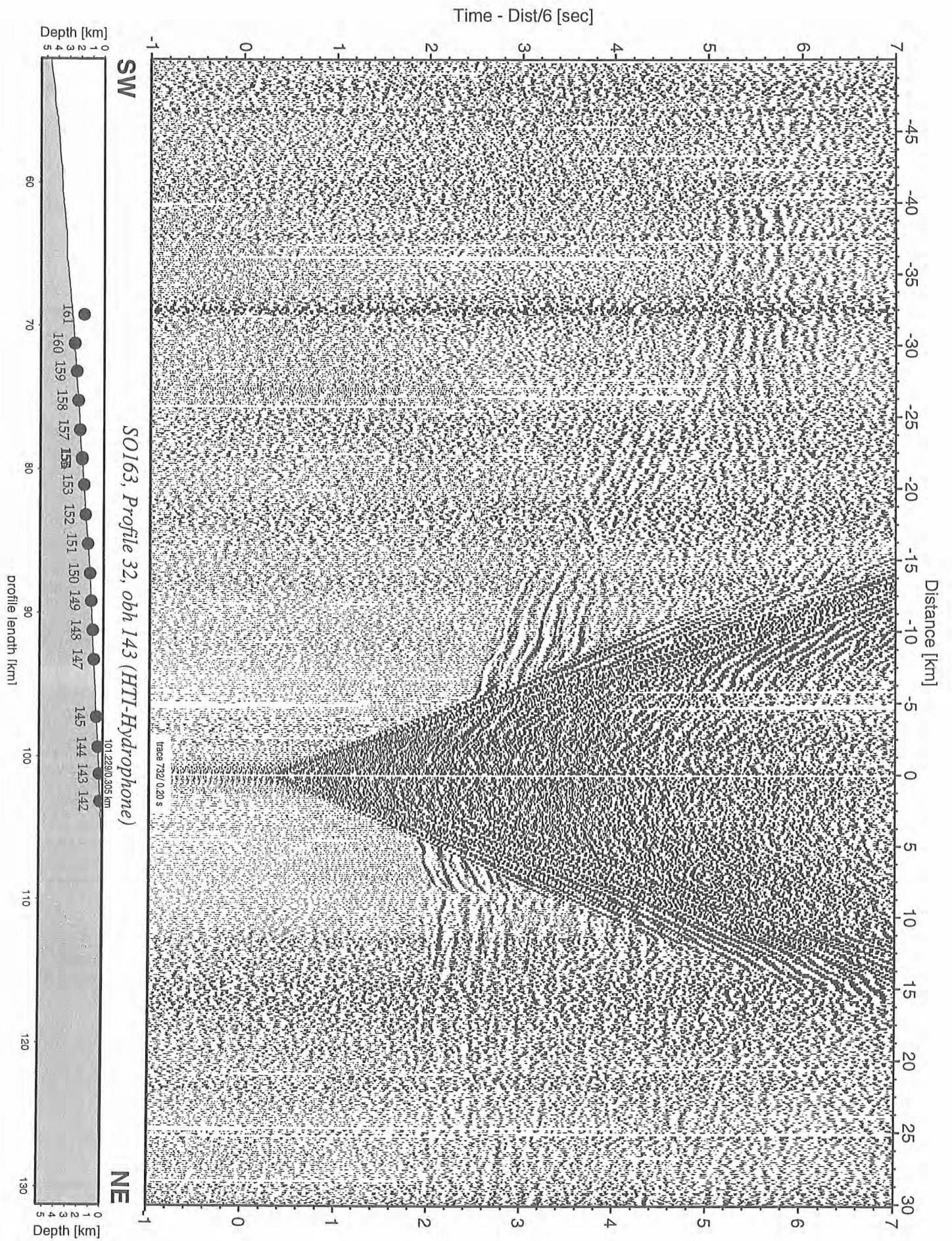


Figure 6.6.2.23 Record section from obh 143 (HTI-Hydrophone), Profile 32.

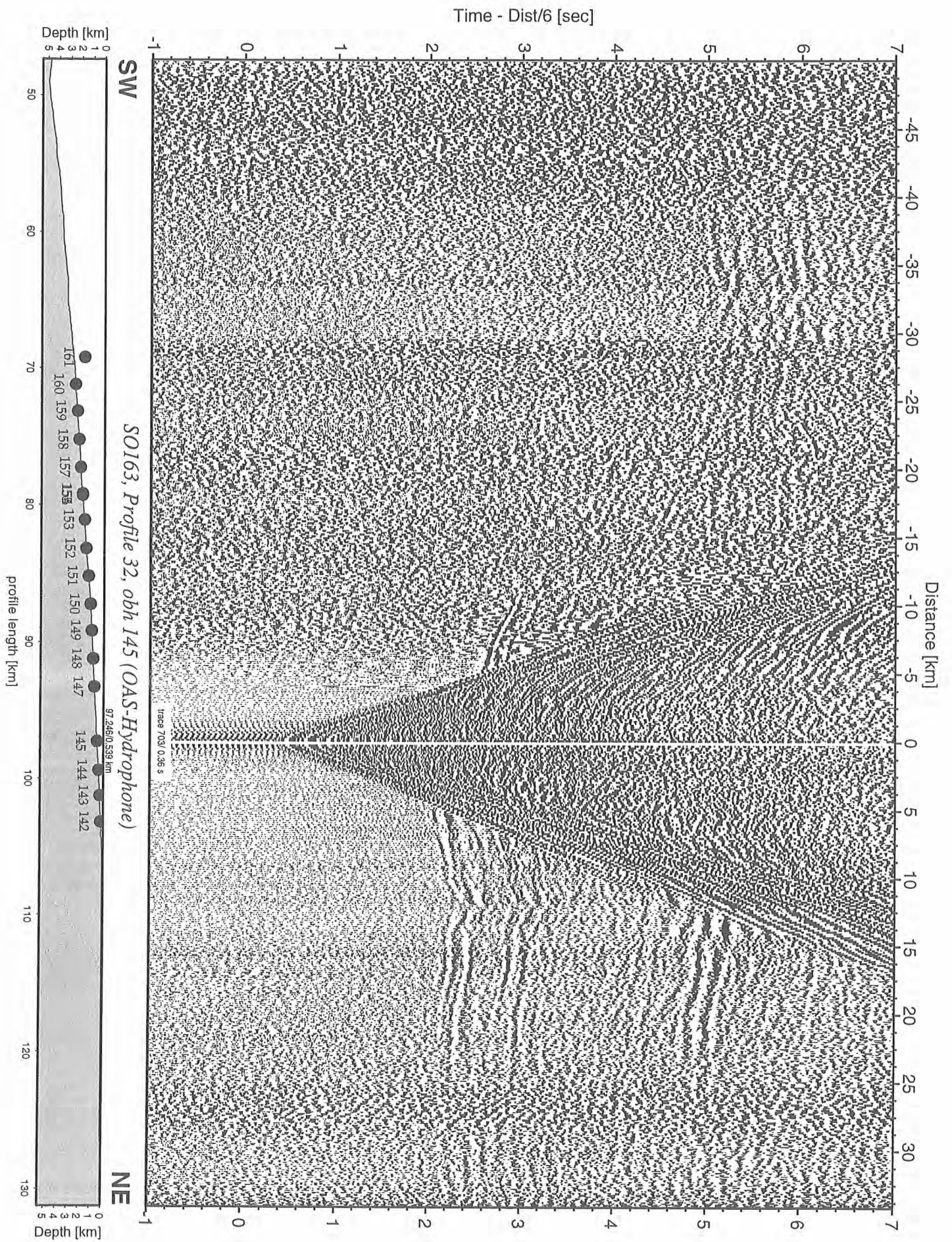


Figure 6.6.2.24 Record section from obh 145 (OAS-Hydrophone), Profile 32.

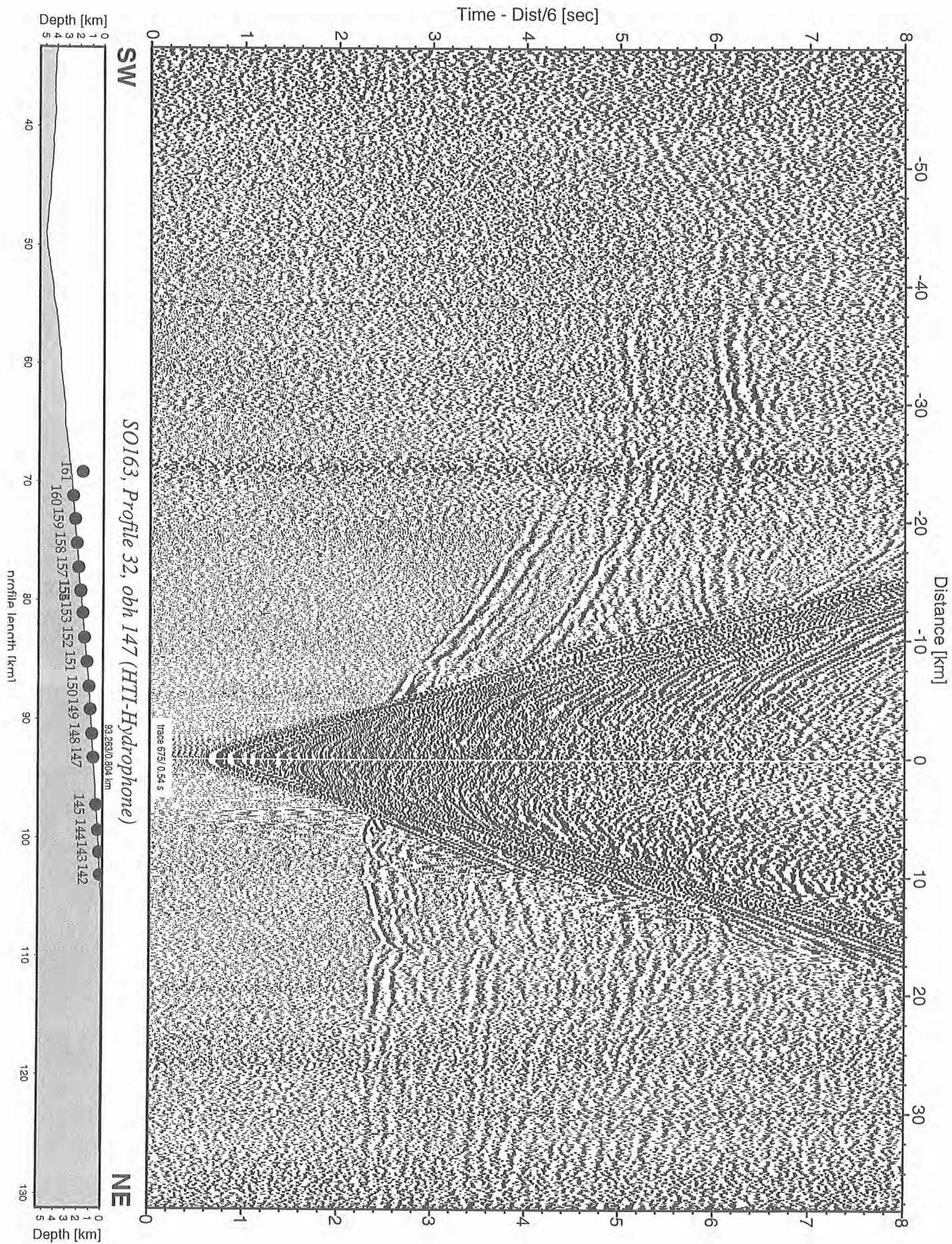


Figure 6.6.2.25 Record section from obh 147 (HTI-Hydrophone), Profile 32.

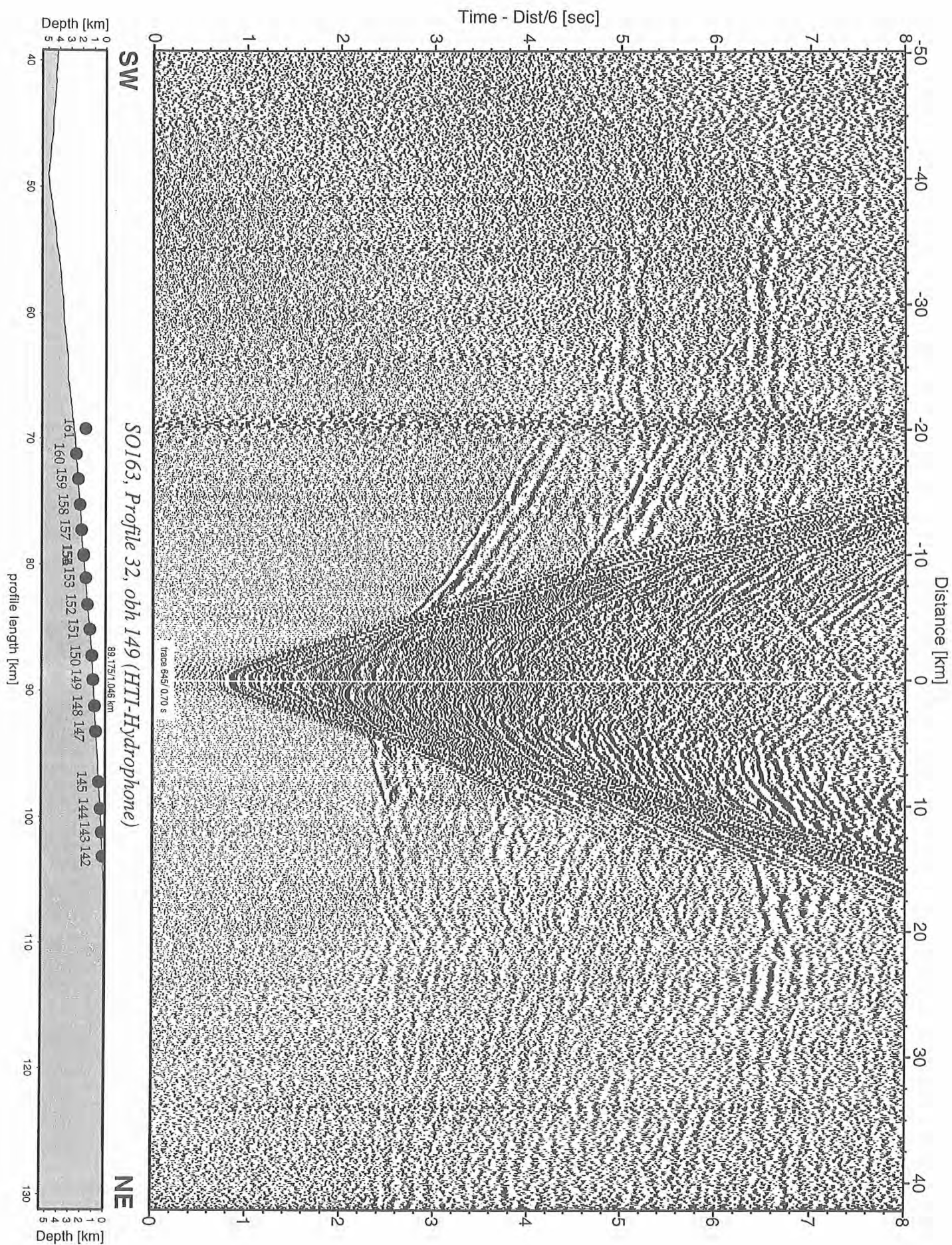


Figure 6.6.2.26 Record section from obh 149 (HTI-Hydrophone), Profile 32.

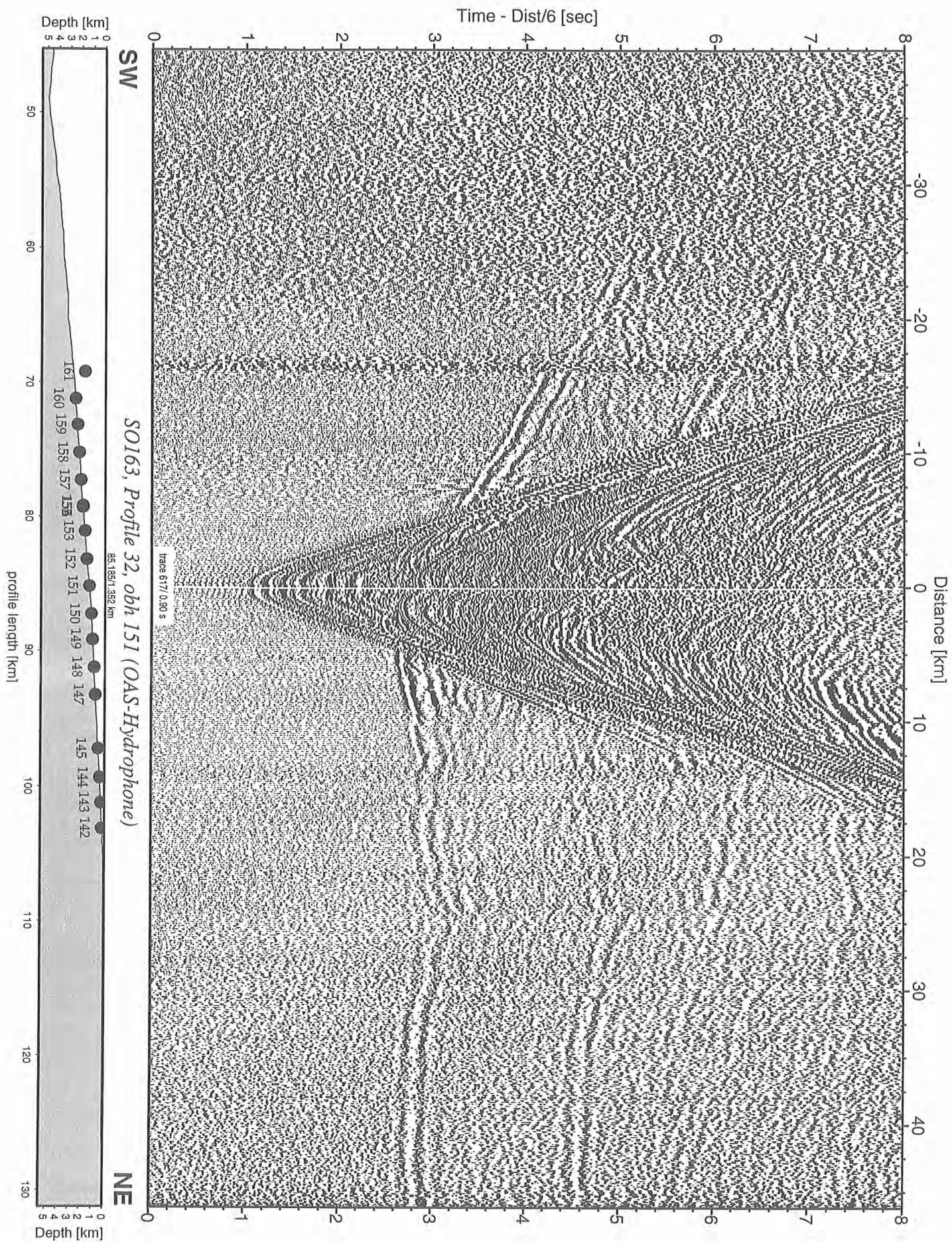


Figure 6.6.2.27 Record section from obh 151 (OAS-Hydrophone), Profile 32.

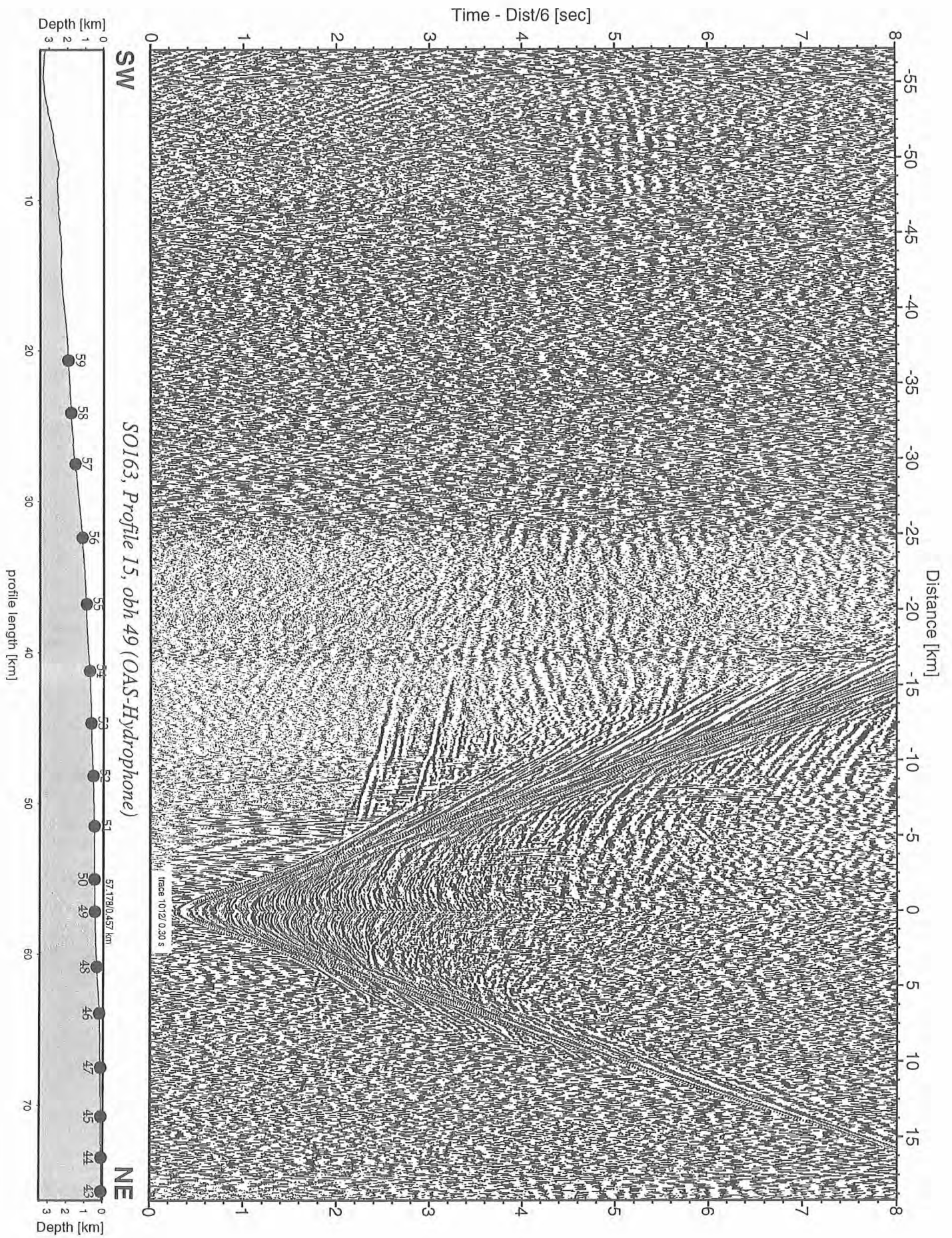


Figure 6.6.3.10: Record section from obh 49 (OAS-Hydrophone), Profile 15.

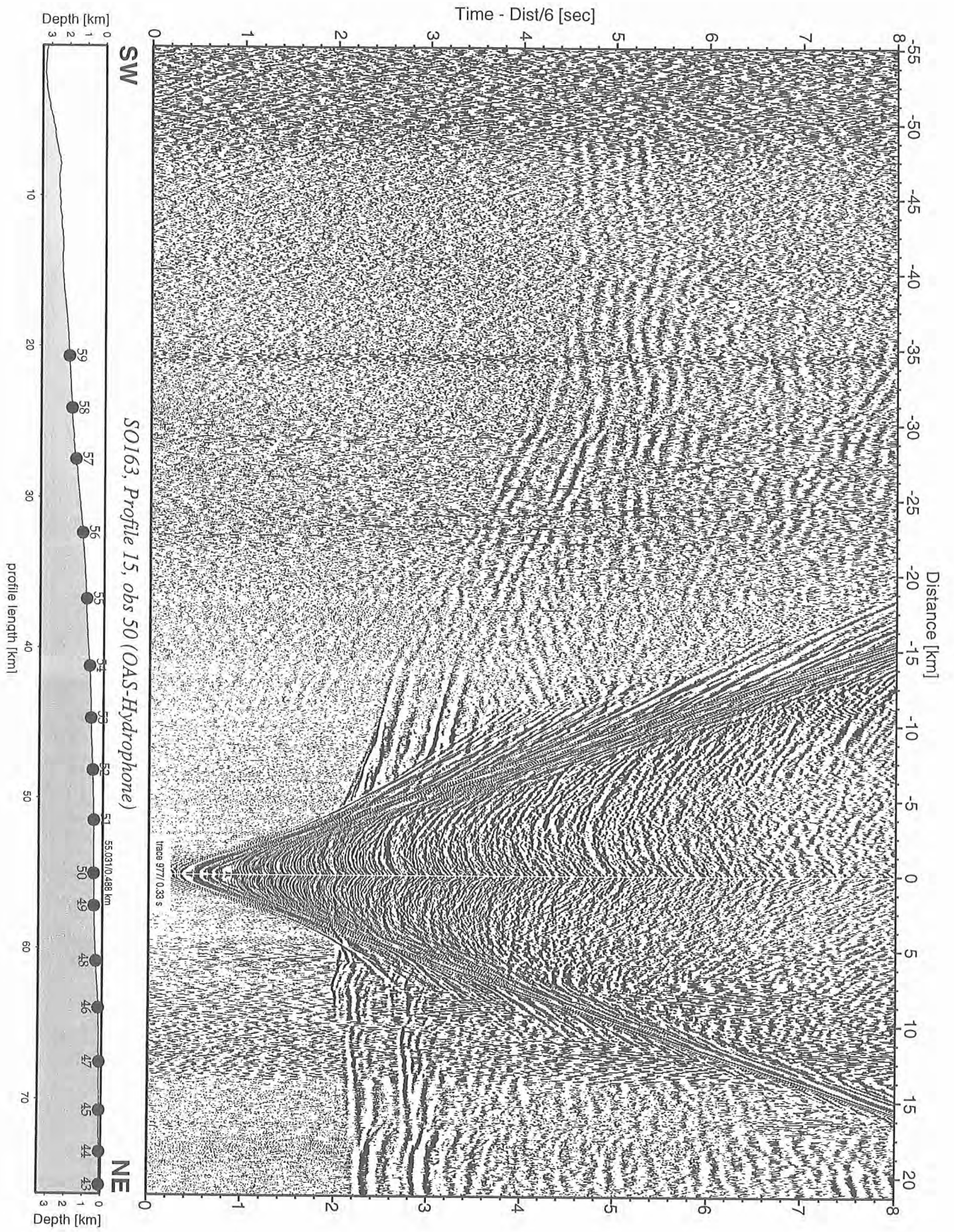


Figure 6.6.3.11: Record section from obs 50 (OAS-Hydrophone), Profile 15.

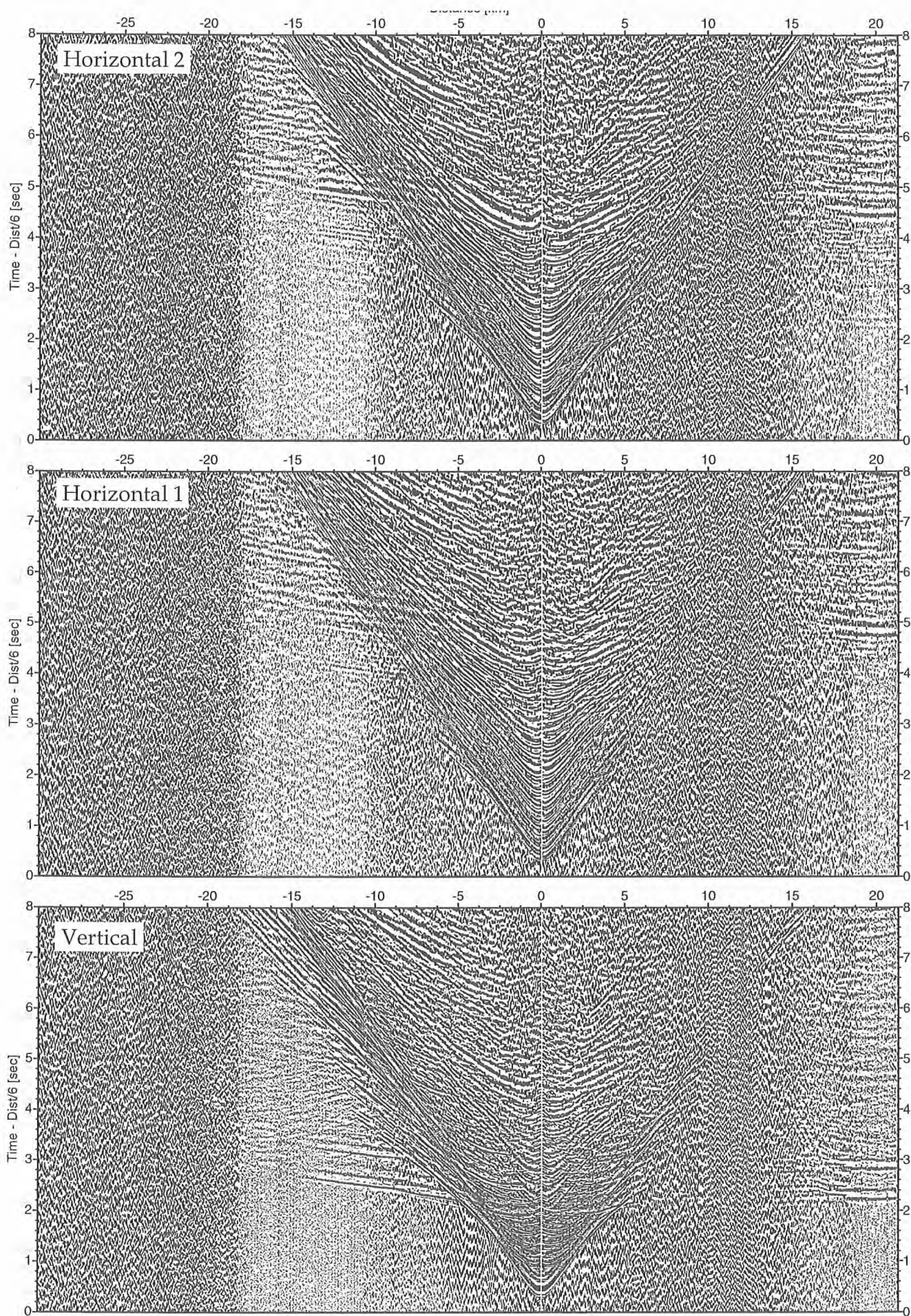


Figure 6.6.3.12: Record sections from obs 50 (Owen-4.5Hz), SO163, Profile 15.

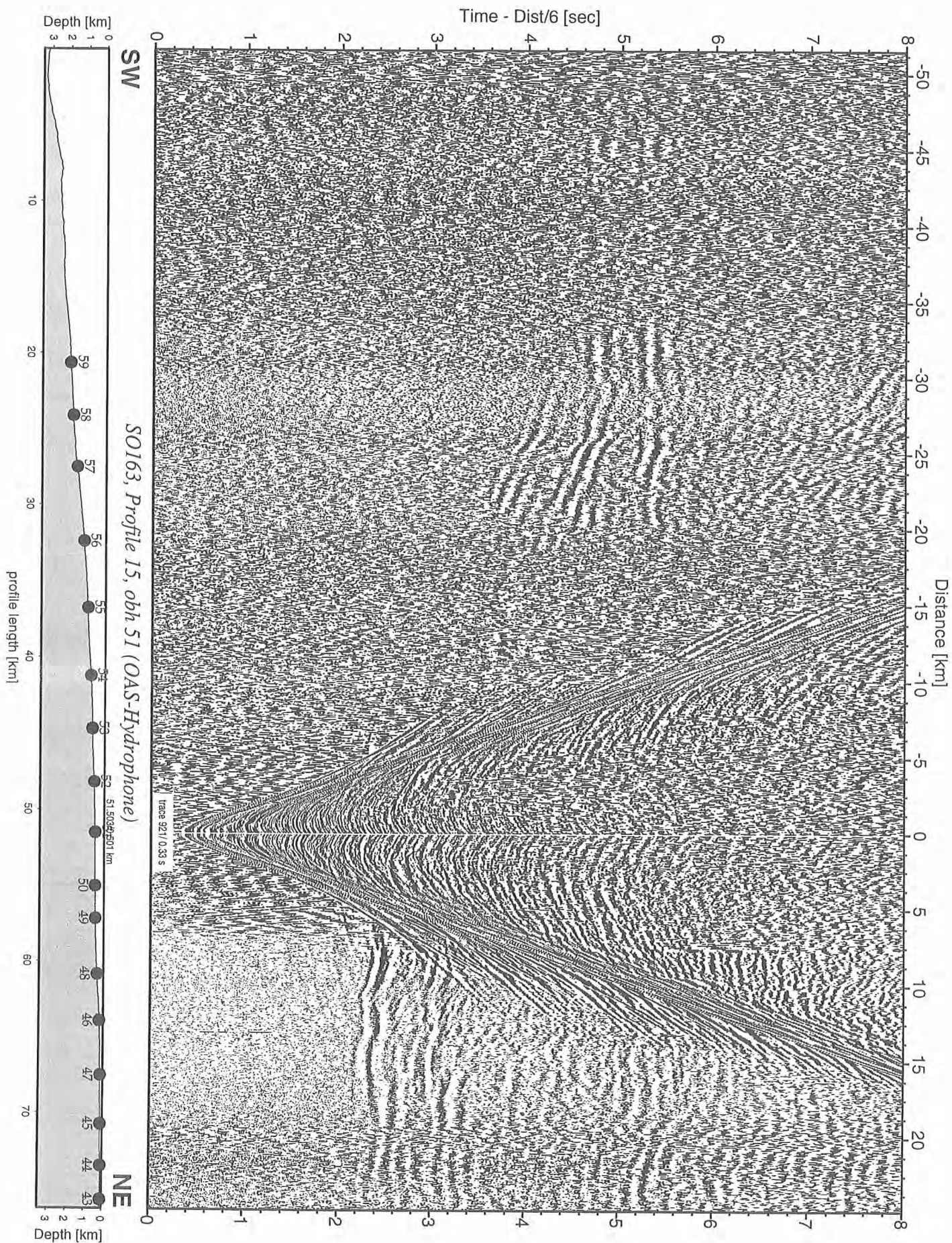


Figure 6.6.3.13: Record section from obh 51 (OAS-Hydrophone), Profile 15.

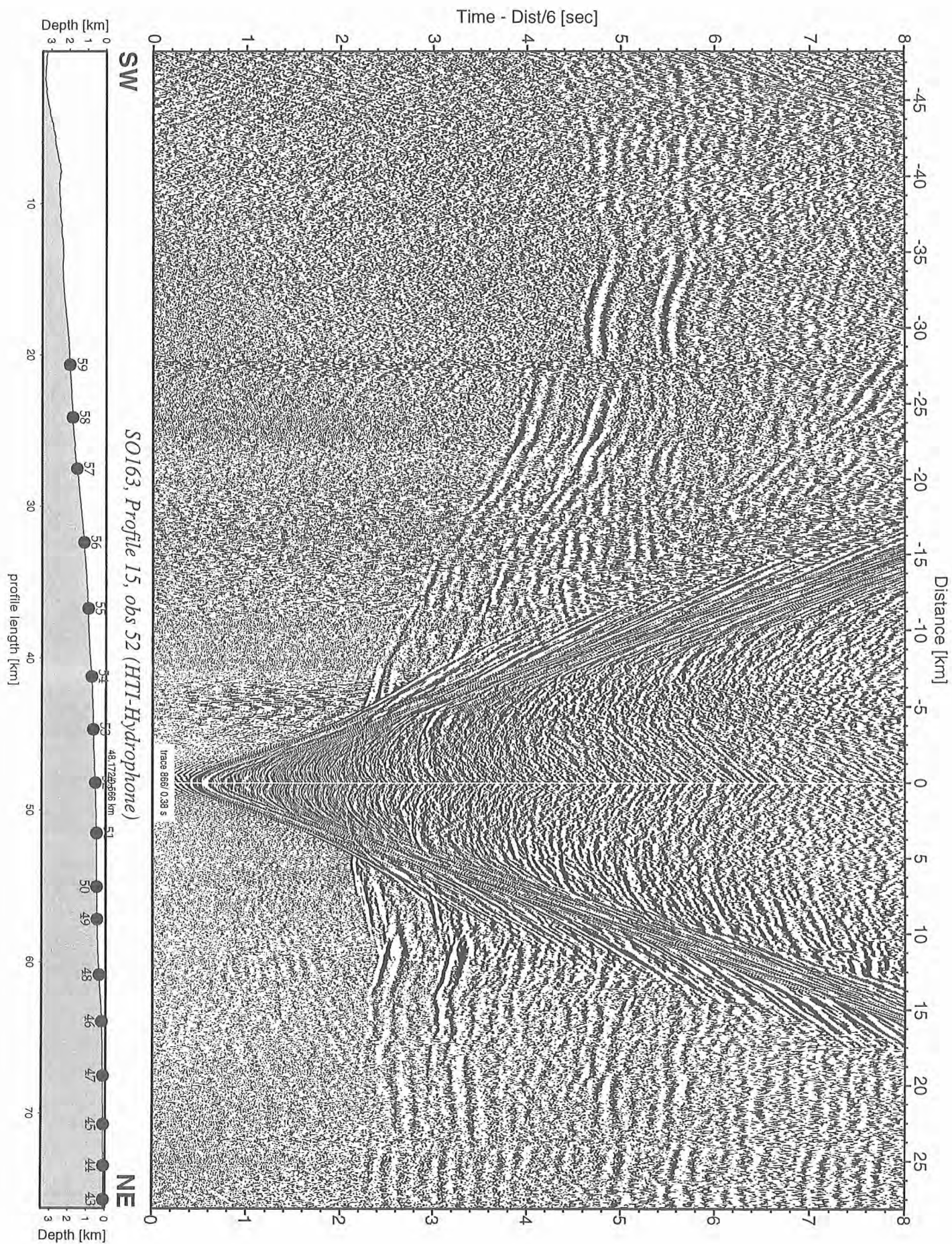


Figure 6.6.3.14: Record section from obs 52 (HTI-Hydrophone), Profile 15.

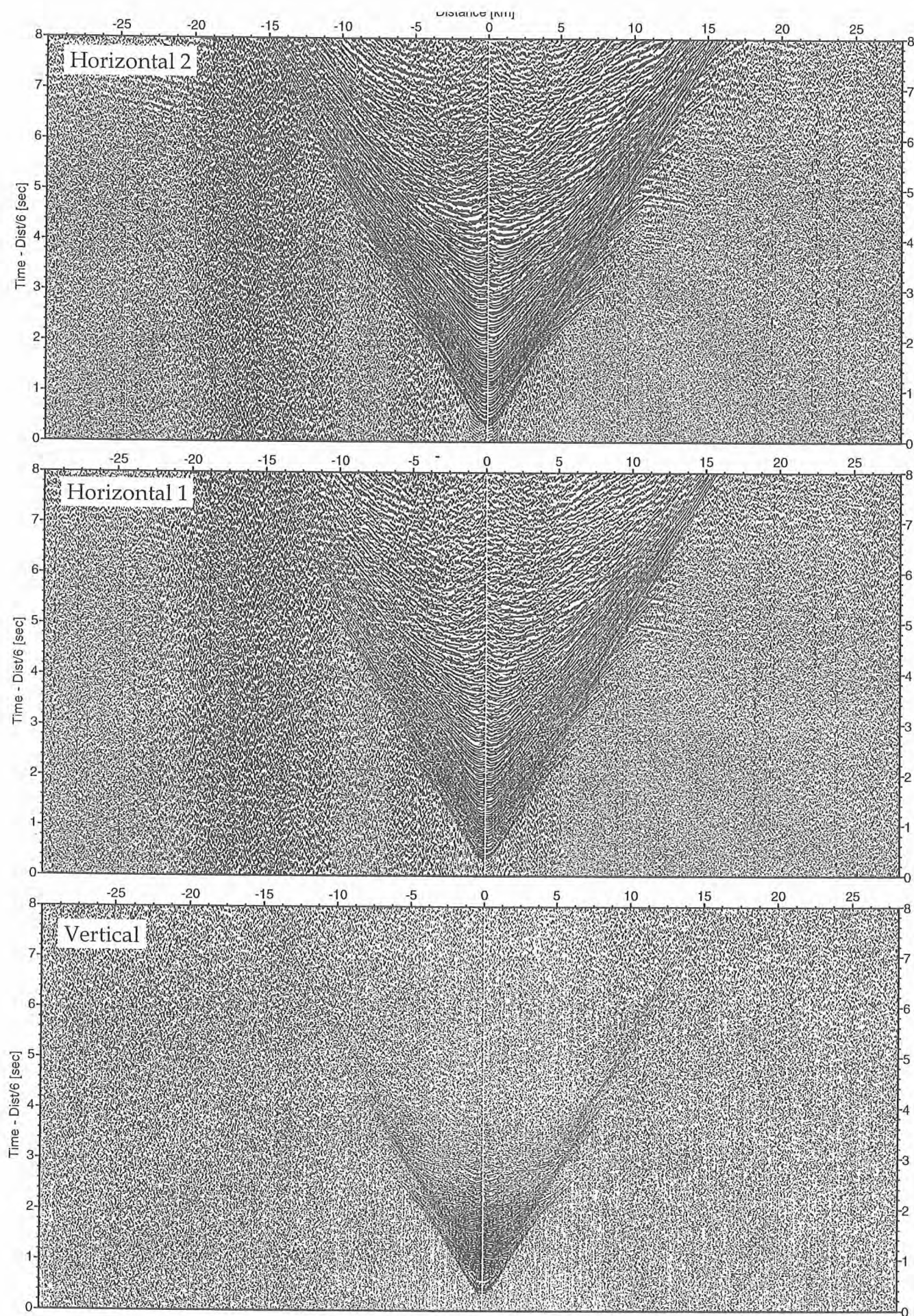


Figure 6.6.3.15: Record sections from obs 52 (Owen-30Hz), SO163, Profile 15.

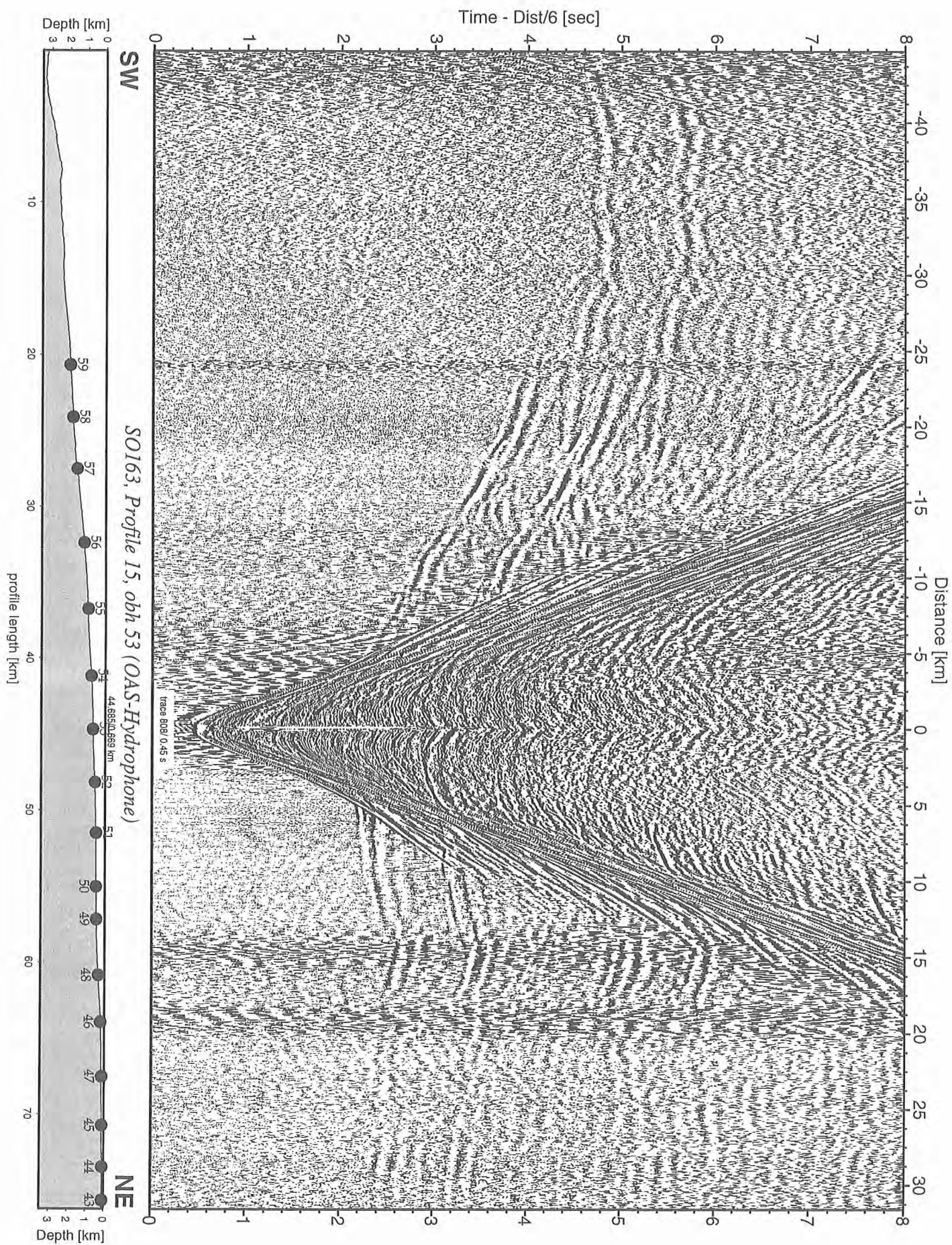


Figure 6.6.3.16: Record section from obh 53 (OAS-Hydrophone), Profile 15.

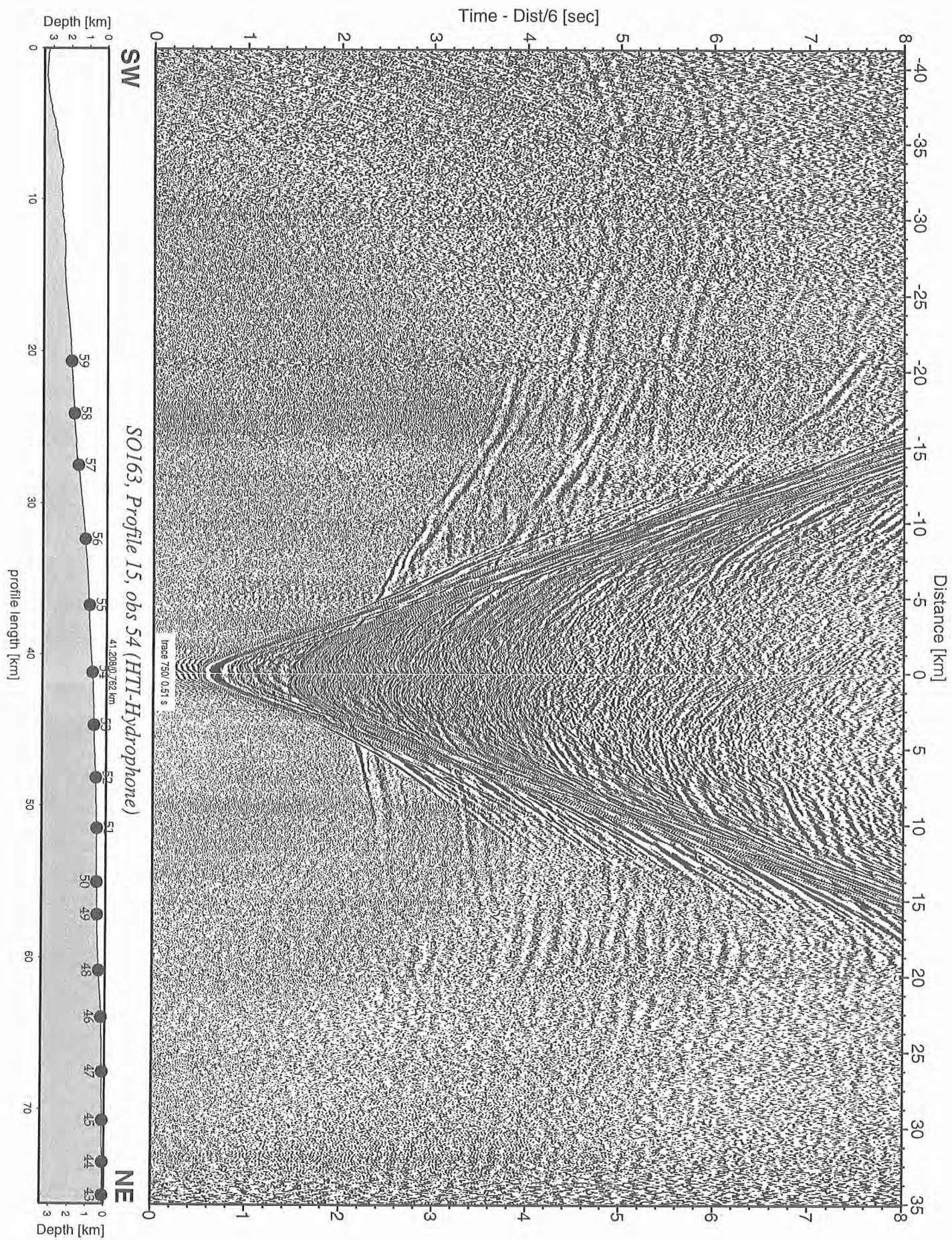


Figure 6.6.3.17: Record section from obs 54 (HTI-Hydrophone), Profile 15.

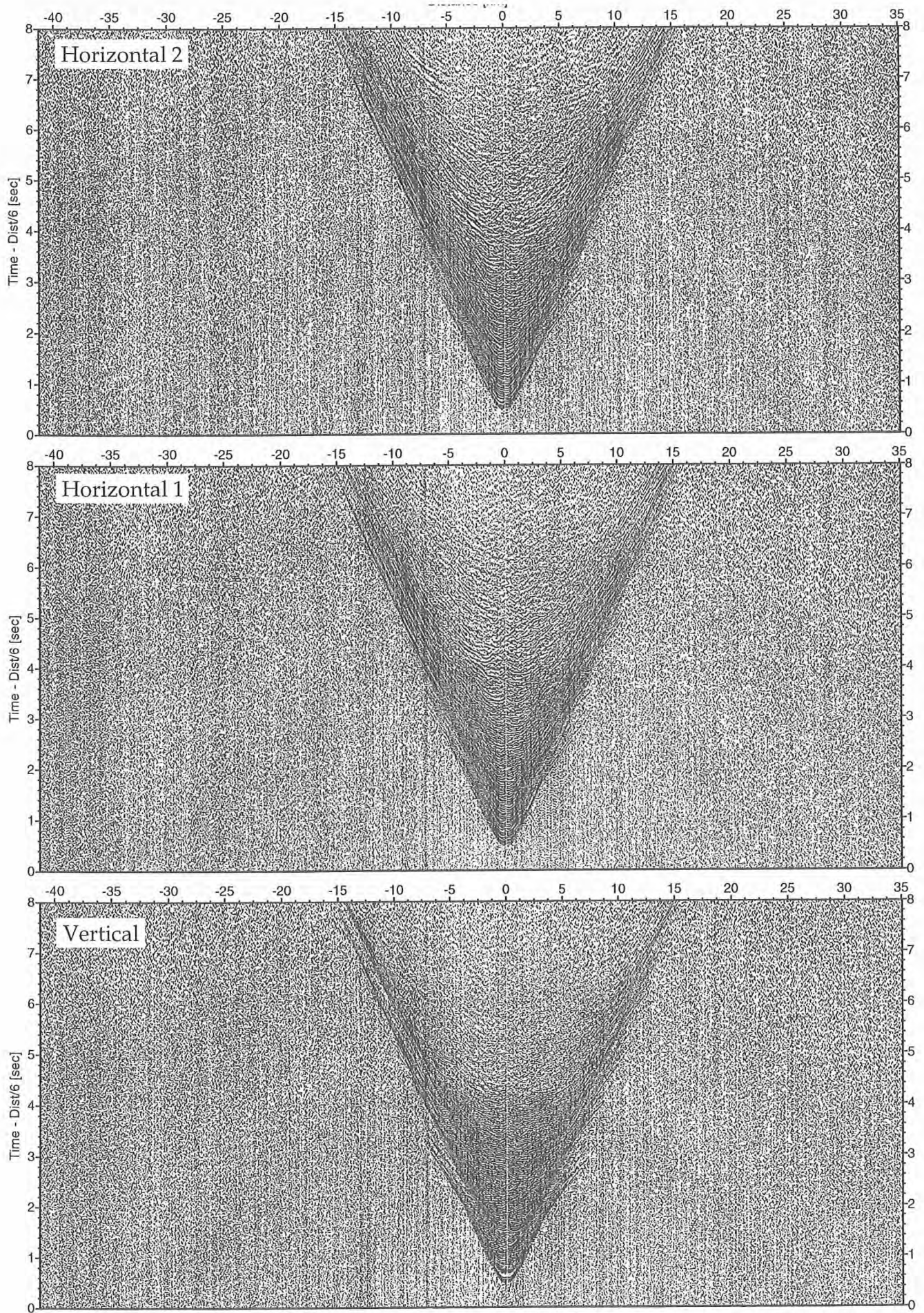


Figure 6.6.3.18: Record sections from obs 54 (Owen-30Hz), SO163, Profile 15.

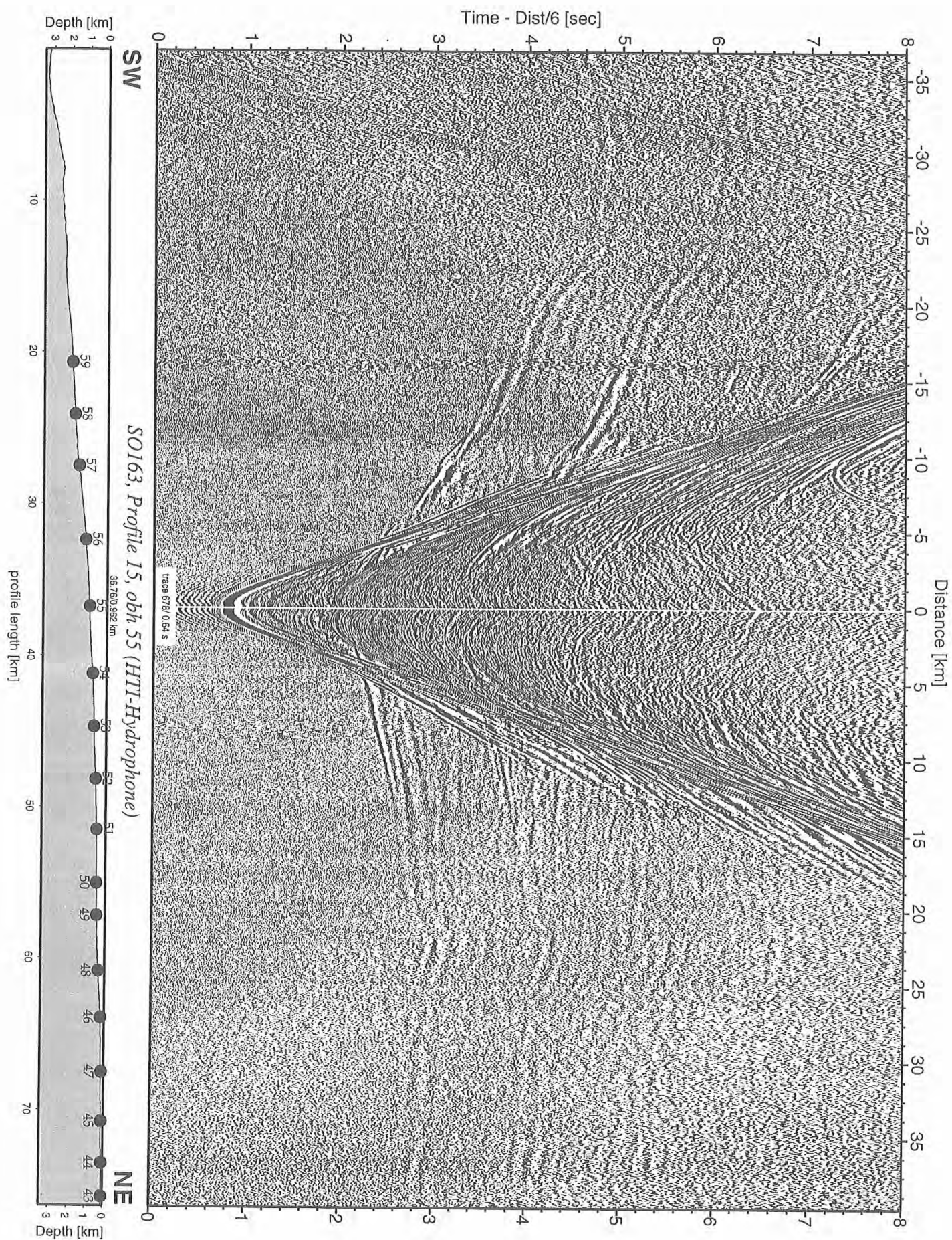


Figure 6.6.3.19: Record section from obh 55 (HTI-Hydrophone), Profile 15.

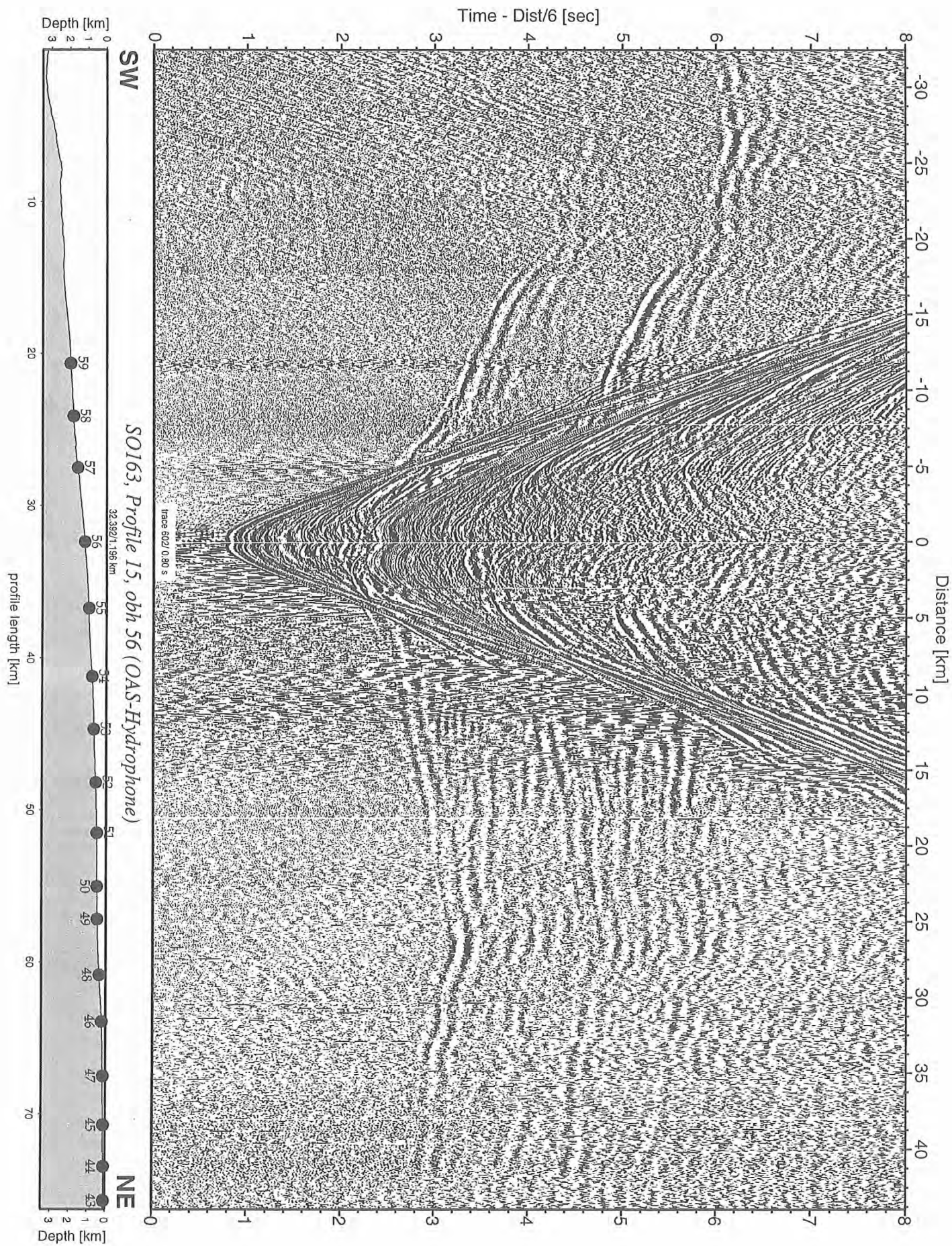


Figure 6.6.3.20: Record section from obh 56 (OAS-Hydrophone), Profile 15.

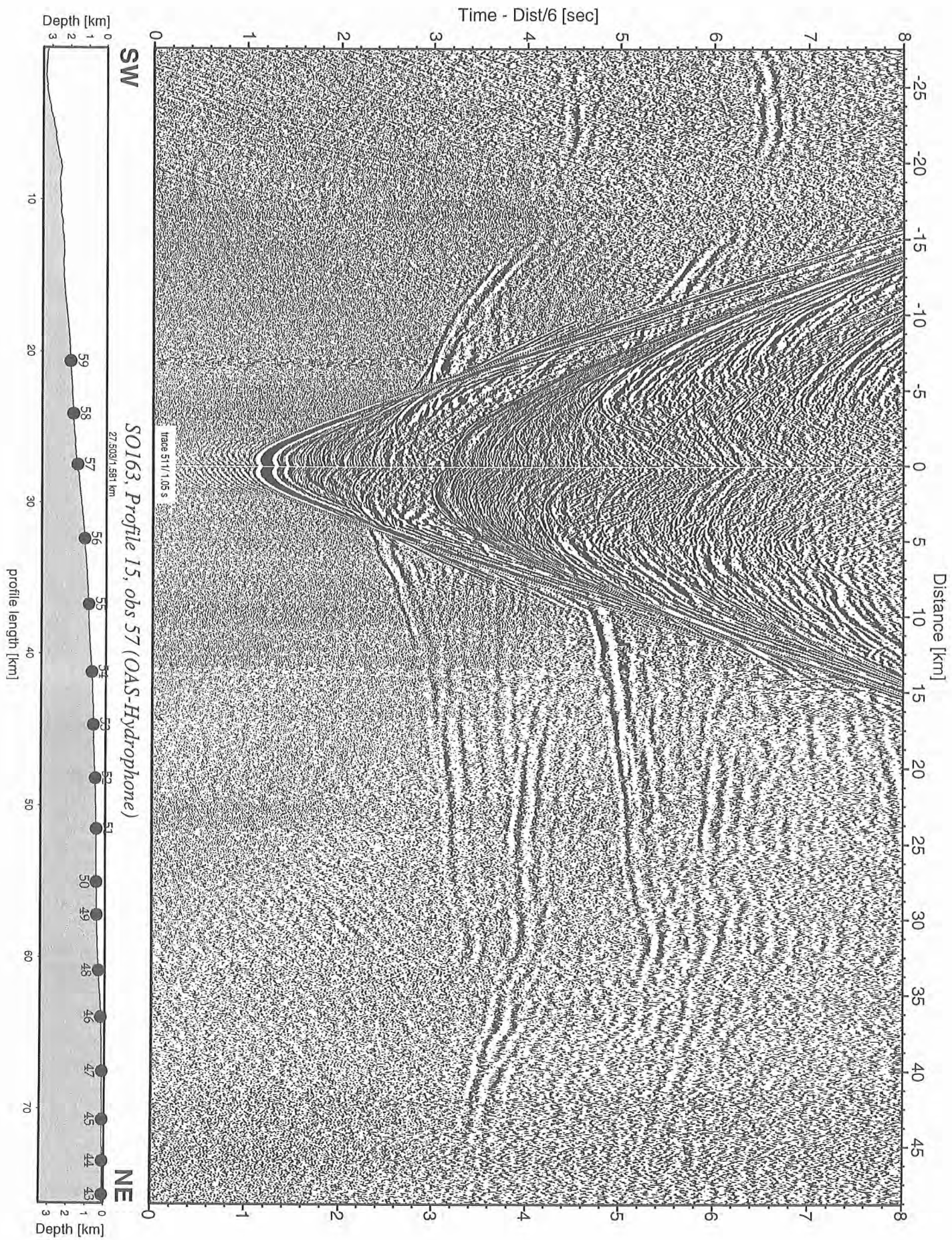


Figure 6.6.3.21: Record section from obs 57 (OAS-Hydrophone), Profile 15.

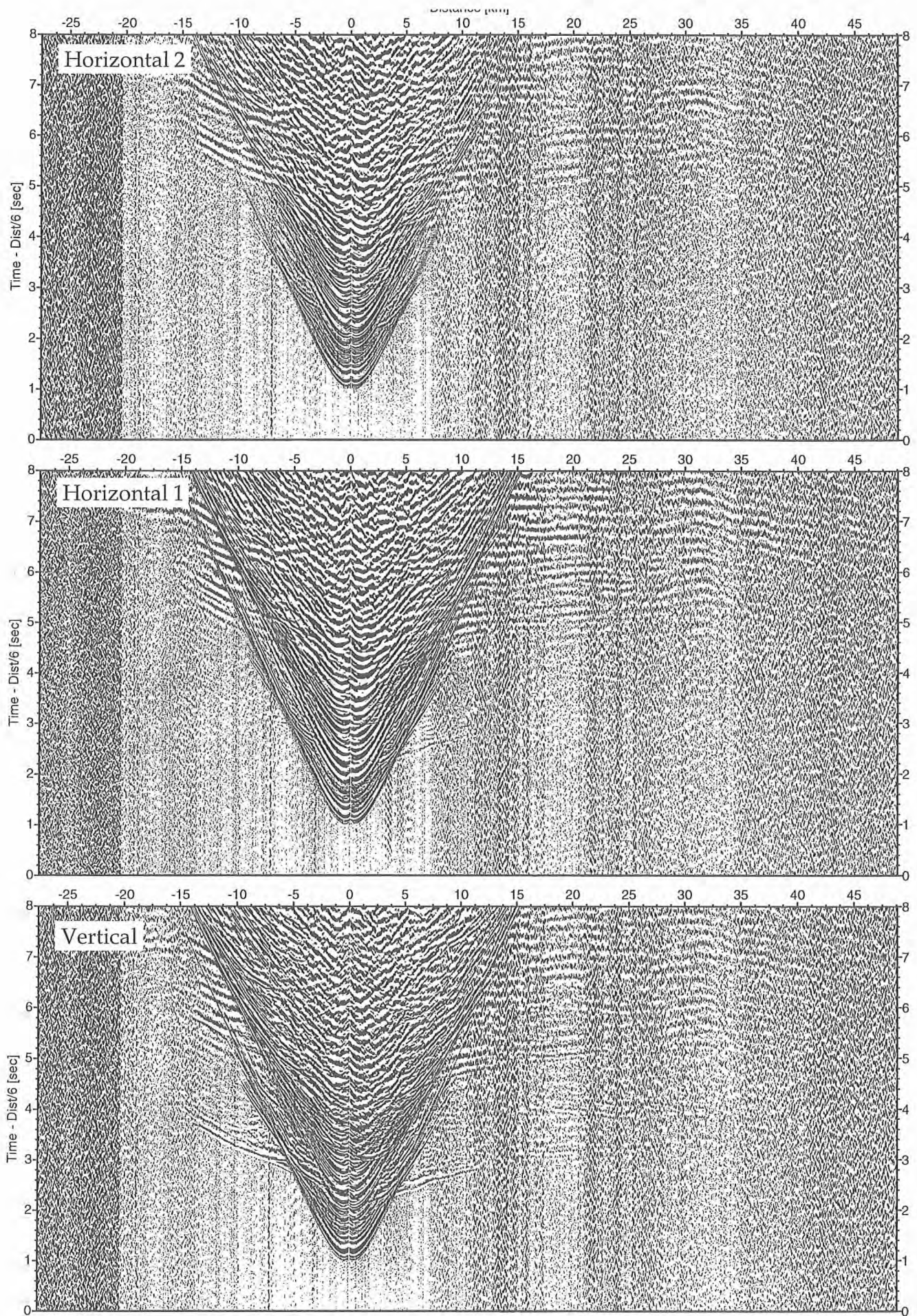


Figure 6.6.3.22: Record sections from obs 57 (LG-4.5HZ), SO163, Profile 15.

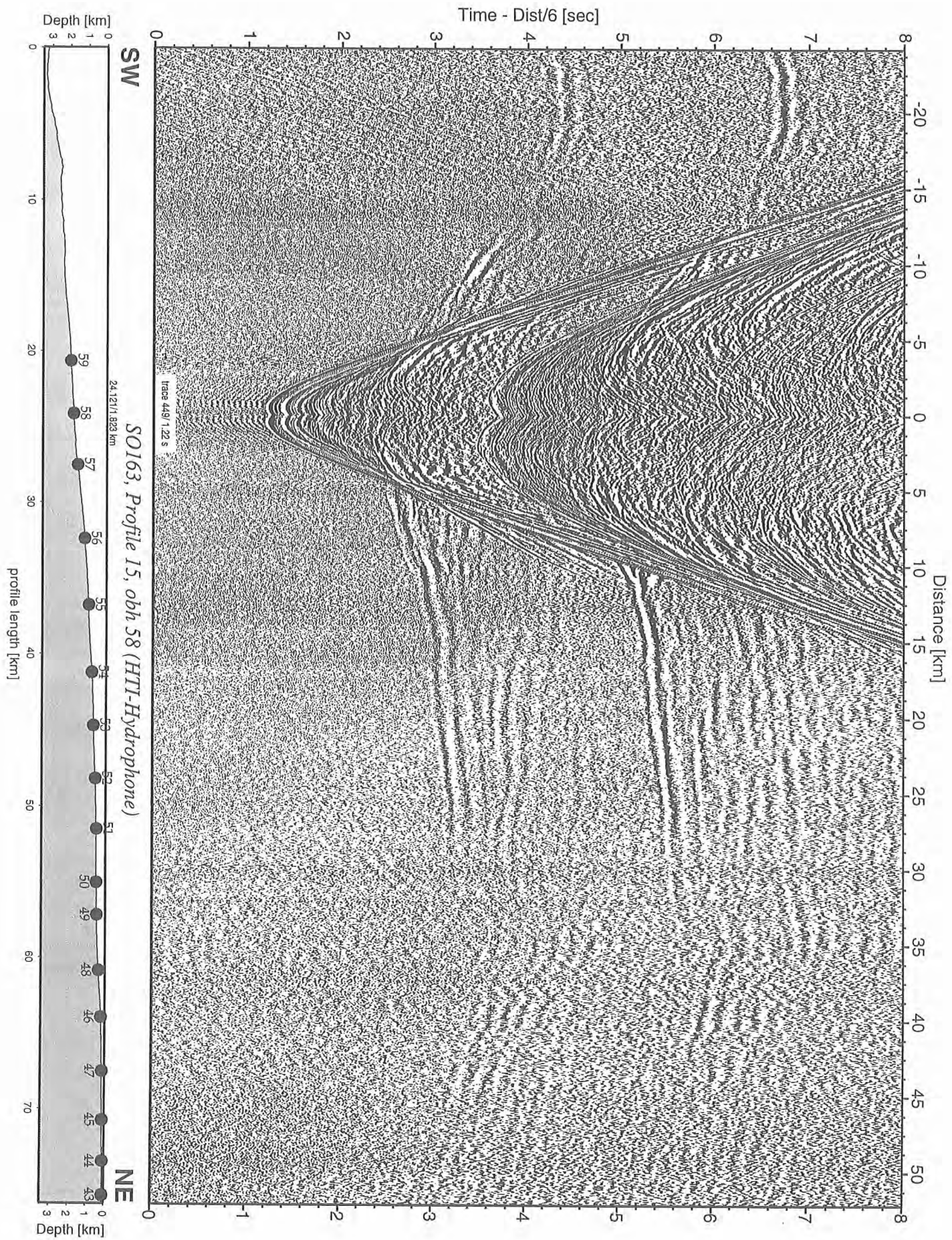


Figure 6.6.3.23: Record section from obh 58 (HTI-Hydrophone), Profile 15.

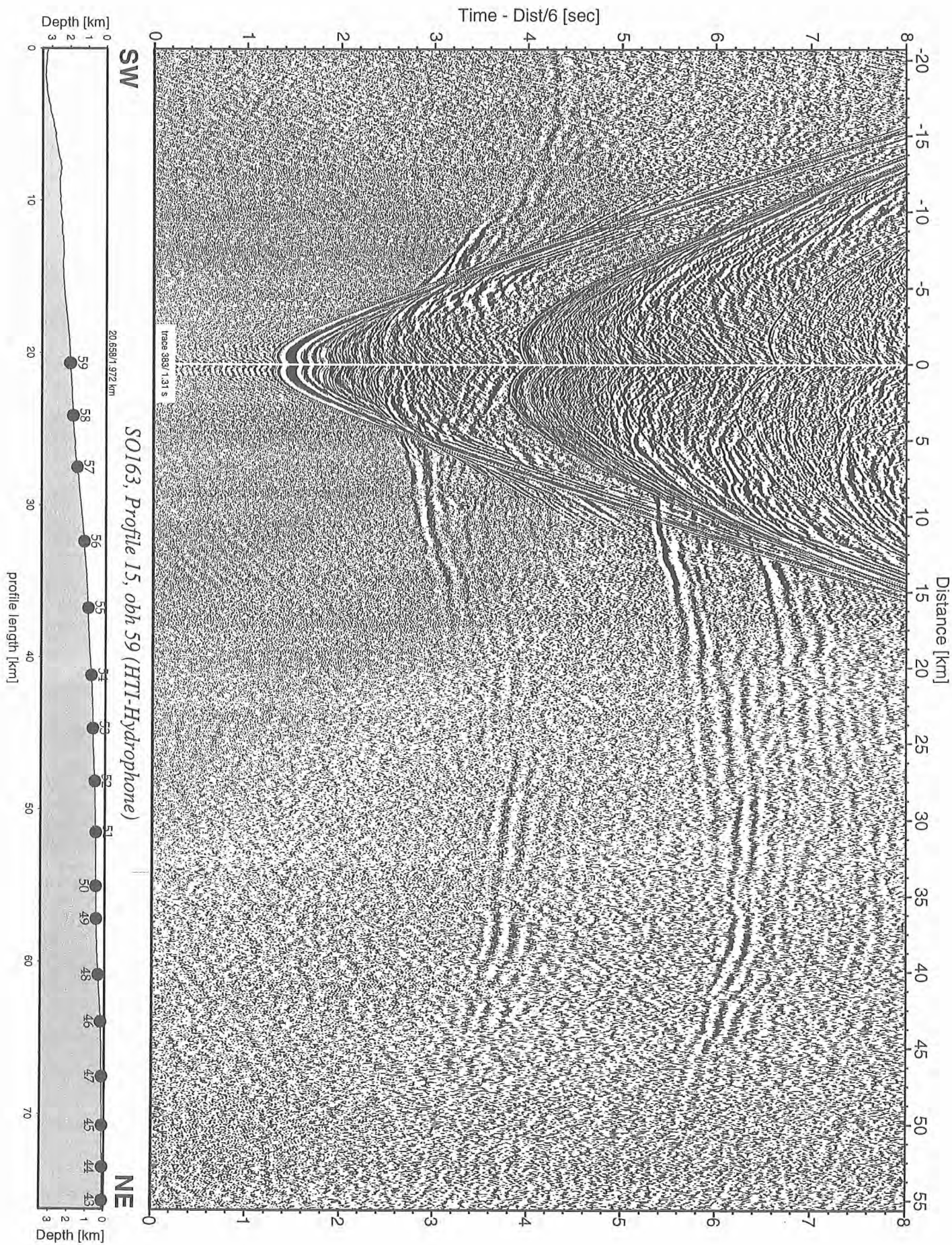


Figure 6.6.3.24: Record section from obh 59 (HTI-Hydrophone), Profile 15.

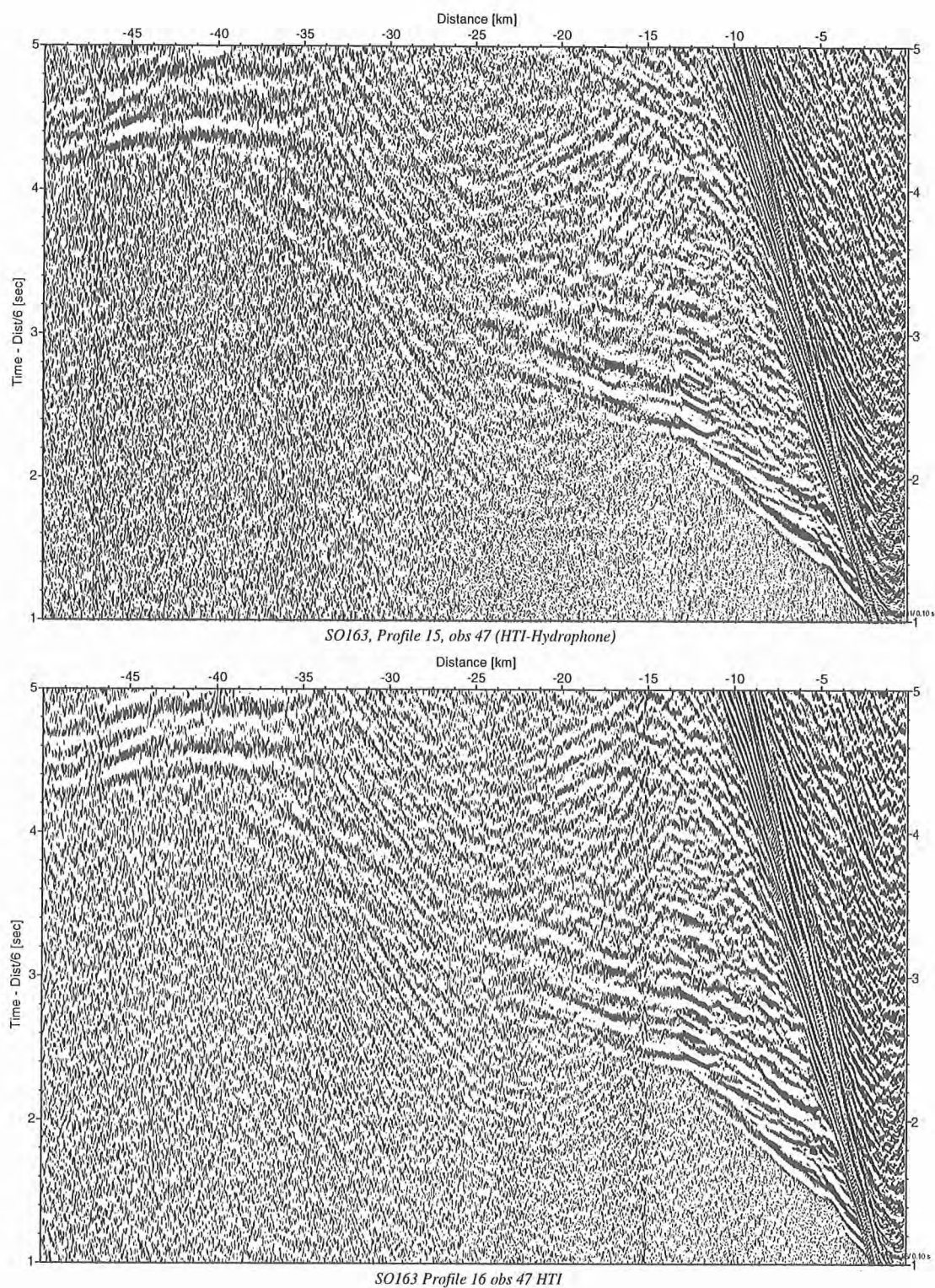


Figure 6.6.3.25.: Comparison between profiles P15 (G-gun) and P16 (Bolt gun) for station OBS 47.

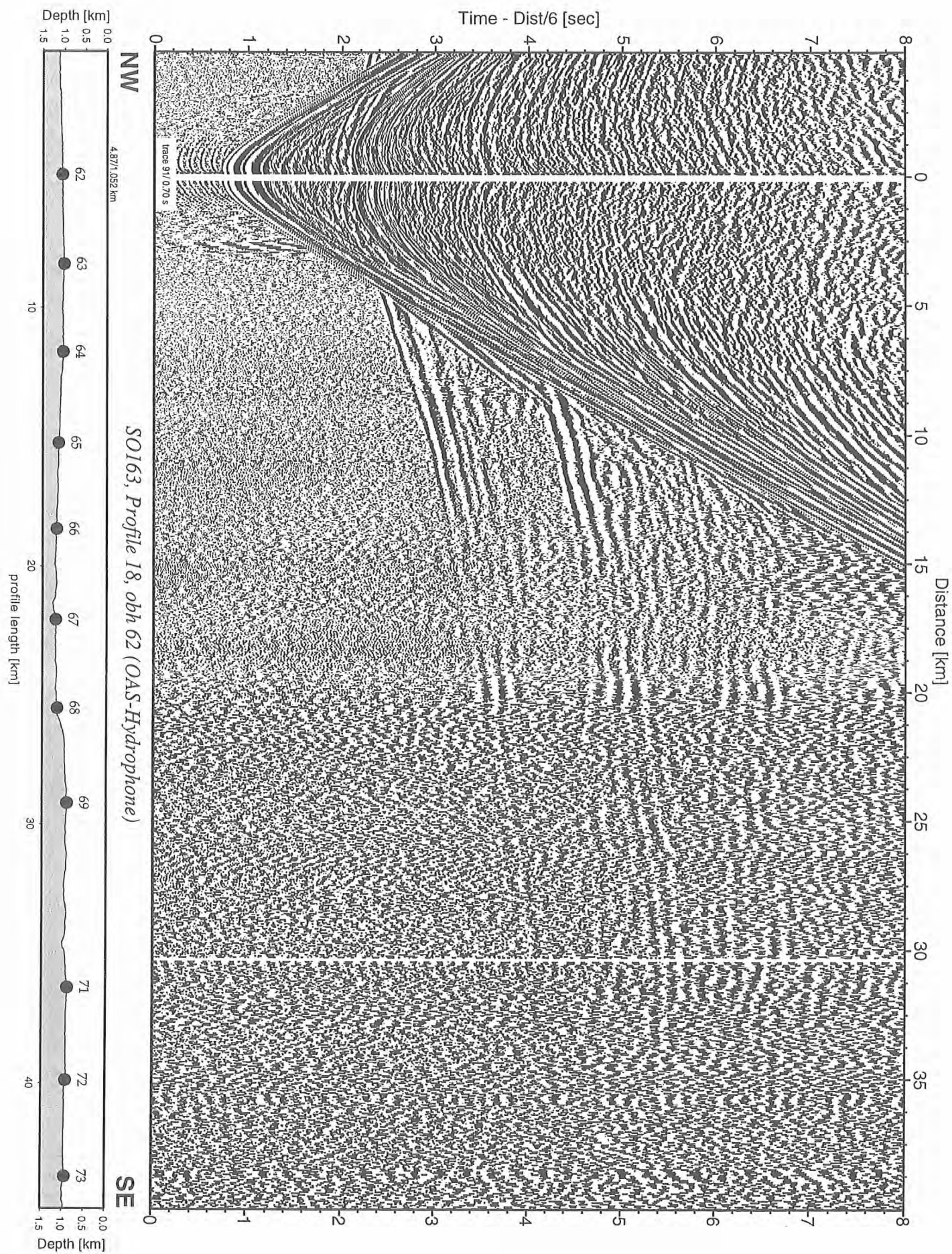


Figure 6.6.3.26: Record section from obh 62 (OAS-Hydrophone), Profile 18.

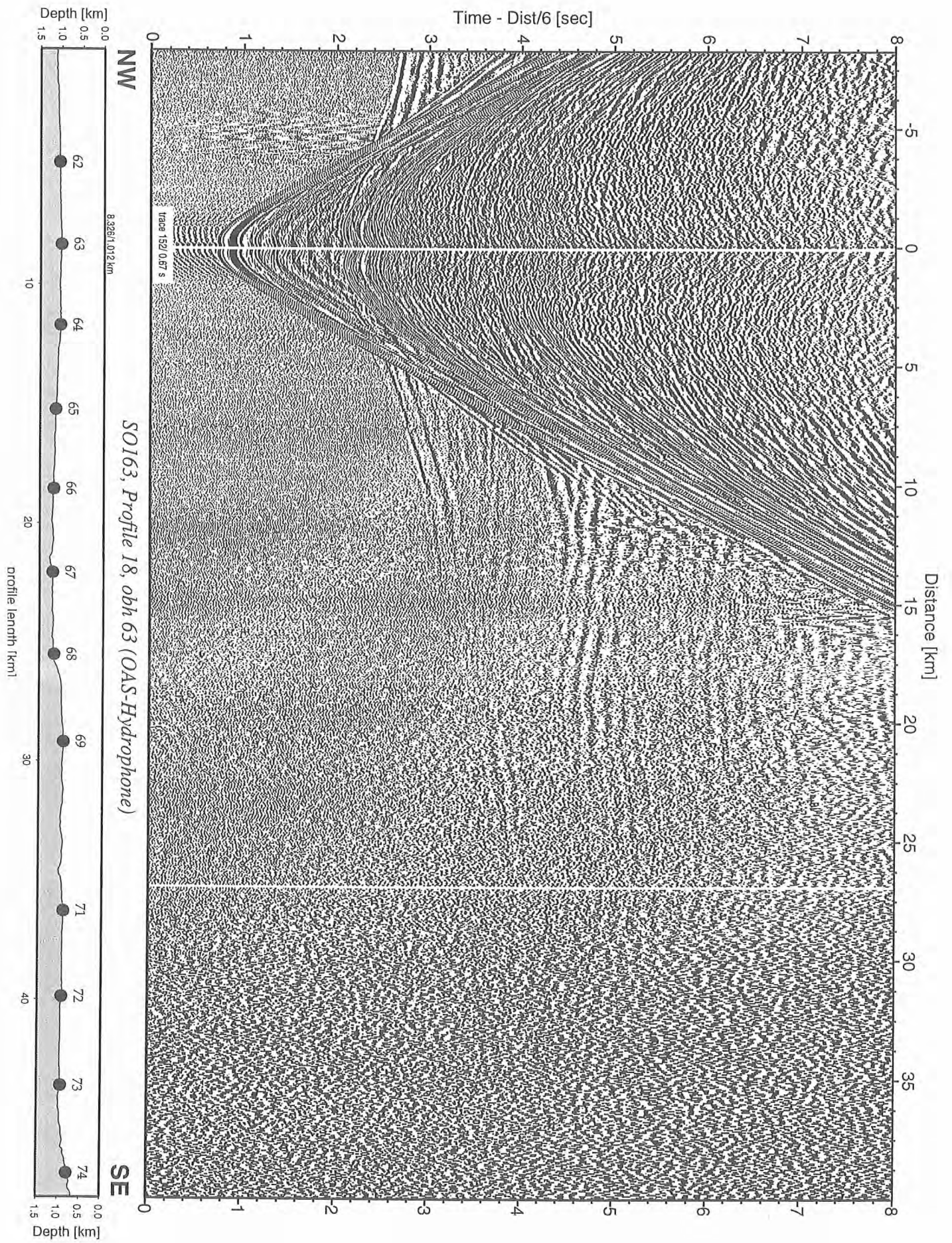


Figure 6.6.3.27: Record section from obh 63 (OAS-Hydrophone), Profile 18.

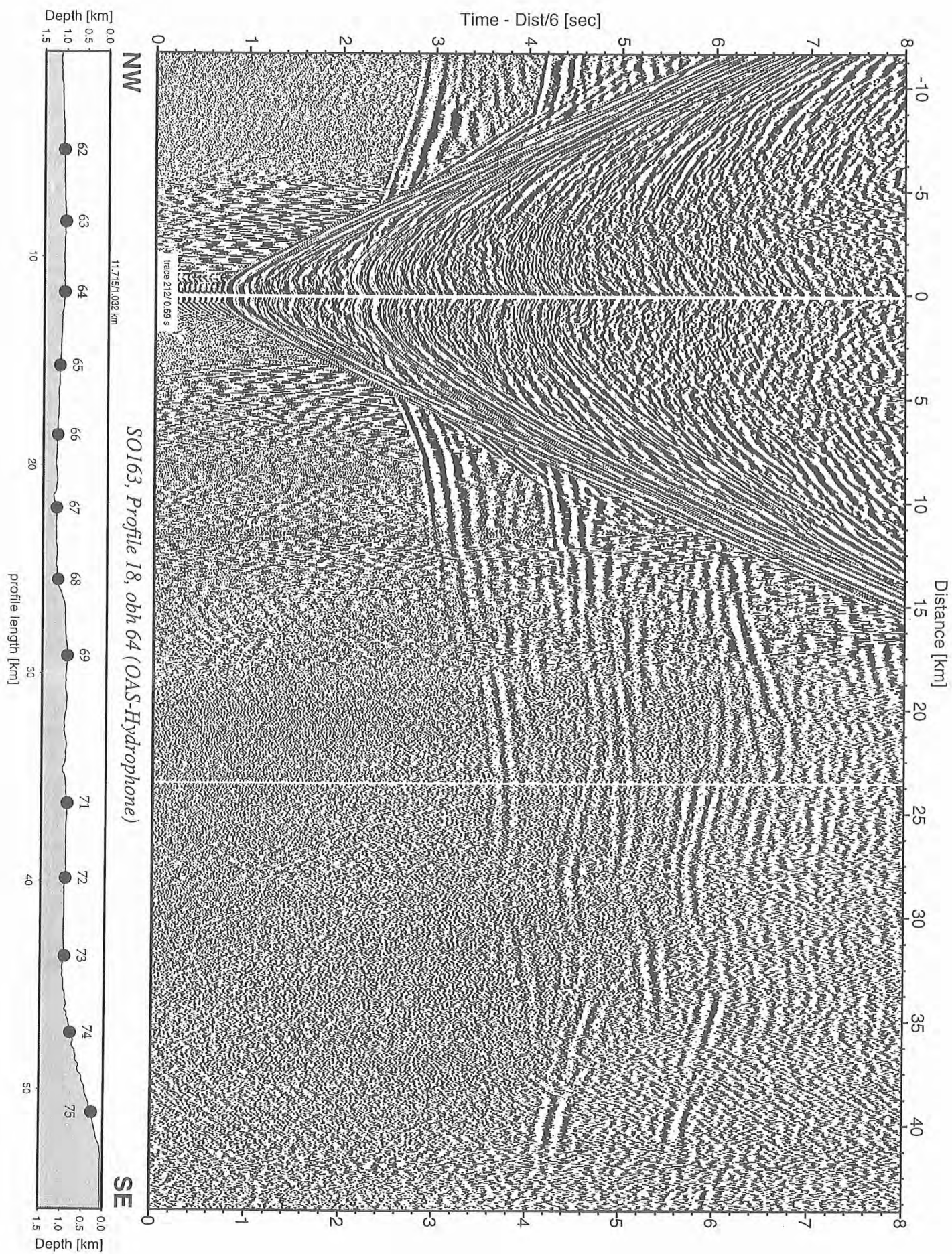


Figure 6.6.3.28: Record section from obh 64 (OAS-Hydrophone), Profile 18.

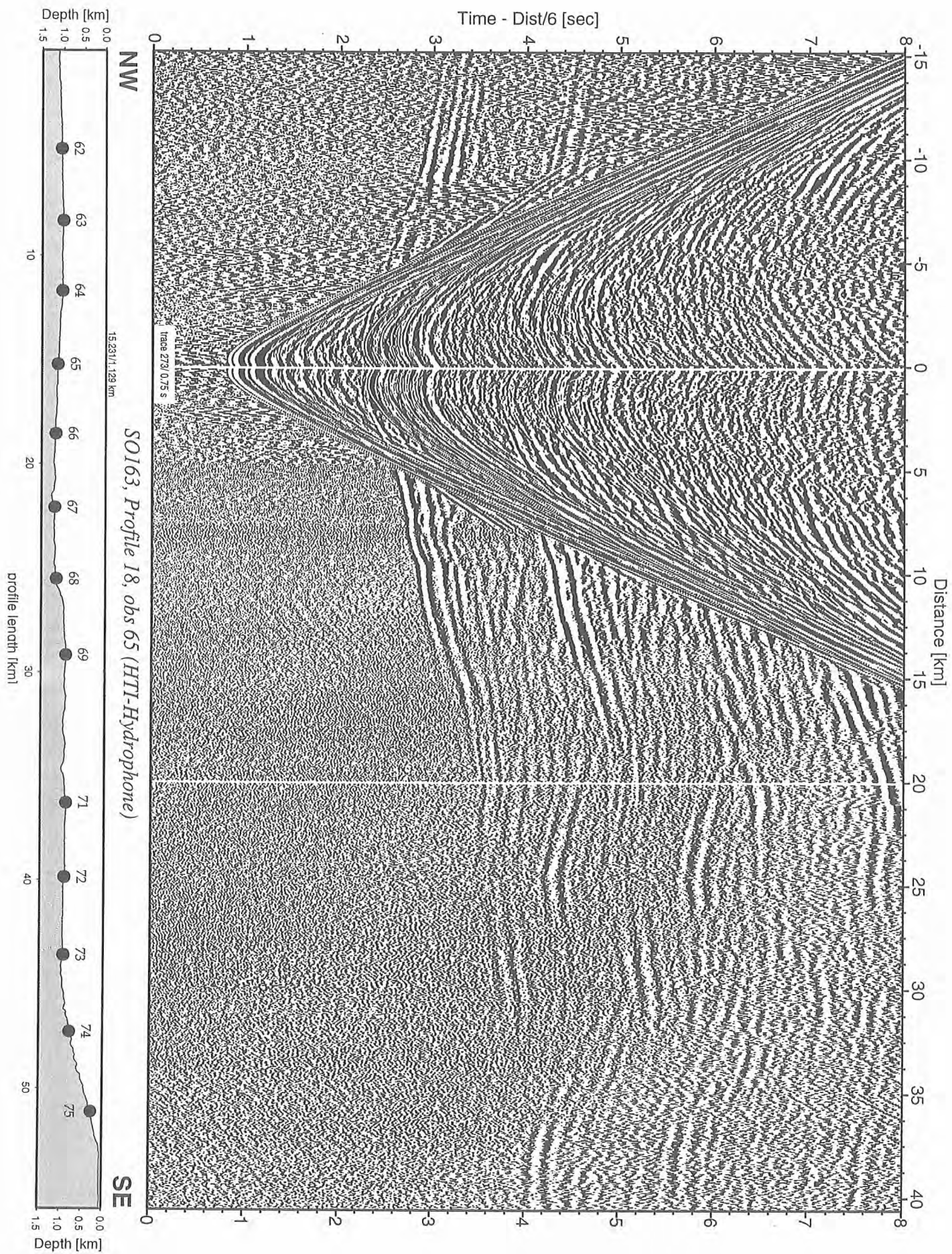


Figure 6.6.3.29: Record section from obs 65 (HTI-Hydrophone), Profile 18.

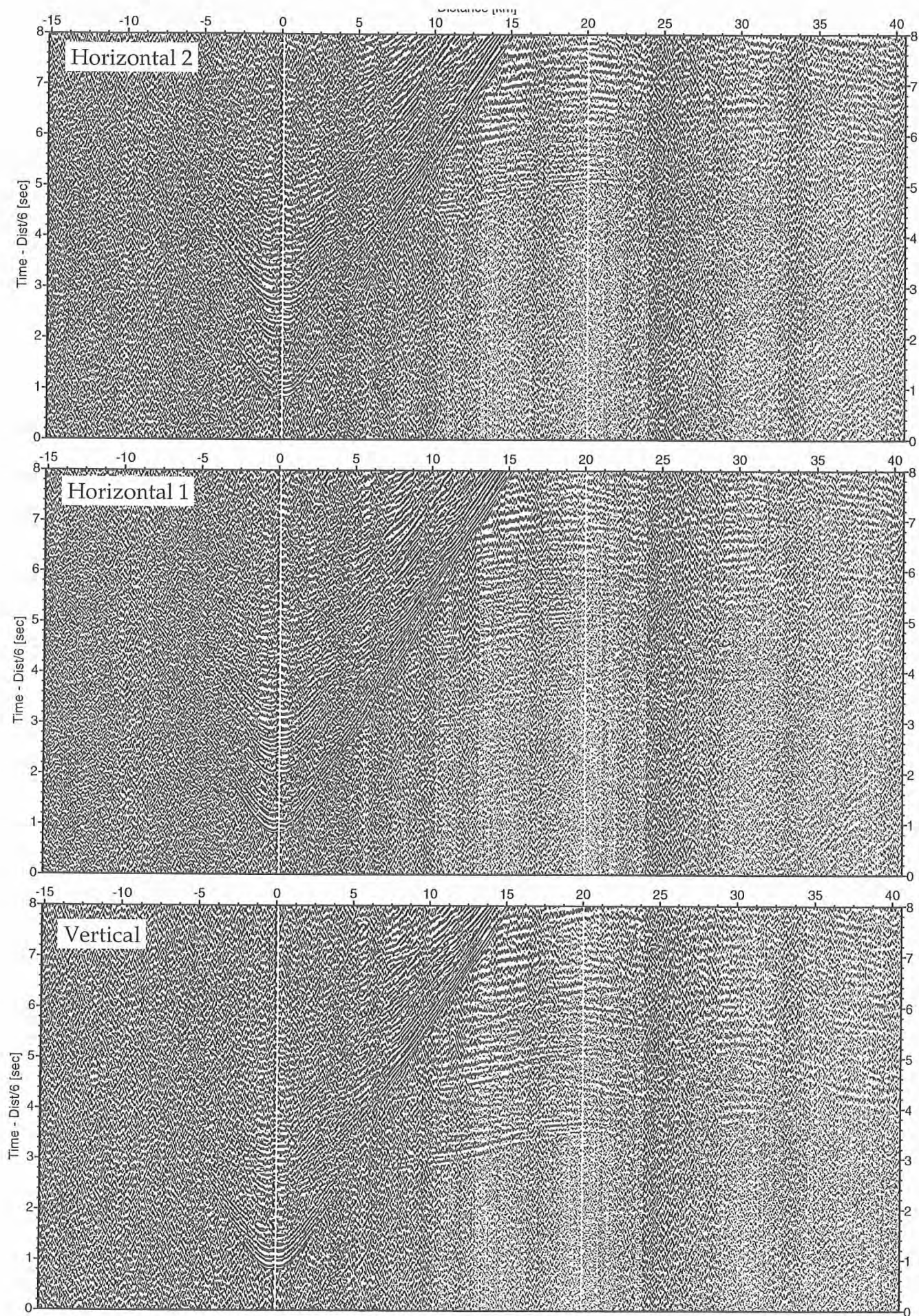


Figure 6.6.3.30: Record sections from obs 65 (Owen-4.5Hz), SO163, Profile 18.

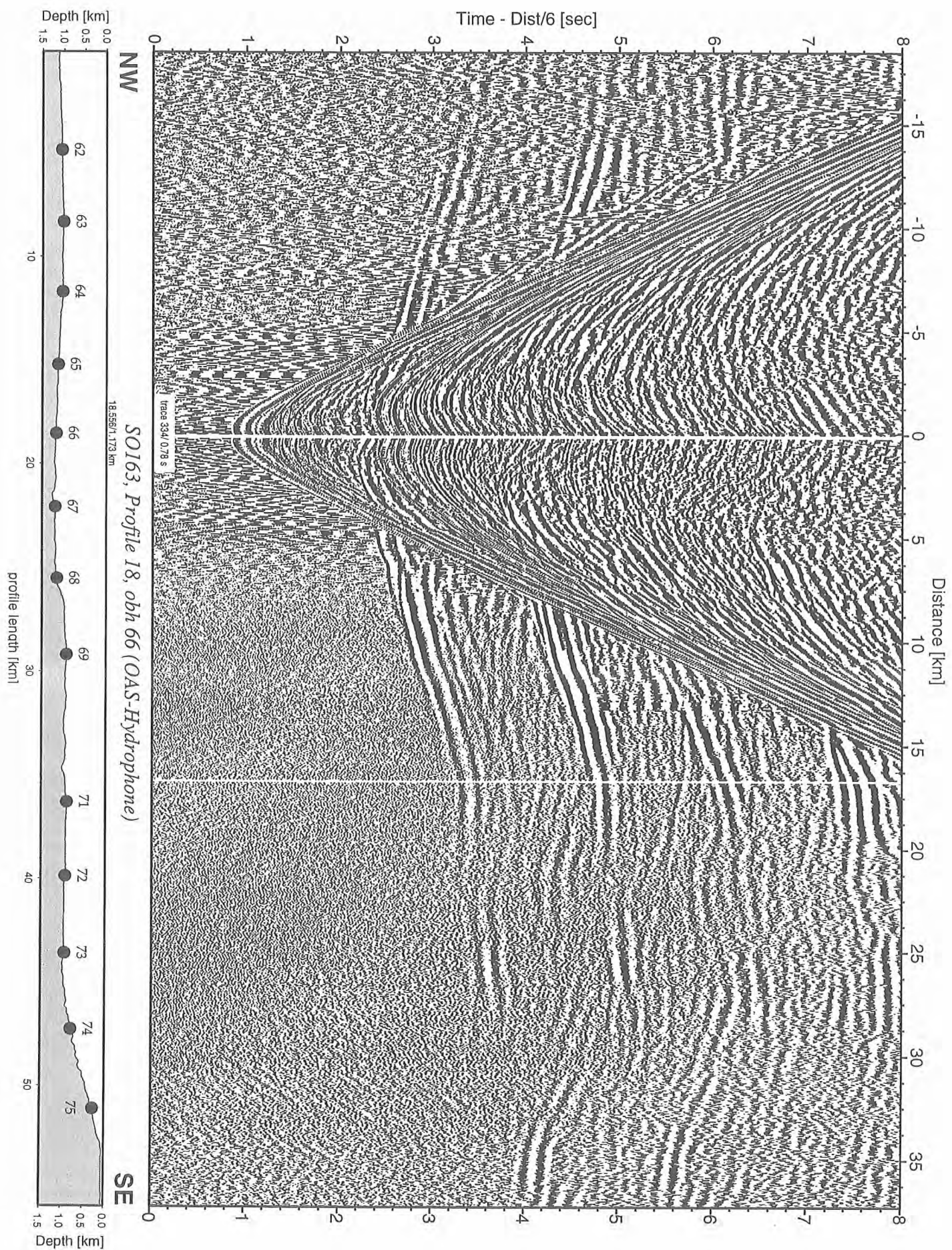


Figure 6.6.3.31: Record section from obh 66 (OAS-Hydrophone), Profile 18.

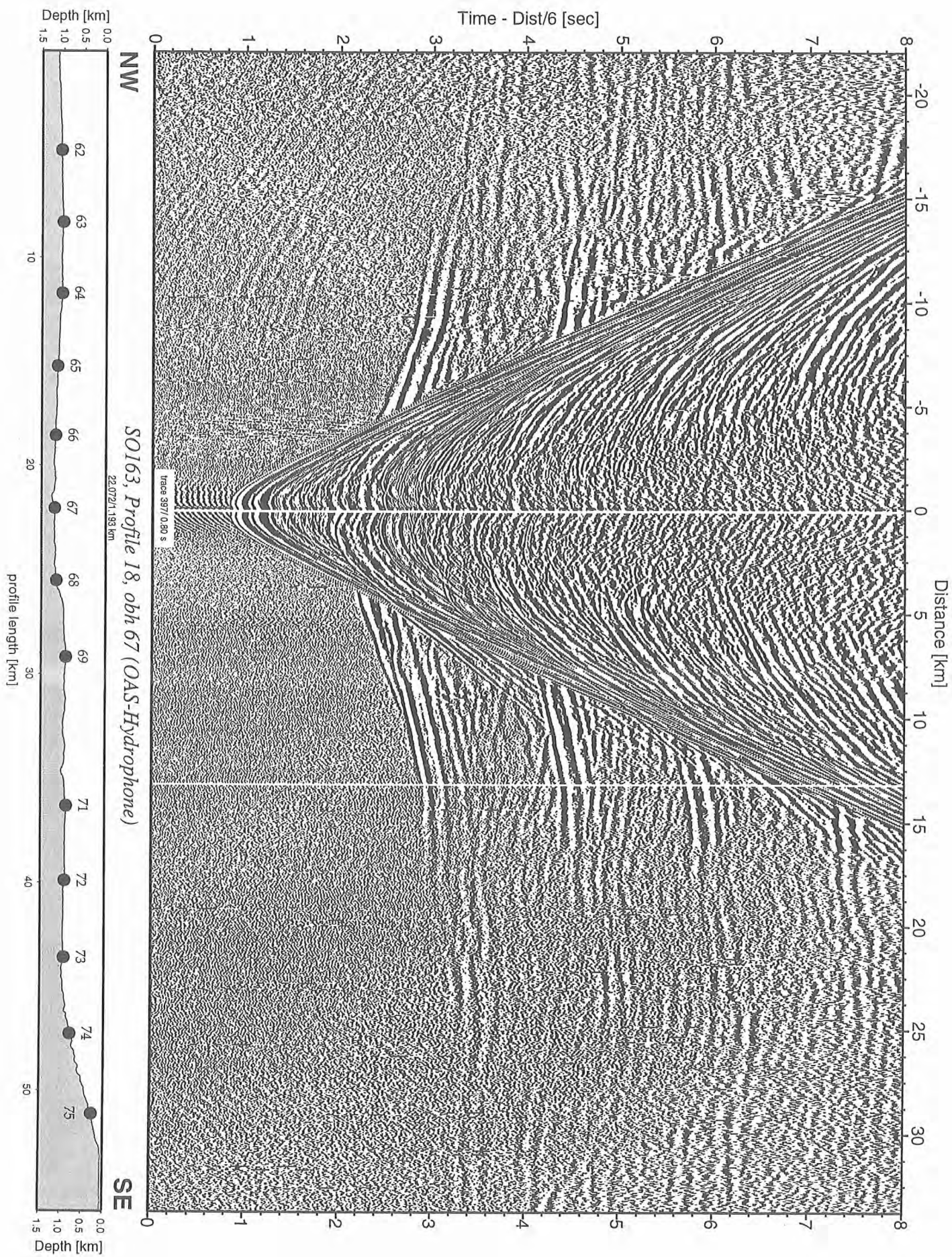


Figure 6.6.3.32: Record section from obh 67 (OAS-Hydrophone), Profile 18.

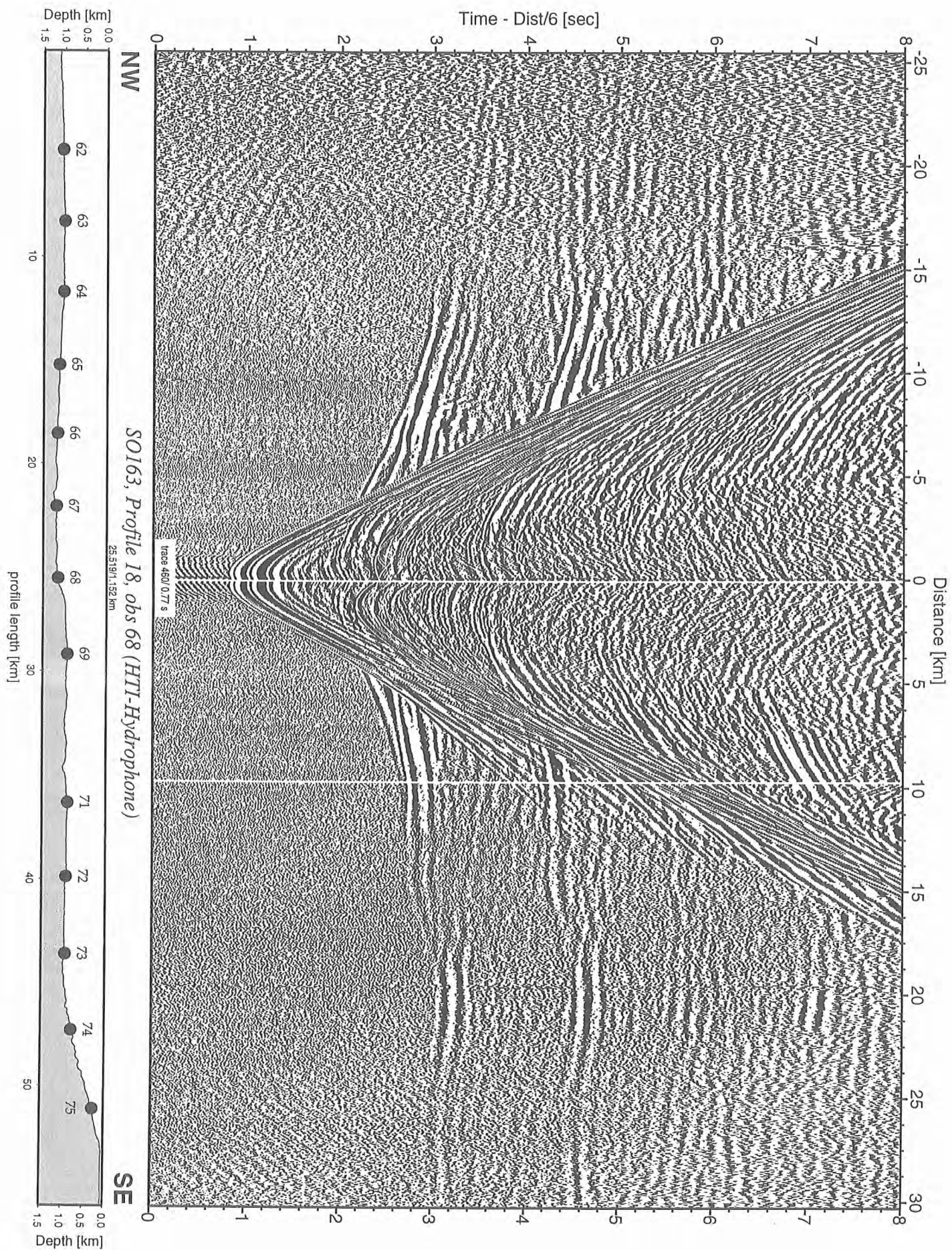


Figure 6.6.3.33: Record section from obs 68 (HTI-Hydrophone), Profile 18.

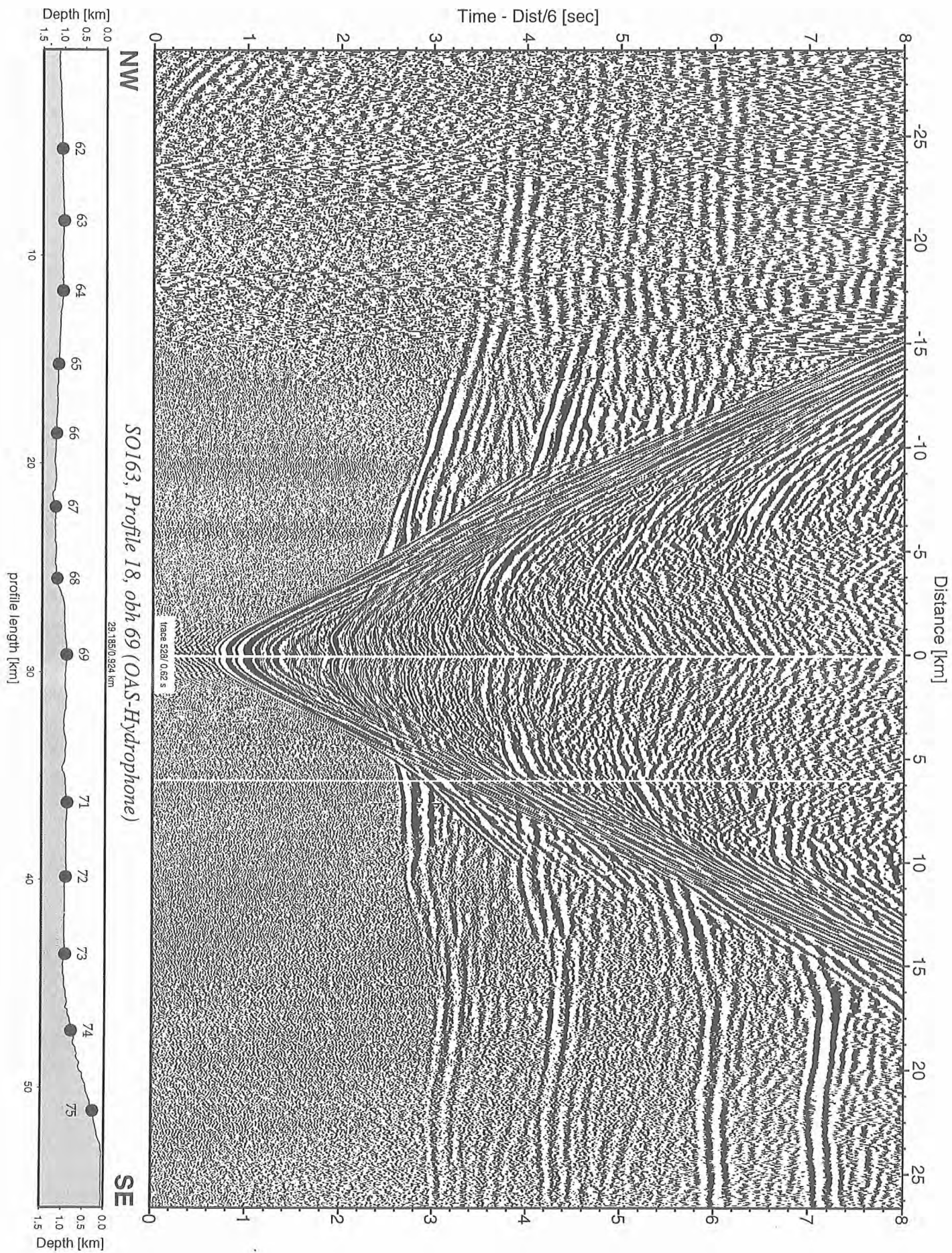


Figure 6.6.3.34: Record section from obh 69 (OAS-Hydrophone), Profile 18.

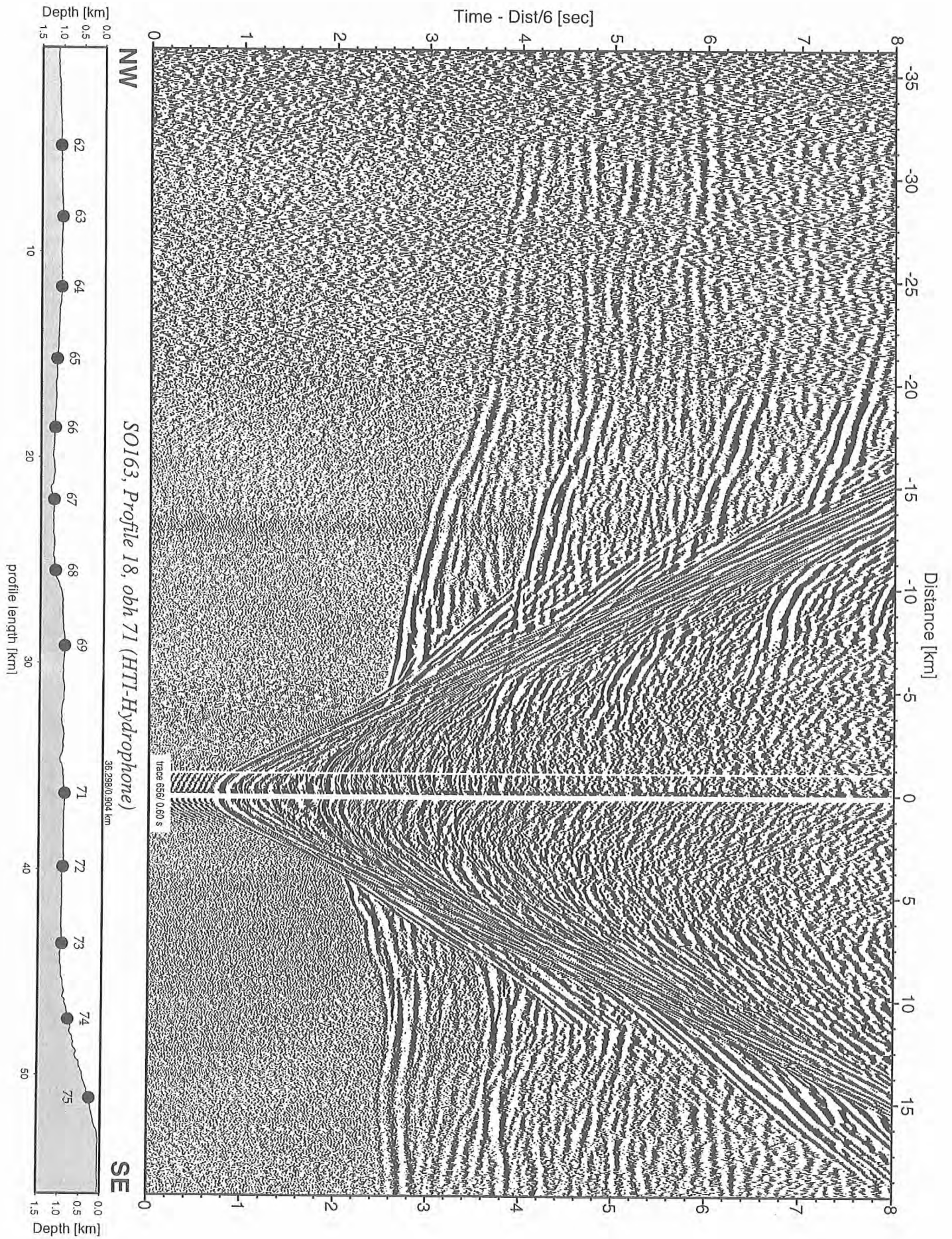


Figure 6.6.3.35: Record section from obh 71 (HTI-Hydrophone), Profile 18.

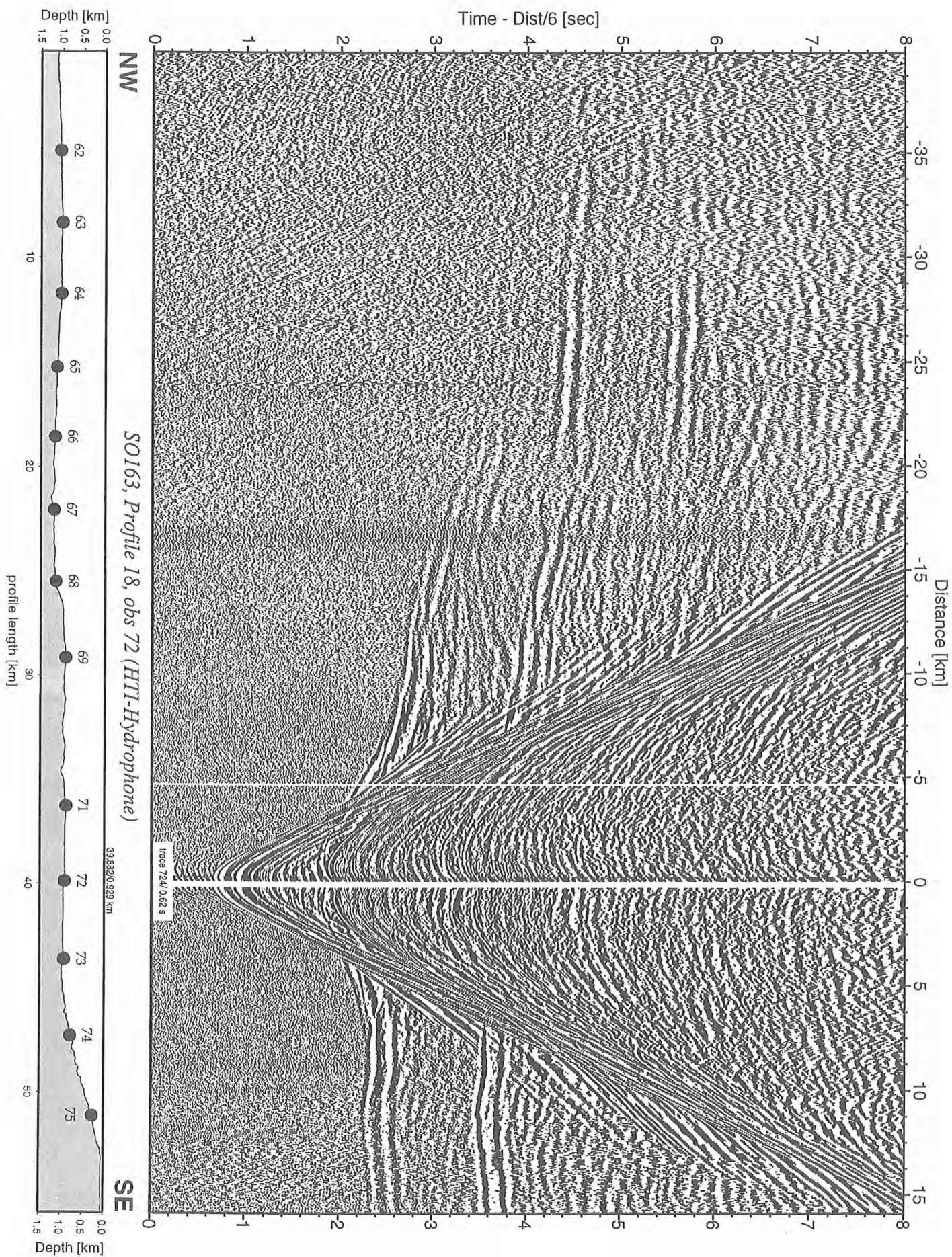


Figure 6.6.3.36: Record section from obs 72 (HTI-Hydrophone), Profile 18.

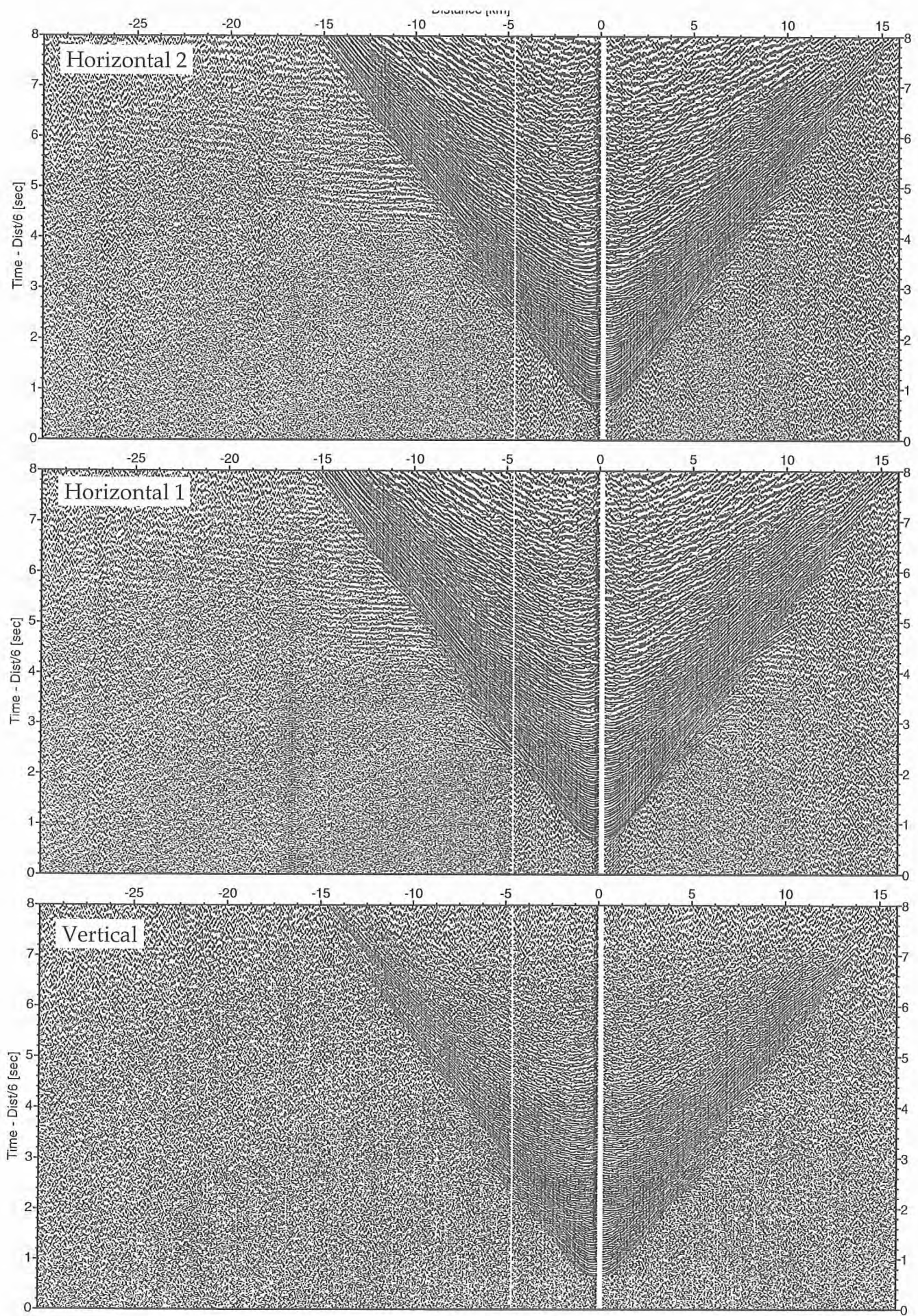


Figure 6.6.3.37: Record sections from obs 72 (Owen-30Hz), SO163, Profile 18.

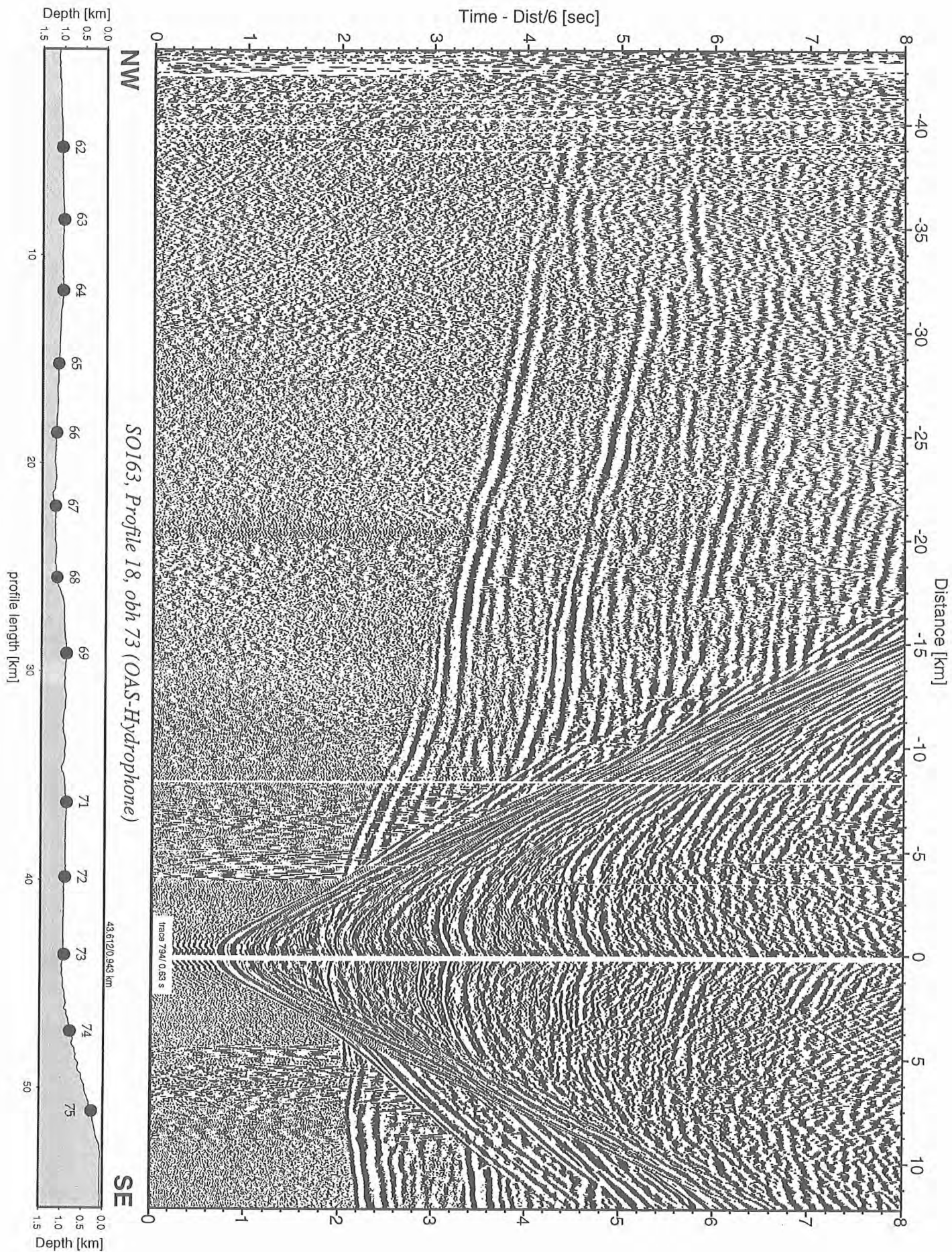


Figure 6.6.3.38: Record section from obh 73 (OAS-Hydrophone), Profile 18.

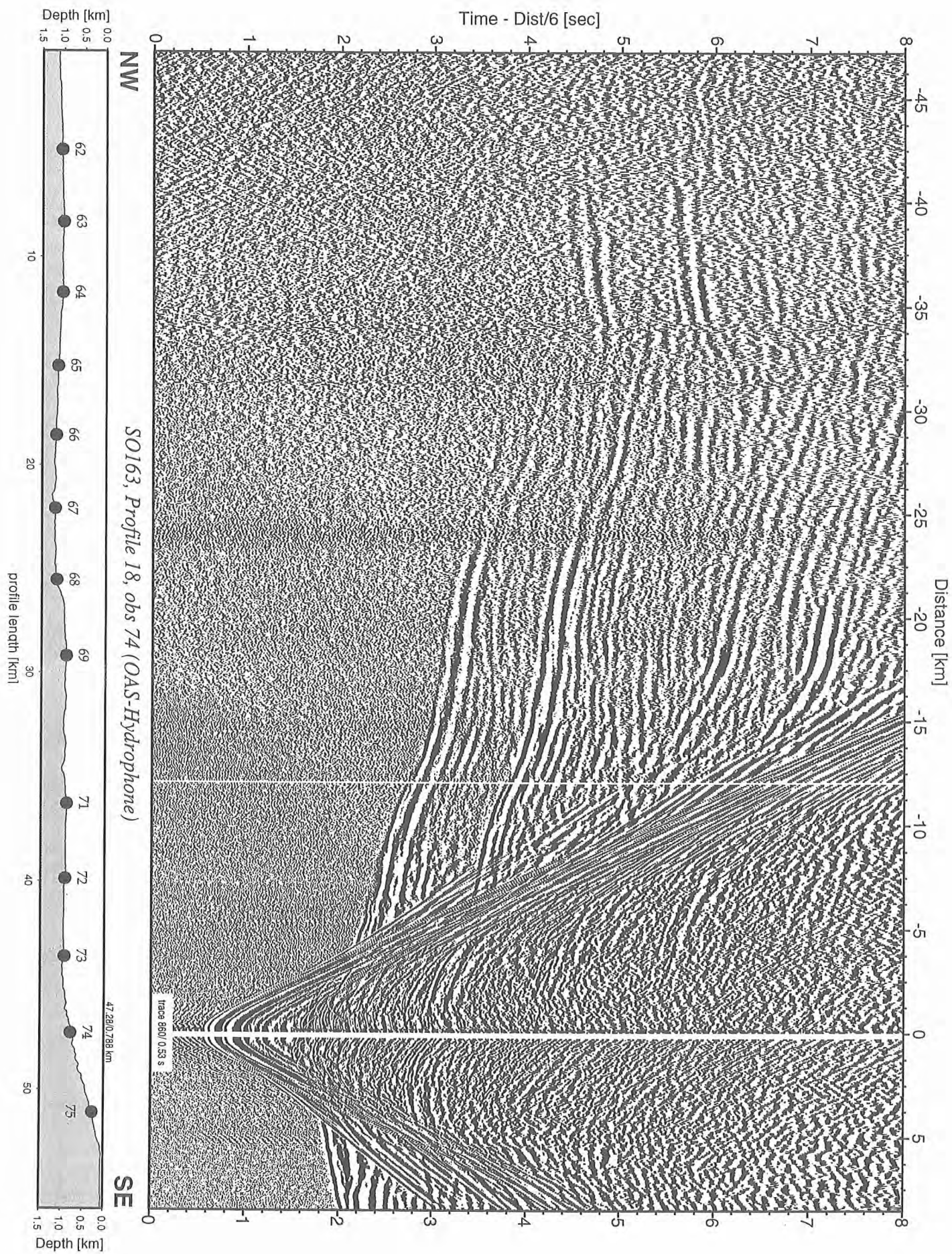


Figure 6.6.3.39: Record section from obs 74 (OAS-Hydrophone), Profile 18.

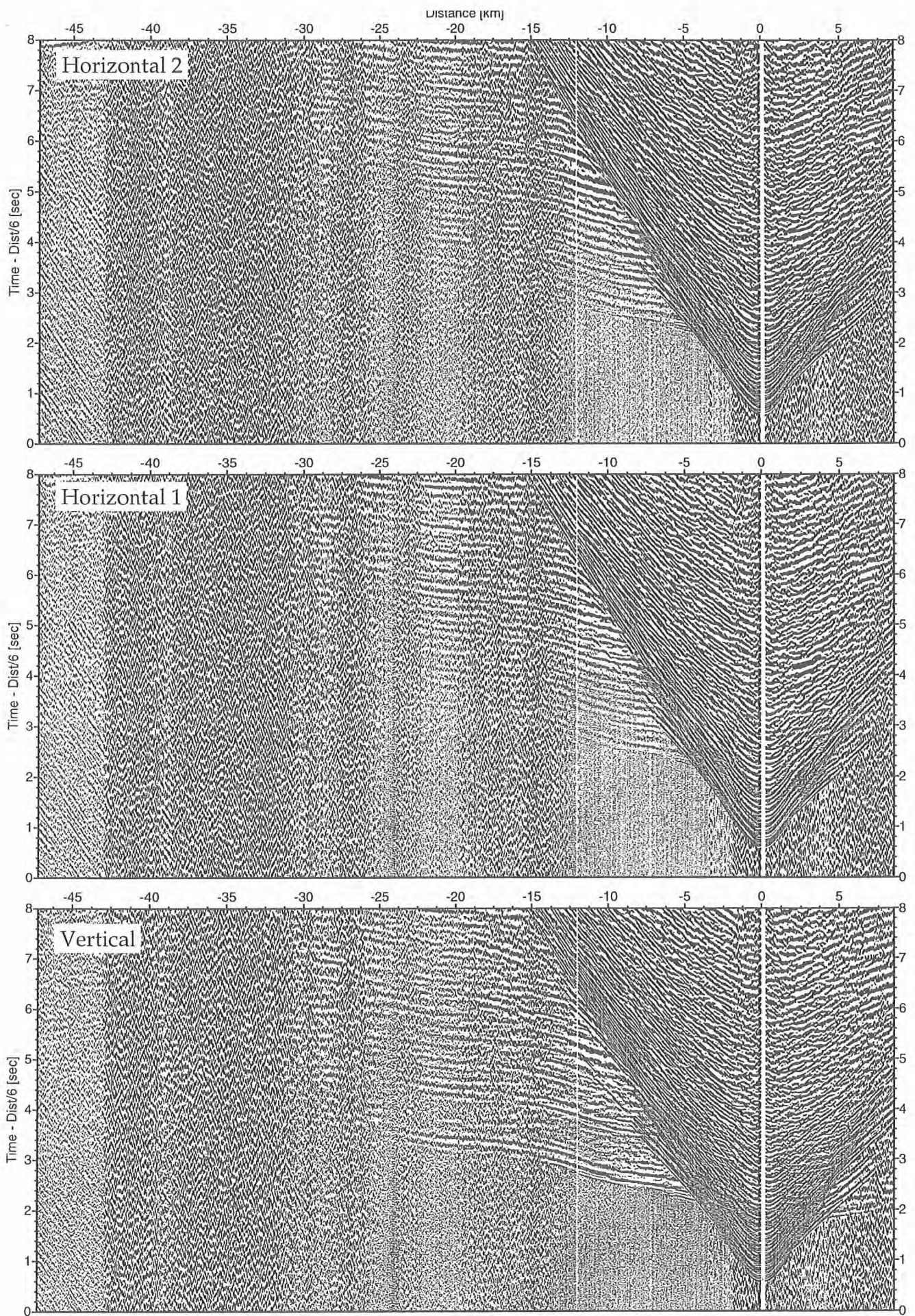


Figure 6.6.3.40: Record sections from obs 74 (LG-4.5Hz), SO163, Profile 18.

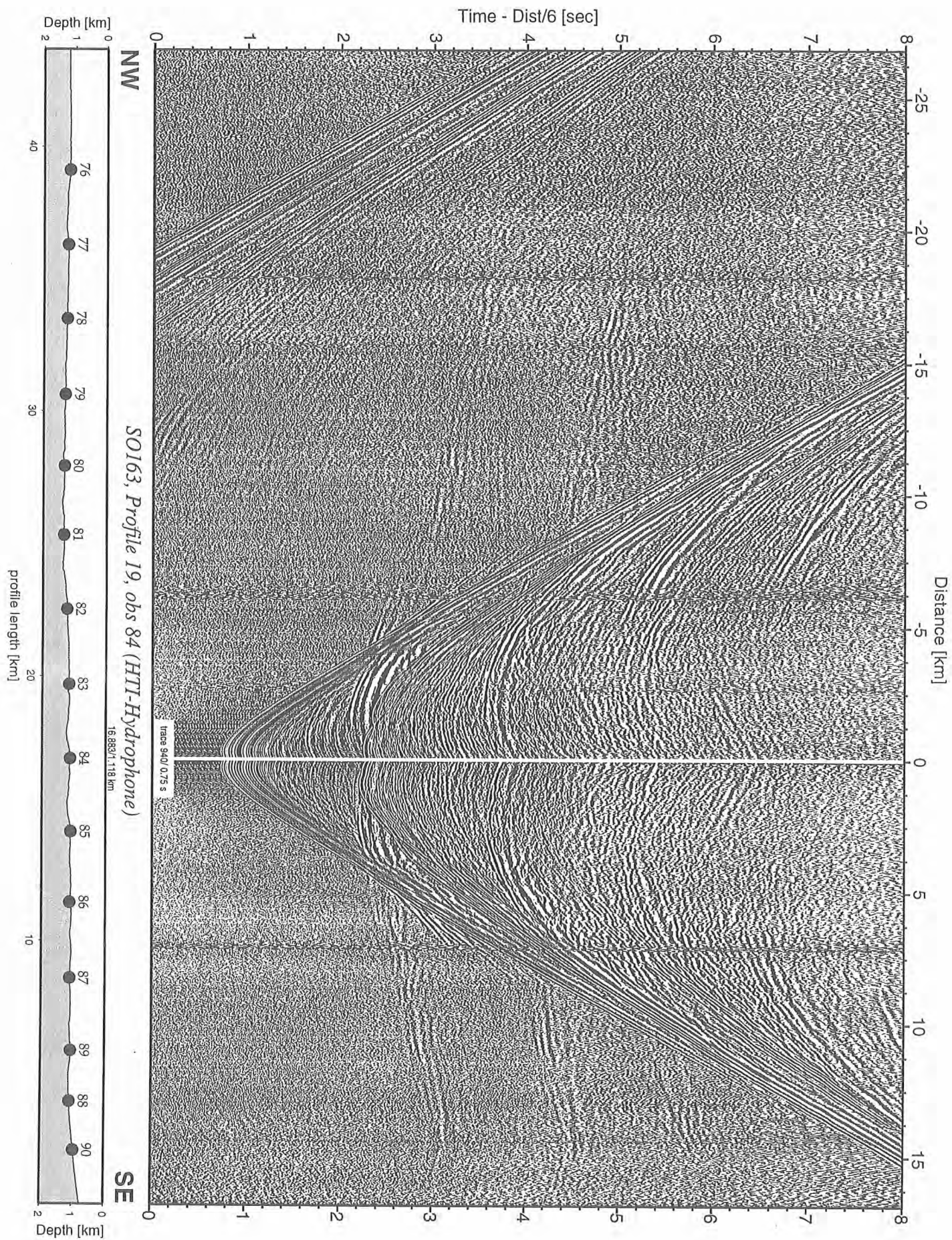


Figure 6.6.3.43: Record section from obs 84 (HTI-Hydrophone), Profile 19.

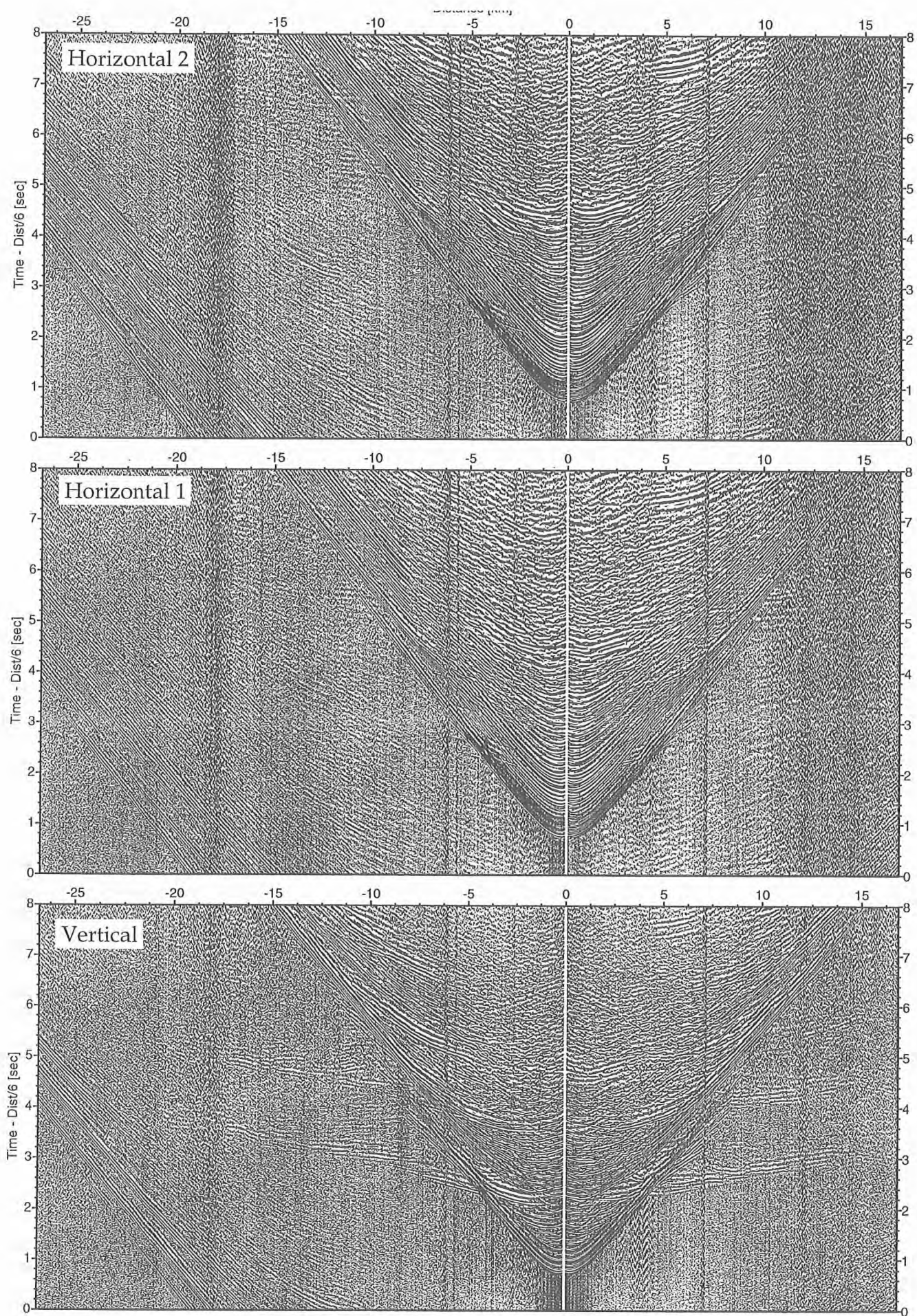


Figure 6.6.3.44: Record sections from obs 84 (Owen-4.5Hz), SO163, Profile 19.

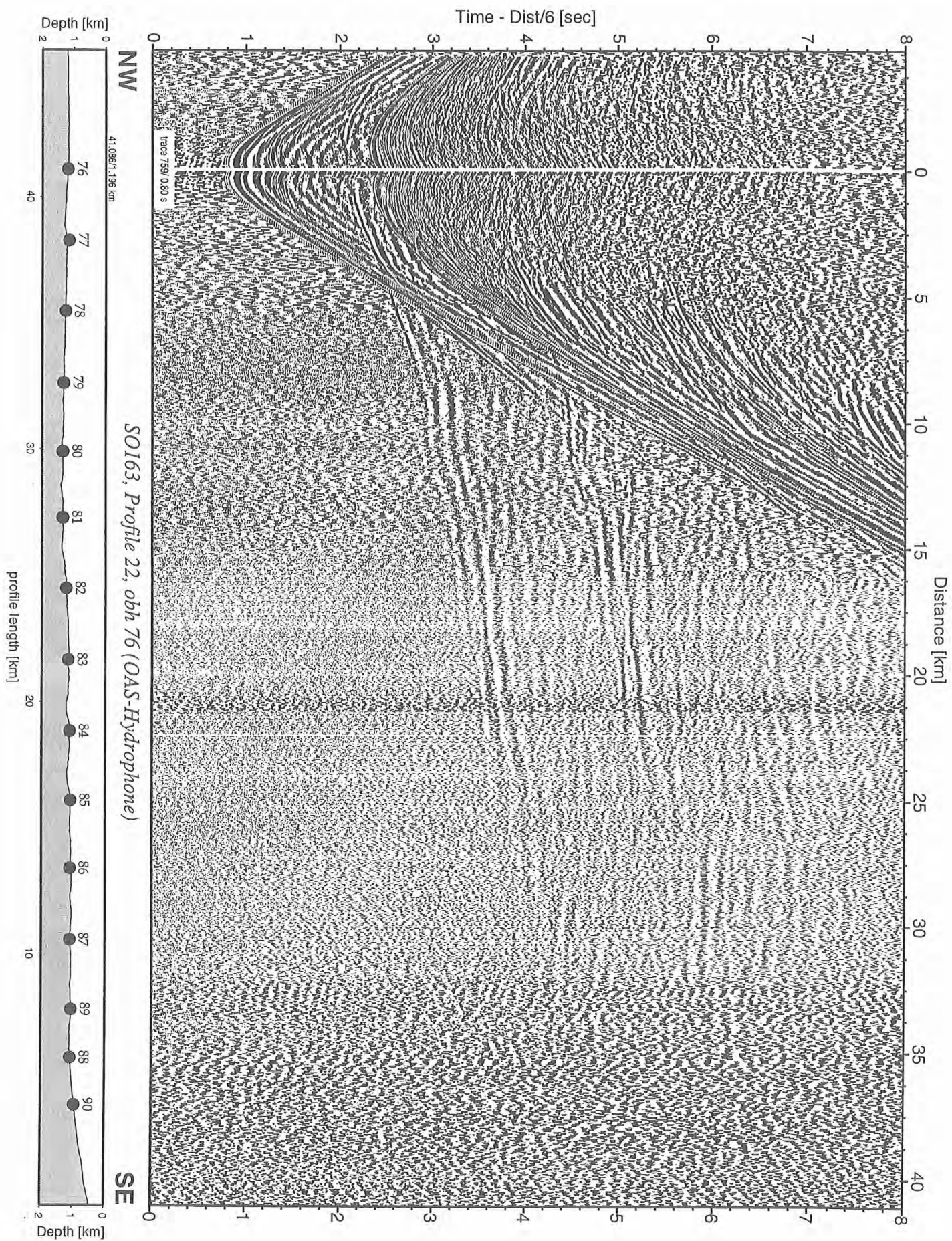


Figure 6.6.3.47: Record section from obh 76 (OAS-Hydrophone), Profile 22.

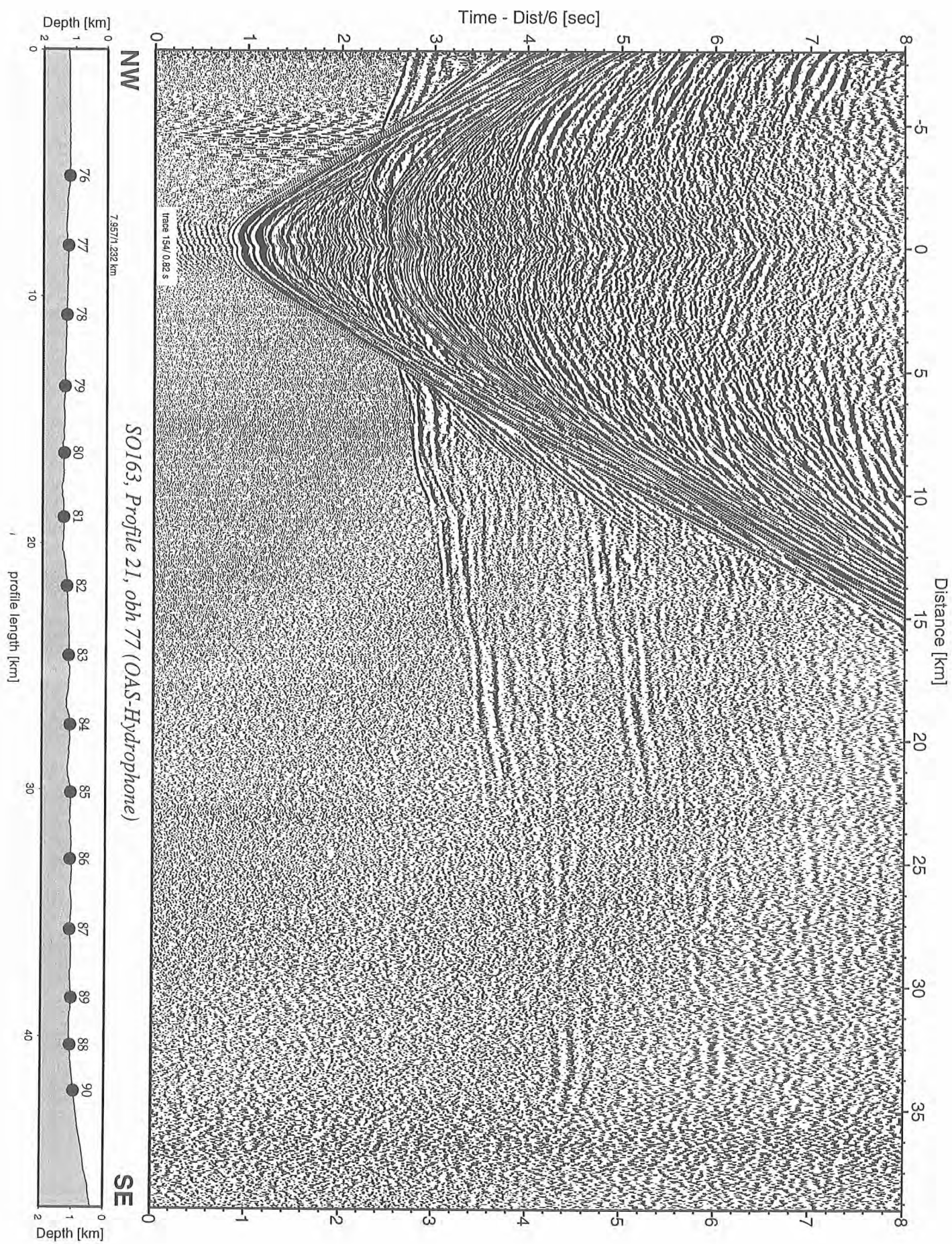


Figure 6.6.3.48: Record section from obh 77 (OAS-Hydrophone), Profile 21.

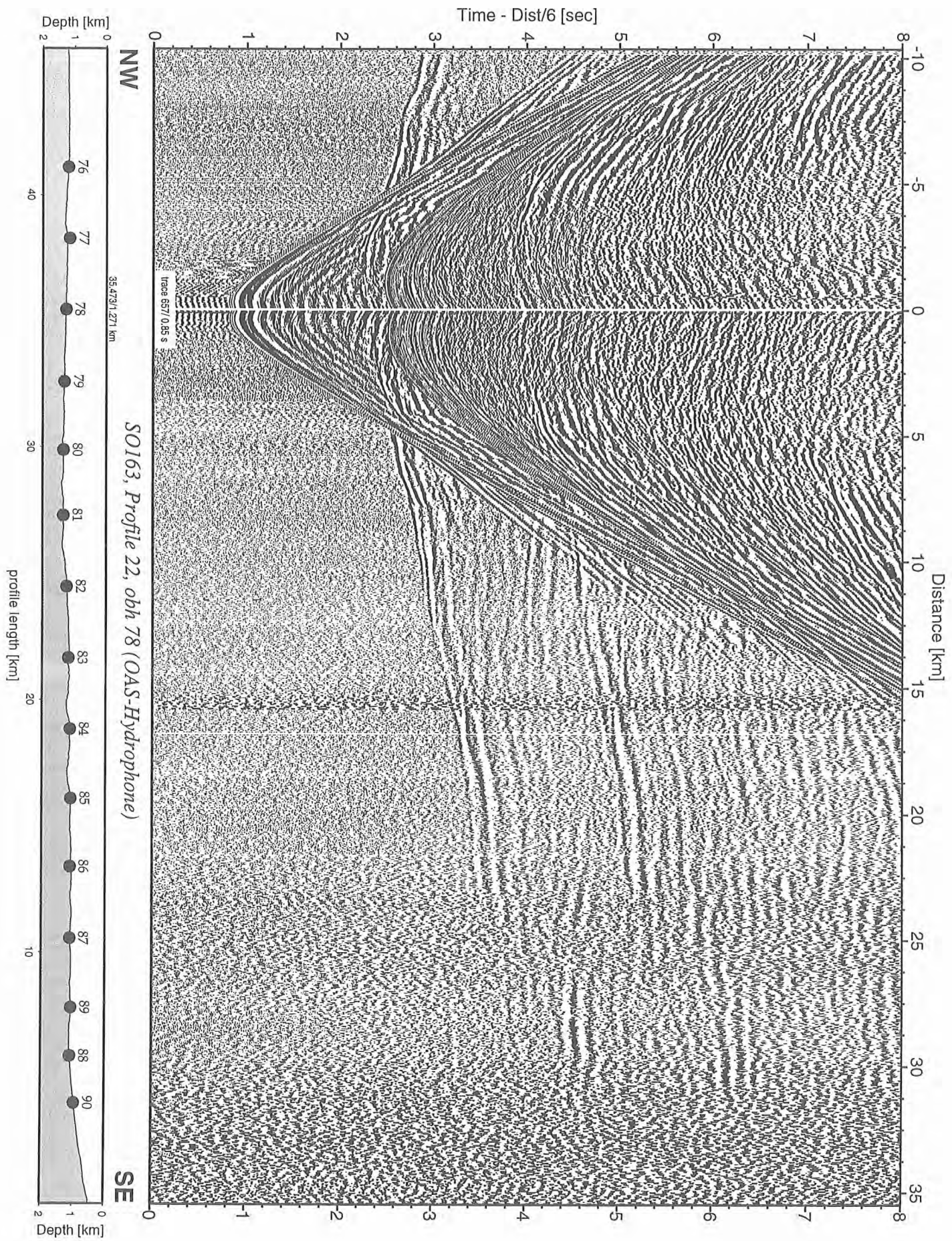


Figure 6.6.3.49: Record section from obh 78 (OAS-Hydrophone), Profile 22.

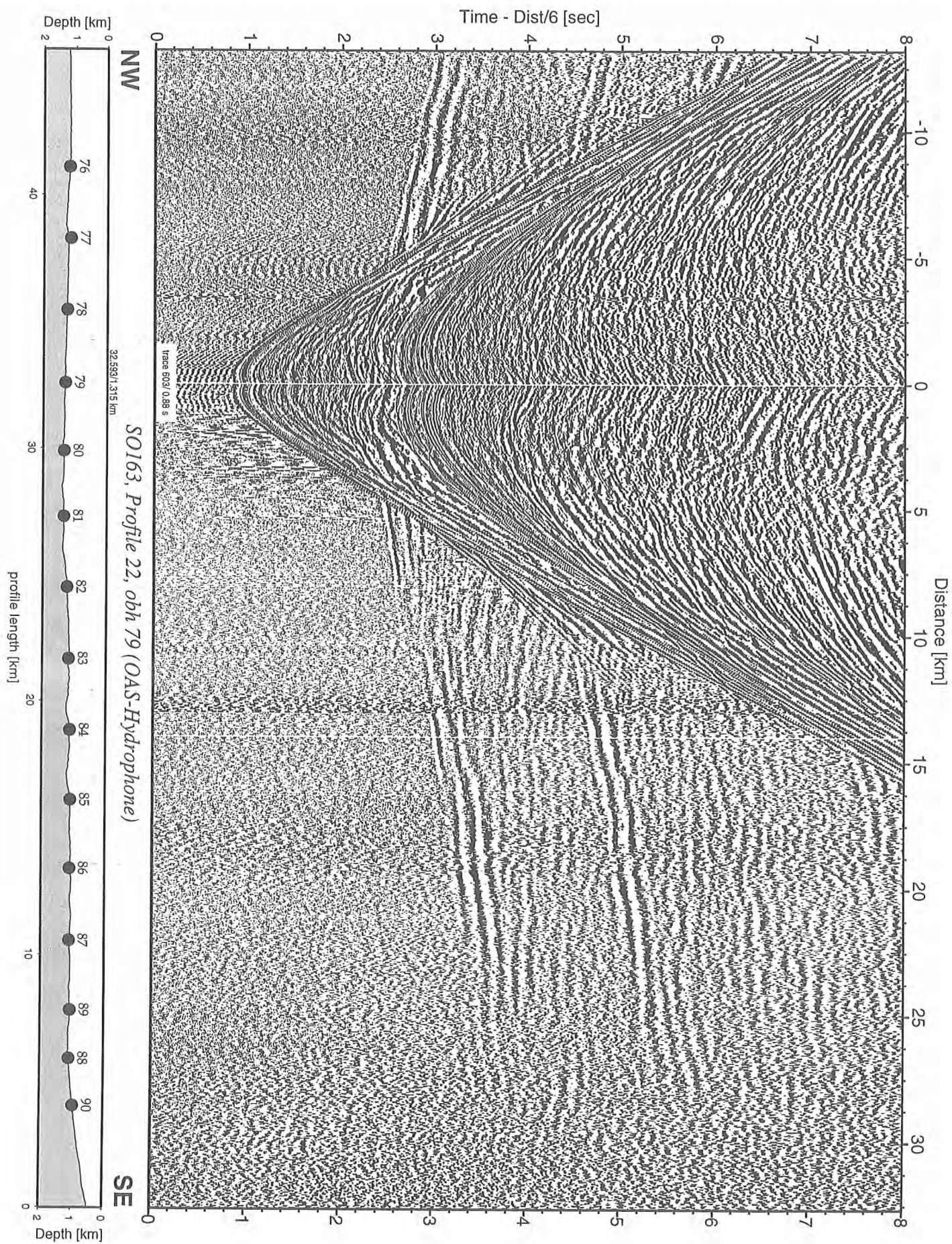


Figure 6.6.3.50: Record section from obh 79 (OAS-Hydrophone), Profile 22.

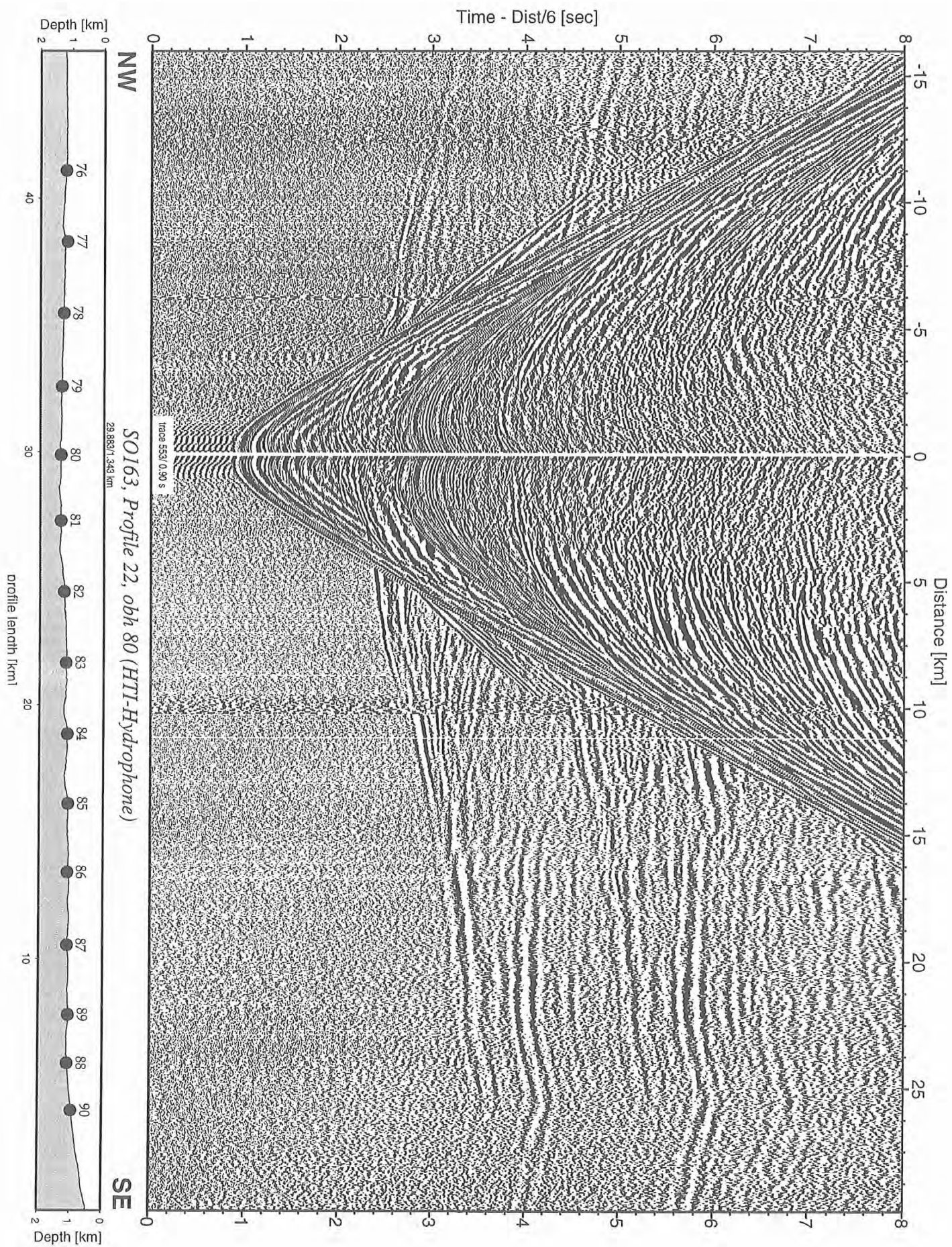


Figure 6.6.3.51: Record section from obh 80 (HTI-Hydrophone), Profile 22.

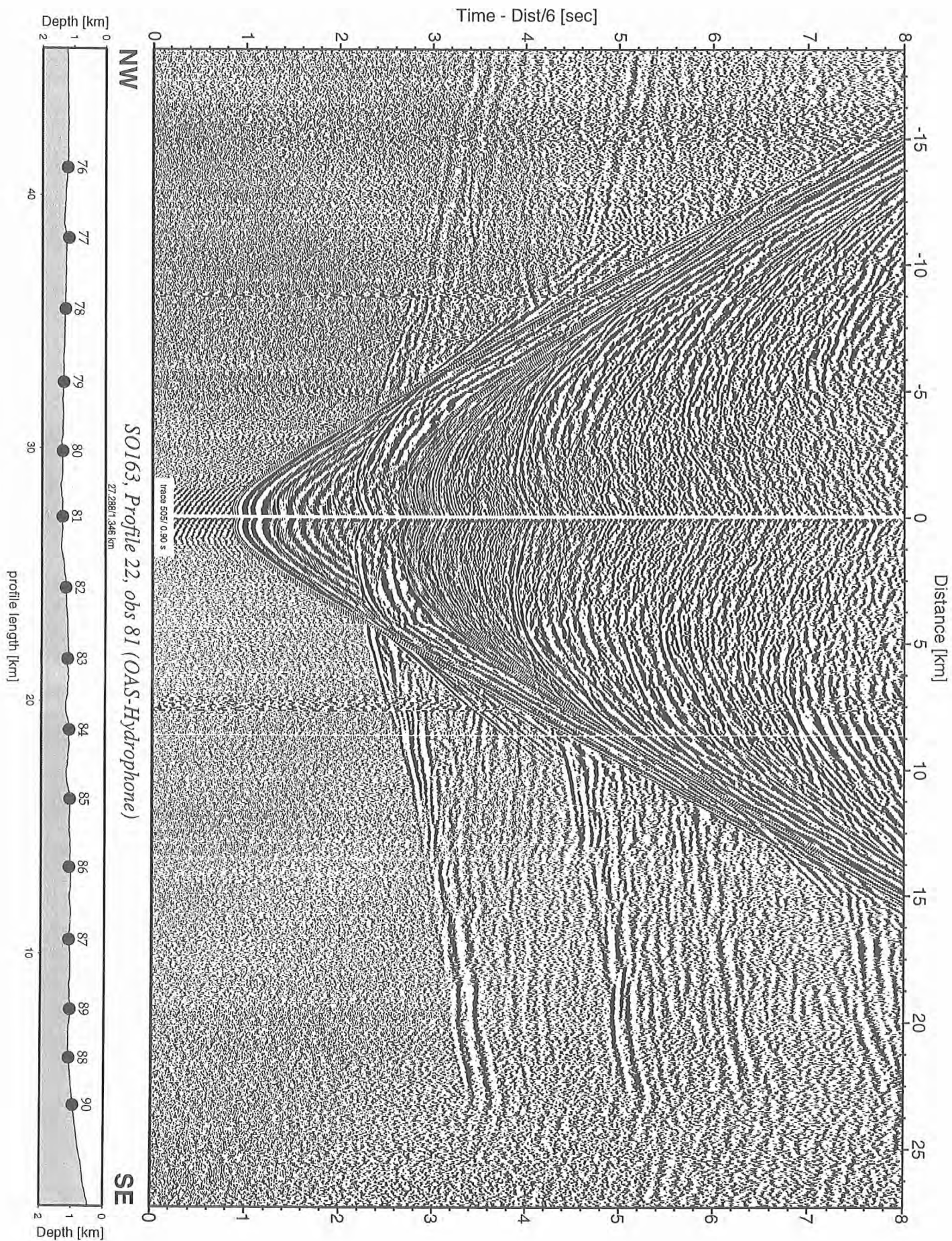


Figure 6.6.3.52: Record section from obs 81 (OAS-Hydrophone), Profile 22.

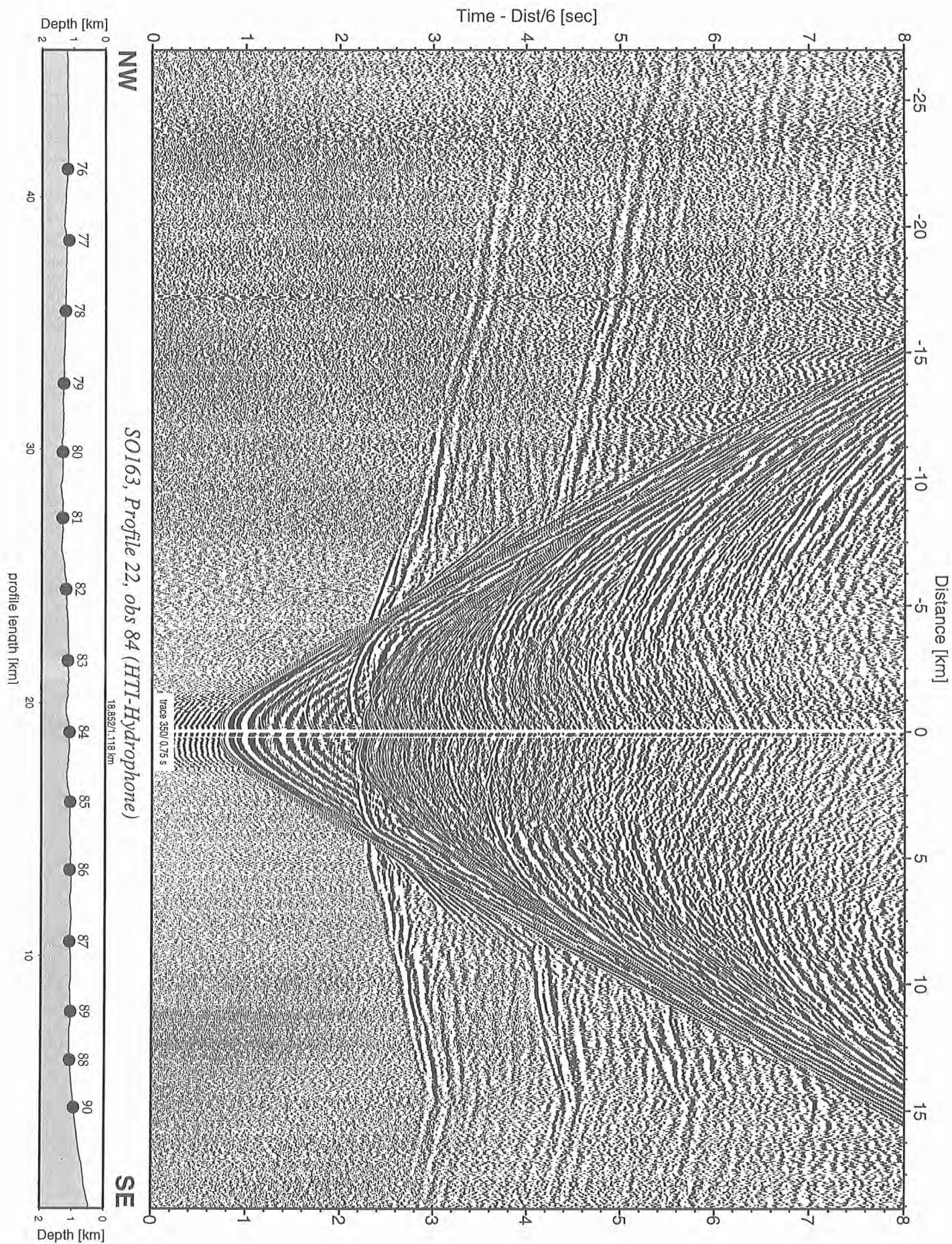


Figure 6.6.3.55: Record section from obs 84 (HTI-Hydrophone), Profile 22.

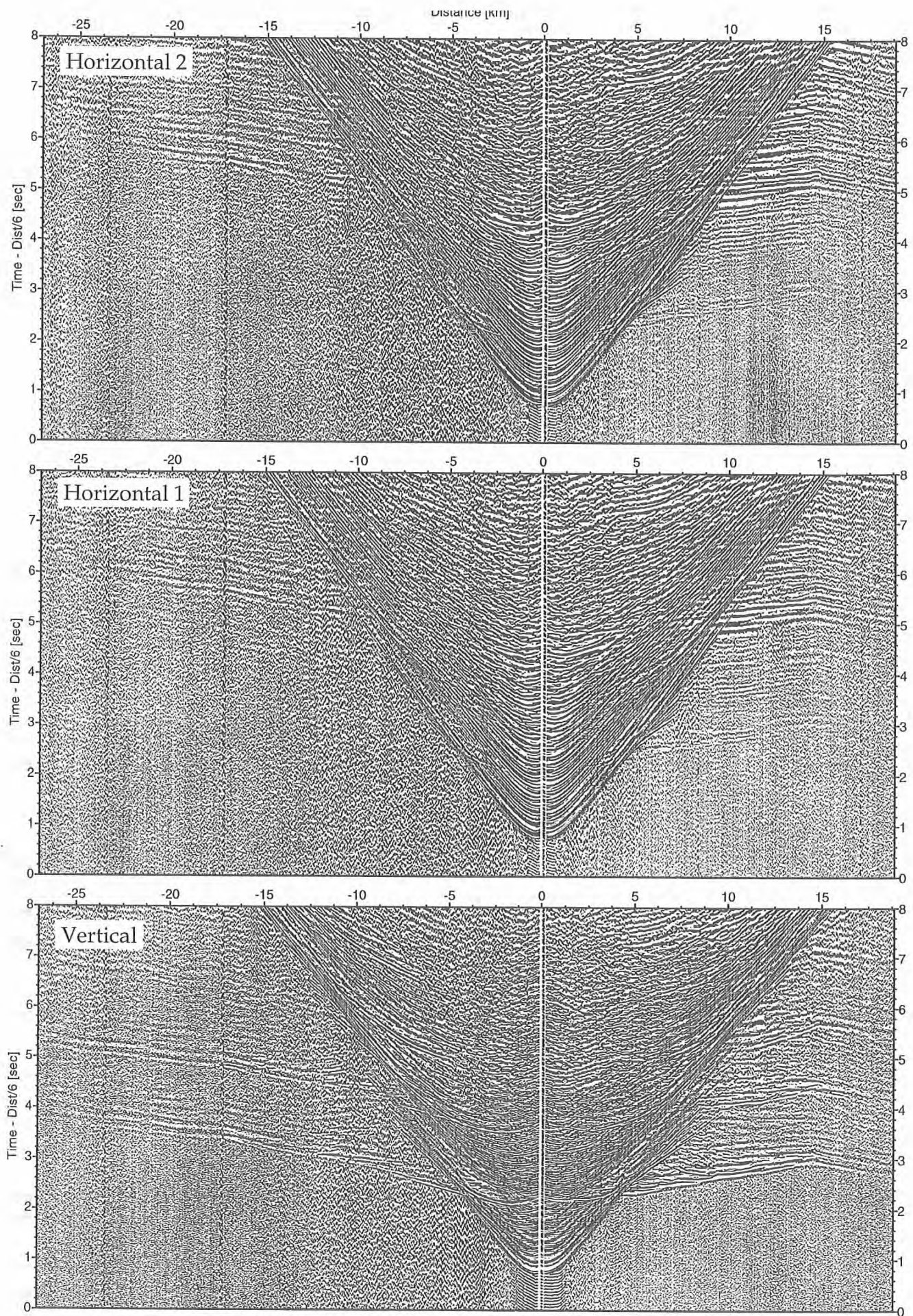


Figure 6.6.3.56: Record sections from obs 84 (Owen-4.5Hz), SO163, Profile 22.

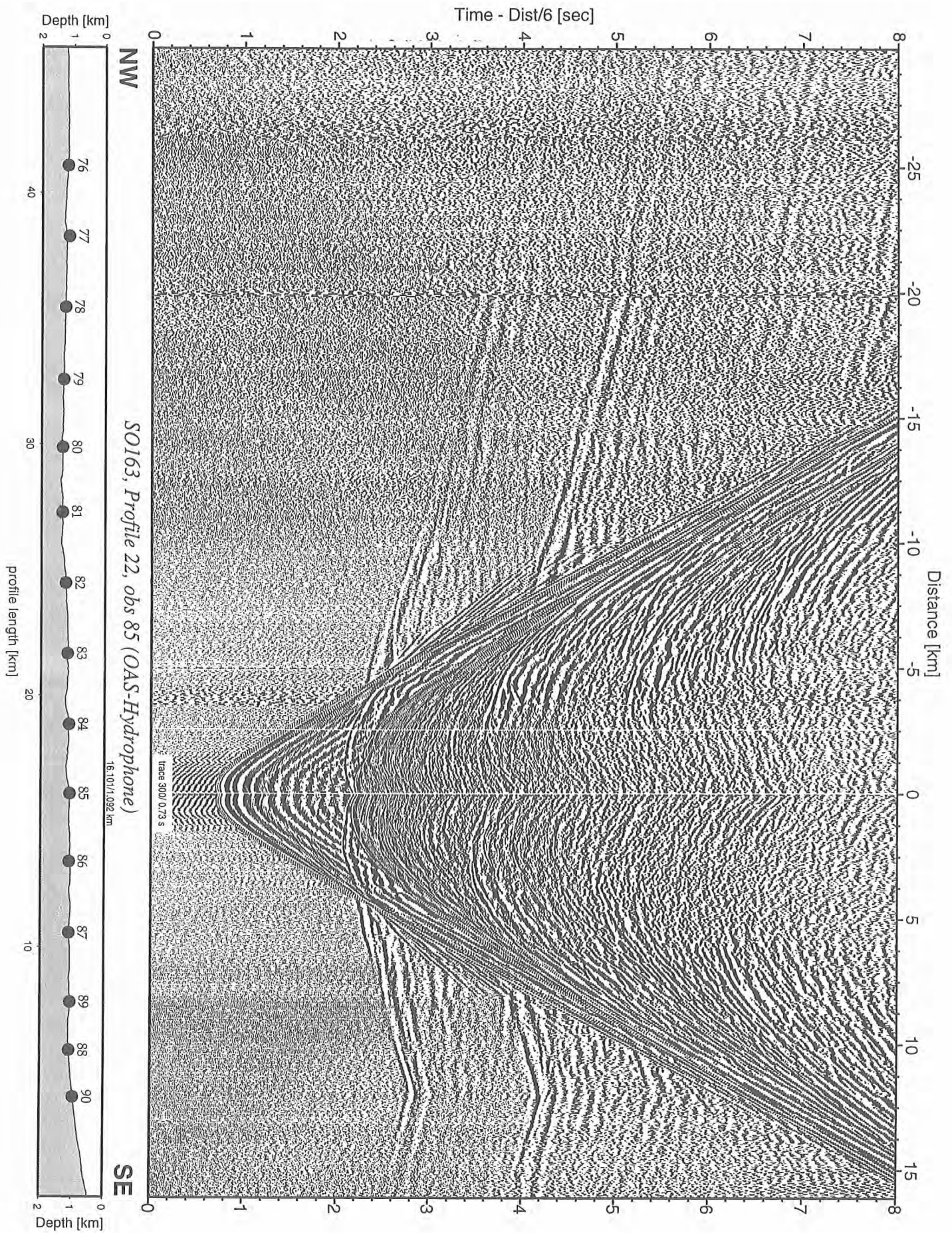


Figure 6.6.3.57: Record section from obs 85 (OAS-Hydrophone), Profile 22.

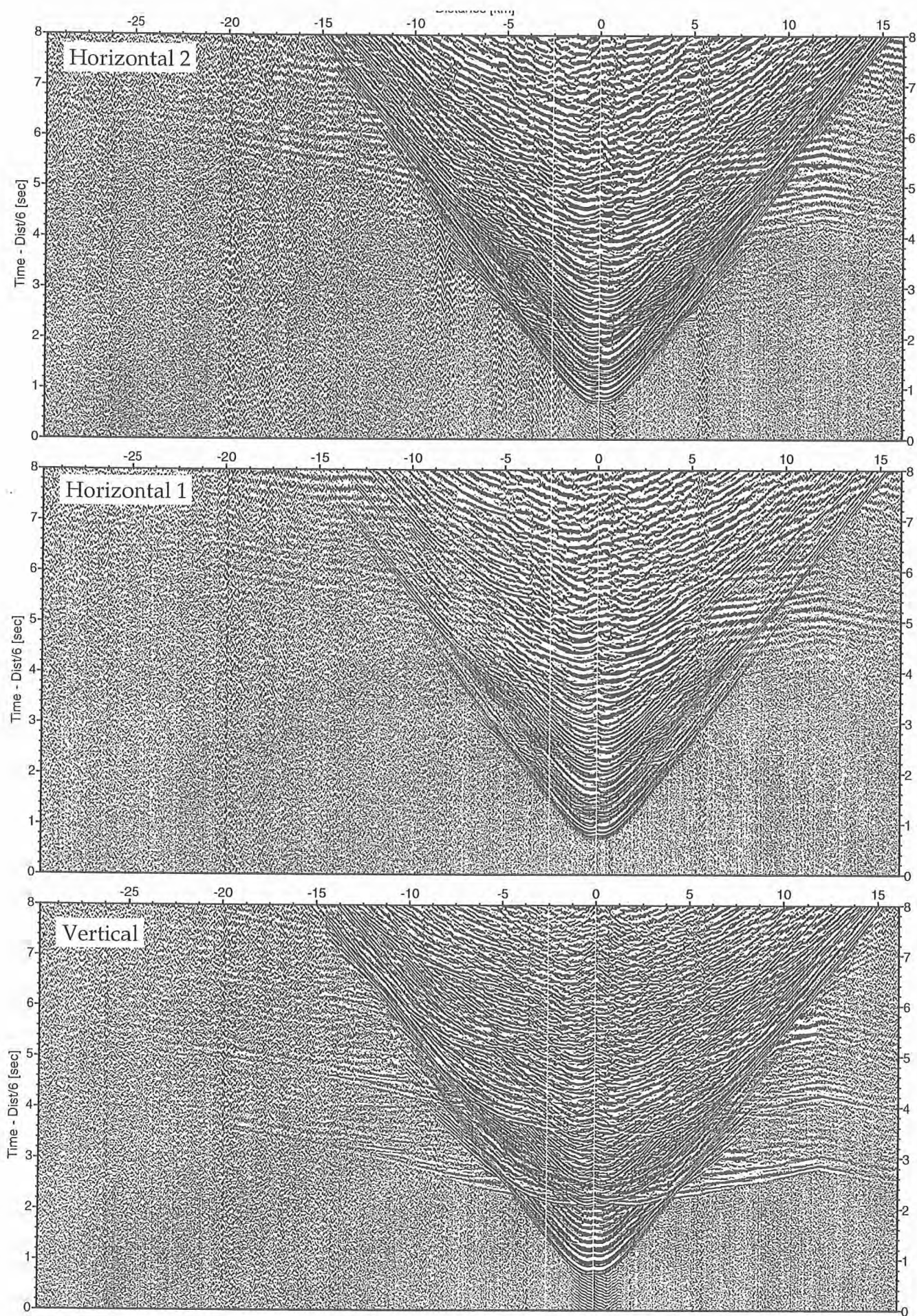


Figure 6.6.3.58: Record sections from obs 85 (LG-4.5Hz), SO163, Profile 22.

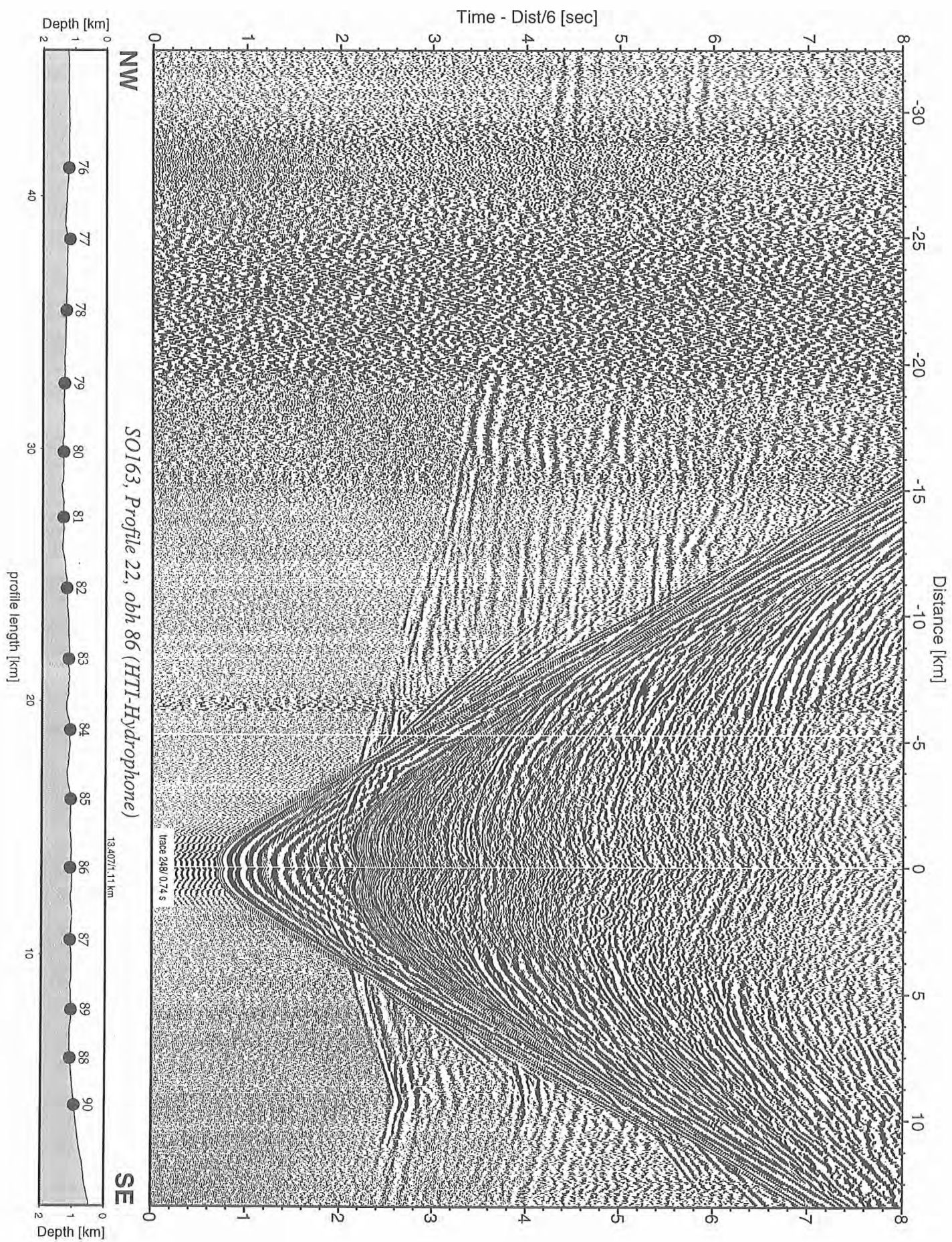


Figure 6.6.3.59: Record section from obh 86 (HTI-Hydrophone), Profile 22.

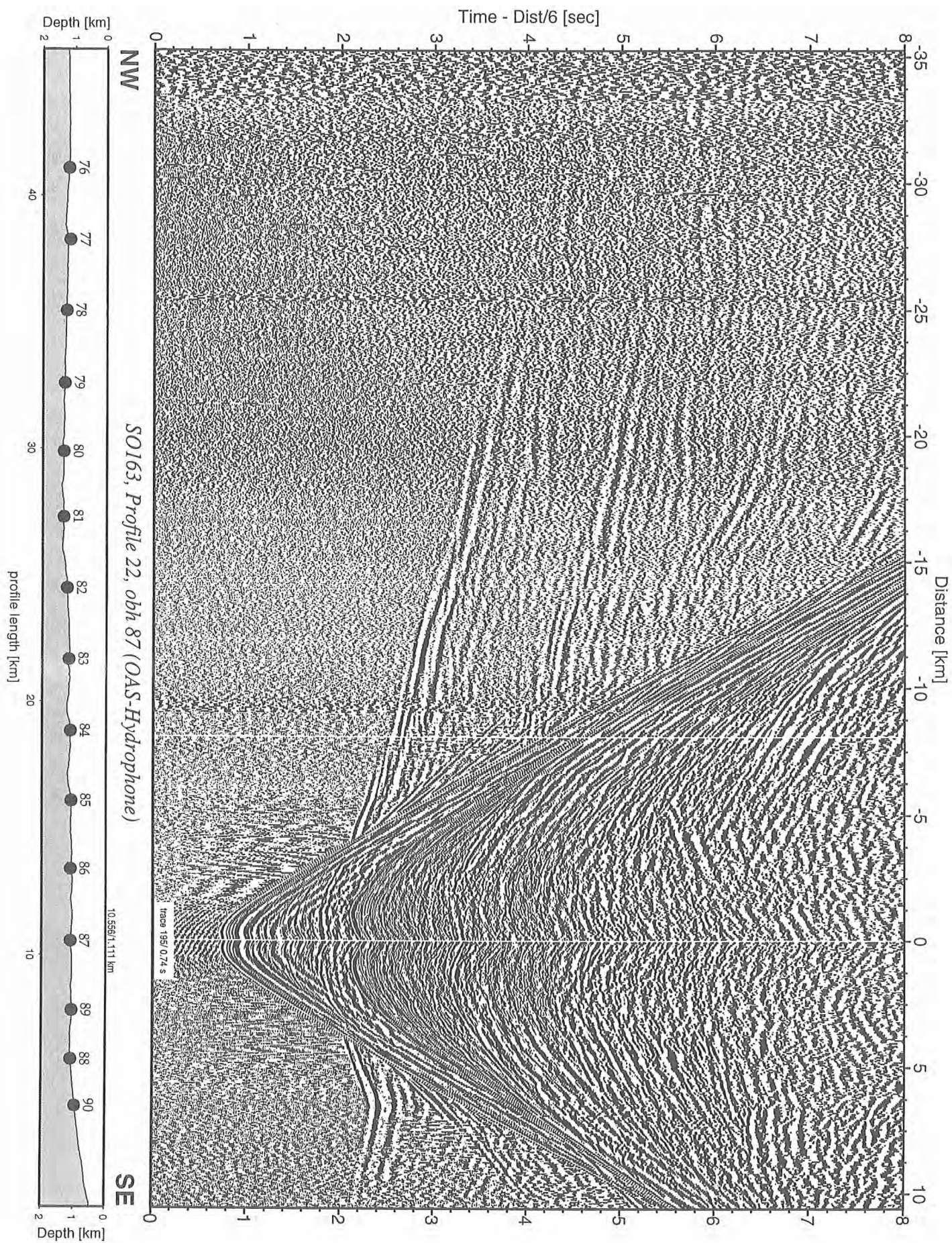


Figure 6.6.3.60: Record section from obh 87 (OAS-Hydrophone), Profile 22.

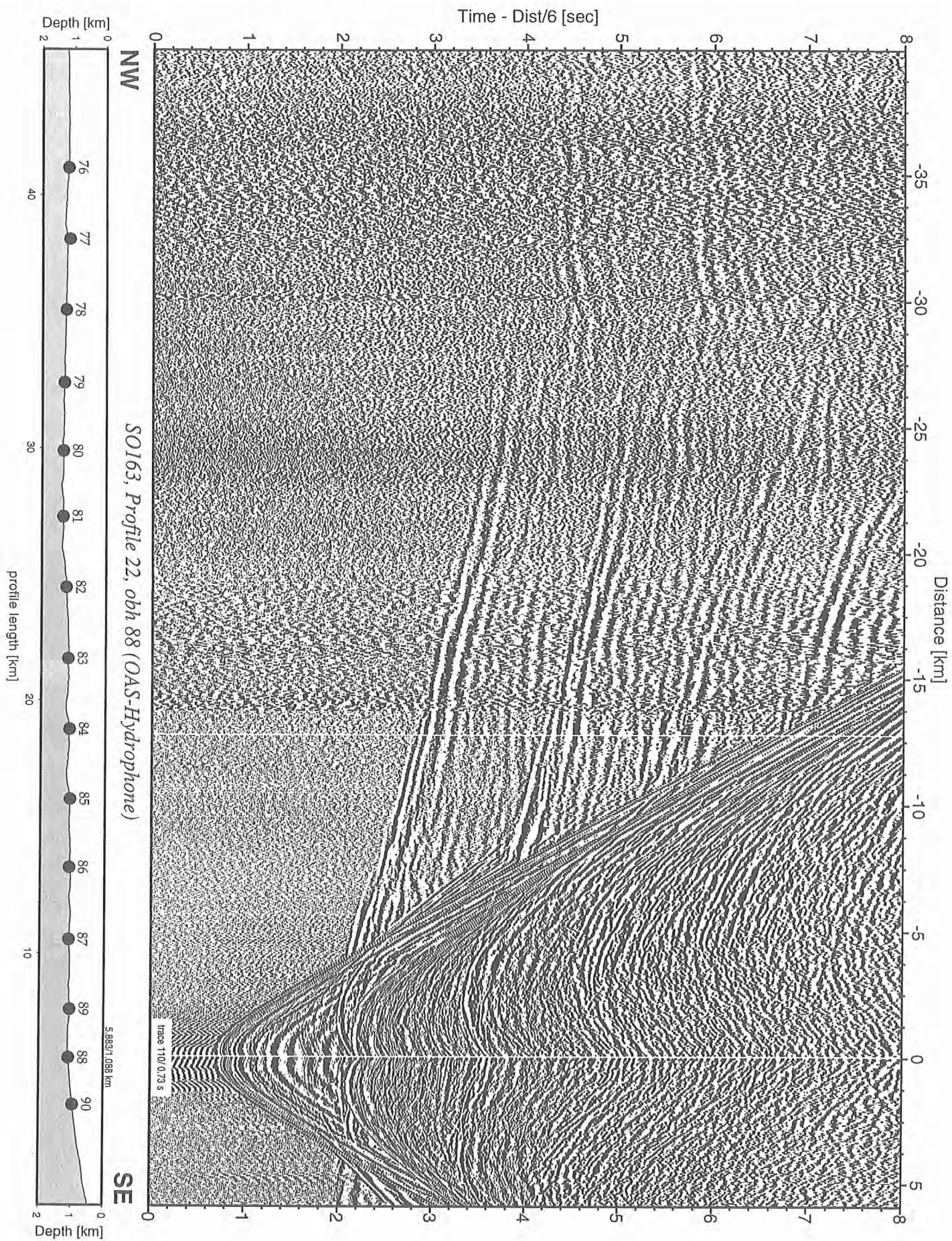


Figure 6.6.3.61: Record section from obh 88 (OAS-Hydrophone), Profile 22.

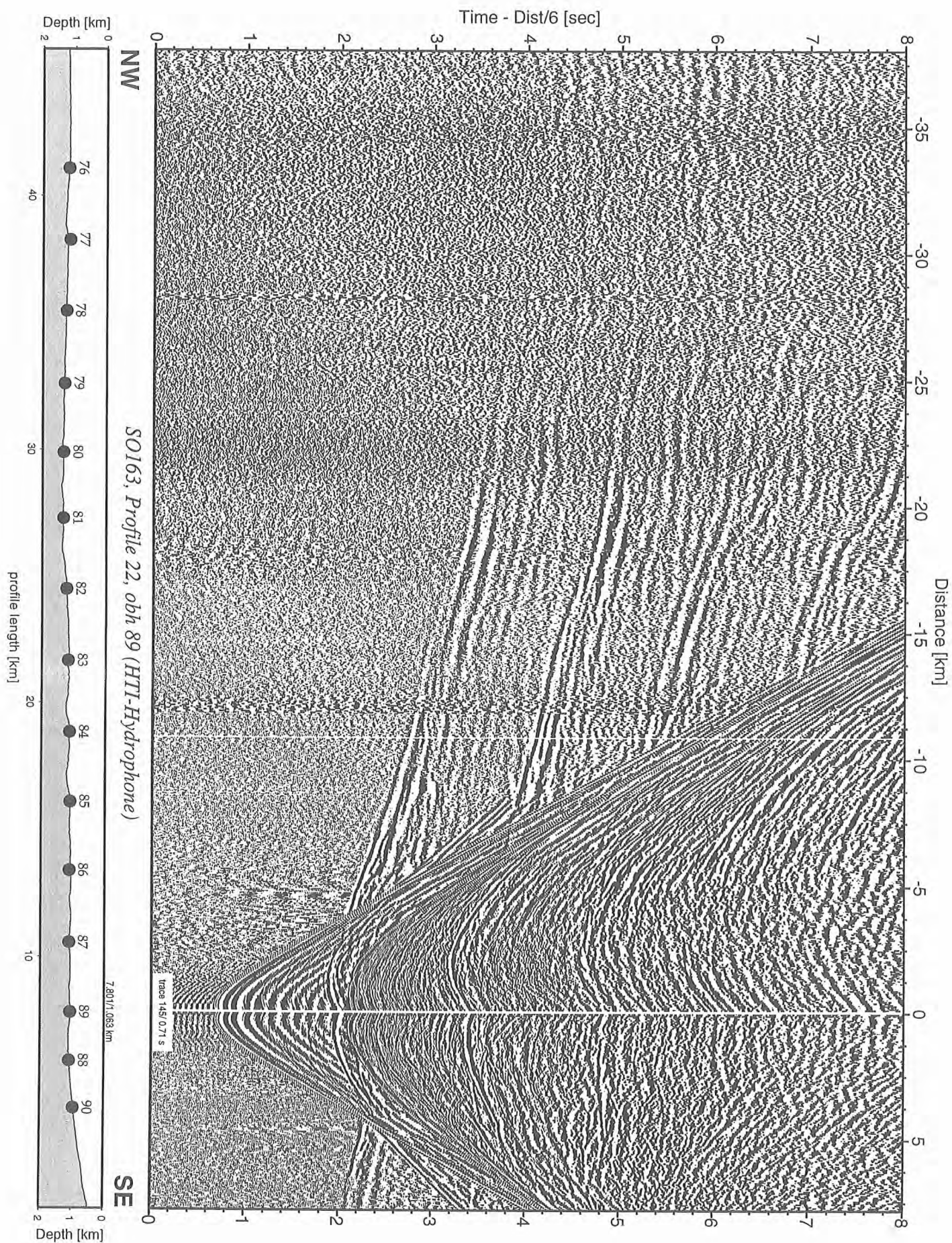


Figure 6.6.3.62: Record section from obh 89 (HTI-Hydrophone), Profile 22.

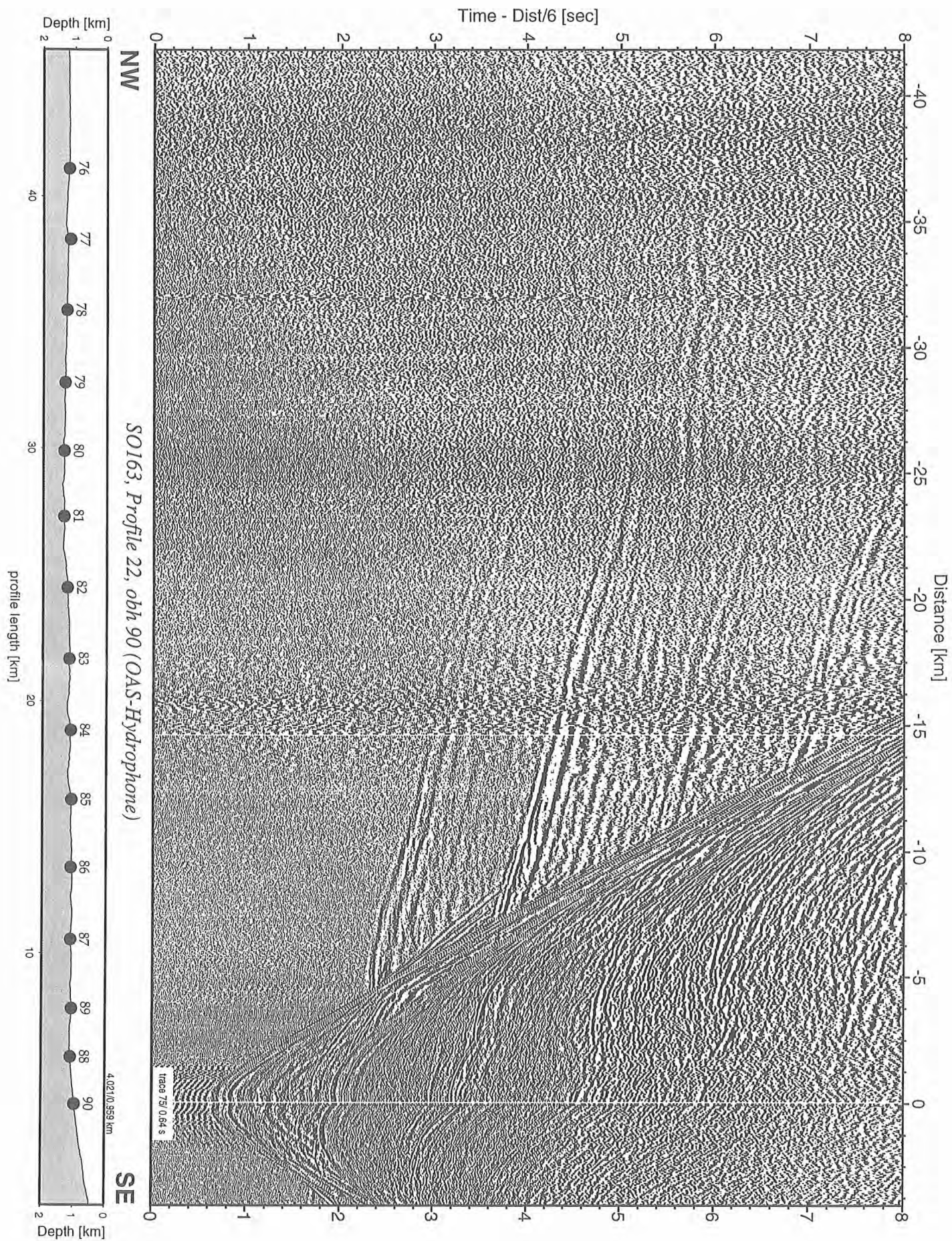


Figure 6.6.3.63: Record section from obh 90 (OAS-Hydrophone), Profile 22.

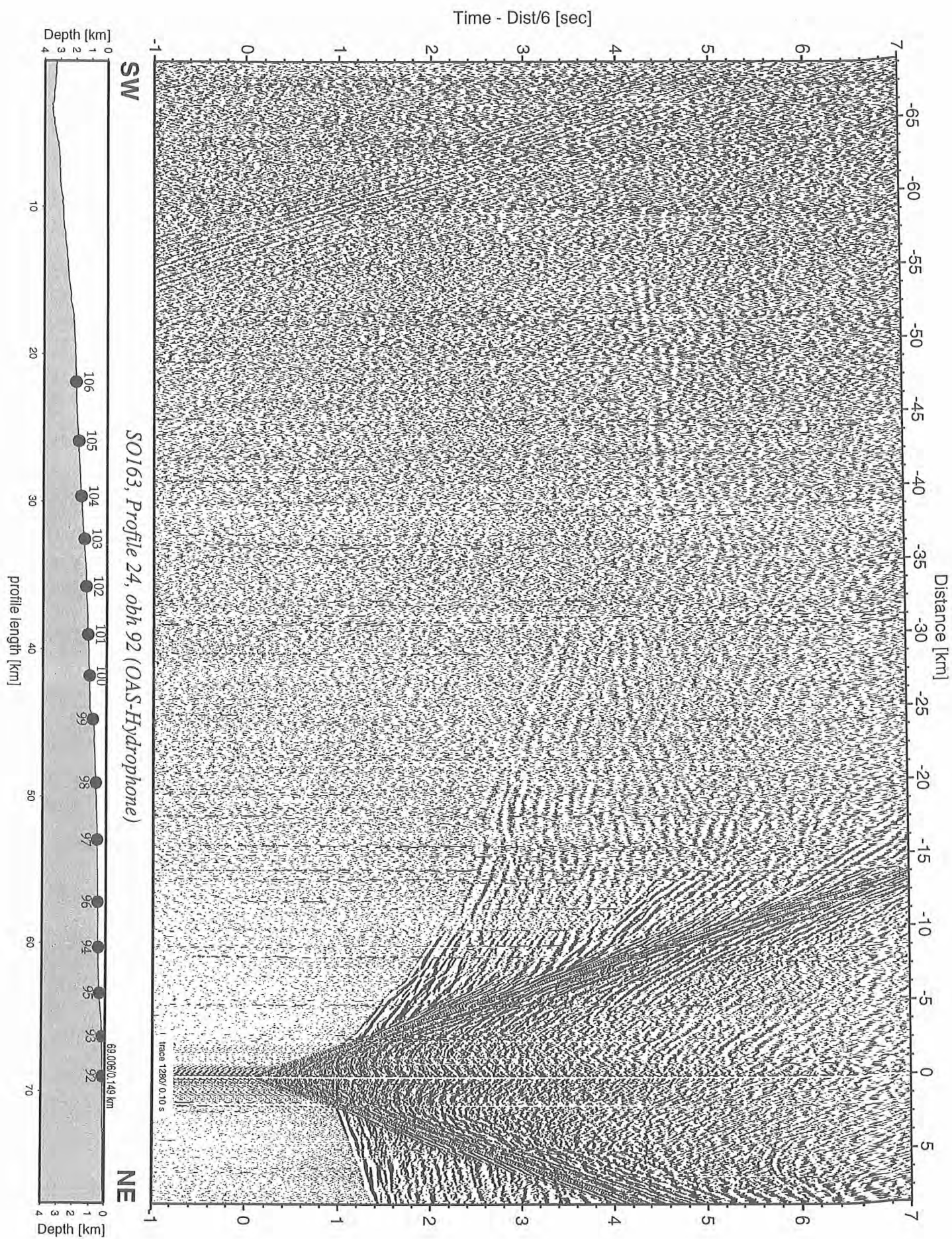


Figure 6.6.3.64: Record section from obh 92 (OAS-Hydrophone), Profile 24.

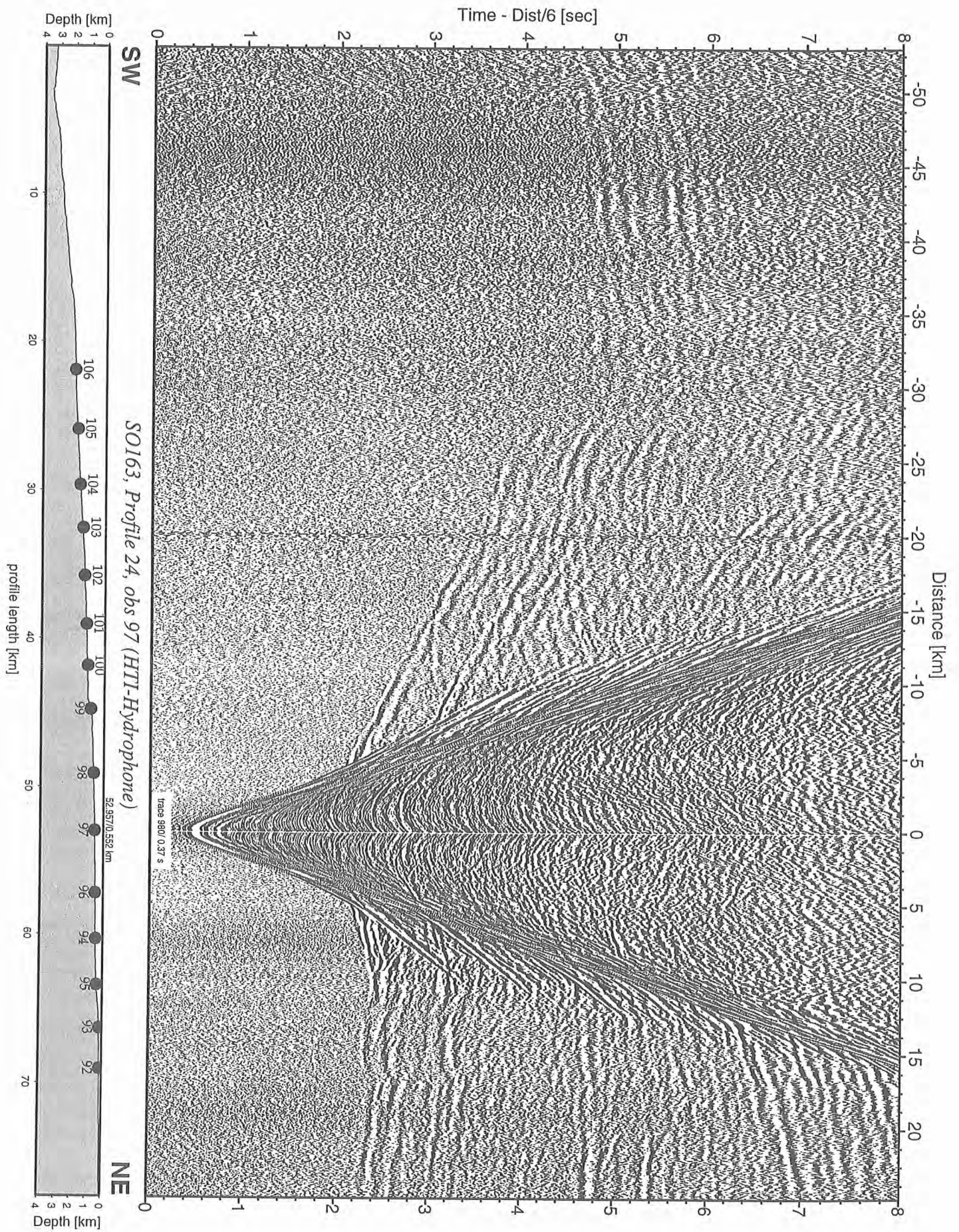


Figure 6.6.3.69: Record section from obs 97 (HTI-Hydrophone), Profile 24.

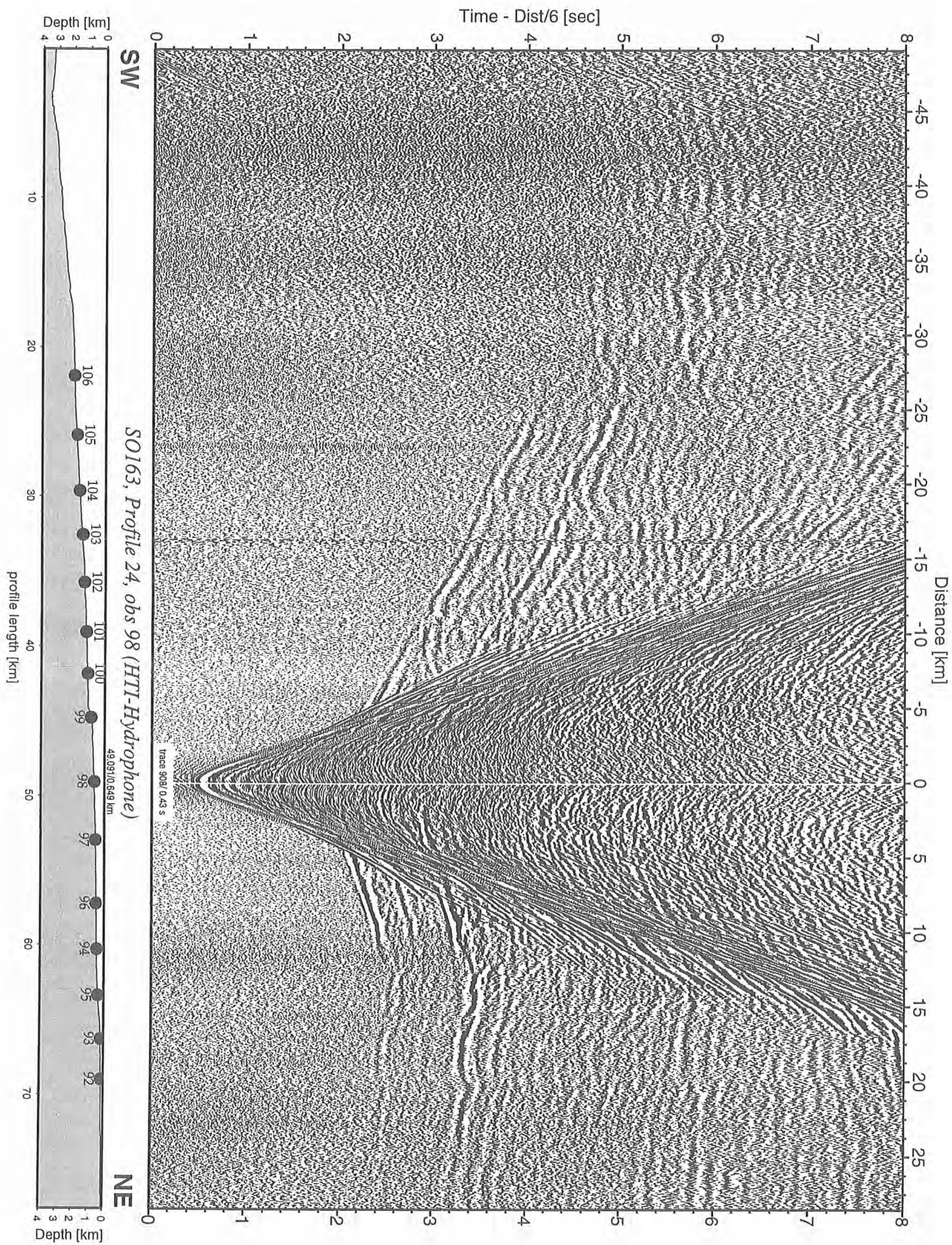


Figure 6.6.3.70: Record section from obs 98 (HTI-Hydrophone), Profile 24.

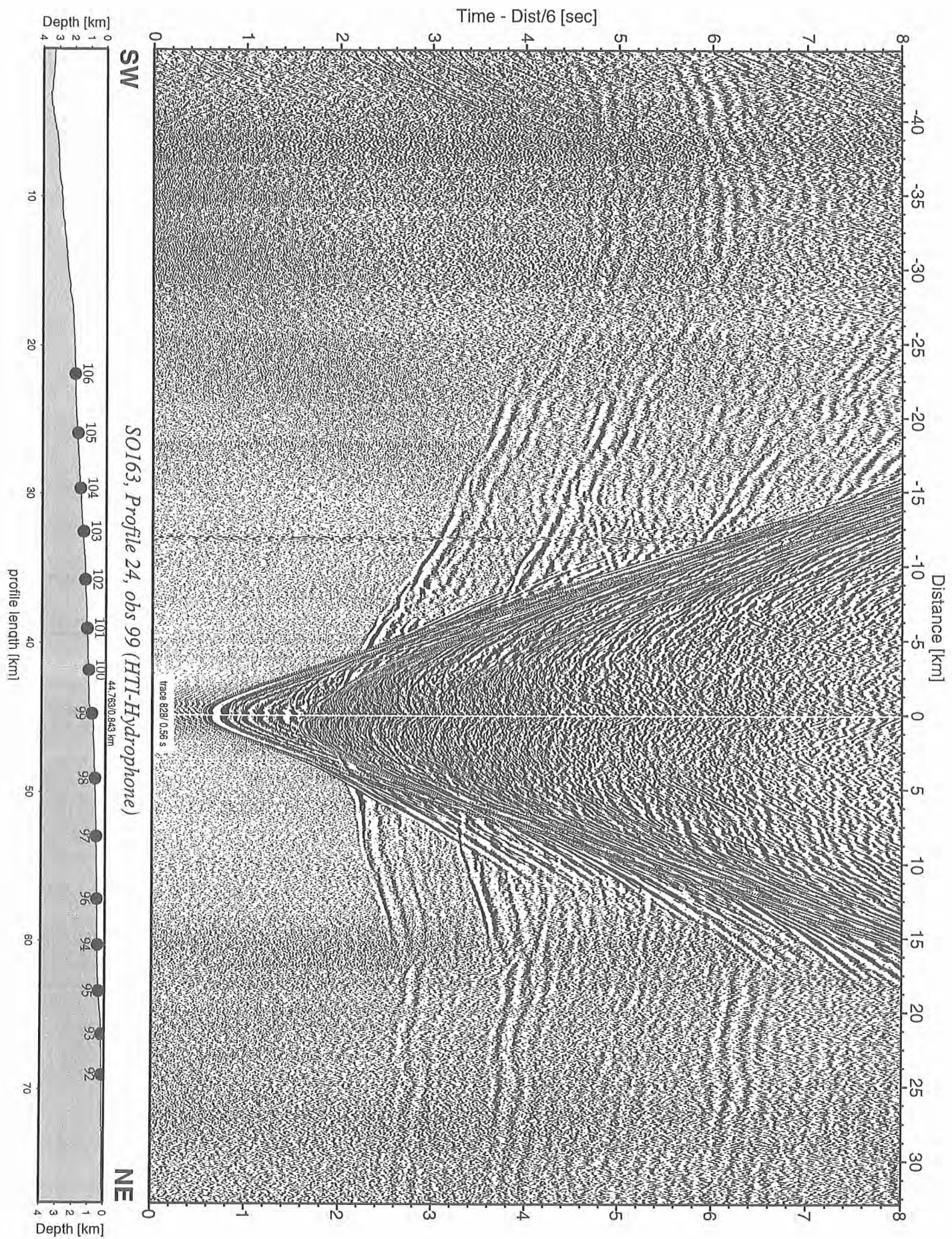


Figure 6.6.3.71: Record section from obs 99 (HTI-Hydrophone), Profile 24.

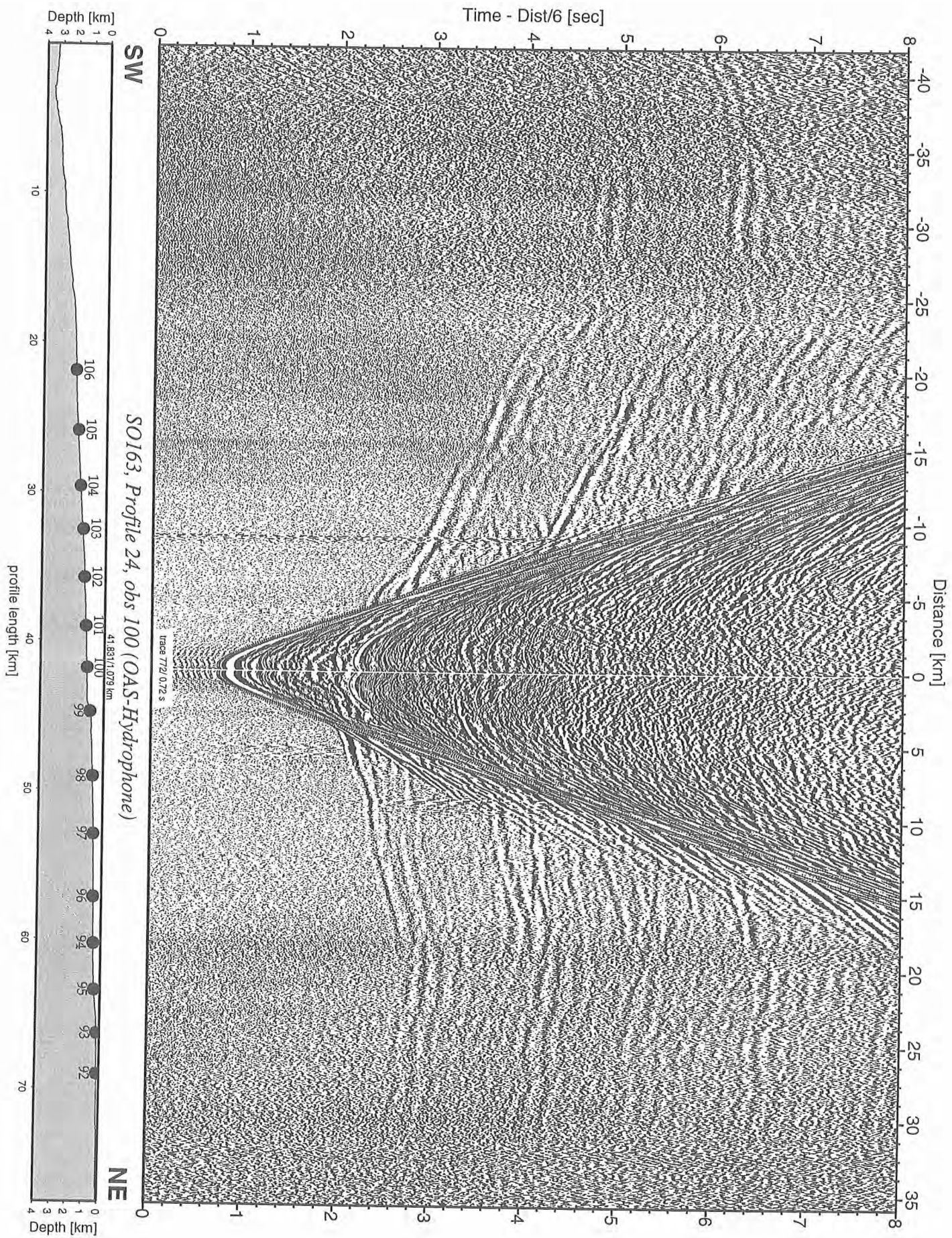


Figure 6.6.3.72: Record section from obs 100 (OAS-Hydrophone), Profile 24.

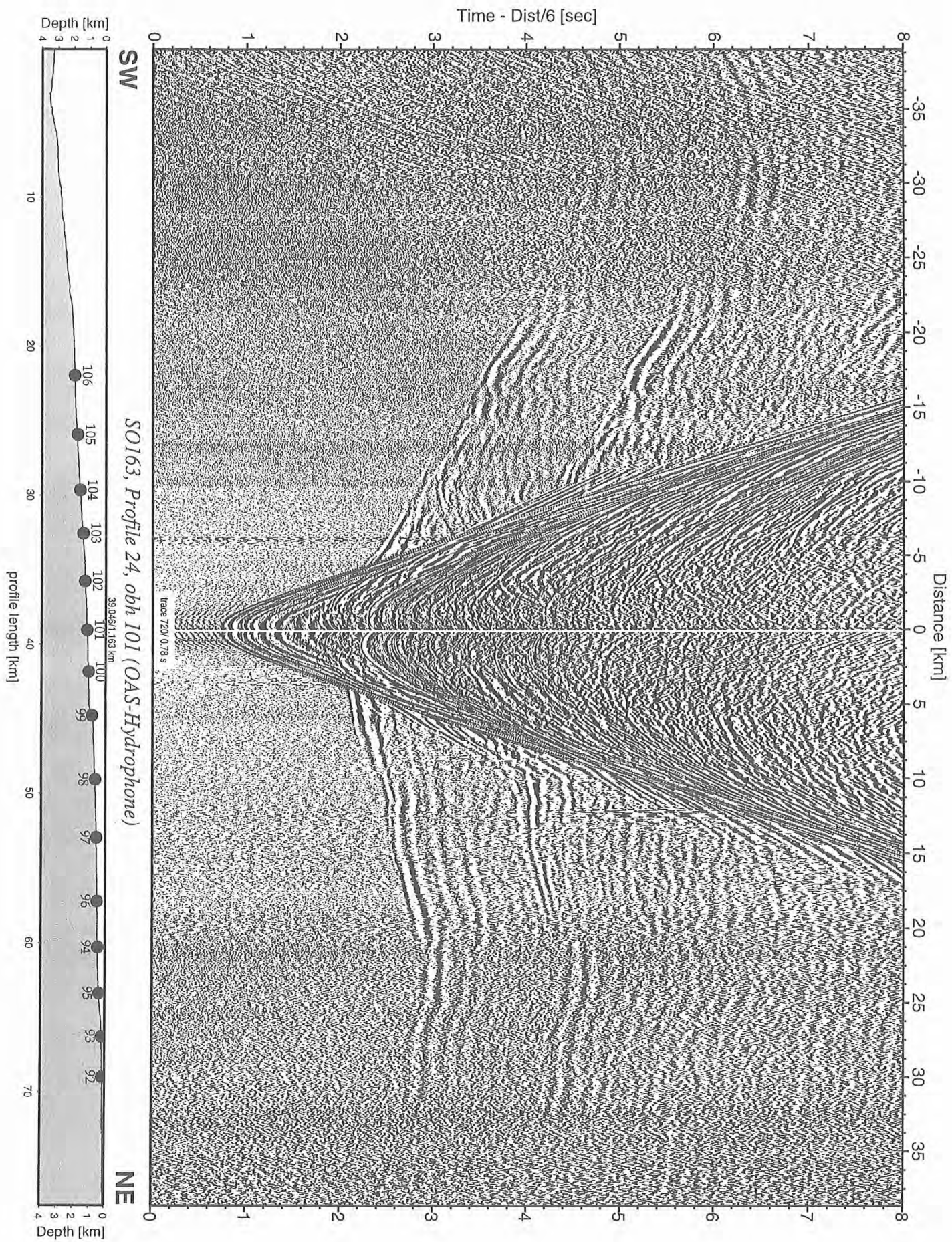


Figure 6.6.3.73: Record section from obh 101 (OAS-Hydrophone), Profile 24.

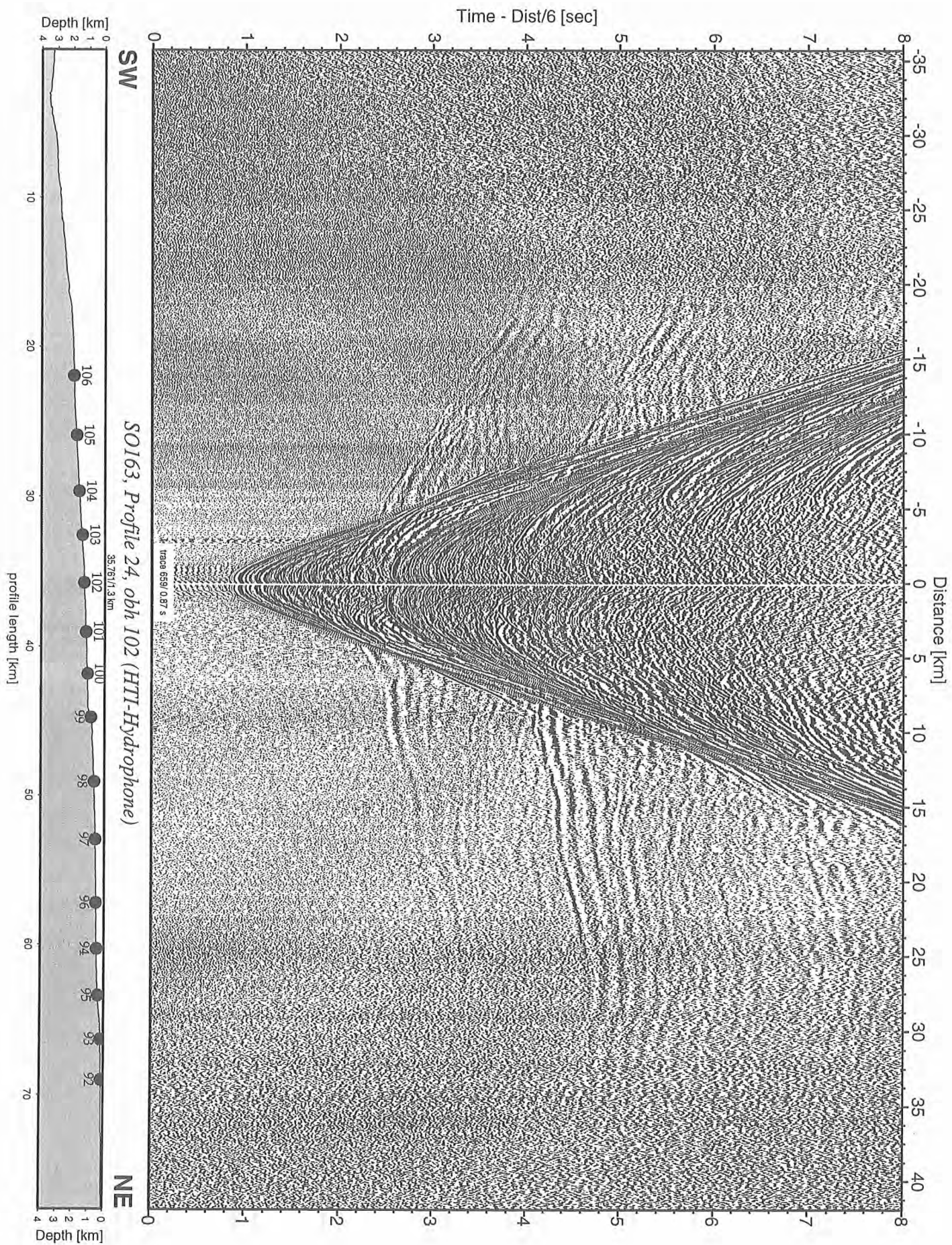


Figure 6.6.3.74: Record section from obh 102 (HTI-Hydrophone), Profile 24.

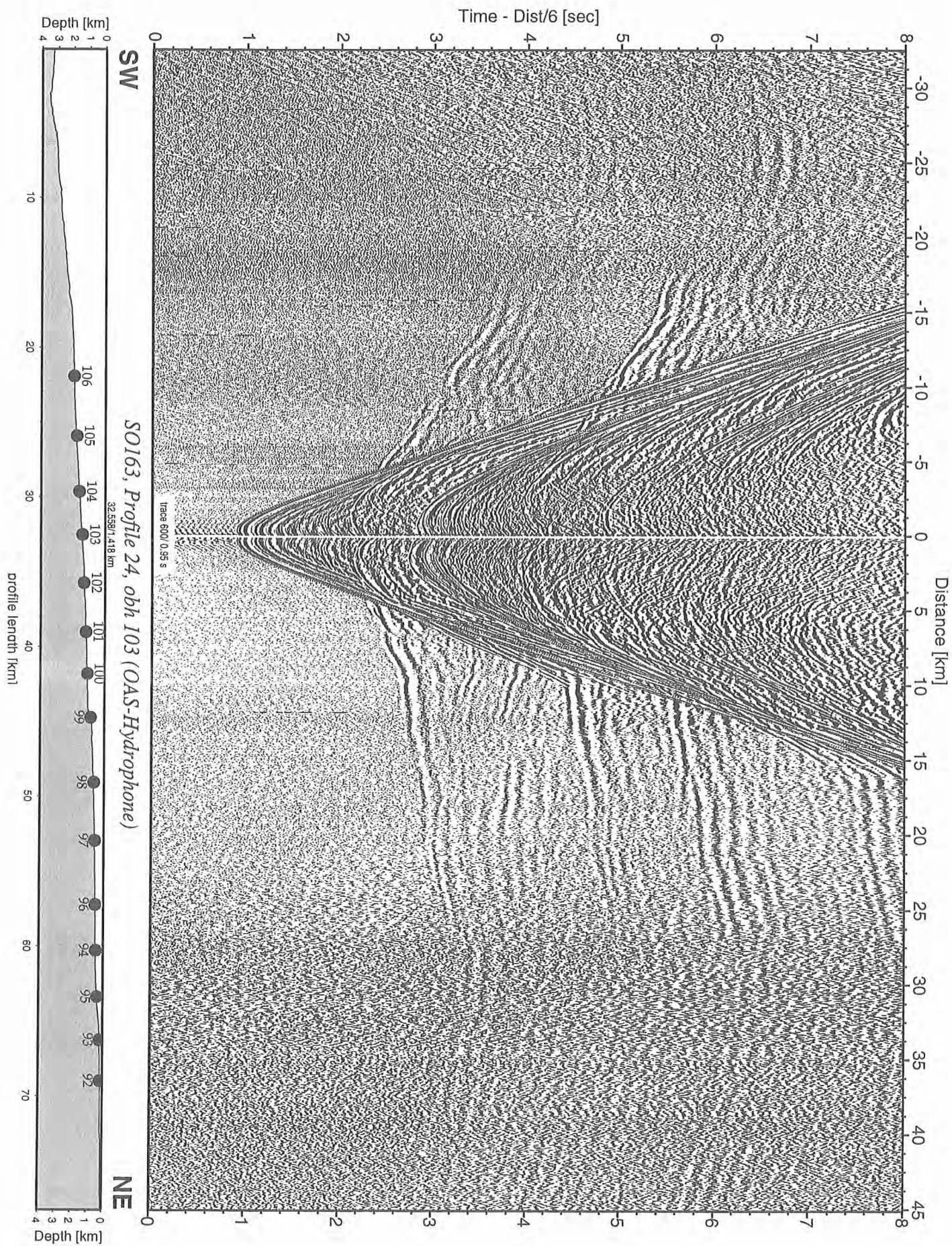


Figure 6.6.3.75: Record section from obh 103 (OAS-Hydrophone), Profile 24.

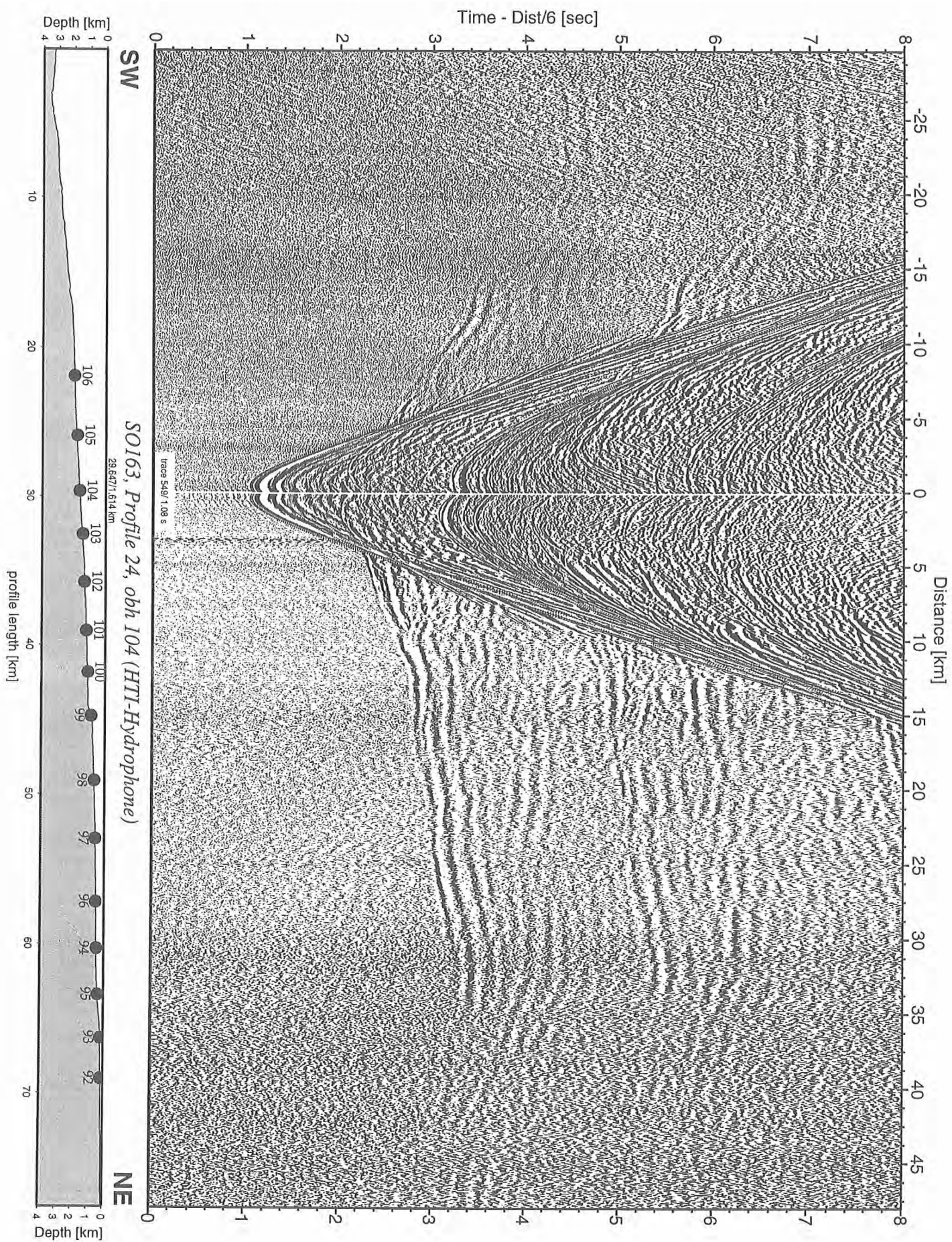


Figure 6.6.3.76: Record section from obh 104 (HTI-Hydrophone), Profile 24.

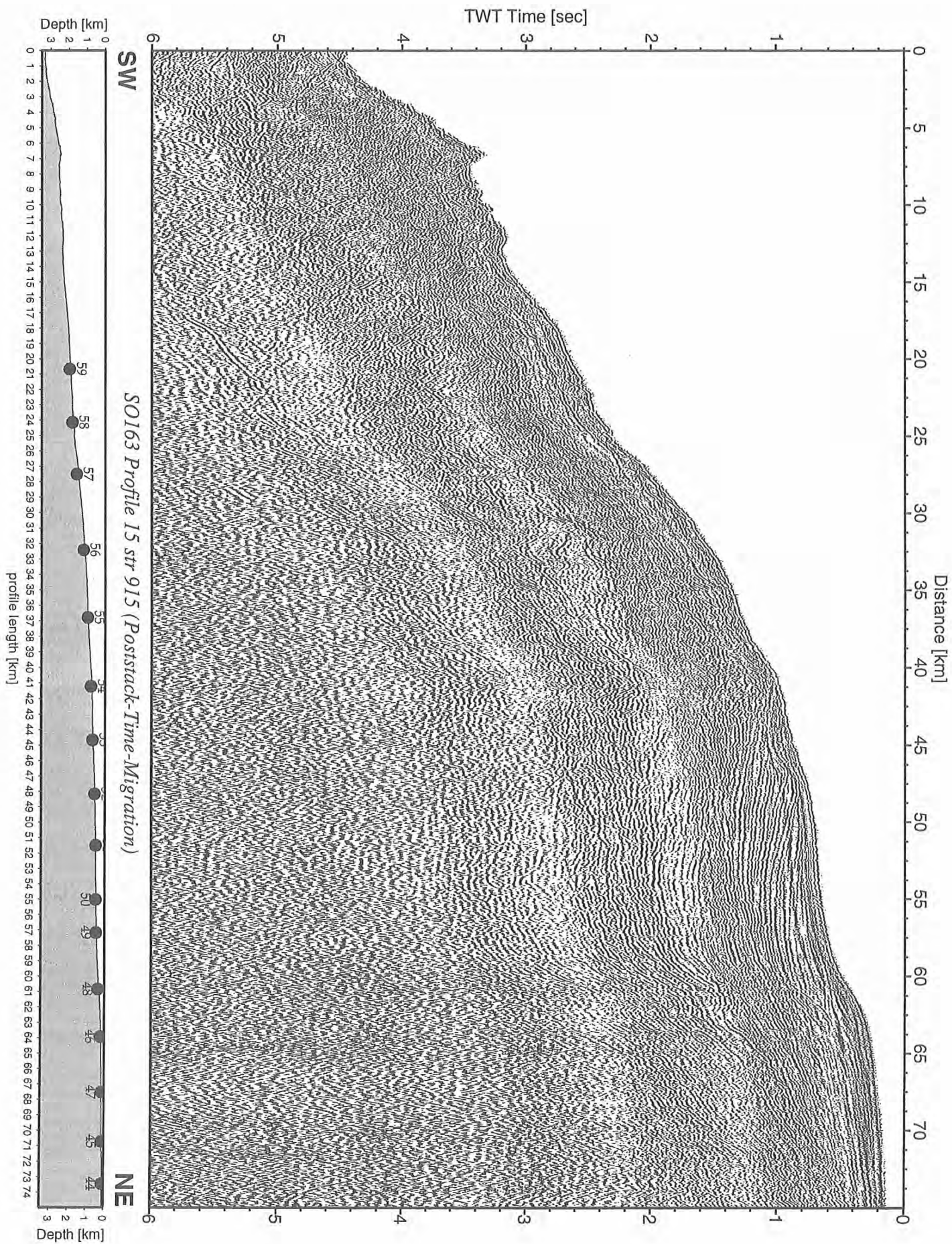


Figure 6.6.3.79: Record section from str 915 (Poststack-Time-Migration), Profile 15.

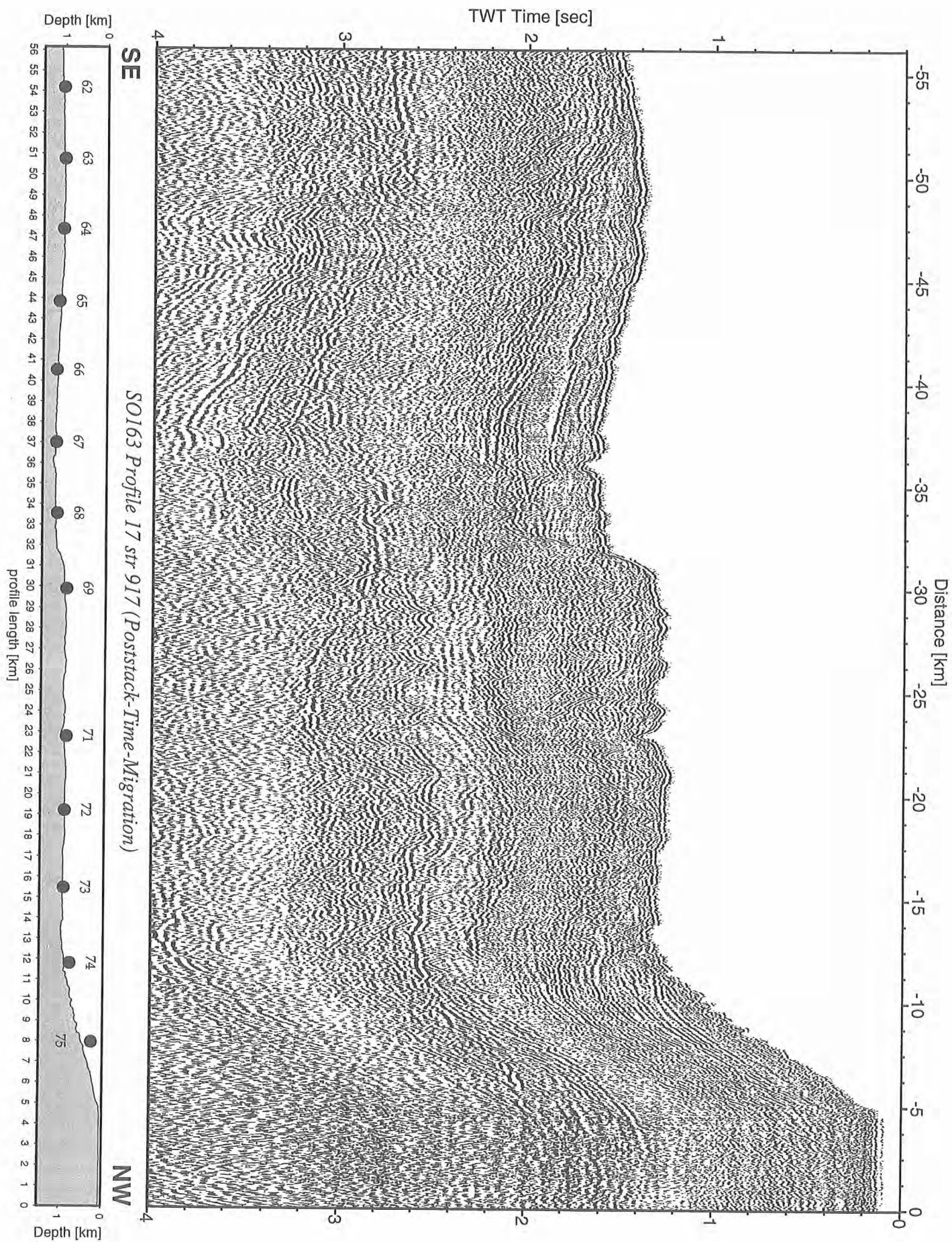


Figure 6.6.3.80: Record section from str 917 (Poststack-Time-Migration), Profile 17.

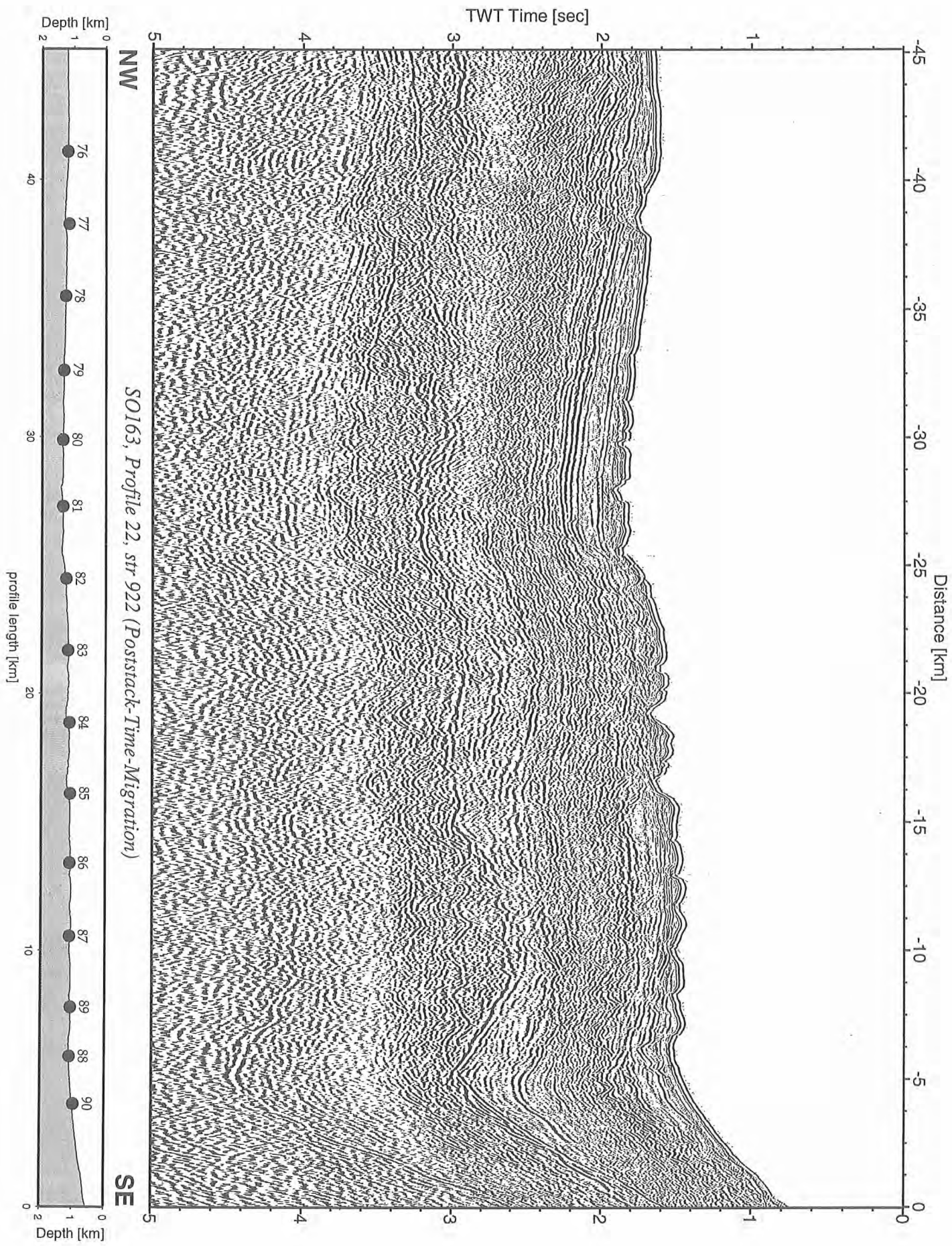


Figure 6.6.3.81: Record section from str 922 (Poststack-Time-Migration), Profile 22.

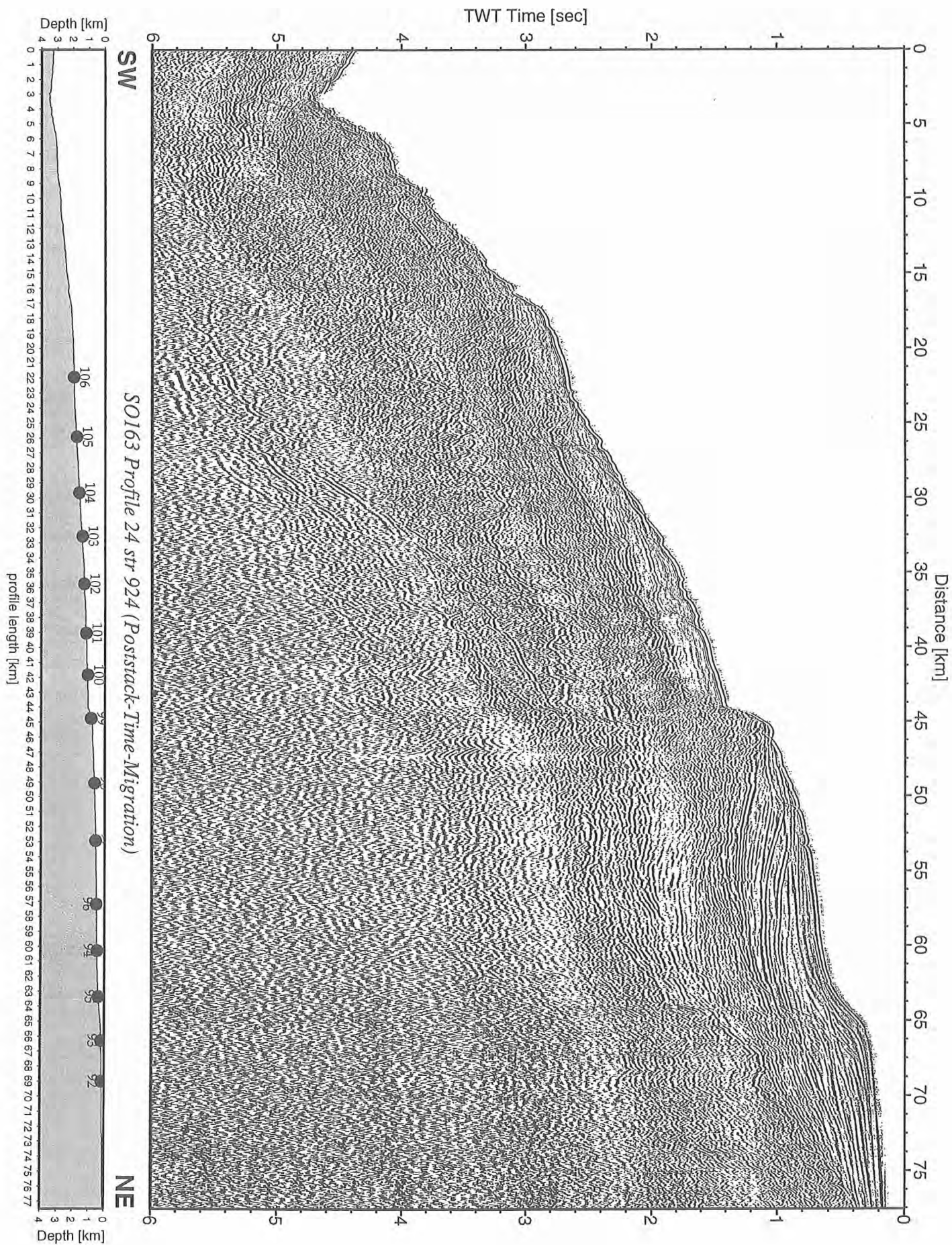


Figure 6.6.3.82: Record section from str 924 (Poststack-Time-Migration)S, Profile 24.

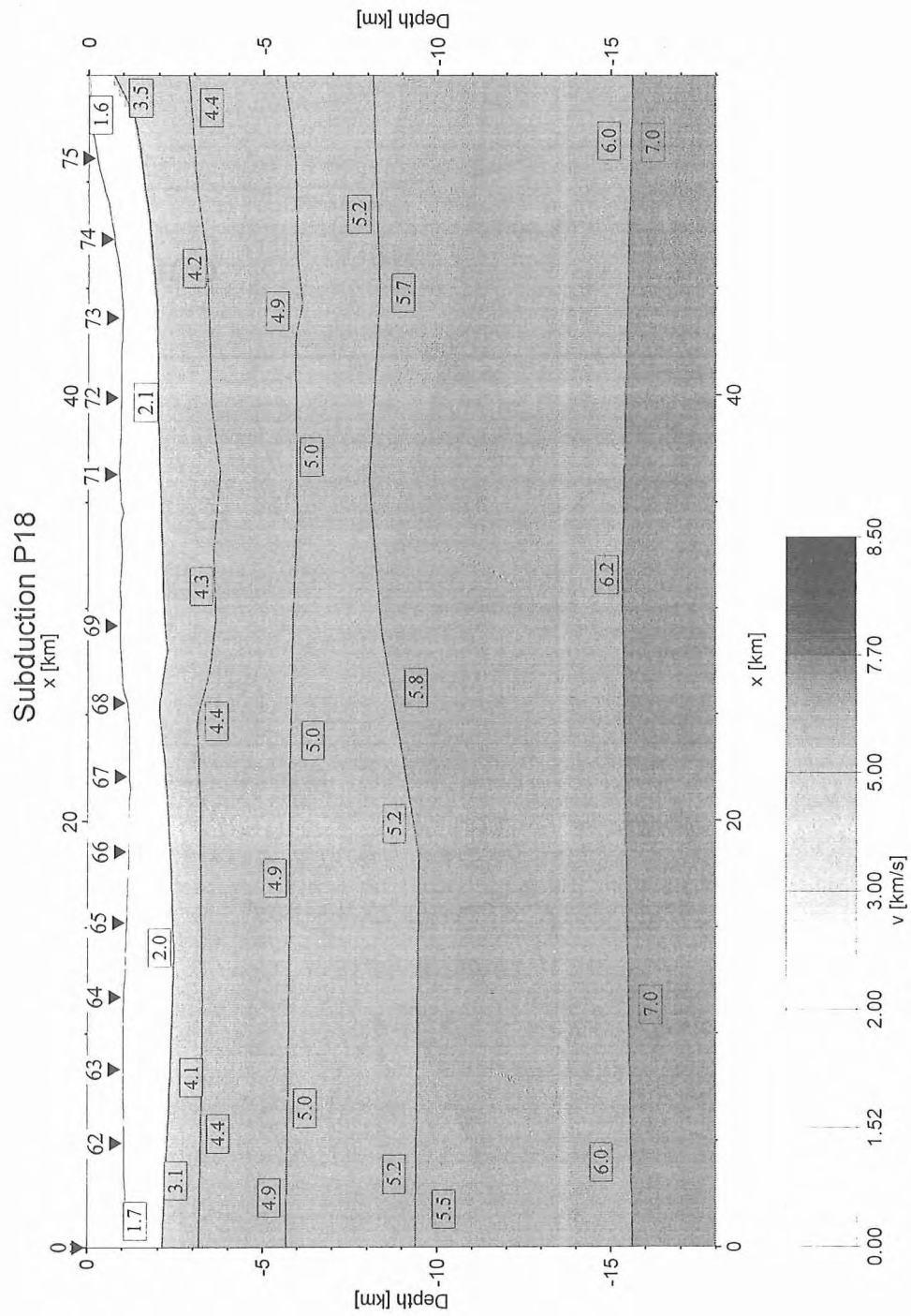


Figure 6.6.3.85: Velocity-depth model of profile P18, SO 163-2. Velocities are give in km/s.

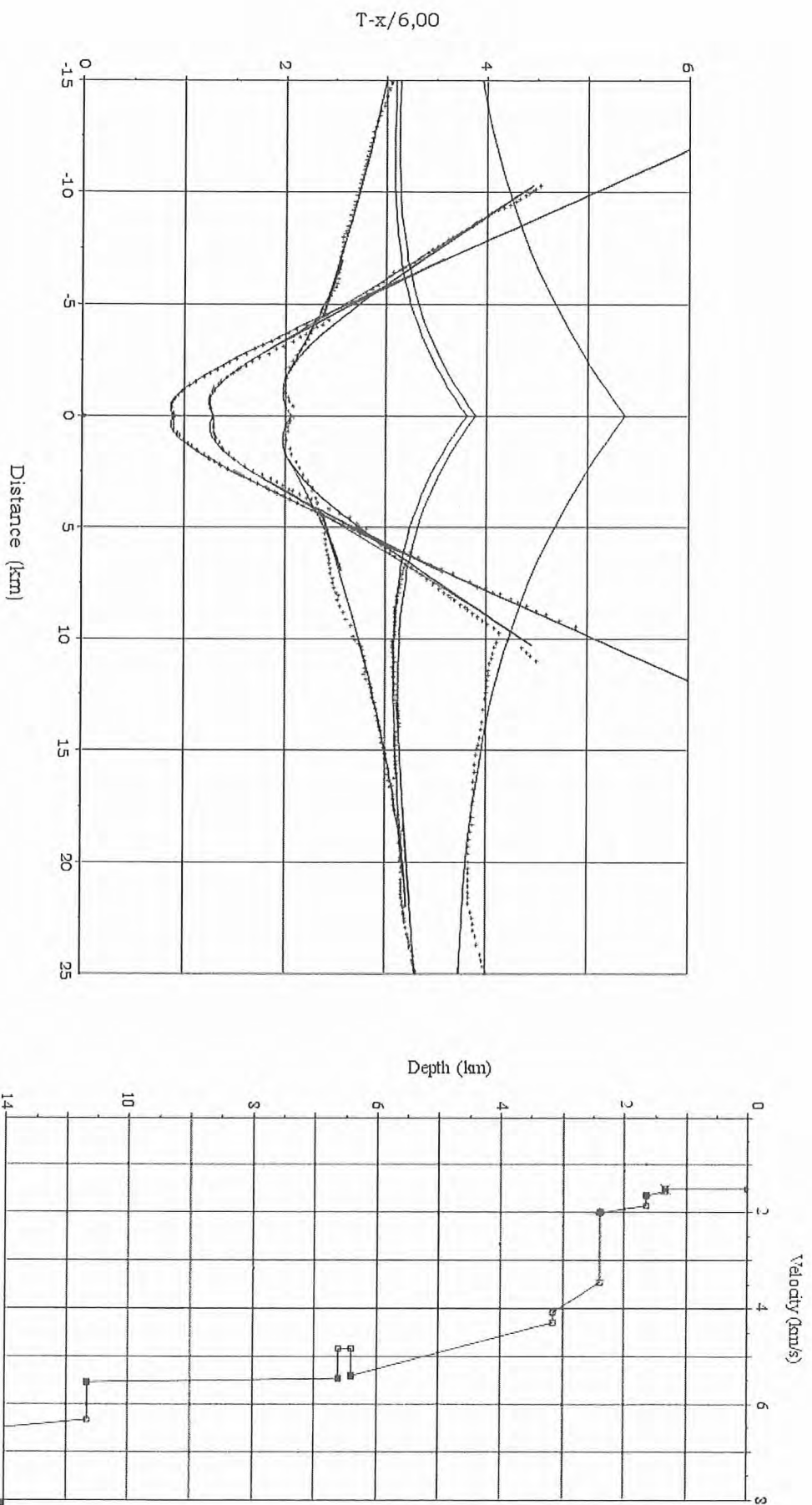


Figure 6.6.3.86: Pickings (crosses) from the record of the OBH80, Profil 22, and predicted arrivals according to the 1-D velocity model shown to the right.

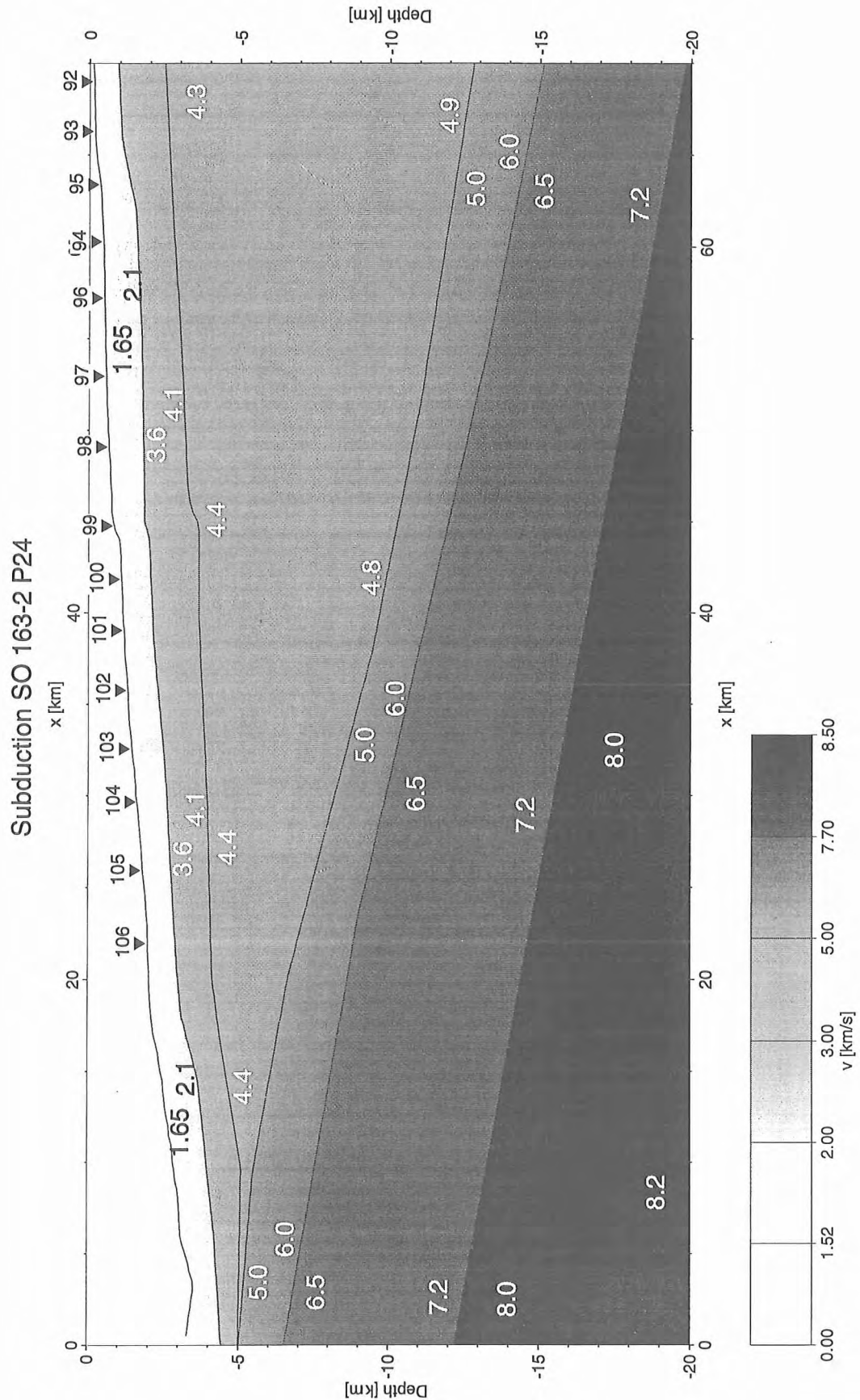


Figure 6.6.3.87: Preliminary Velocity-depth model of profile P24, SO 163-2.
Velocities are given in km/s.

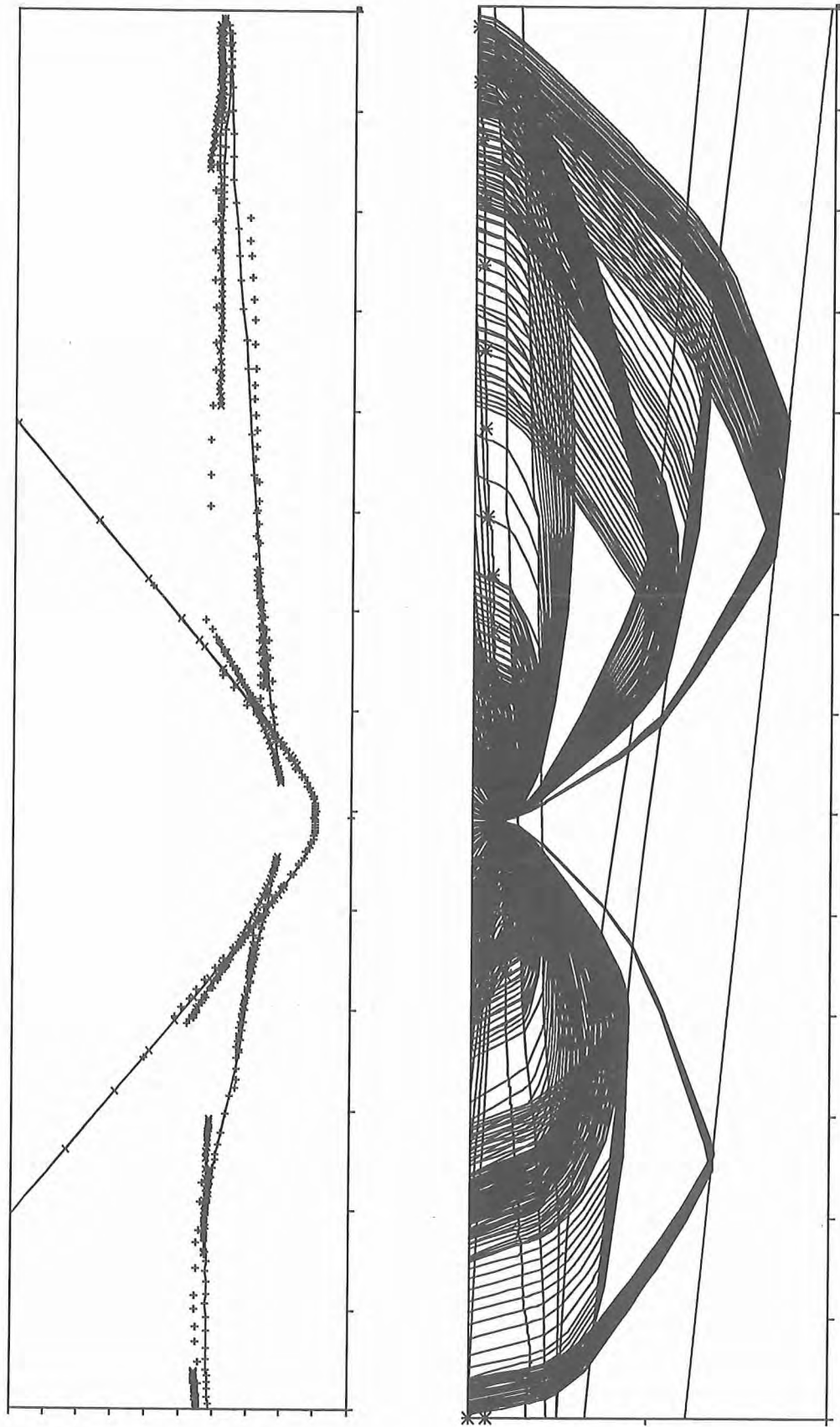


Figure 6.6.3.88: Travel times and ray penetration for OBH 106 of profile P24, SO163-2.

6.6.4 The décollement

(M. Schnabel)

The main objectives within the subproject A2 are to determine the key factors that lead to accretion or erosion at convergent margins. The coupling and mass-transfer between upper and lower plate should be studied and seismic images of the resulting tectonic units should be delivered. One target-area is the décollement, which separates the upper and lower plate.

The décollement zone plays an important role in the volatile budget of an subduction zone. It is the focus of mechanical compaction, which releases pore water. New volatile reservoirs are generated by the alteration of sediments. Off Costa Rica, the décollement has in places a high reflectivity, which implies a massive dewatering within about 4 km of the deformation front (Shipley & Moore, 1986). In this zone, porosity decreases from perhaps 50 to 25%. Based on ODP Leg 170, Kimura et al. (1997) showed that the décollement reflection is a result of a very sharp boundary in physical properties, with higher porosities below than above, indicating marked mechanical decoupling. The plate boundary décollement off Costa Rica is structurally divisible into an upper brittle-fracture-dominated domain overlying a lower, ductile domain (Tobin et al., 2001).

During SO163, we wanted to collect new seismic wide-angle data, which should enable us to describe the décollement reflections in greater detail. We are especially interested in the lithological properties, which may be used to analyse and quantify the amount of sediment incorporated in the subduction channel and indicate potential locations of porosity changes and fluid losses. Further on, we hope to deliver important informations for the subproject A3, which is concerned about diagenetic reactions and alterations of sediment.

Along profile line SO81-5, which showed one of the best reflections of the décollement in this region (Hinz et al., 1996), we deployed a profile with 10 OBH and 5 OBS across the lower slope (Figure 6.6.4.1). After a first look at the streamer recording of this profile, we decided to deploy two strike-profiles, the first 2 km, and the second 3.2 km behind the trench. Selected record sections are given in Figures 6.6.3.2 to 6.6.3.41.

The data quality on these profiles were very high on profile 26 up to 29. All hydrophones were recording and showed strong reflections. Only 2 seismometer delivered weak data.

On profile 30 and 31, there were problems with two stations. OBH 124 had not enough battery power, and OBH 130 recorded just electronic noise. But for the rest of the stations, the data quality was very satisfying. No onboard velocity modelling was done with these stations.

Figures 6.6.4.39 up to 6.6.4.42 show streamer recordings. All three profiles were shot with the G-Gun array and a shot intervall of 20 seconds. The data is poststack time migrated.

On profile 28, the décollement is clearly imaged up to 5 km behind the trench. Below the detachment, a graben-like-structure can be seen.

On the strike-profiles, the décollement can be found at 4.4 seconds TWT. The oceanic crust shows again large fault planes. These fault planes might be pathways for fluids rising from the deeper crust, explaining the high reflectivity of the décollement above. A further analysis of the amplitudes might give answer to this question.

SO-163-2: Profiles P26-P31 (Decollement)

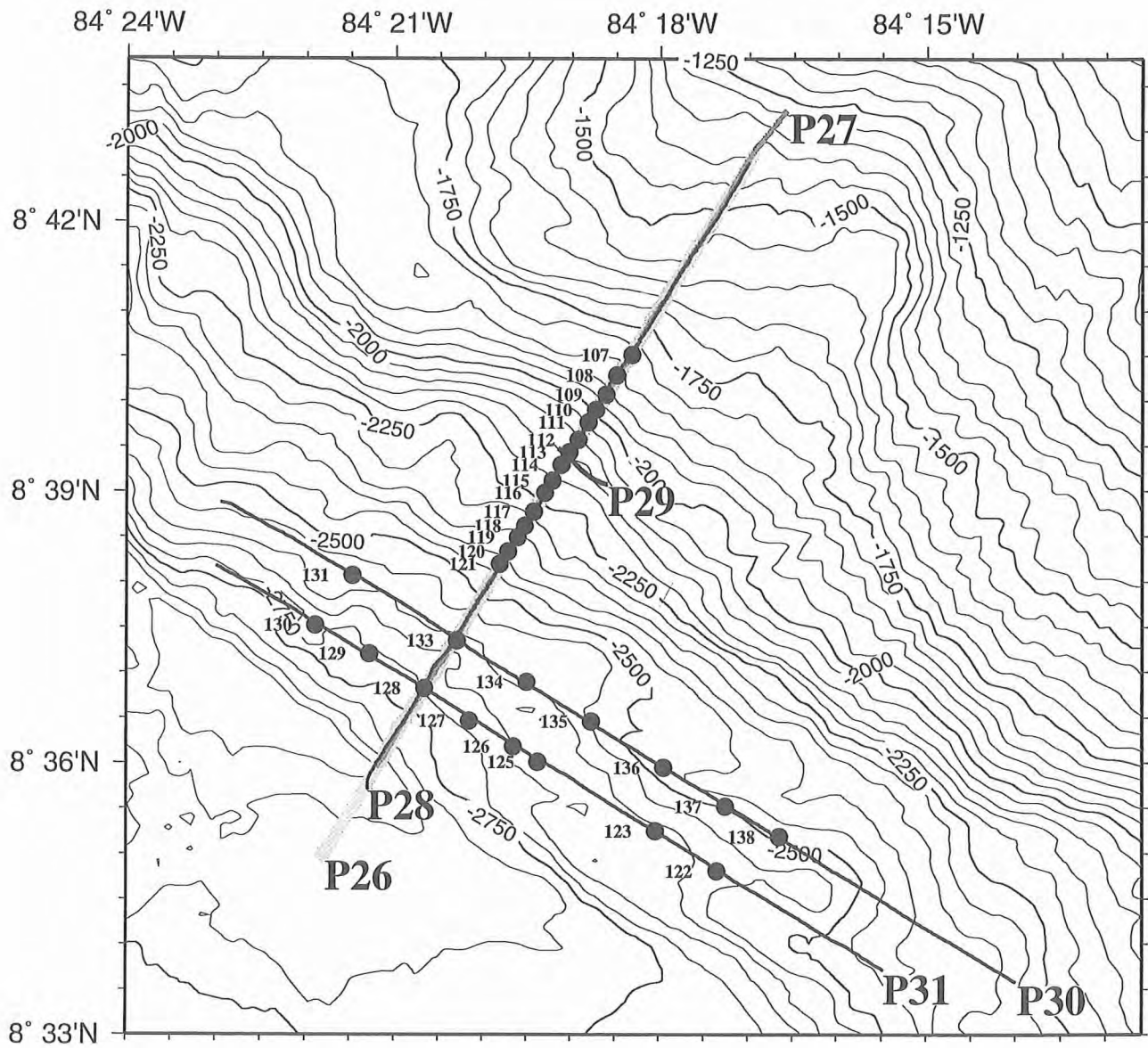


Figure 6.6.4.1: Location map of profiles P26-P31 with 50 m isolines. Stations OBH107-OBH121 (lines P26-P29, relocated) and OBH122-OBH138 (lines P30 and P31) are superimposed.

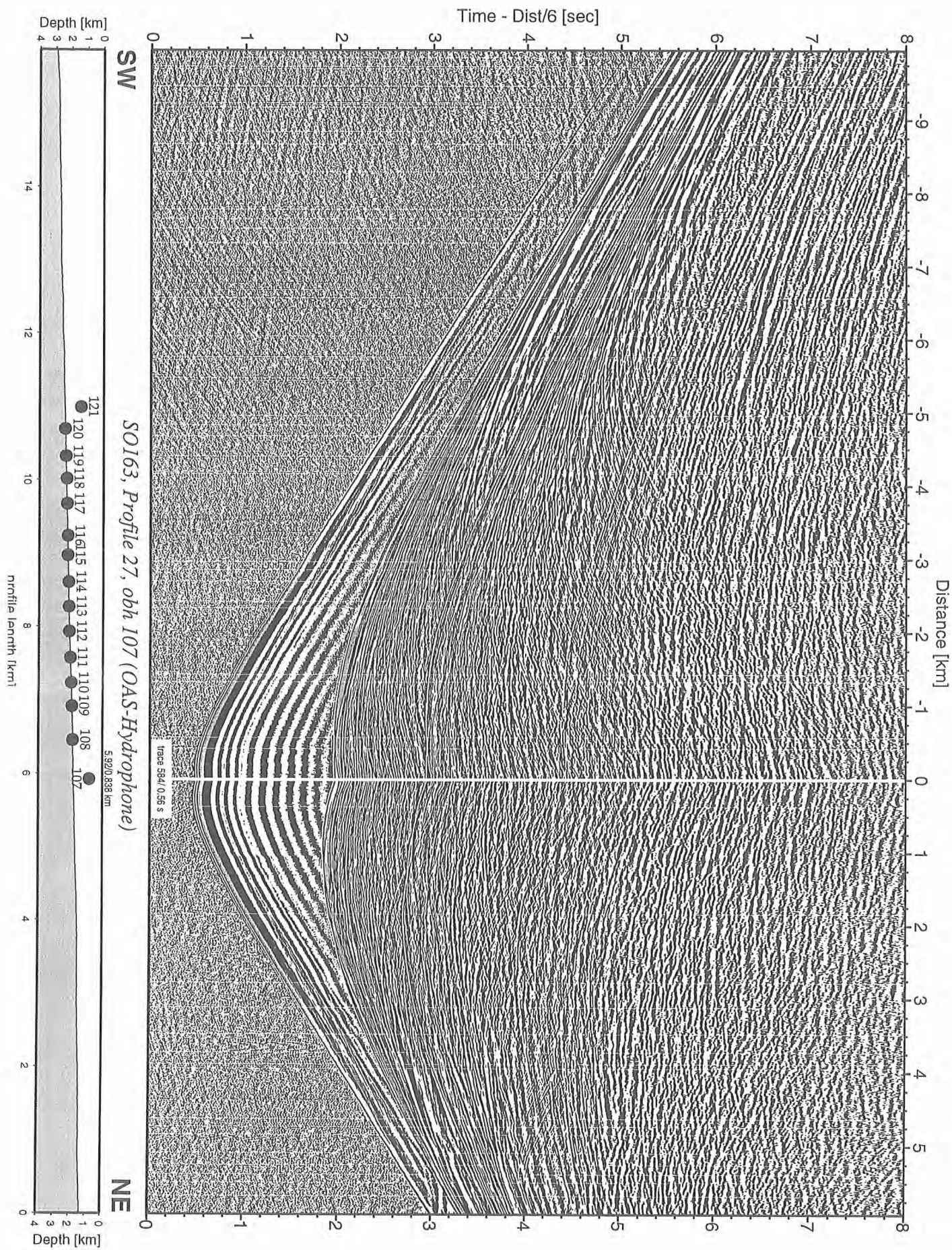


Figure 6.6.4.2: Record section from obh 107 (OAS-Hydrophone), Profile 27.

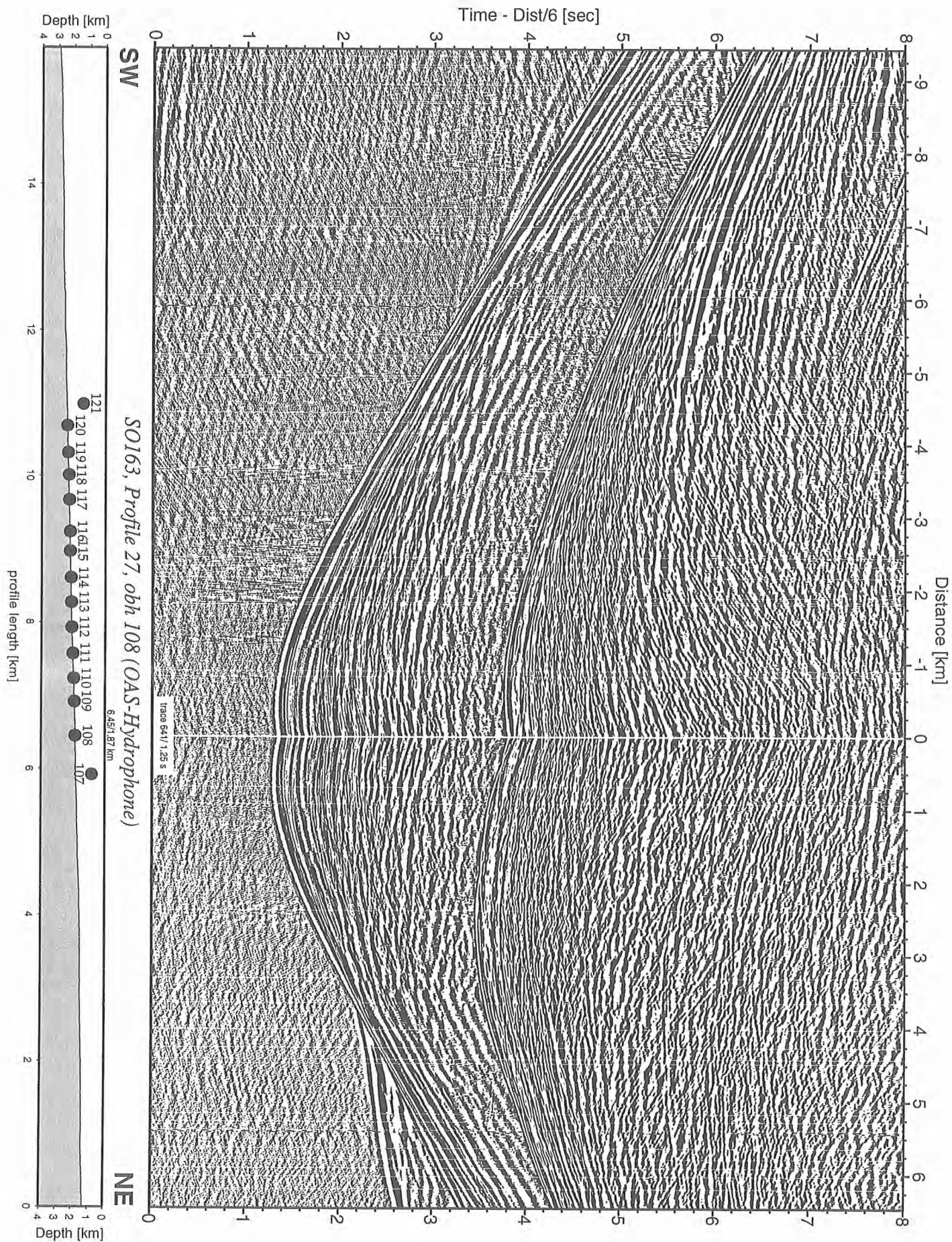


Figure 6.6.4.3: Record section from obh 108 (OAS-Hydrophone), Profile 27.

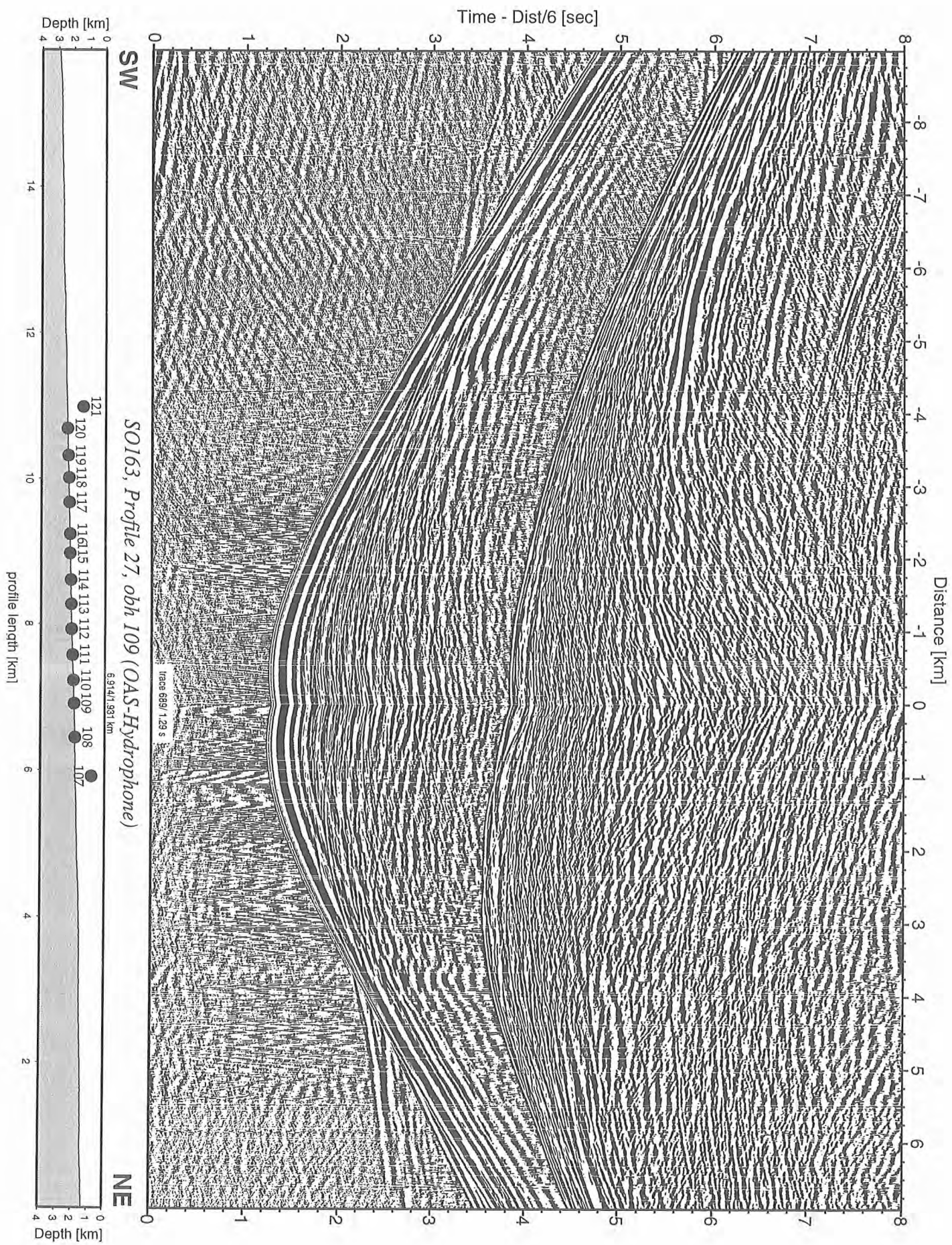


Figure 6.6.4.4: Record section from obh 109 (OAS-Hydrophone), Profile 27.

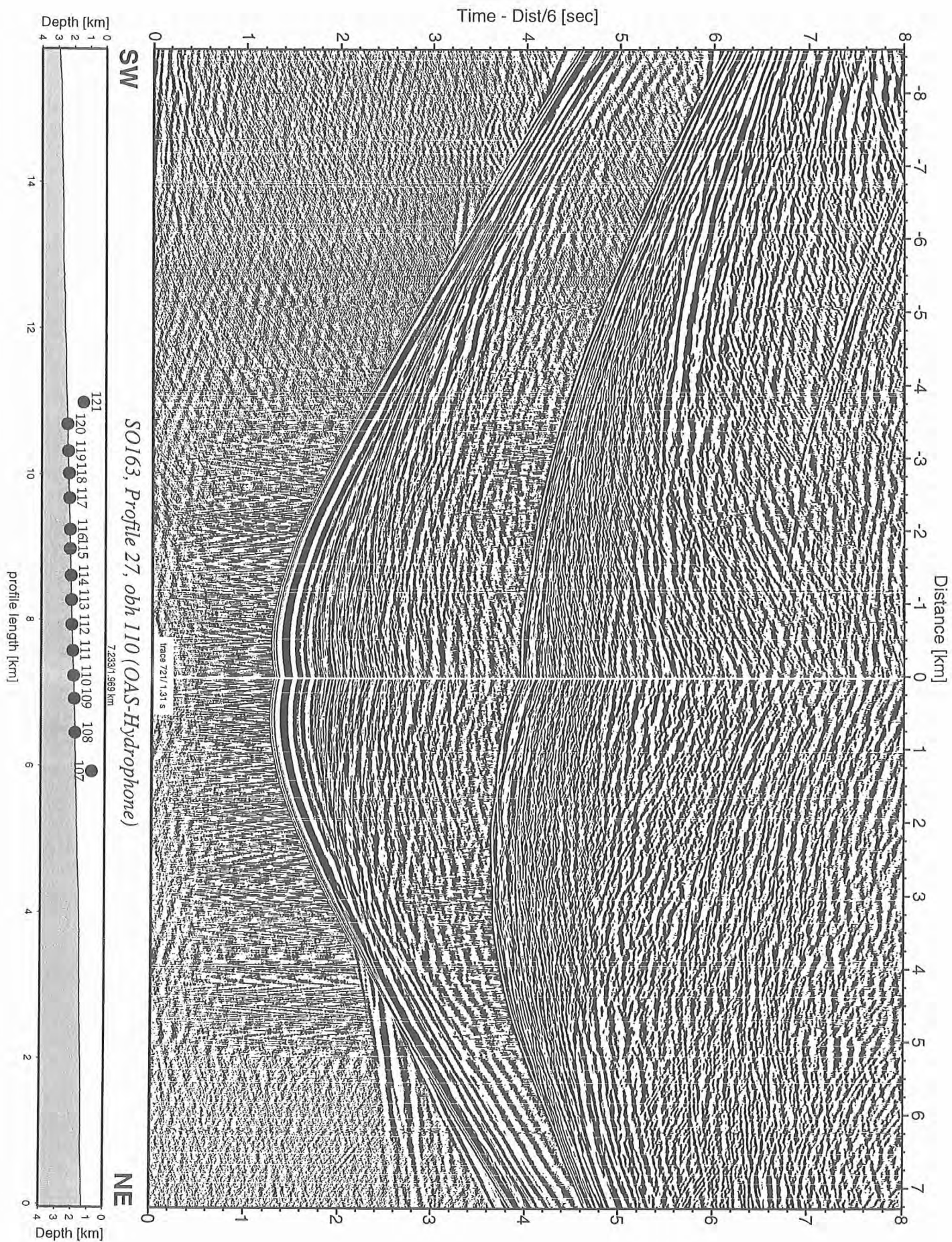


Figure 6.6.4.5: Record section from obh 110 (OAS-Hydrophone), Profile 27.

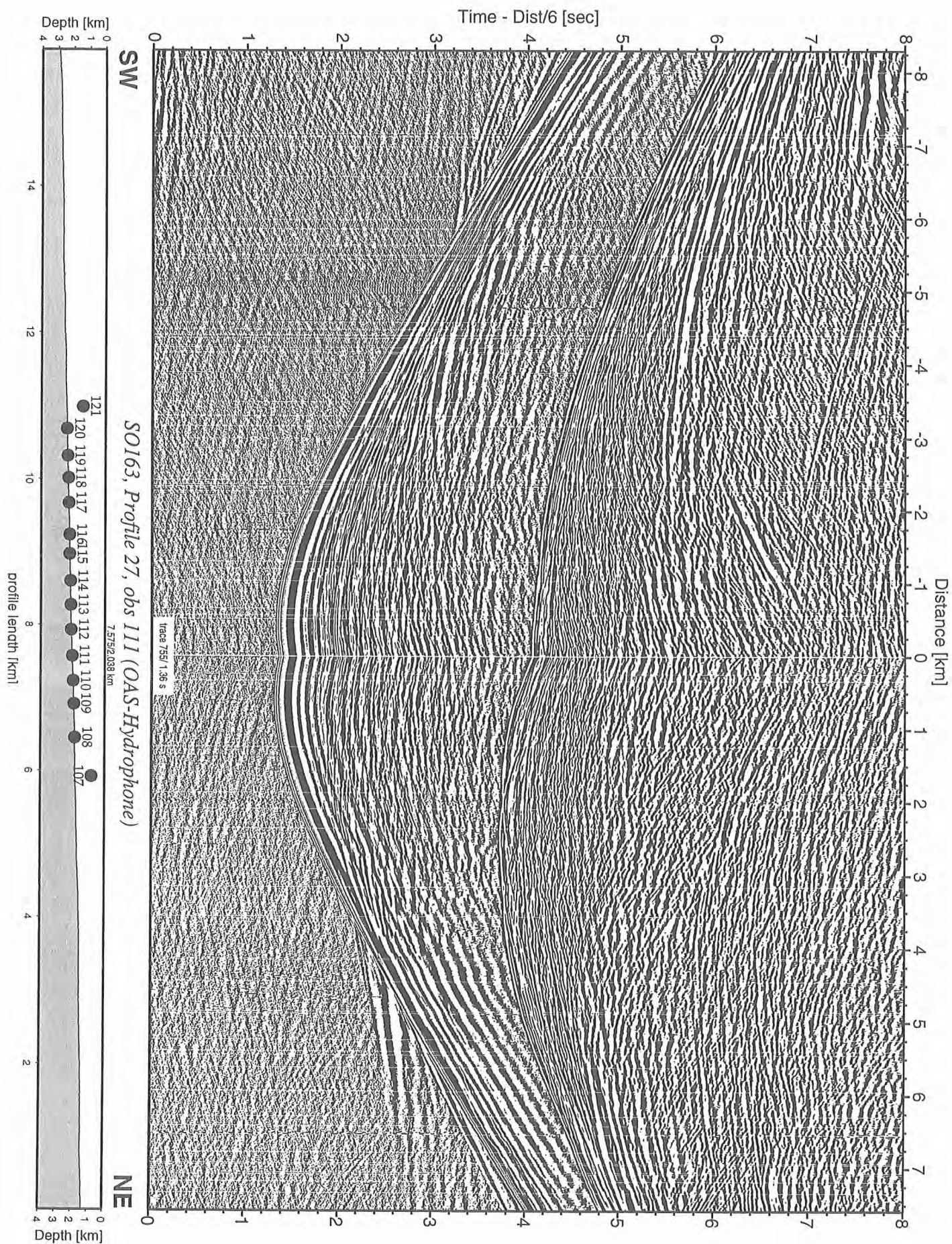


Figure 6.6.4.6: Record section from obs 111 (OAS-Hydrophone), Profile 27.

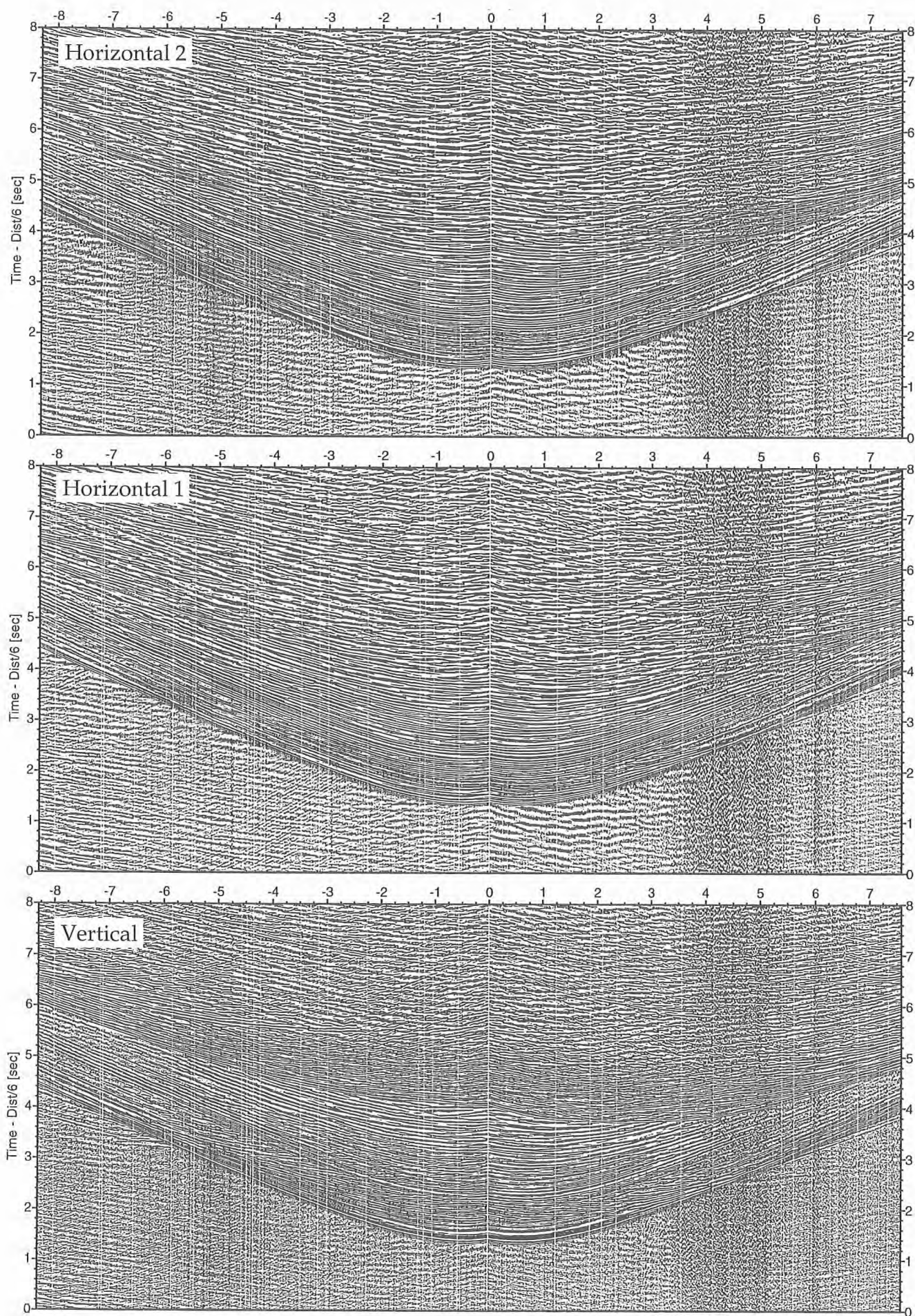


Figure 6.6.4.7: Record sections from obs 111 (Owen-30Hz), SO163, Profile 27.

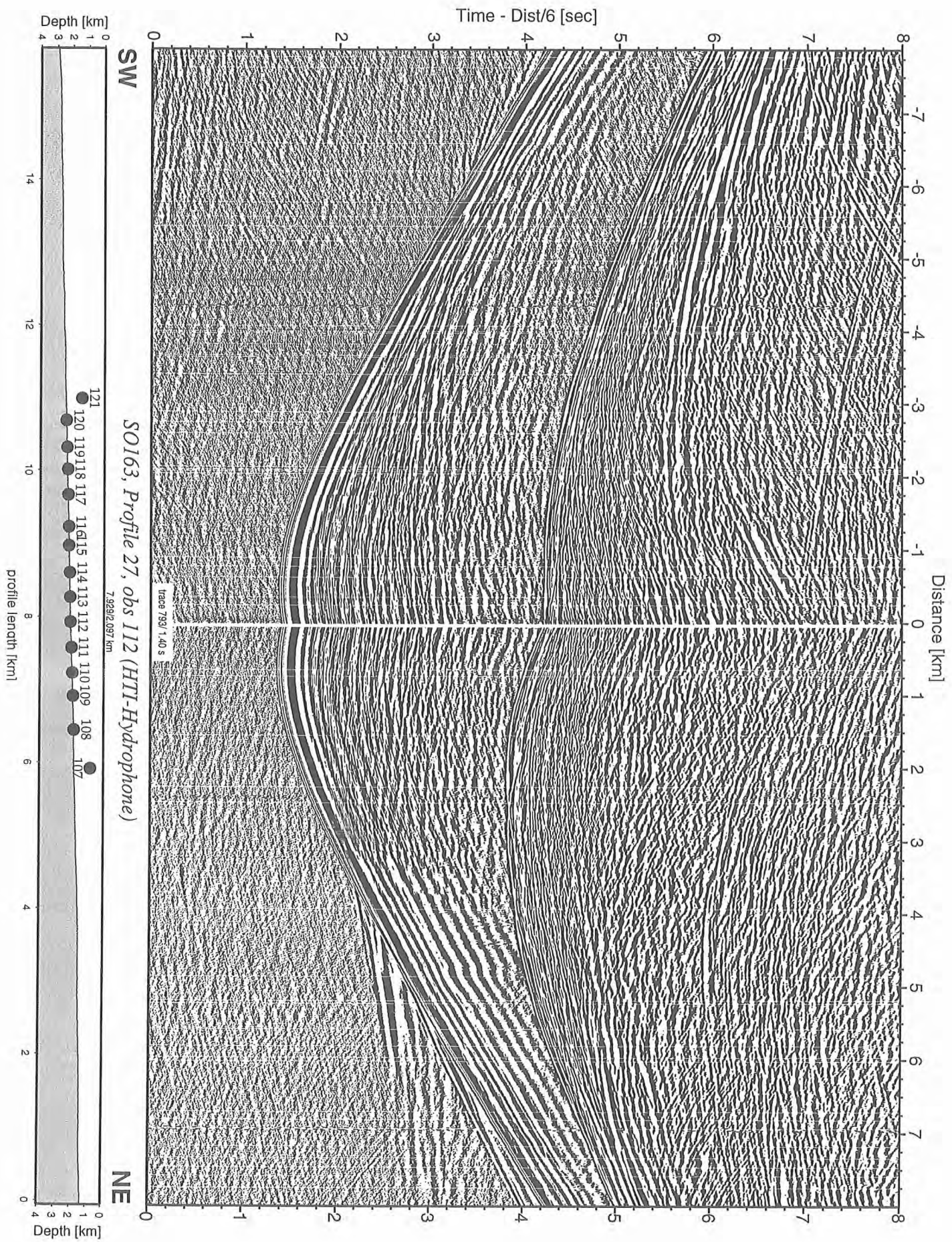


Figure 6.6.4.8: Record section from obs 112 (HTI-Hydrophone), Profile 27.

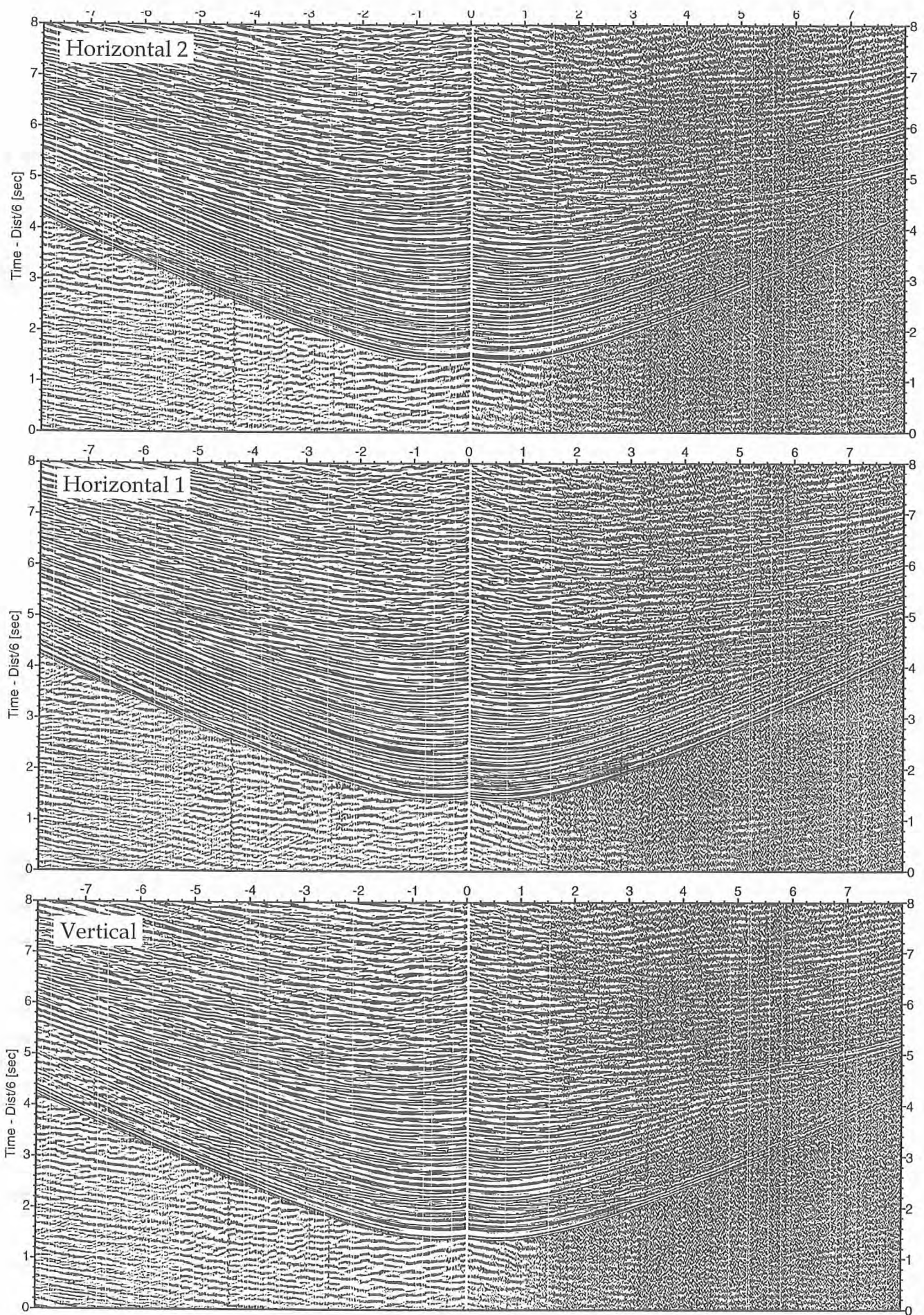


Figure 6.6.4.9: Record sections from obs 112 (Owen-4.5Hz), SO163, Profile 27.

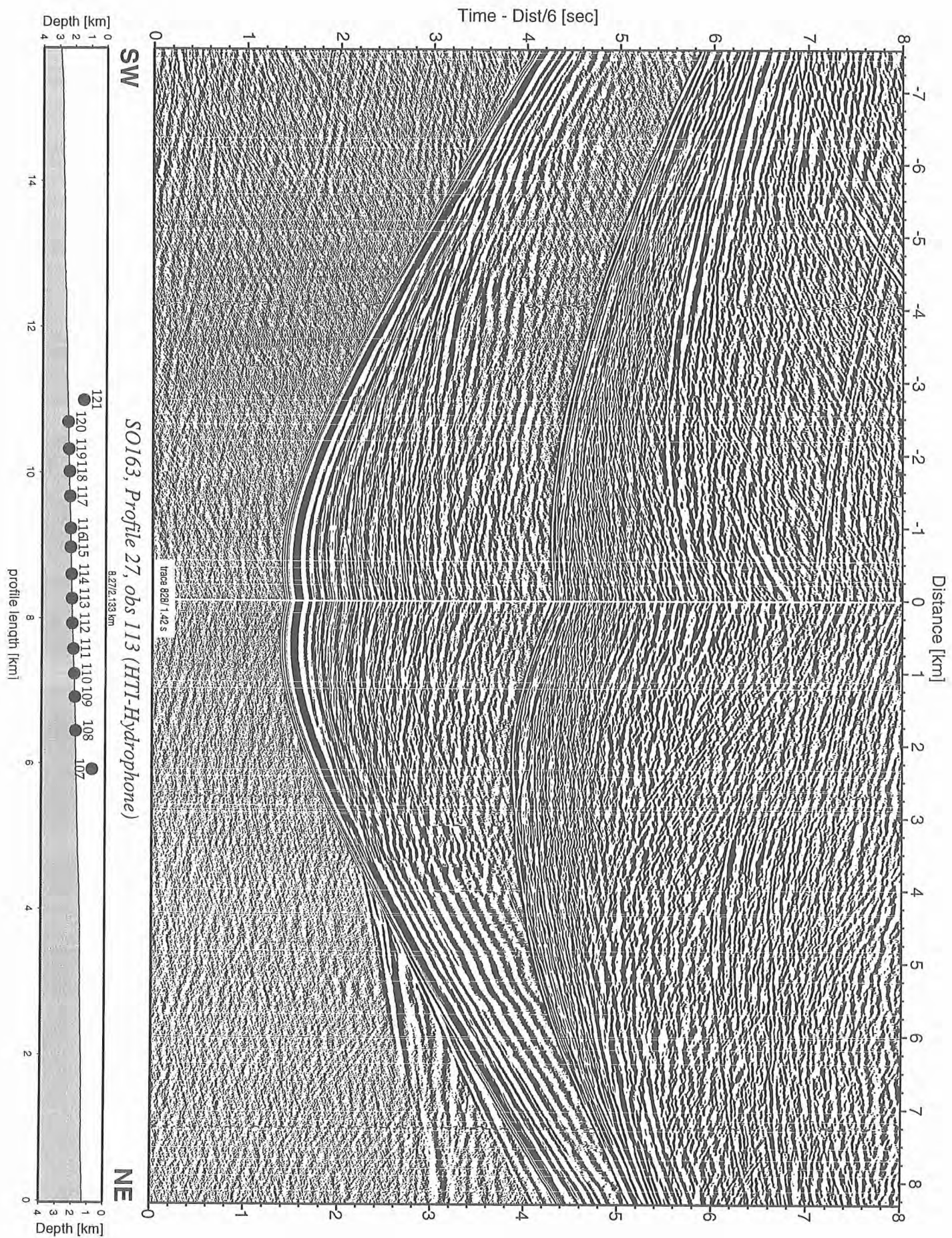


Figure 6.6.4.10: Record section from obs 113 (HTI-Hydrophone), Profile 27.

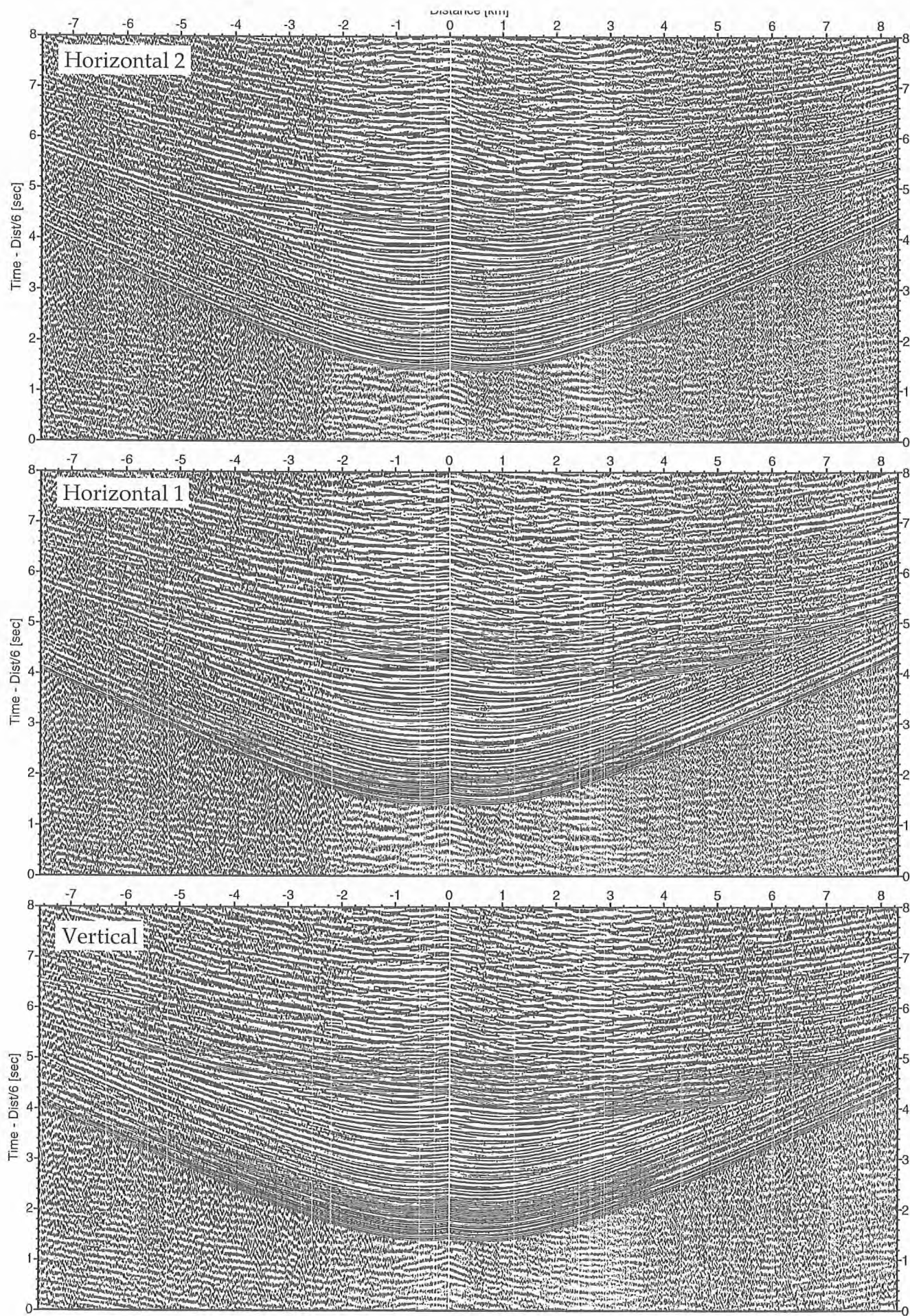


Figure 6.6.4.11: Record sections from obs 113 (Owen-4.5Hz), SO163, Profile 27.

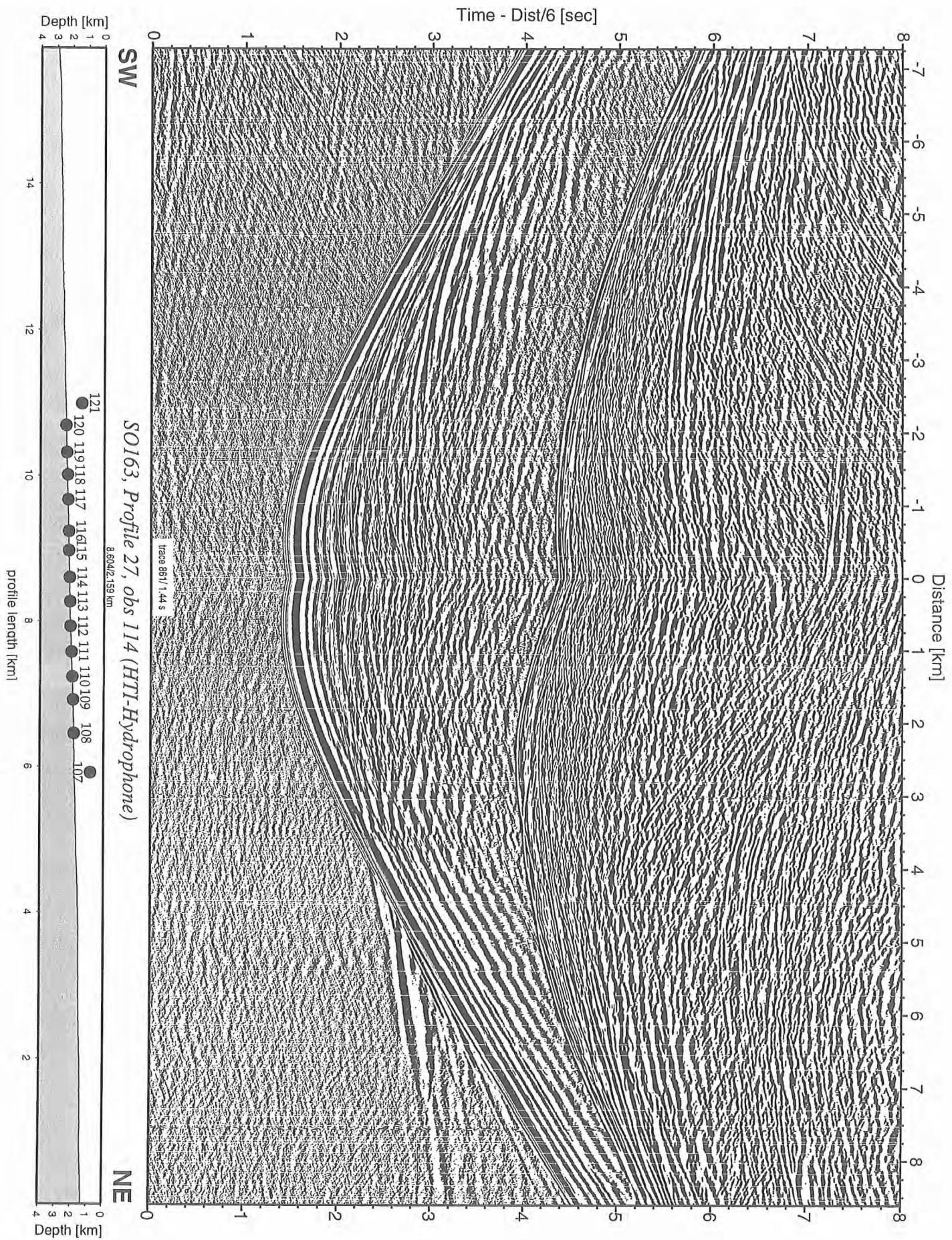


Figure 6.6.4.12: Record section from obs 114 (HTI-Hydrophone), Profile 27.

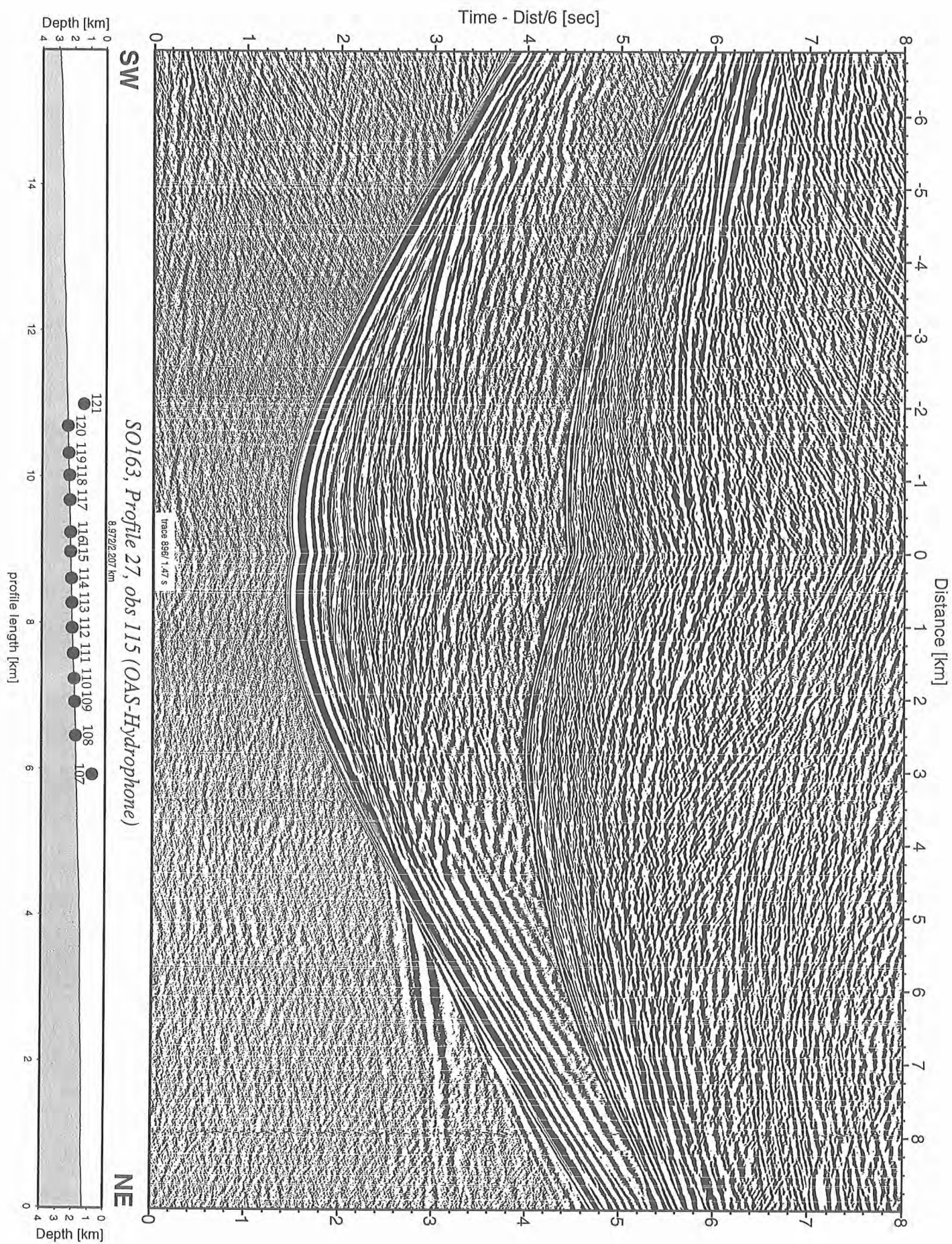


Figure 6.6.4.13: Record section from obs 115 (OAS-Hydrophone), Profile 27.

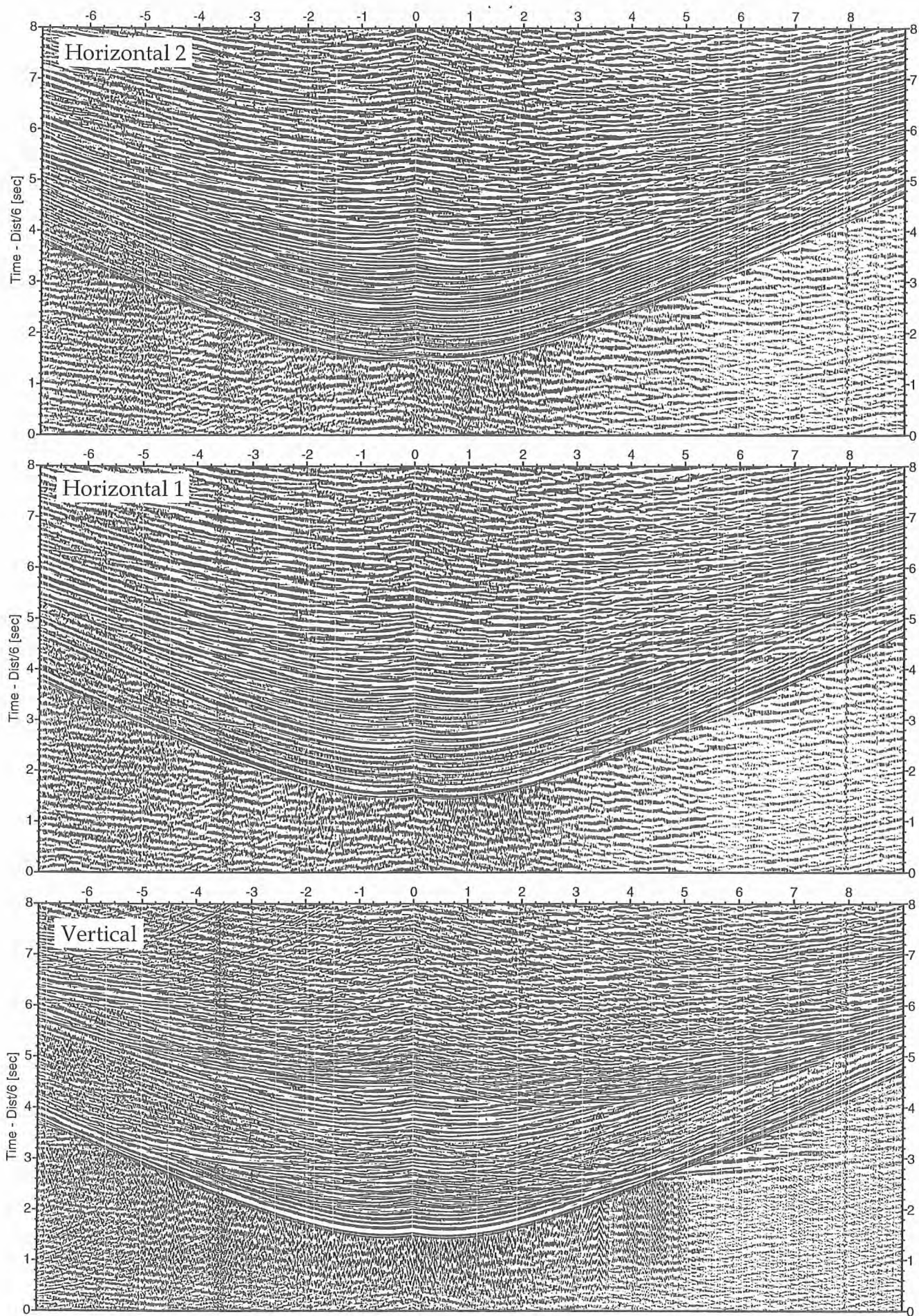


Figure 6.6.4.14: Record sections from obs 115 (Owen-30Hz), SO163, Profile 27.

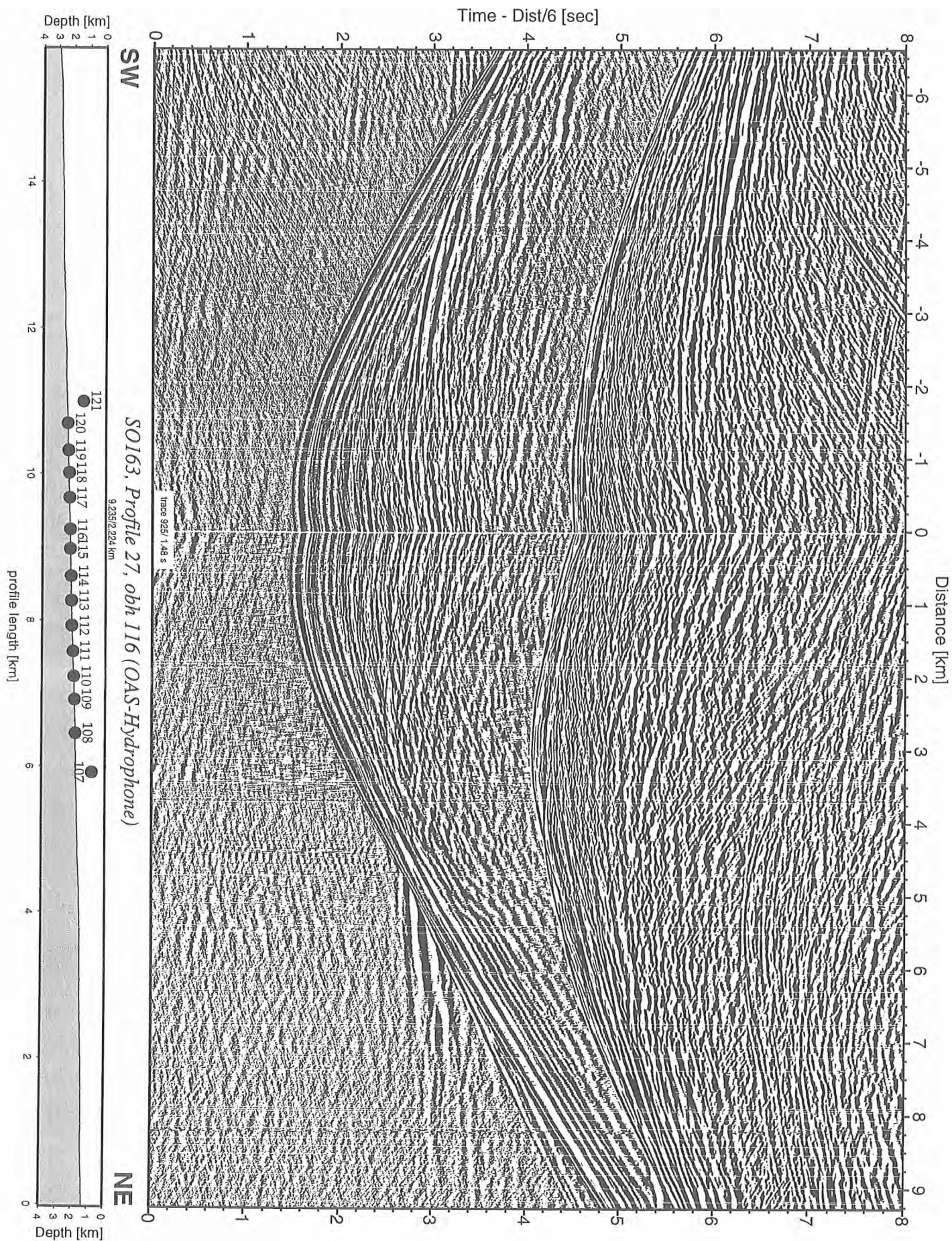


Figure 6.6.4.15: Record section from obh 116 (OAS-Hydrophone), Profile 27.

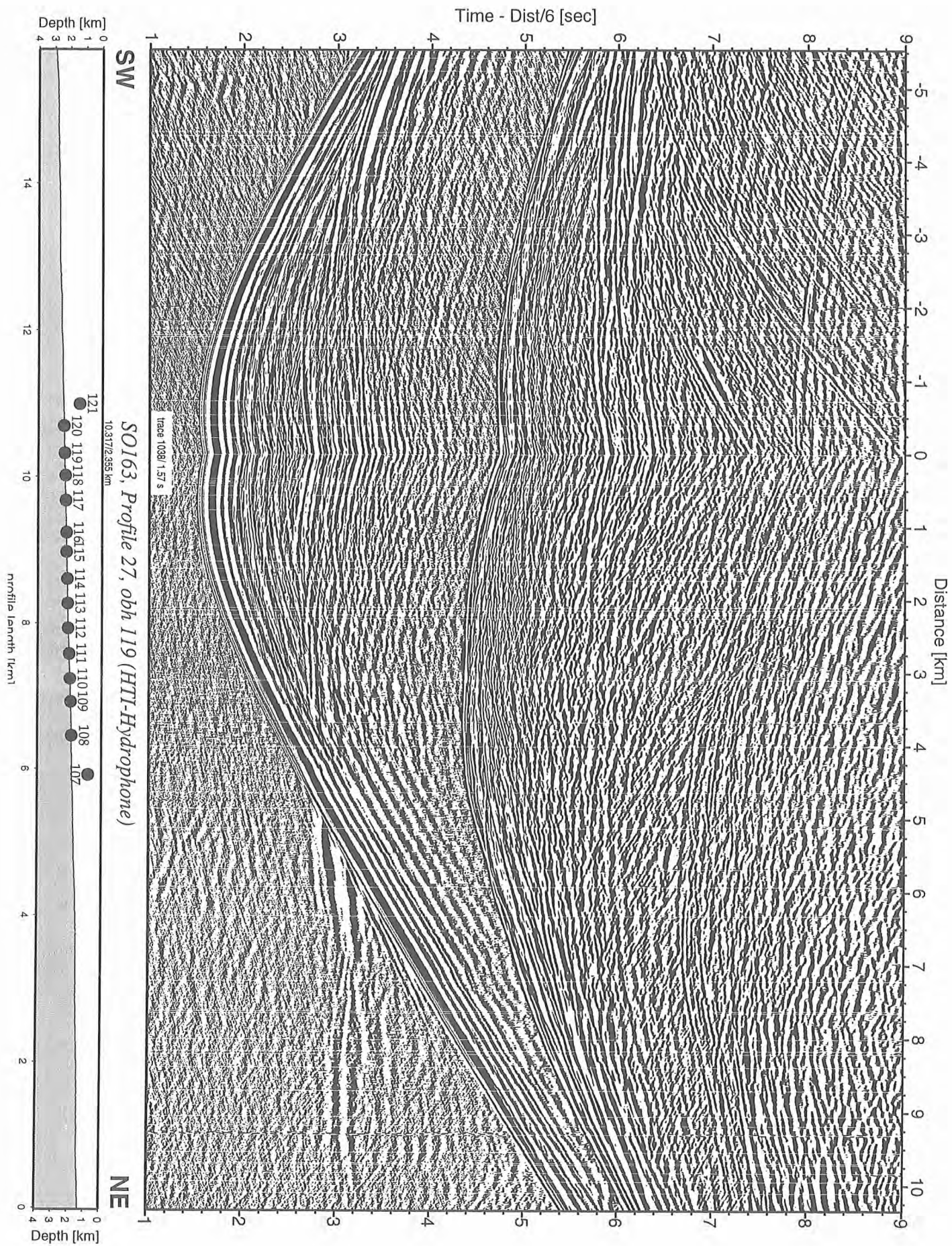


Figure 6.6.4.18: Record section from obh 119 (HTI-Hydrophone), Profile 27.

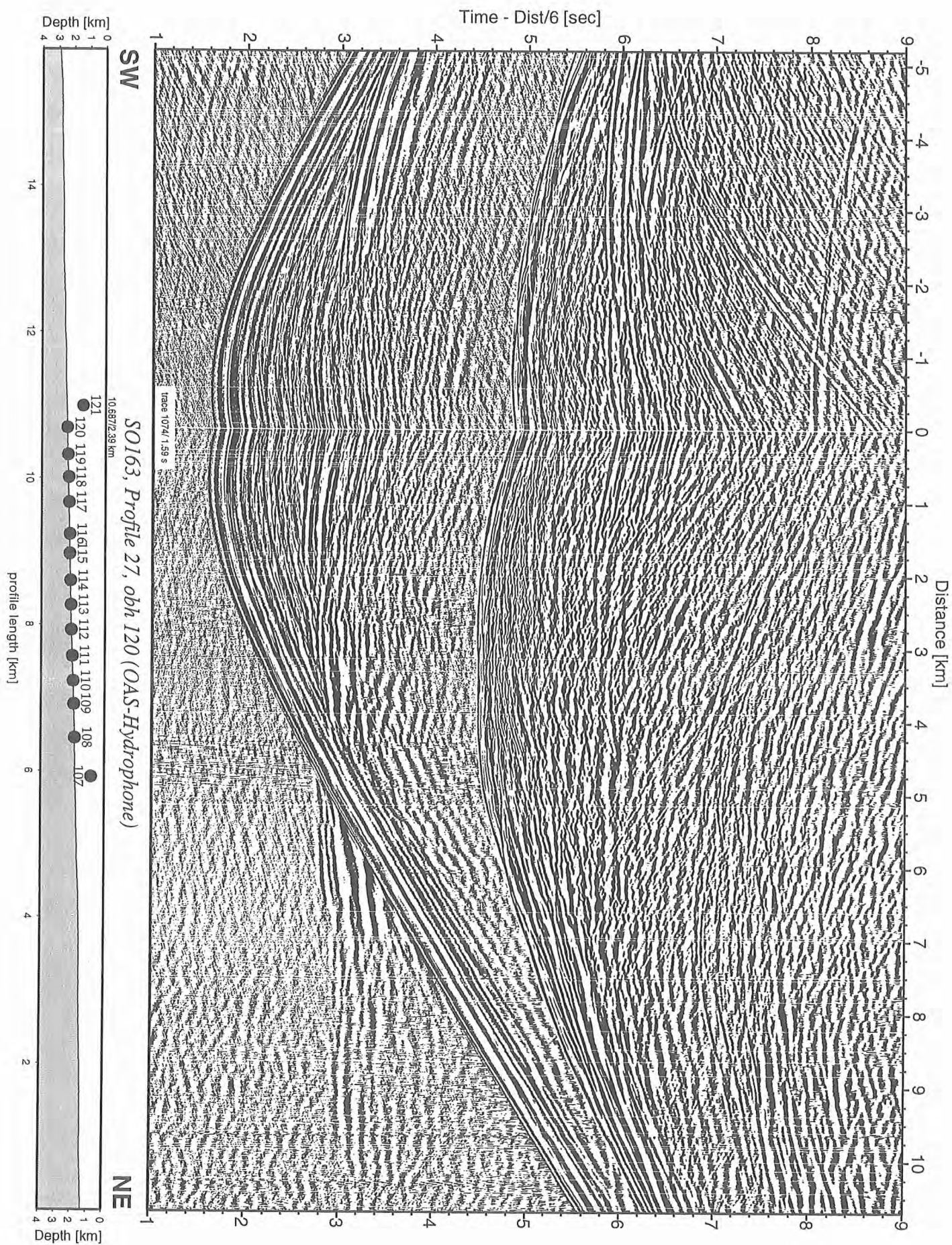


Figure 6.6.4.19: Record section from obh 120 (OAS-Hydrophone), Profile 27.

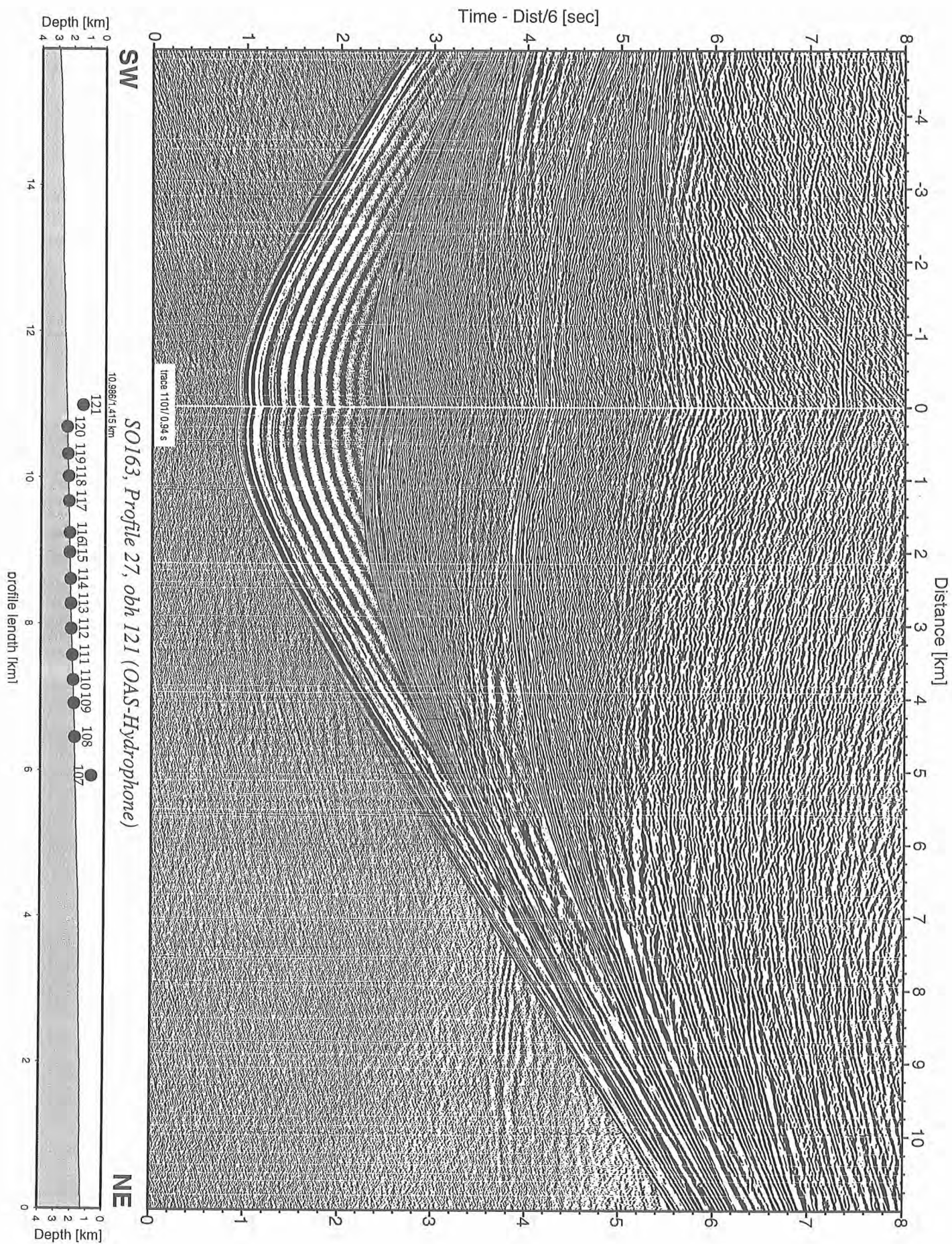


Figure 6.6.4.20: Record section from obh 121 (OAS-Hydrophone), Profile 27.

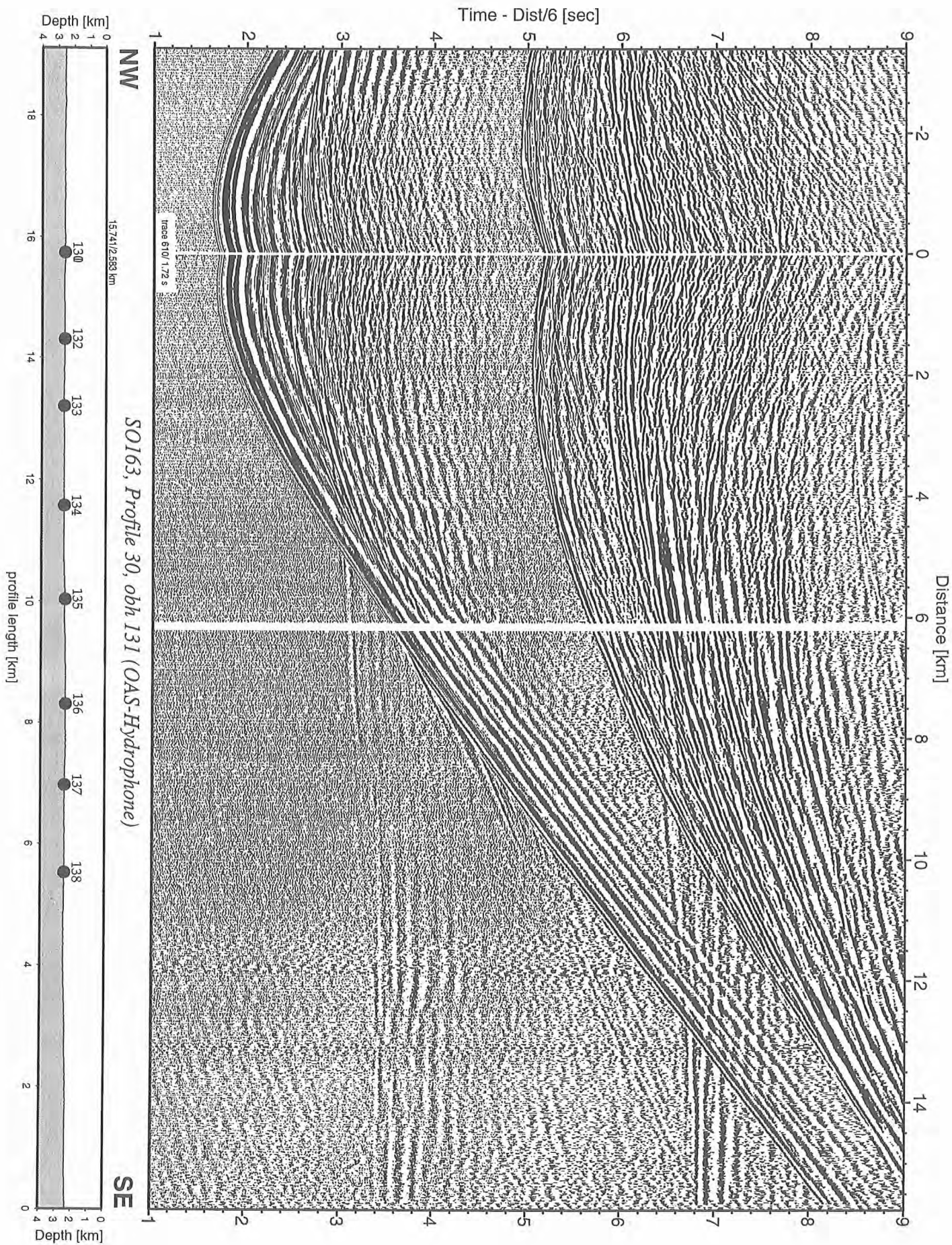


Figure 6.6.4.21: Record section from obh 131 (OAS-Hydrophone), Profile 30.

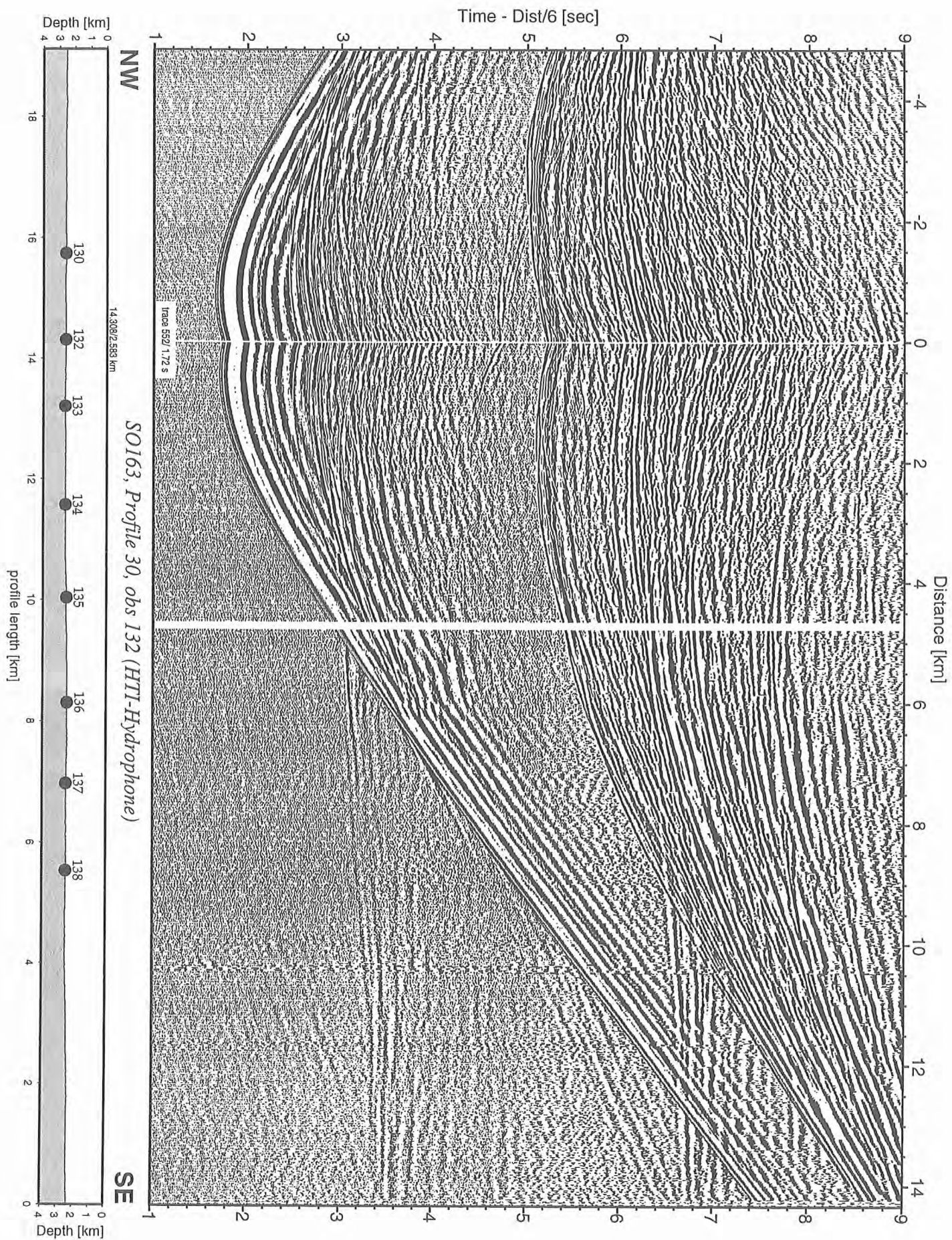


Figure 6.6.4.22: Record section from obs 132 (HTI-Hydrophone), Profile 30.

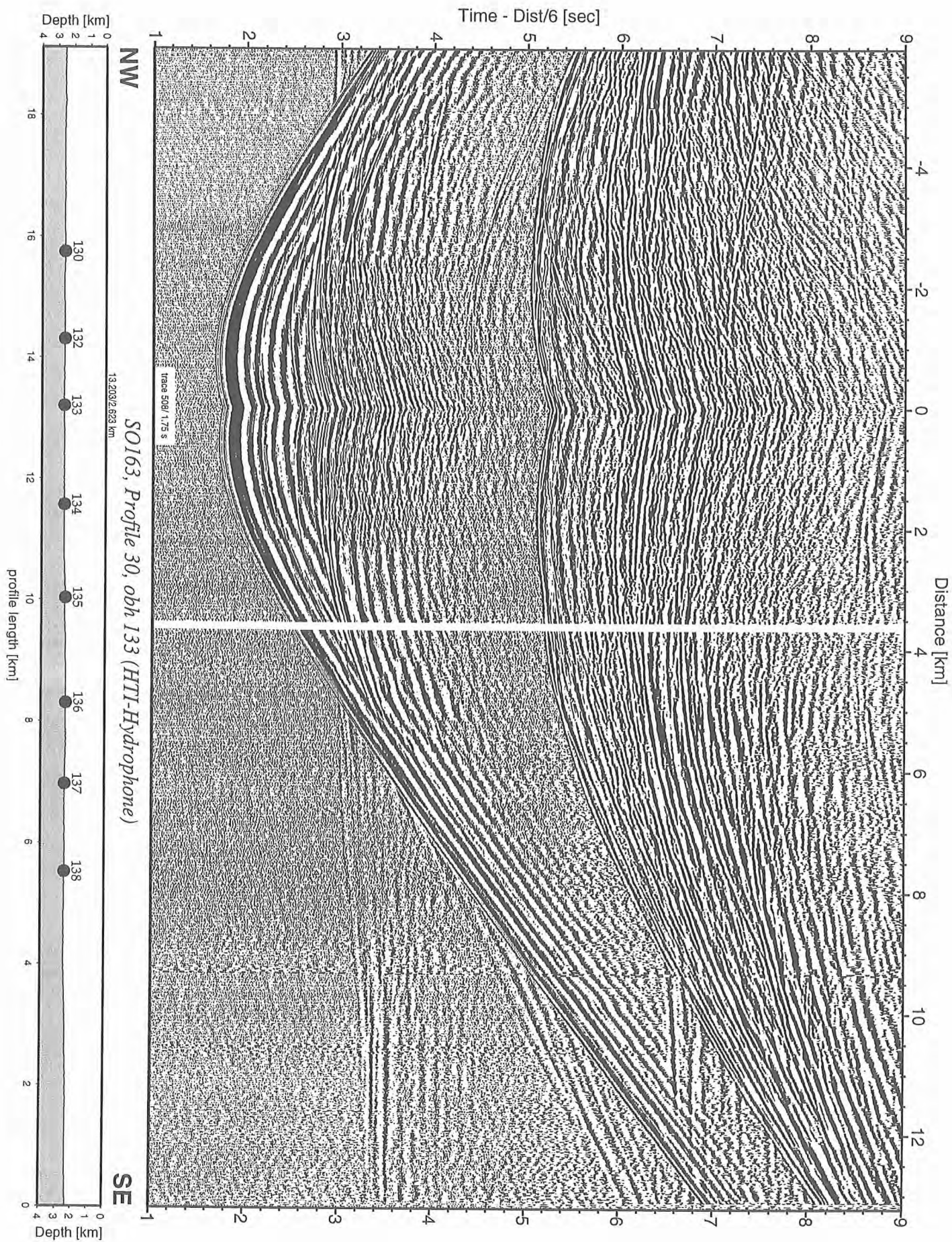


Figure 6.6.4.23: Record section from obh 133 (HTI-Hydrophone), Profile 30.

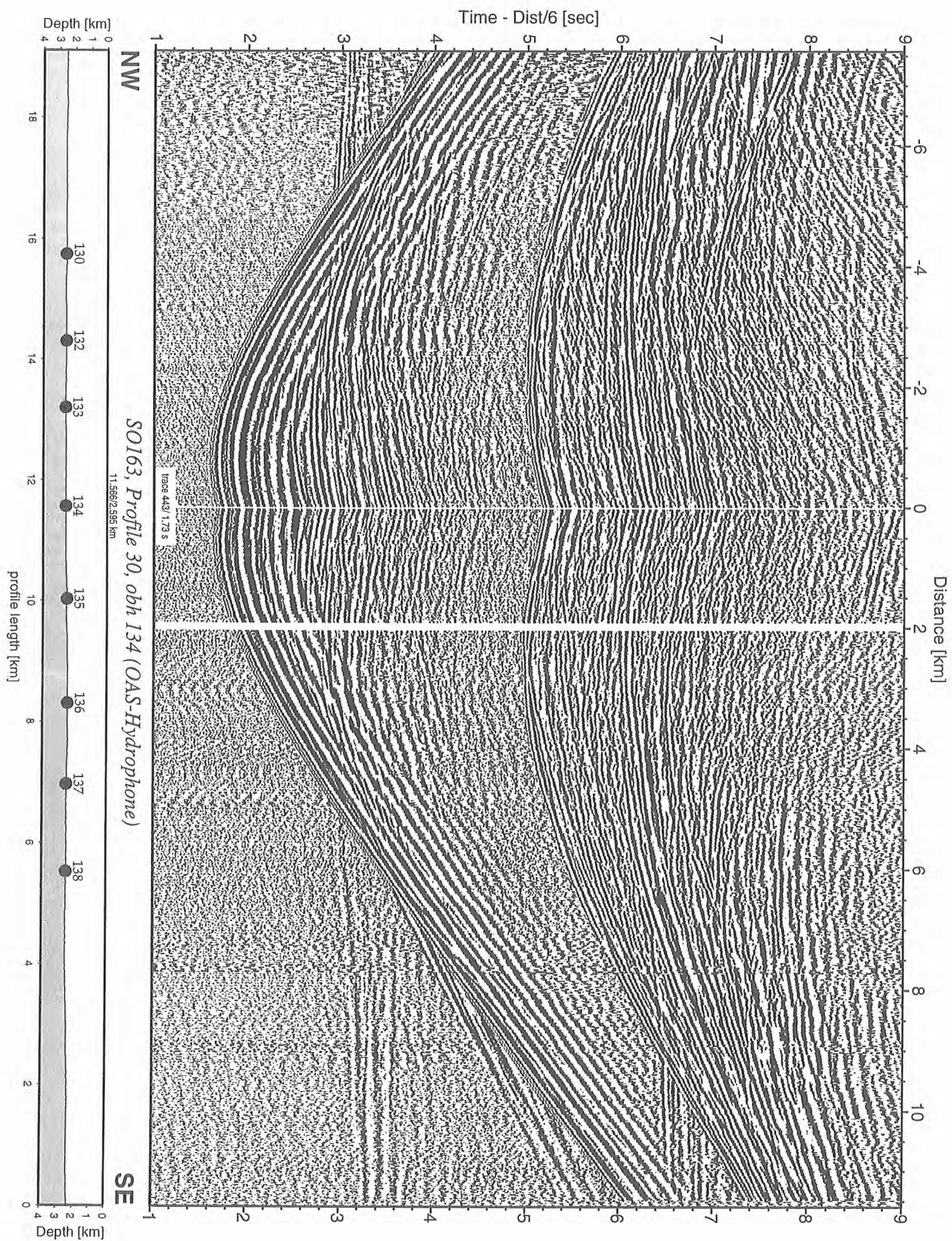


Figure 6.6.4.24: Record section from obh 134 (OAS-Hydrophone), Profile 30.

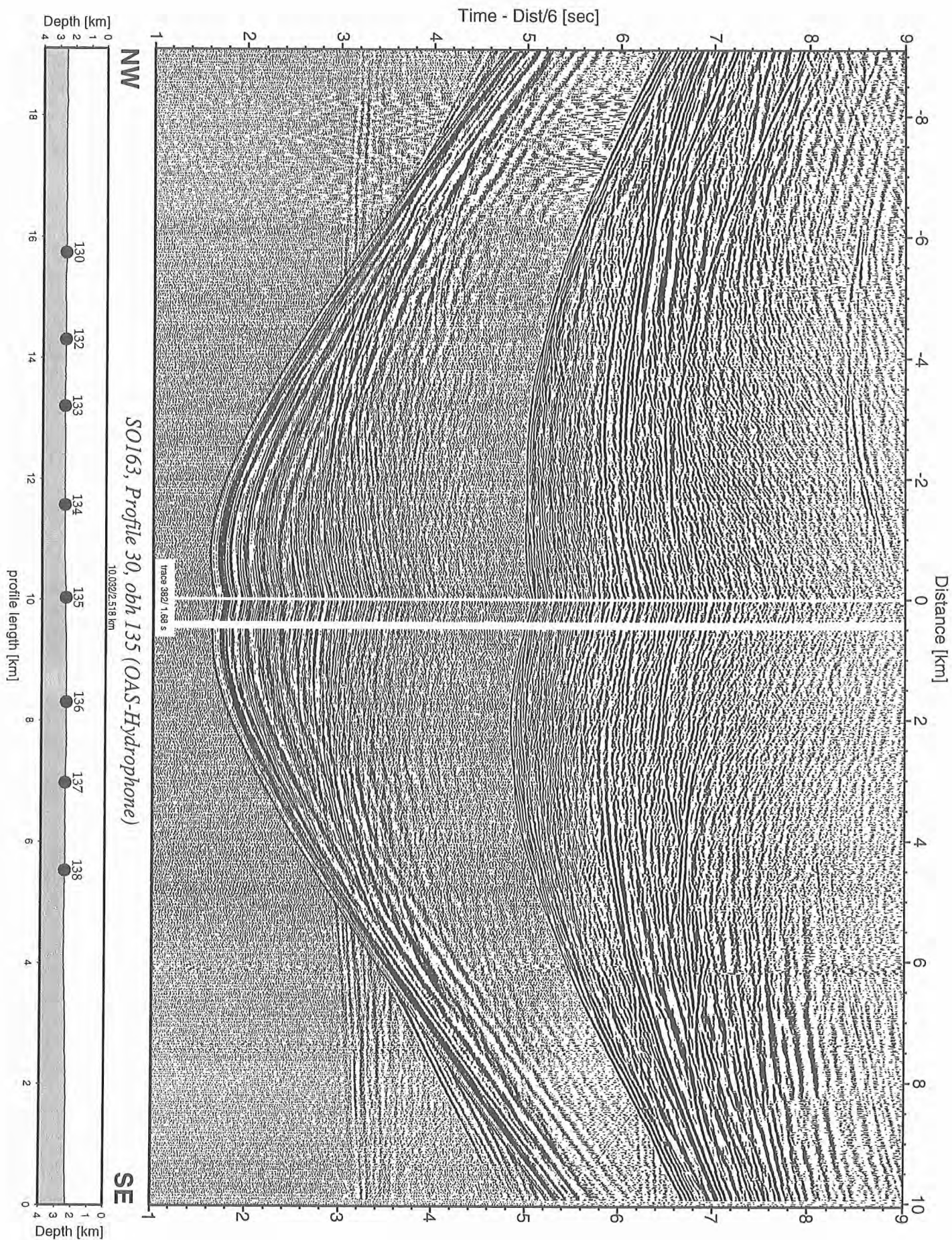


Figure 6.6.4.25: Record section from obh 135 (OAS-Hydrophone), Profile 30.

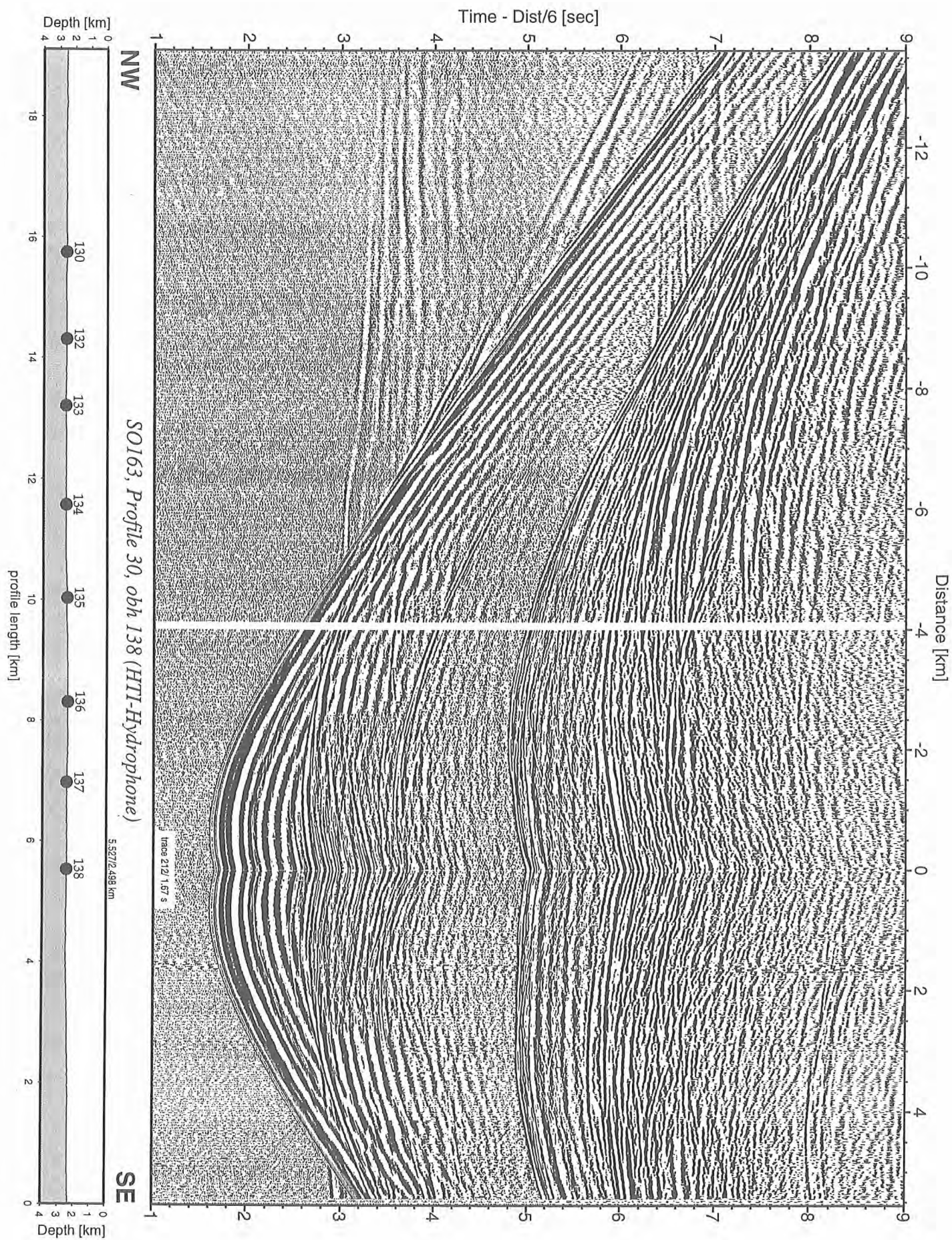


Figure 6.6.4.28: Record section from obh 138 (HTI-Hydrophone), Profile 30.

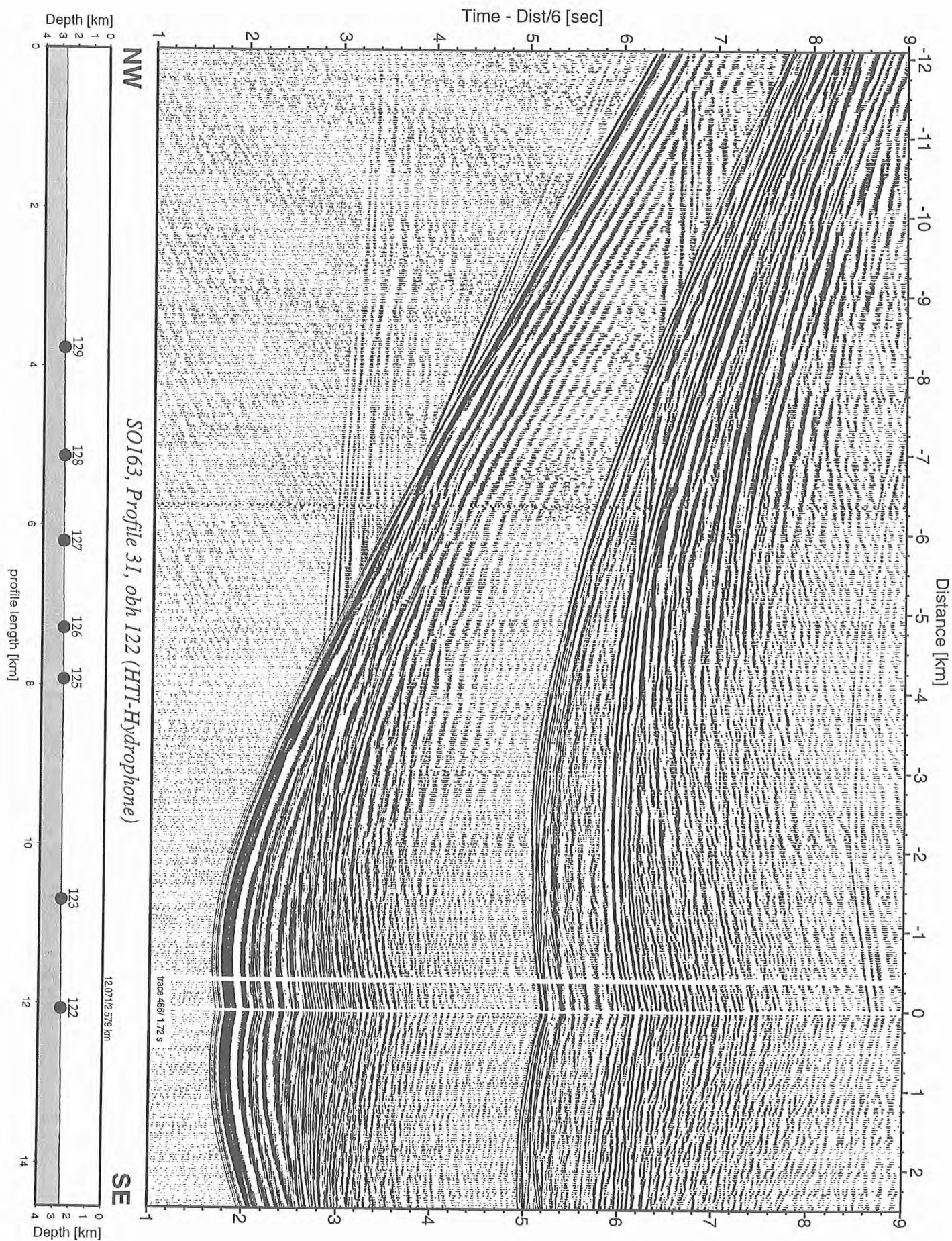


Figure 6.6.4.29: Record section from obh 122 (HTI-Hydrophone), Profile 31.

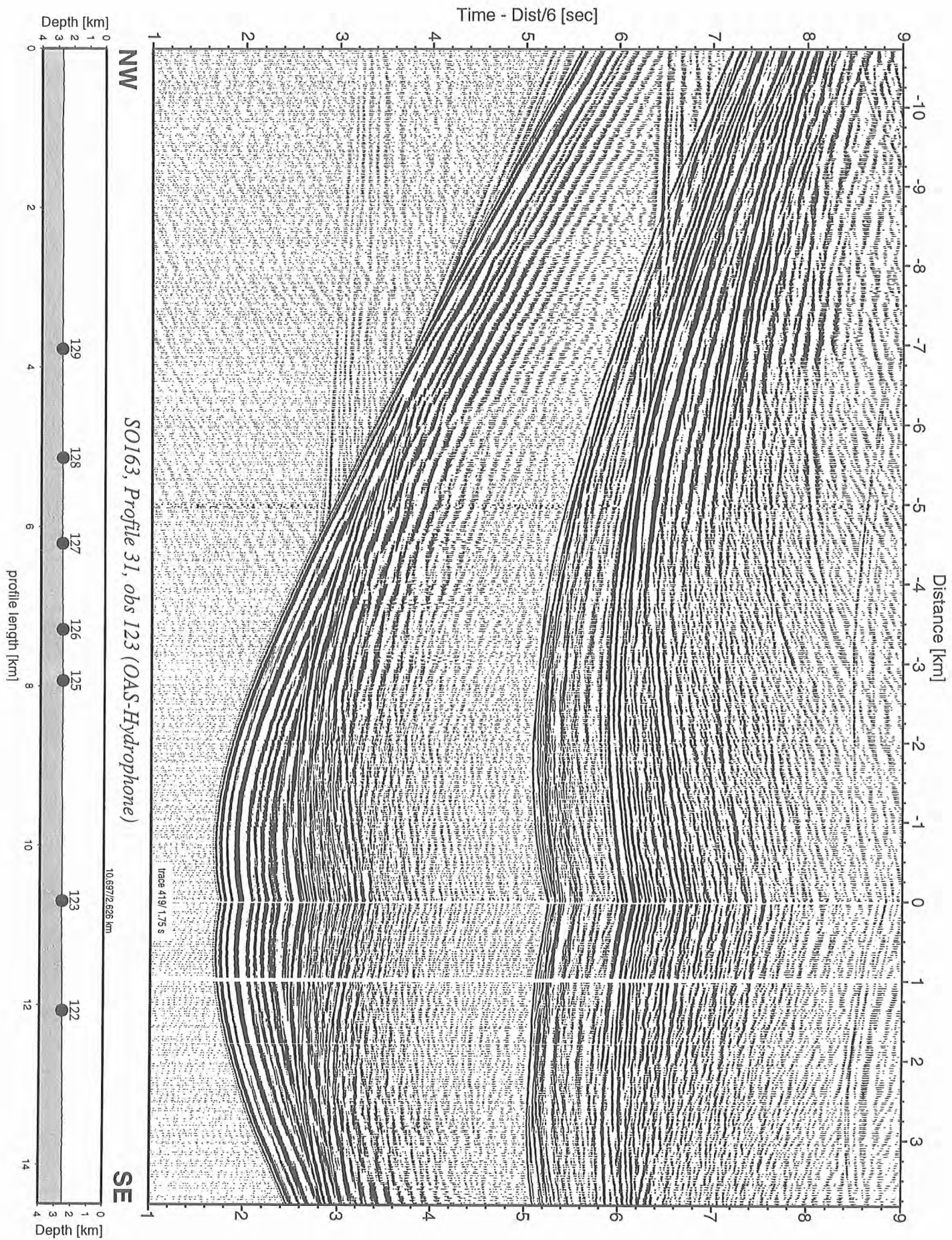


Figure 6.6.4.30: Record section from obs 123 (OAS-Hydrophone), Profile 31.

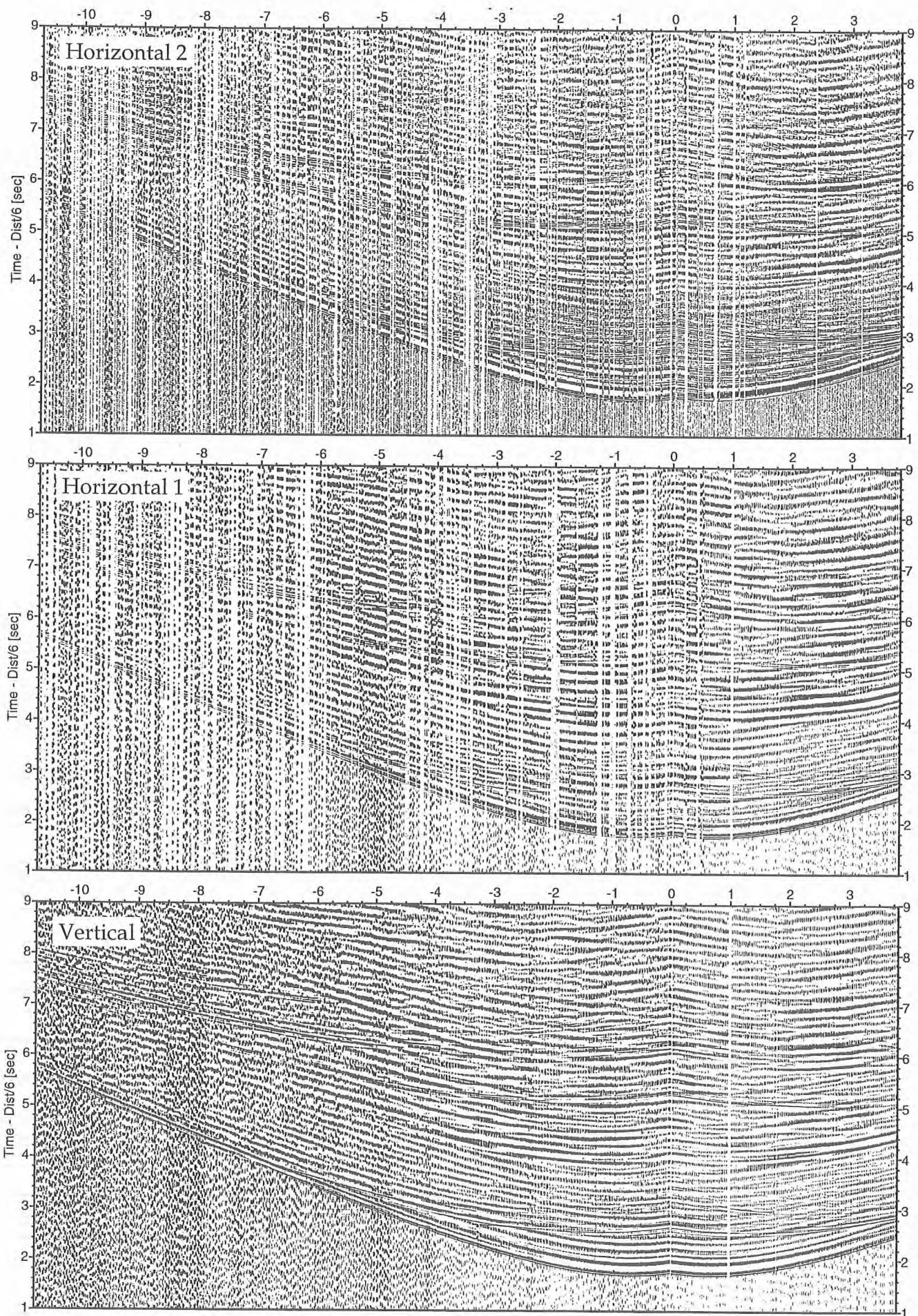


Figure 6.6.4.31: Record sections from obs 123 (PMD), SO163, Profile 31.

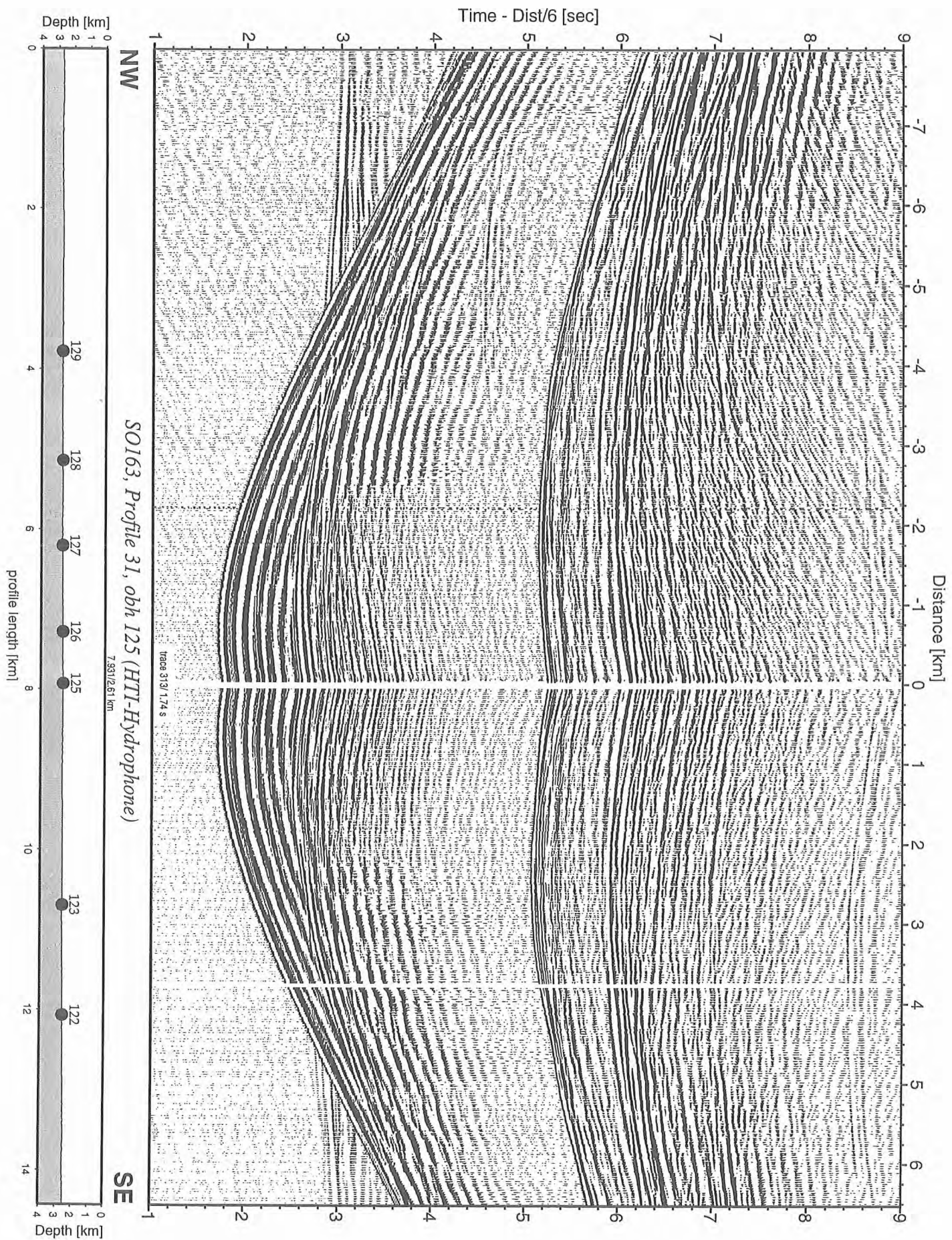


Figure 6.6.4.32: Record section from obh 125 (HTI-Hydrophone), Profile 31.

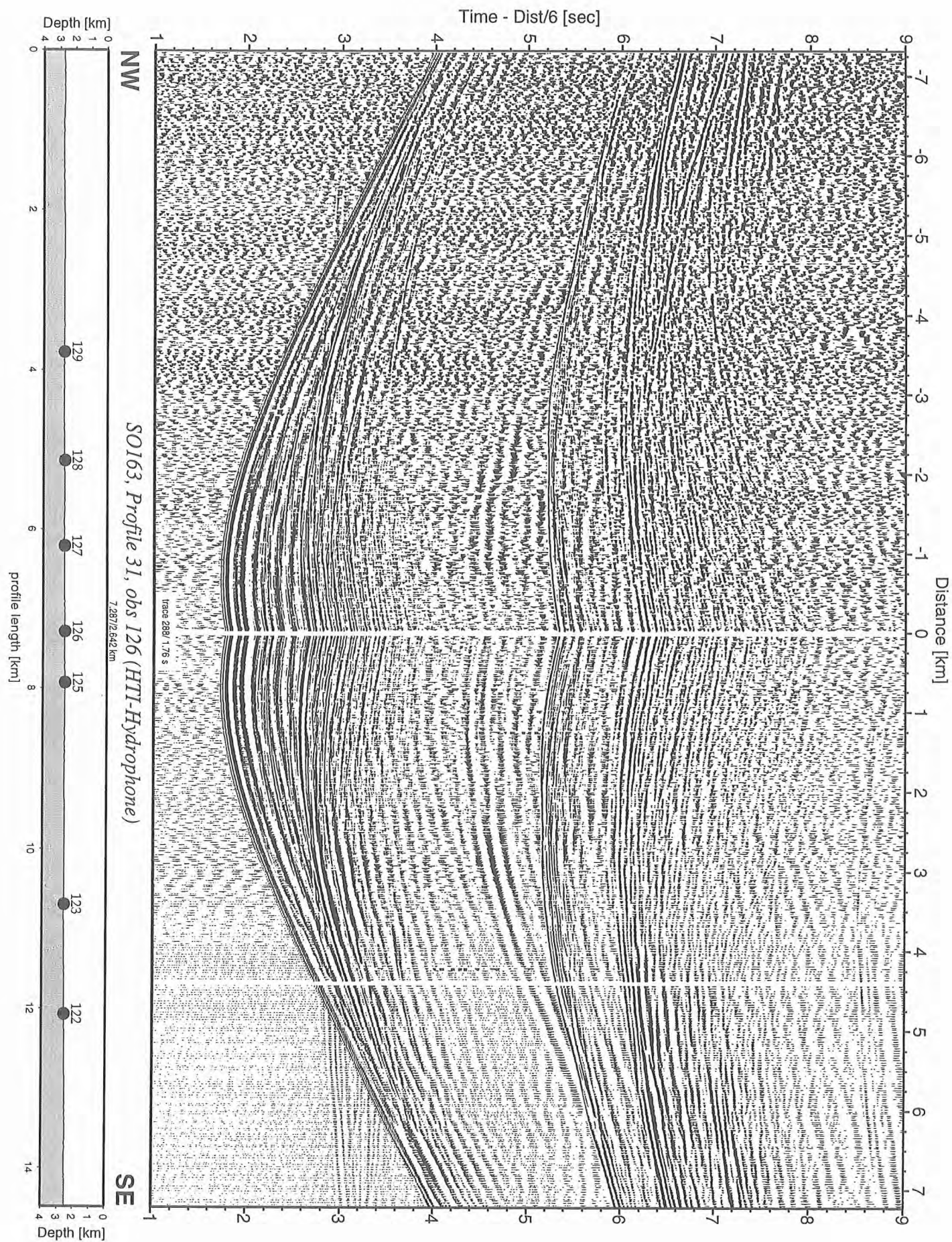


Figure 6.6.4.33: Record section from obs 126 (HTI-Hydrophone), Profile 31.

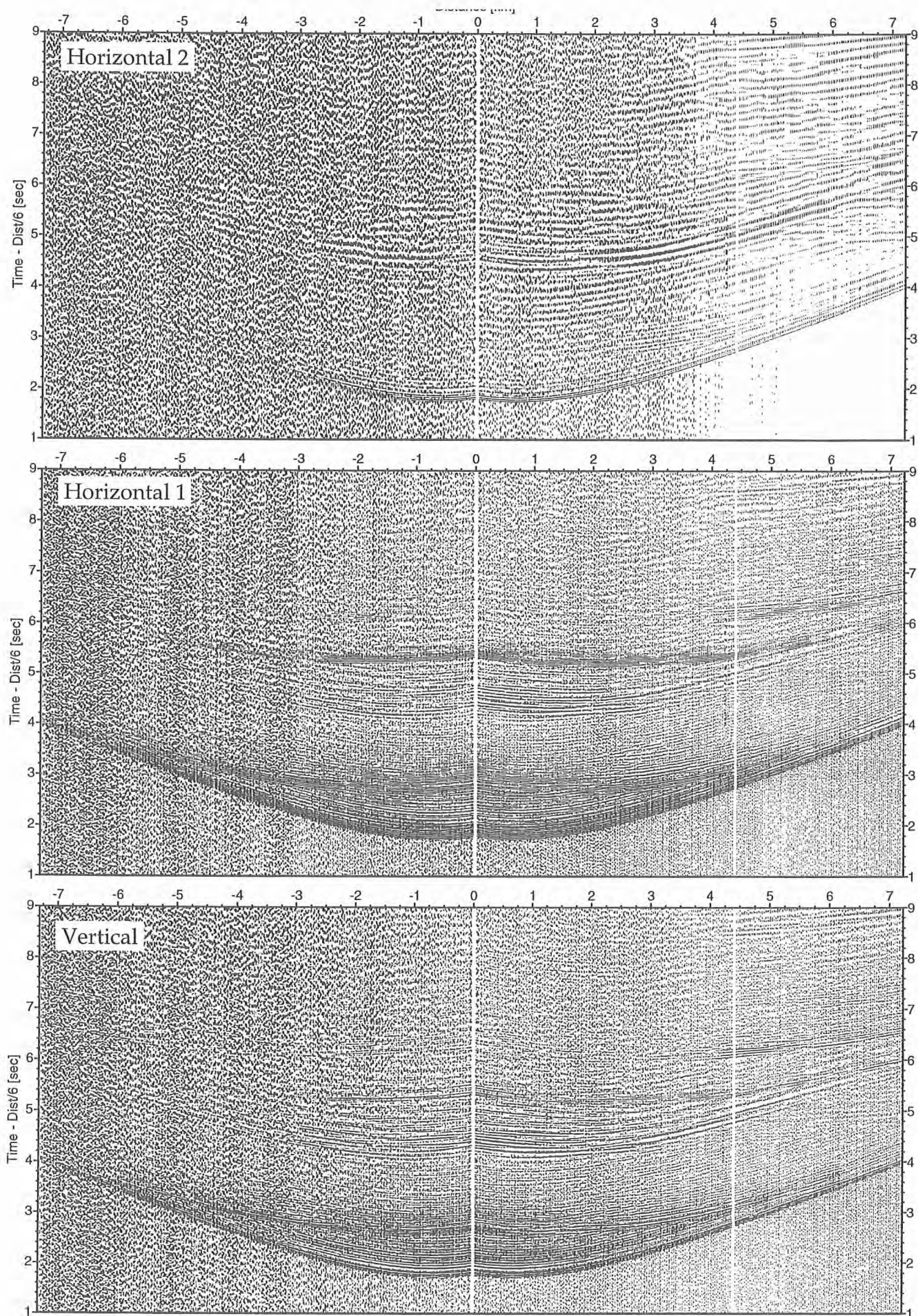


Figure 6.6.4.34: Record sections from obs 126 (Owen-4.5Hz), SO163, Profile 31.

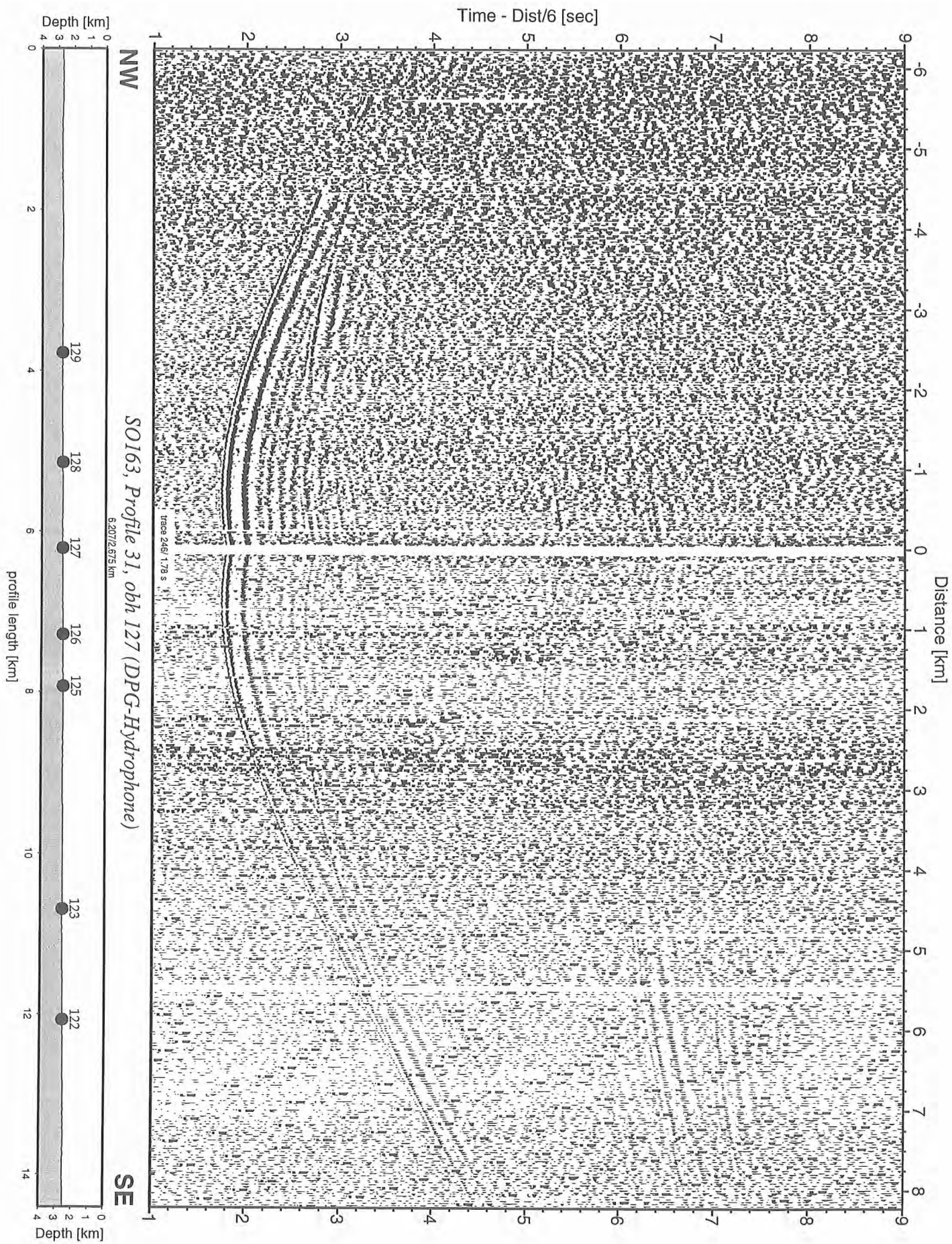


Figure 6.6.4.35: Record section from obh 127 (DPG-Hydrophone), Profile 31.

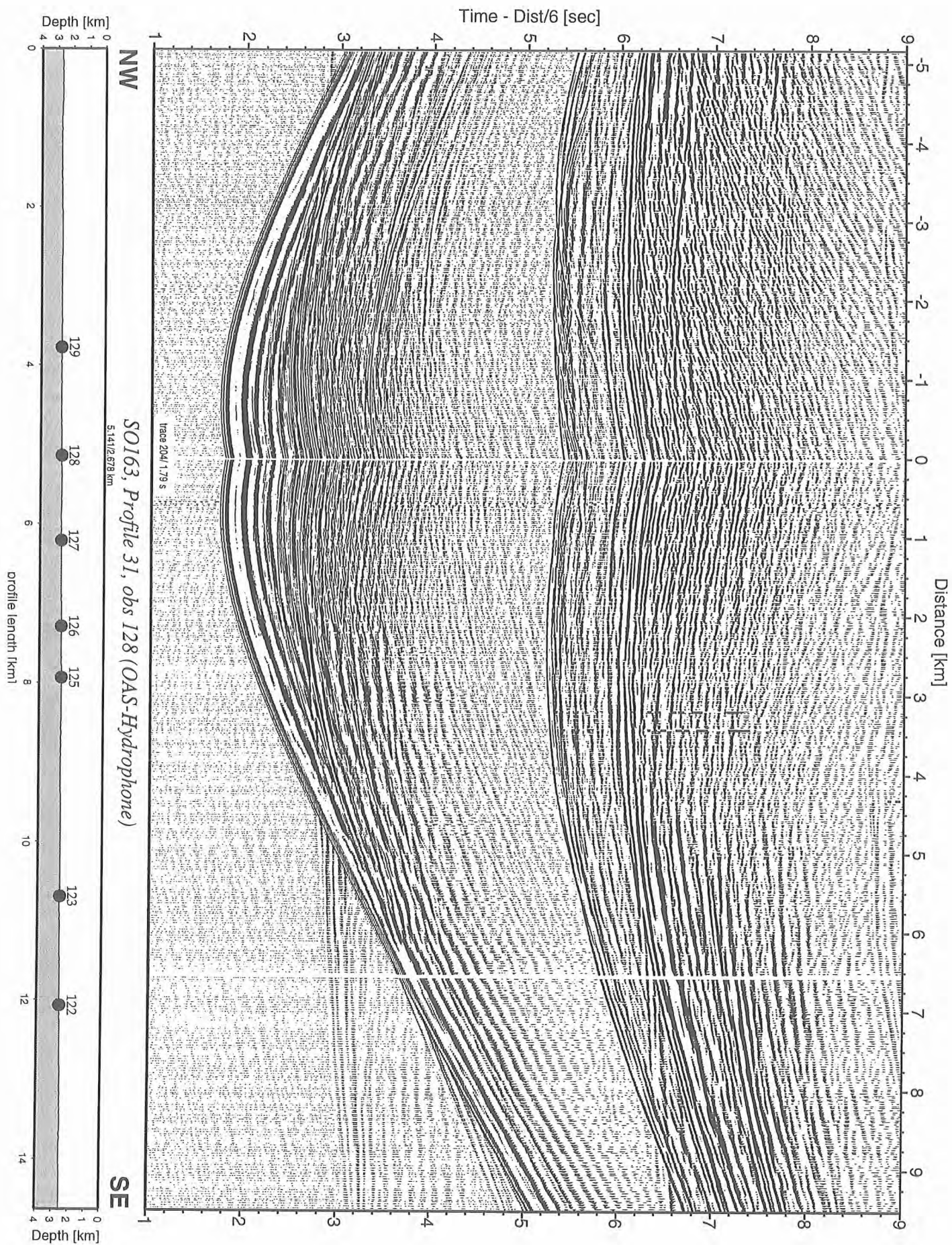


Figure 6.6.4.36: Record section from obs 128 (OAS-Hydrophone), Profile 31.

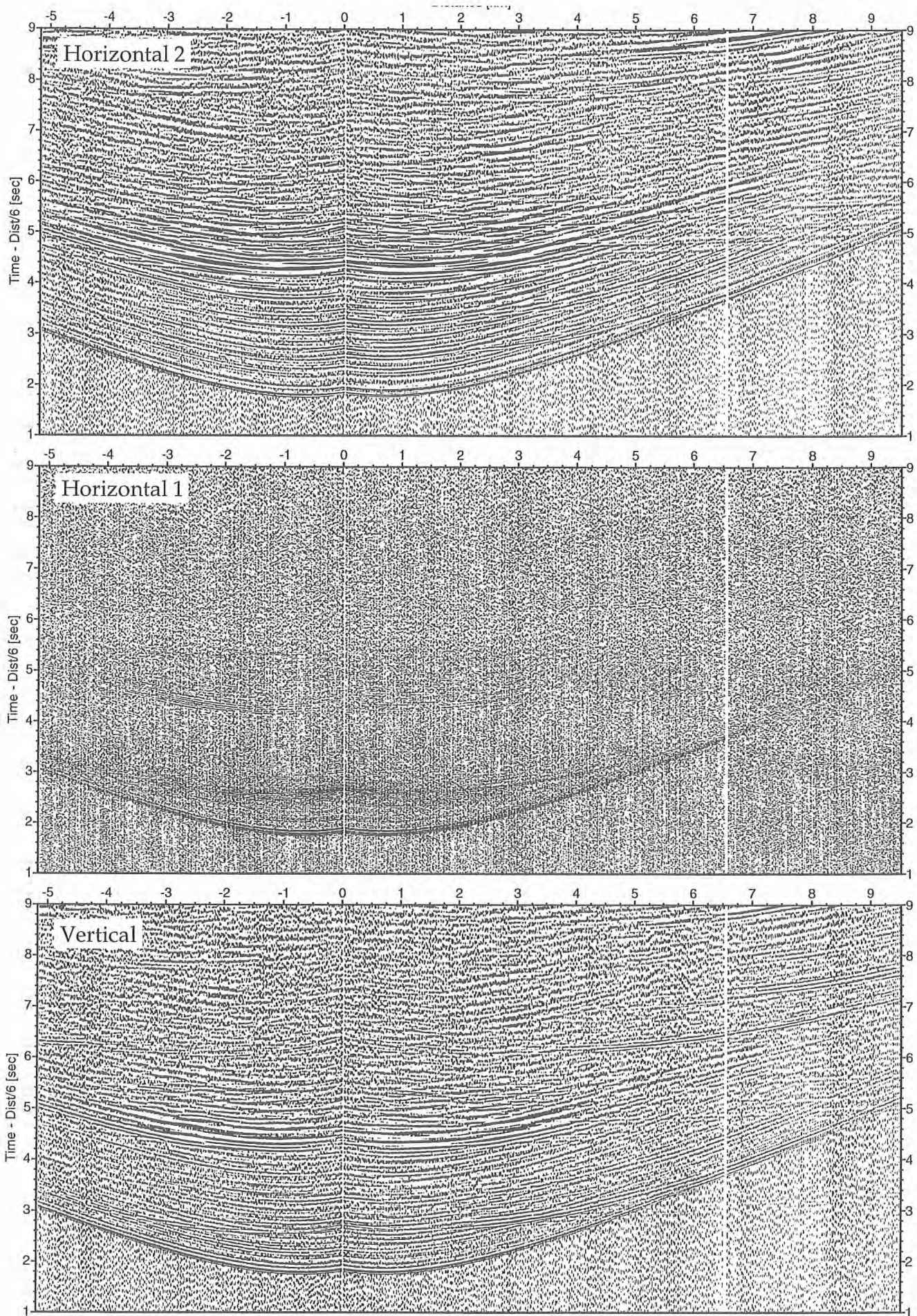


Figure 6.6.4.37: Record sections from obs 128 (Owen-4.5Hz), SO163, Profile 31.

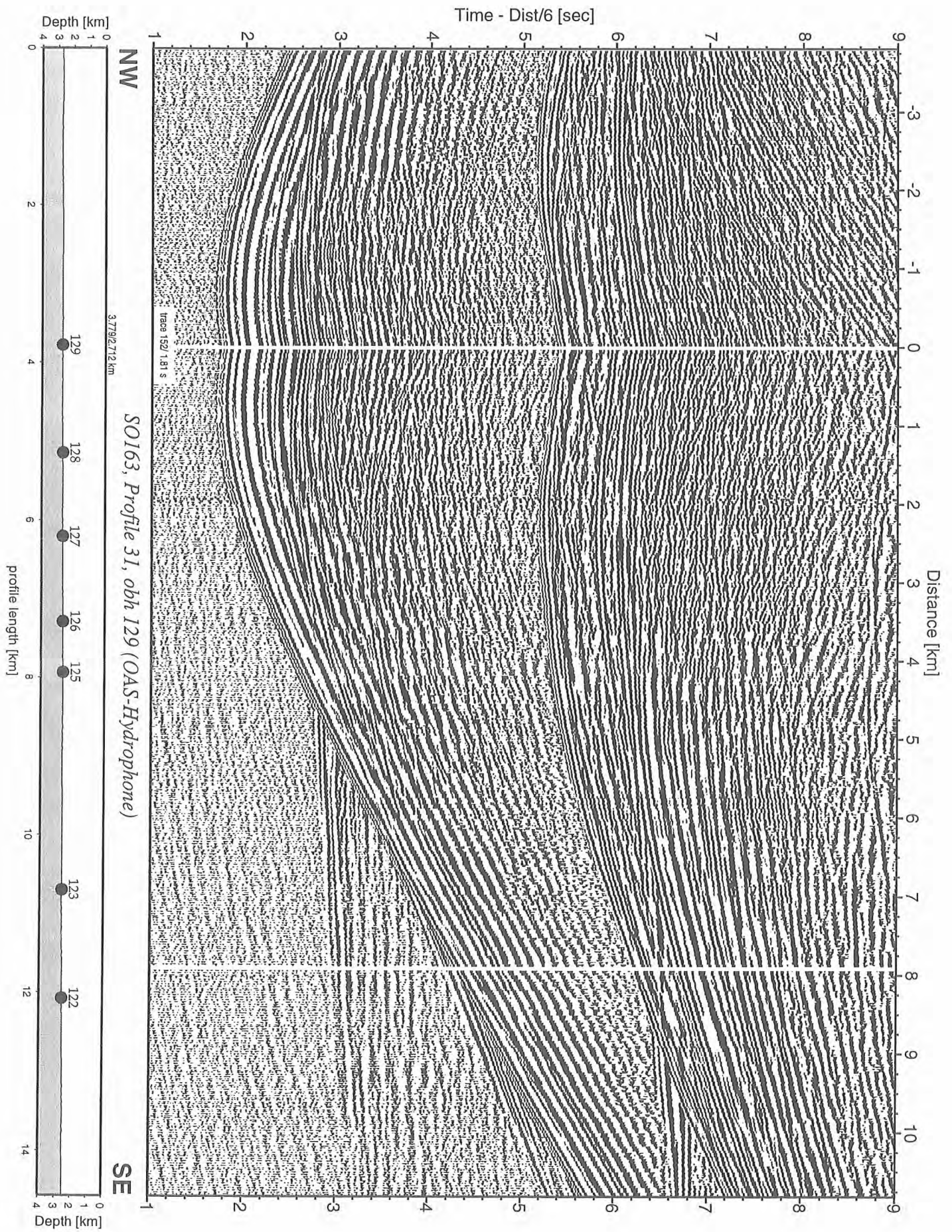


Figure 6.6.4.38: Record section from obh 129 (OAS-Hydrophone), Profile 31.

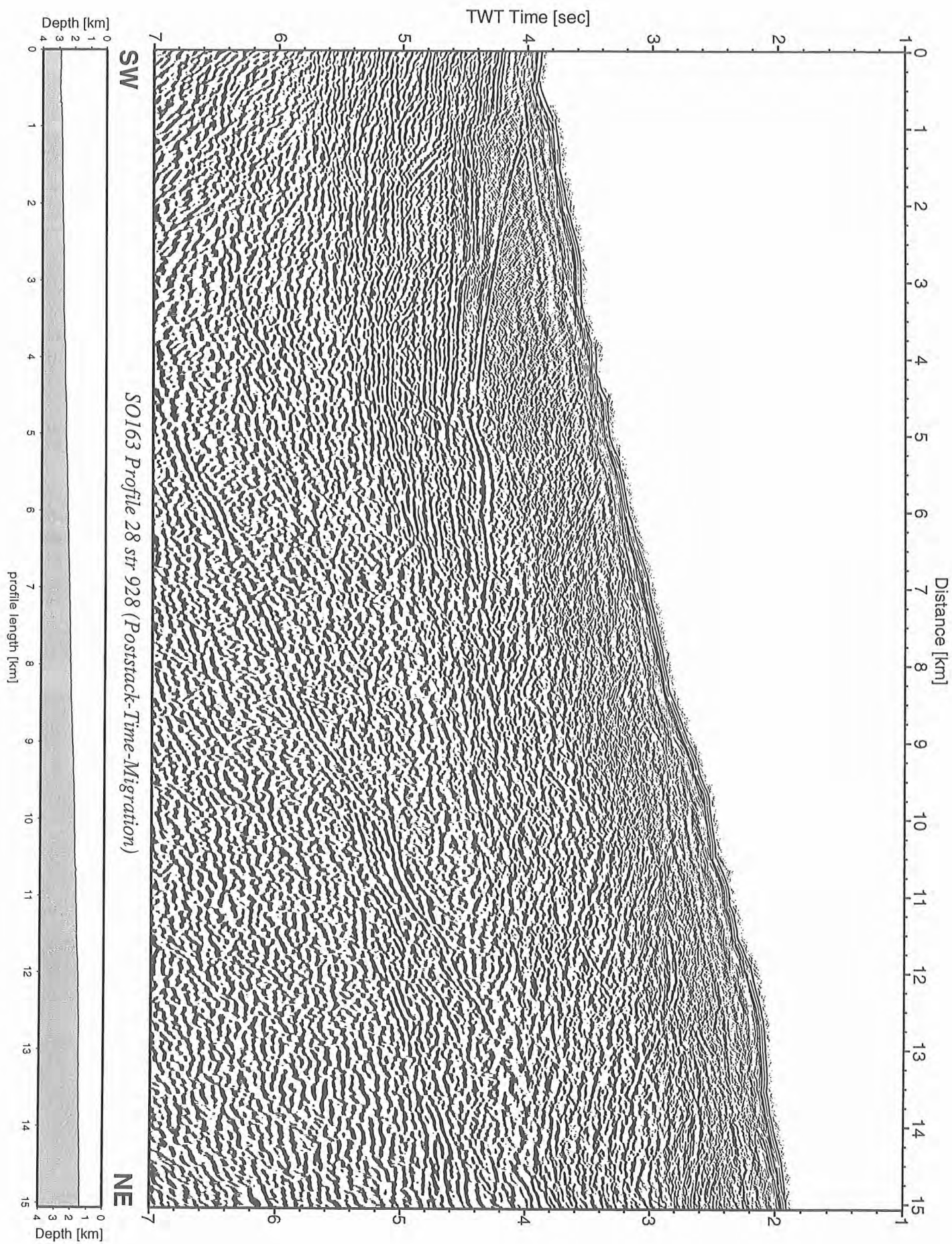


Figure 6.6.4.39: Record section from str 928 (Poststack-Time-Migration), Profile 28.

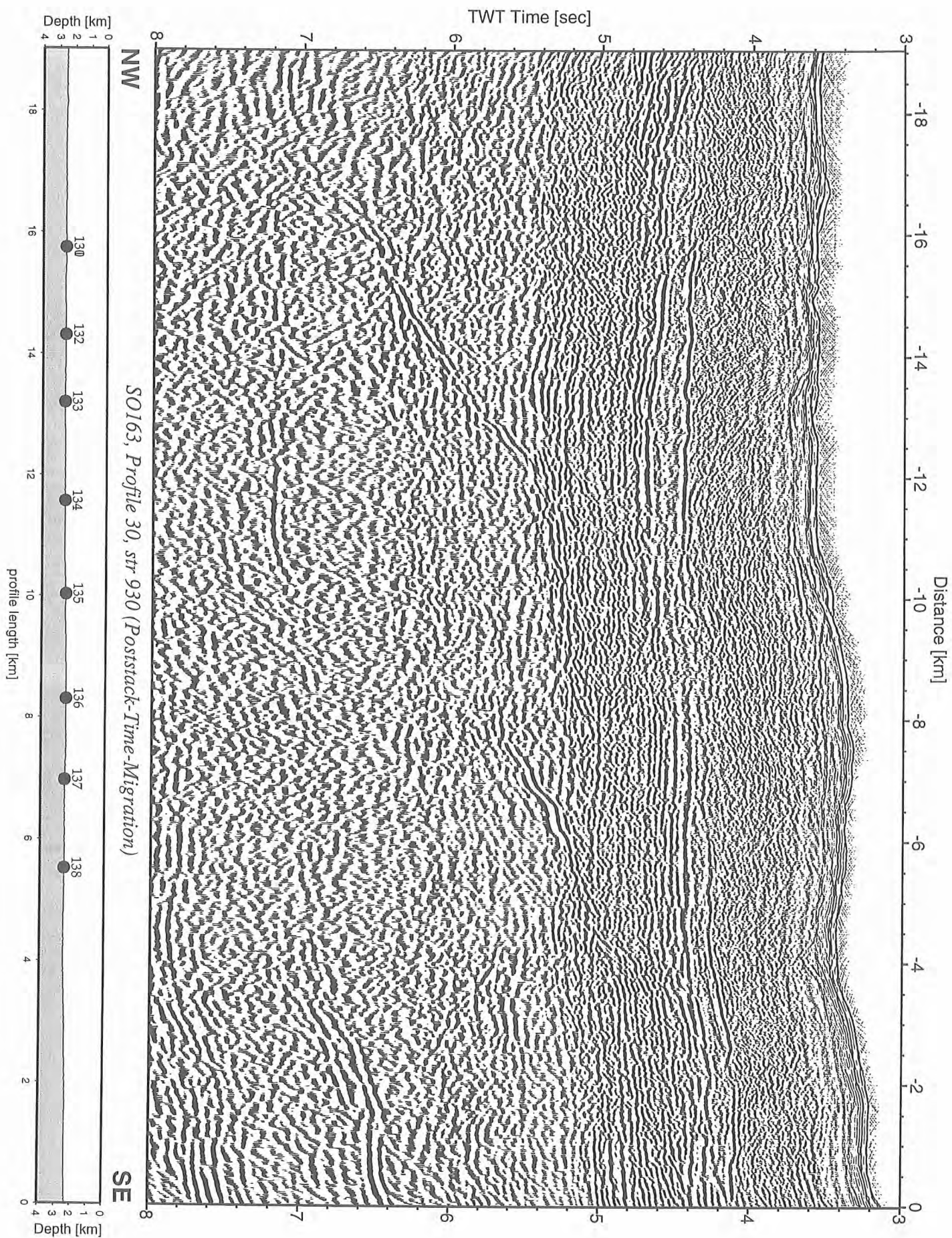


Figure 6.6.4.40: Record section from str 930 (Poststack-Time-Migration), Profile 30.

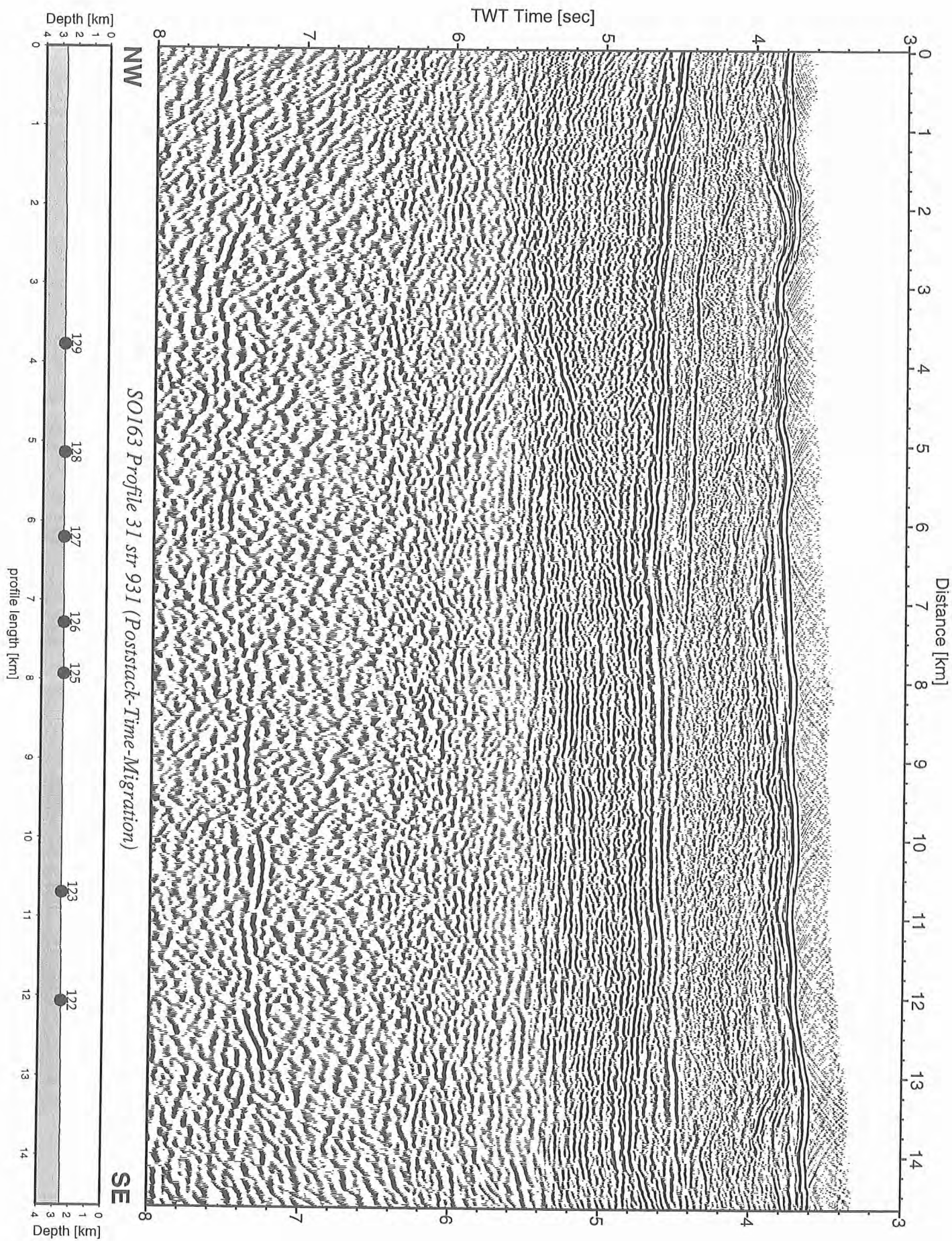


Figure 6.6.4.41: Record section from str 931 (Poststack-Time-Migration), Profile 31.

6.6.5 The Jaco Scar seismological Network

(I. Boschini, J. Gossler, P.O. Thierer)

Description of the network

A larger network than installed on OFOS 10 mudvolcanoe was deployed centered on Jaco Scar. Here a total of 23 instruments were deployed and shall be recovered during METEOR Leg M54/3B in early October 2002. Another five OBH have been deployed east of the network on Quepos Mound, which may be included into the Jaco Scar network. The Quepos Mound stations will be recovered in July. In a first portion, seven ocean bottom instruments were deployed. To the end of SO163-2, we recovered the instruments and redeployed them after checking the quality of the obtained data (Figure 6.6.5.1).

Contemporaneously with the ocean bottom deployment of the ten OBH, the Instituto Costarricense de Electricidad (ICE) together with GEOMAR and the University of Kiel are running a land based local earthquake network. The additional data from the land stations will be used for further constrain earthquake locations and should provide more data for focal mechanism analysis of events located nearby the marine ocean bottom array (Figure 6.6.5.1).

Additionally to the passive seismological recordings, active seismic shootings were carried out following the baselines of the network. Shooting was done using two G-Gun arrays and an additional single Bolt gun of 32 liters volume.

The Jaco Scarp region is a structural place where a seamount is subducting underneath the Caribbean Plate. As seen in the multibeam bathymetry, the subducting seamount produces an uplift of the upper part of the margin right beneath. Moreover, numerous recent fractures are cutting the Jaco uplift area at that place.

The main objectives of this passive seismological experiment is to detect and register the seismicity induced by the convergent dynamics between the subducting seamount and the Caribbean Plate. The spatial dimensions of the joined marine and land network are designed to register local tectonic events of the subducting plate.

Day of month	April														Remarks
	22	23	24	25	26	27	28	29	30	1	2	3	4	5	
Julian day				115					120					125	
OBS14	s	o	o	o	o	o	o	o	o	o					Hydrophone channel out of range, only vertical seismometer component was fairly working. Because of power problems recording stopped after 10 days
OBH15	s	-	-	-	-	-	-	-	-	-	-	-	-	s	DPG pressure sensor did not work
OBH16	s	x	x	x	x	x	x	x	x	x	x	x	x	s	Hydrophone
OBH18	s	x	x	x	x	x	x	x	x	x	x	x	x	s	Hydrophone
OBS19	s	o	o	o	o	o	o	o	o	o	o	o	o	s	Hydrophone sometimes out of range, only one seismometer channel was fairly working. Because of power low recording stopped after 13 days
OBH20	s	-	-	-	-	-	-	-	-	-	-	-	-	s	recorder problem
OBH21	s	x	x	x	x	x	x	x	x	x	x	x	x	s	Hydrophone

x operating
s deployment/recovery
o data with specific problems but at least one component usable (see comment column)
- inaccurate recording, so that data cannot be used
blank not recording

Table 6.6.5.1: The station status of the first deployed JACO net stations, that have been already recovered during RV sonne cruise SO 163-2.

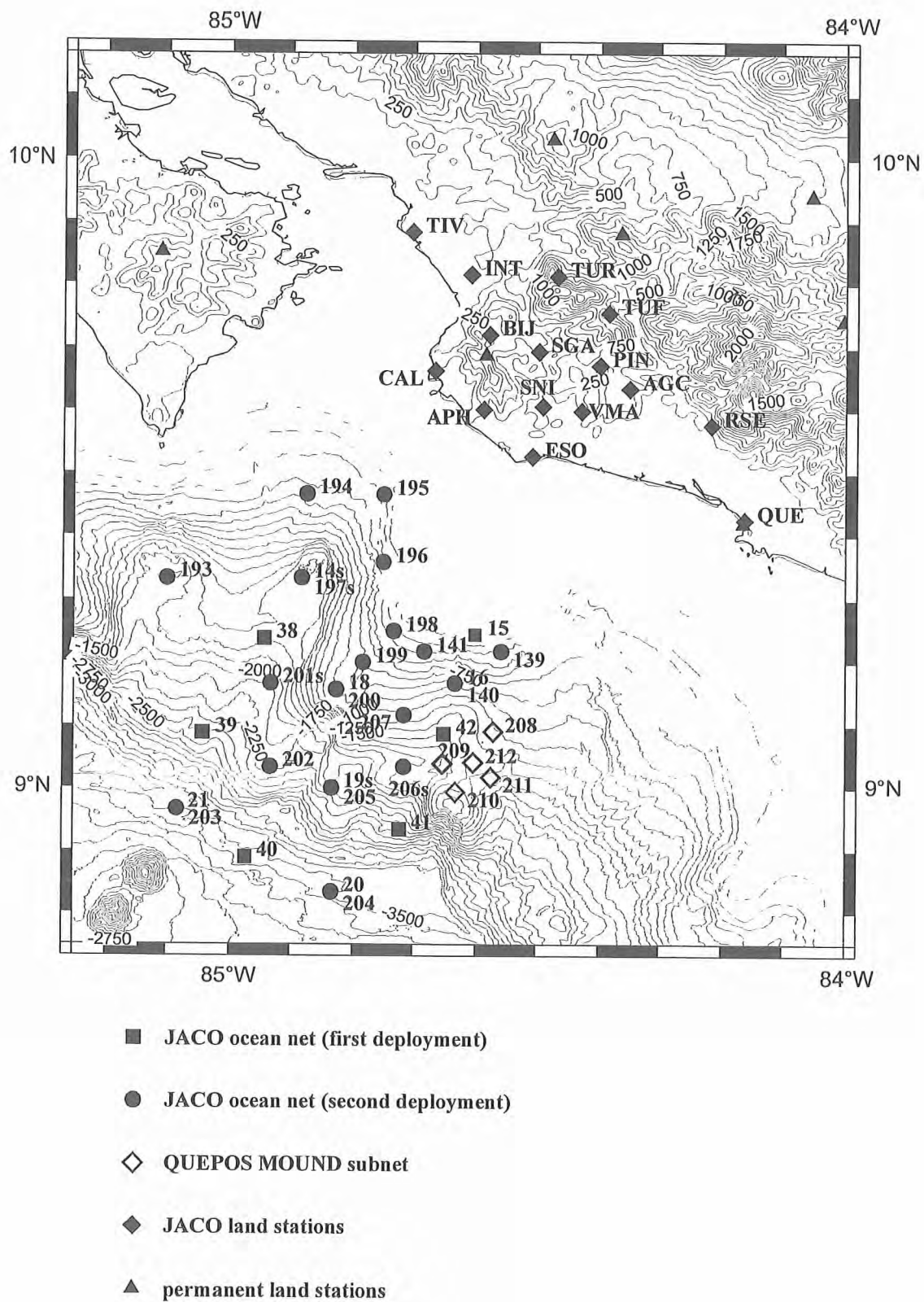


Figure 6.6.5.1: The combined land and seafloor seismic network of the JACO area. The land stations have been installed in April 2002. The ocean bottom instruments have been deployed during RV Sonne cruise SO 163-1 and SO 163-2. The traces of the subducting seamounts are clearly visible as large depressions in the continental margin.

The interpretations on both datasets, the seismic and seismological data, must be done interactive. Active shooting allows to develop a precise velocity-depth model which is elementary for further seismological investigations. Furthermore, a distinct relocalisation of the instruments could be done. Hypocenter determination and calculations of fault plane solutions of the registered earthquake events is as well planned

Data Quality and Noise characteristics

As mentioned above, seven stations of the network were recovered successfully and have been deployed again. Half of the stations operated as intended and recorded high quality data of local and regional earthquakes that occur during their deployment period of 17 days, as well as the airgun source from MCS profiles shot during leg 2 of SO163 (Table 6.6.5.1). Data recovery was not complete because of a number of factors.

1. The PMD seismometers installed at obs14 and obs19 seemed to draw more power than specified such that those stations ran out of power prematurely. Furthermore these instruments jumped often between the minimum and maximum offset, so that only a small portion of their data could be used. Because this failure could not be fixed during the cruise, the PMD sensors have not been deployed again, but will be shipped to the manufacturer for repair.
2. The DPG pressure sensor at station obh15 also did not work correctly and produced a lot of noise and spikes, so that only stronger events could be registered successfully. This sensor also has not been deployed any more.
3. Obh20 failed due to recorder problems.
4. Some of the stations are affected by spikes, and more mysteriously, by bursts of almost monochromatic noise with a fundamental frequency close to 5 Hz, as already reported on RV Sonne cruise SO 161-3 often supplemented by overtones. The noisy periods can last up to several hours. Because of its relative monochromaticity this type of noise can be relatively easily removed by filtering.
5. No problem occur with the new 1 Gbyte flash disks, so that they could be read out completely. Hopefully that will be the same in October when the stations will be recovered and the amount of data is much larger than on the first deployment.

Combination of MLS recorders with Owen-type 4.5 Hz OBS

In Figure 6.6.5.2 we compare the signals of a hydrophone (upper seismogram) with the signals of a seismometer during a shooting session together with a coinciding earthquake. While the shot amplitudes on the hydrophone are larger than the biggest earthquake signals, it is opposite on the seismometer channels (lower seismograms). There, the shot peaks can only hardly be recognized besides the earthquake signal.

Because hydrophones are pressure sensitive, while naturally the seismometer records mainly ocean bottom displacement, the different behaviour is clear.

Though, the foreseen PMD broadband seismometer could not be used due to malfunction, we tested the 4.5 Hz Owen-type sensor in combination with a MLS recorder. Usually these sensors are only used together with MBS recorders, which cannot stay longer than a few weeks at the sea floor, without being recovered for changing the batteries. Although that kind of sensor is not as sensitive to lower frequencies as the broadband sensor, it is sufficient for the detection of local earthquakes.

Testing the direct connection of the sensor to the recorder seismometer connector failed due to insufficient amplification of the sensor's signals. In a quick action a 10-times pre-amplifier has been constructed. In the laboratory testing was successful, so that we decided to use three 4.5 Hz

seismometers for the long-term survey, though we could not test it because deployment has to be done immediately.

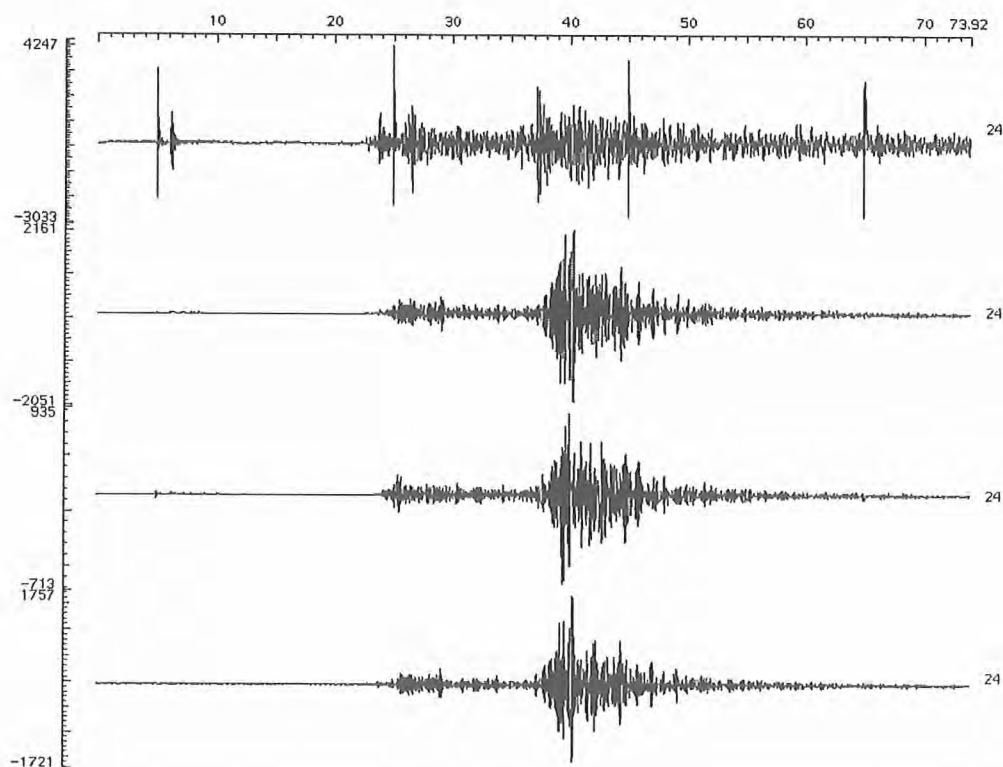


Figure 6.6.5.2: Comparison of hydrophone (upper trace) with seismometer signals during the coincidence of airgun shooting and a local earthquake (lower three traces) on May 8, 2002 19:58 UTC. The shot signals are equally spaced every 20 seconds, with the first shot at the 5th second of the registration window. The earthquake first arrival is 23 seconds after trace start.

The JACO network ocean bottom instruments

A total of 20 OBH and 3 combined OBH/OBS stations are deployed in the JACO seismological network. The instruments sites have been chosen to cover the area of three prominent subducted seamounts, promising a lot of seismicity due to tension when the seamounts pass through the continental lithospheric plate.

Furthermore, a small network of five OBH is closely placed east of the JACO net, to investigate the so called Quepos Mound, a fluid vent site on the subducting plate. These instruments will be recovered in July, earlier than the JACO net stations. Nevertheless, their signals can be combined with the JACO data set, too.

Table 6.6.5.2 specifies the instruments and their deployment date. The first 8 stations including a tiltmeter station are from the first deployment, and have already been recovered. Some of the stations of the second deployment occupy their sites. Data from the already recovered stations will be presented in the following section.

SFB 574 Jaco Earthquake Net (ocean bottom instruments)												
STATION	LAT (N)		LON (W)		DEPTH (m)	DEPLOYMENT (Date)	RECOVERY (Date)	RELEASER CODE	ANTENNA CH.	RECORDER S/N	SENSORS	
	D	M	D	M							HYDROPHONE	TILTMETER
obs14	9	19.985	84	52.996	1741	22. Apr	09. Mai	0399+0355	C	706	OAS 34	PMD 540
obh15	9	14.661	84	36.046	114	22. Apr	09. Mai	0398+0355	A	711	DPG 95	
obh16	9	9.975	84	38.015	944	22. Apr	09. Mai	0397+0355	C	10405	HTI 302	
obh17	9	9.405	84	49.577	824	22. Apr	28. Apr	03BC+0355	C	20303		4948
obh18	9	9.397	84	49.588	827	23. Apr	09. Mai	4A44	C	10402	OAS 18	
obs19	8	59.996	84	50.055	2206	23. Apr	09. Mai	03B9+0355	C	10401	OAS 31	PMD 509
obh20	8	50.006	84	49.996	3473	23. Apr	09. Mai	03B7+0355	D	991238	HTI 602	
obh21	8	58.032	85	5.046	3602	23. Apr	09. Mai	03BB+0355	C	991243	HTI 102	
JACO network (until October 2002)												
obh38	9	14.271	84	56.637	1752	28. Apr		03BD+0355	D	713	DPG 93	
obh39	9	5.280	85	2.560	2496	28. Apr		03B6+0355	C	991249	DPG 75	
obh40	8	53.328	84	58.432	3420	28. Apr		03B5+0355	C	991237	DPG 91	
obh41	8	56.043	84	43.428	2615	28. Apr		03B3+0355	C	991242	DPG 87	
obh42	9	5.173	84	39.092	1481	28. Apr		03B8+0355	C	991250	DPG 92	
obh139	9	13.023	84	33.508	379	12. Mai		03B7+0355	C	991247	OAS 02	
obh140	9	10.017	84	37.993	939	12. Mai		03BB+0355	A	10405	OAS 04	
obh141	9	13.016	84	41.022	351	12. Mai		03B9+0355	A	10402	OAS 29	
obh193	9	19.989	85	6.063	1880	16. Mai		3659	C	991238	HTI 502	
obh194	9	28.015	84	52.503	560	16. Mai		039A+0355	D	707	OAS 18	
obh195	9	28.001	84	44.997	324	16. Mai		0387+0355	B	991256	OAS 18	
obh196	9	21.531	84	45.034	413	16. Mai		3624	C	10401	HTI 302	
obs197	9	20.022	84	53.006	1745	17. Mai		03BA+0355	C	10404	HTI 702	Owen 22 (4.5 Hz)
obh198	9	14.975	84	44.001	291	16. Mai		3674	C	991243	OAS 27	
obh199	9	12.016	84	46.990	734	16. Mai		03B1+0355	D	20303	HTI 802	
obh200	9	9.419	84	49.601	824	16. Mai		03B2+0355	C	20302	HTI 602	
obs201	9	10.013	84	56.020	2080	17. Mai		0397+0355	D	991252	OAS 31	
obh202	9	1.992	84	56.010	2237	17. Mai		0386+0355	D	991236	HTI 102	
obh203	8	57.974	85	5.004	3561	17. Mai		3614	D	706	OAS 18	4949
obh204	8	50.000	84	49.995	3474	17. Mai		3629	D	991246	OAS 38	4948
obh205	8	59.985	84	49.996	2207	17. Mai		3669	B	991241	HTI 902	
obs206	9	1.996	84	43.006	1726	17. Mai		03BC+0355	D	10403	OAS 46	
obh207	9	7.000	84	43.003	1159	17. Mai		3609	C	711	OAS 22	Owen 21 (4.5 Hz)
Quepos Mound Net (until July 2002)												
obh208	9	5.376	84	34.248	1155	17. Mai		0399+0355	D	991248	HTI 202	
obh209	9	2.296	84	39.205	1615	17. Mai		3664	A	991244	OAS 75	
obh210	8	59.579	84	37.907	1202	17. Mai		03B4+0355	D	708	HTI 101	
obh211	9	1.004	84	34.461	1411	17. Mai		6334	C	616	OAS 09	
obh212	9	2.387	84	36.203	1326	17. Mai		4A44	B	971202	OAS 28	

Table 6.6.5.2: The seafloor stations of the JACO network and the Quepos Mound Net

The JACO network earthquake dataset from the first deployment

During the two and a half week period of the first deployment from April 23 to May 8, 2002 a total of 34 earthquakes has been detected by the JACO seismological network. Preliminary hypocenter locations have been determined for 20 events (Table 6.6.5.3).

Date	Time		Latitude	Longitude	Depth	Magnitude
2002 4 23	13 44	7.0 L				
2002 4 25	6 58	52.0 L				
2002 4 26	16 24	49.7 L				
2002 4 26	21 51	12.1 L	9.365	-84.862	24.2	2.1C
2002 4 26	22 2	17.3 L	9.332	-84.869	1.6	2.3C
2002 4 26	22 27	51.1 L	9.370	-84.857	31.2	2.4C
2002 4 26	23 14	42.4 L	9.334	-84.867	1.3	
2002 4 27	0 10	8.1 L	9.329	-84.888	1.6	1.9C
2002 4 27	0 29	38.5 L	9.349	-84.874	1.3	2.0C
2002 4 27	1 30	55.7 L	9.352	-84.867	1.3	2.2C
2002 4 27	1 37	31.9 L	9.340	-84.854	1.7	2.1C
2002 4 27	2 29	42.5 L	9.340	-84.877	1.2	2.0C
2002 4 27	3 26	38.2 L	9.308	-84.886	1.9	2.2C
2002 4 27	10 52	23.3 L	9.368	-84.856	34.7	2.3C
2002 4 28	1 52	44.0 L				
2002 4 28	13 55	47.2 L	9.247	-84.455	53.9	2.3C
2002 4 28	19 4	11.2 R				
2002 4 29	2 44	30.2 L	9.087	-84.669	11.0	
2002 4 29	5 15	21.1 L	10.104	-84.580	134.1	2.9C
2002 4 30	18 54	7.0 L				
2002 5 1	5 21	18.2 L	9.752	-84.458	121.7	
2002 5 2	14 35	52.0 L				
2002 5 3	8 14	58.0 L				
2002 5 4	4 33	37.0 L				
2002 5 5	0 6	18.0 L				
2002 5 5	2 36	38.0 L				
2002 5 5	3 29	3.0 L				
2002 5 5	3 34	3.0 L				
2002 5 5	13 14	25.3 L	10.096	-84.685	96.0	3.2C
2002 5 5	13 21	4.2 L	8.662	-85.335	7.3	
2002 5 6	1 29	43.6 L	9.062	-84.738	7.1	2.6C
2002 5 8	1 24	19.1 L	9.121	-84.895	3.4	1.9C
2002 5 8	14 12	28.7 L	9.053	-84.740	7.1	3.0C
2002 5 8	23 50	56.8 L				

Table 6.6.5.3: Preliminary earthquake catalogue from the JACO seismic ocean bottom network. Origin times are only valid for localized events, all others are the start time of the time window, which starts three minutes before the triggered event. L means local, R regional event, latitude and longitude are in ° (deg), depth is in km, magnitude is the coda-magnitude due to missing receiver transfer functions.

Half of the earthquakes have been located exactly at the place where a big seamount has already entered into the continental lithosphere (Figure 6.6.5.3). Eleven of those events belongs to a swarm that occurred between April 26, 21:51 and April 27, 03:26. The other subducted seamounts did not produce heavy seismicity in the observed time interval. One event could be detected on the track subducted seamounts draw below the Caribbean plate. Three earthquakes have been observed beneath the continent at depth around 100 km, which are originating from the subducting oceanic plate (Benioff Zone).

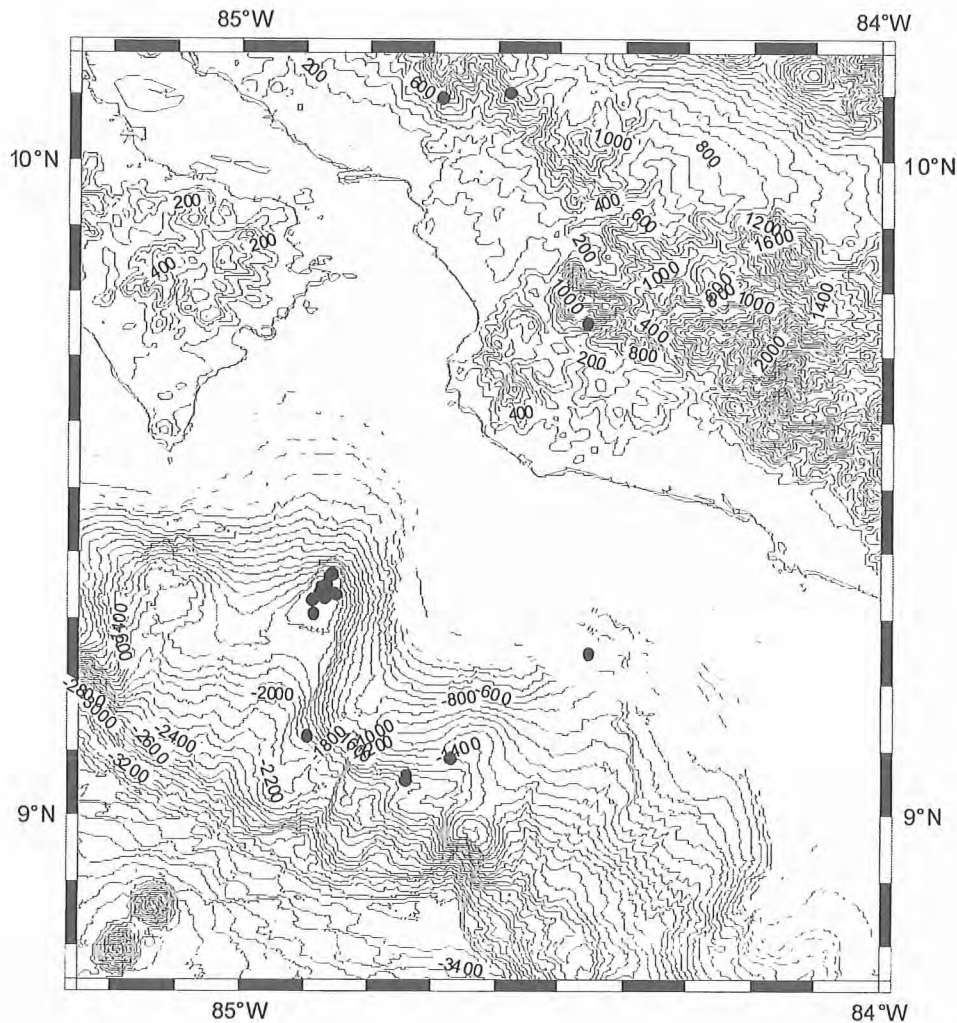


Figure 6.6.5.3: Earthquake epicenters that have been located with the JACO seismic network during April 23 and May 8, 2002 (first deployment)

The following pages show examples of earthquakes recorded at the JACO net array. They serve to illustrate the signal quality to be expected from the different instruments. In general, P phases can be clearly identified on all instruments in the unfiltered records, however, seismometer records are often dominated by long period microseismic noise. Fortunately, the dominant frequencies of this noise (<1 Hz) are sufficiently removed from the dominant band excited by local earthquakes (>3-4 Hz, see Figures 6.6.5.4-9) to be easily filtered out, or, for high precision onset and polarity picking (which can be affected adversely by filtering), to only appear as a trend or DC shift after zooming in on an arrival with a window size of only tenths of a second. For a first inspection, a 10-49 Hz bandpass filter was found to be the most suitable.

S-waves could occasionally be picked, too. As expected, these phases were quite clear on the seismometer data, particularly on the horizontal components, but much harder or impossible to see in the hydrophone records.

In the following plots, the event origin time and hypocentral co-ordinates as determined by phase picks on the ocean bottom stations are indicated by the heading above the seismograms. Component label BHH is used to indicate hydrophone channels, while BHZ, BH1, and BH2 are the vertical and two horizontal seismometer components. Amplitude and polarity of all data are 'raw', i.e., uncorrected for instrumental response, preamplifier gain, and polarity flips. All times are UTC. First, an event from the subducting plate is shown, then examples of oceanic events.

JACO 2002-05-01-0518-38S.JACO_010
 Plot start time: 2002 5 1 5:21 23.166

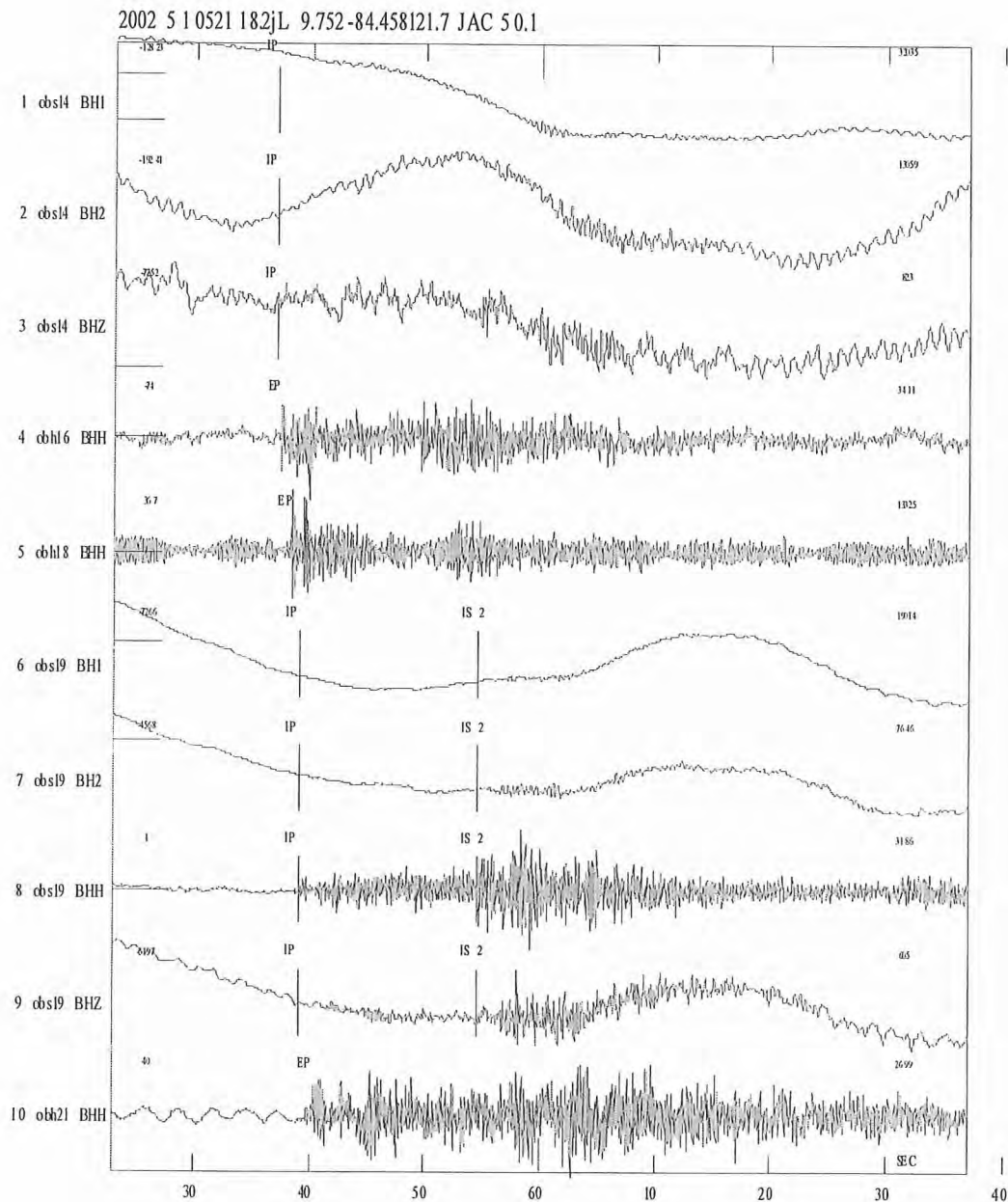


Figure 6.6.5.4: Earthquake from May 21, 2002 at 5:21 UTC from the subducting plate at a depth of about 122 km located at the north-western edge of the land stations network. All ocean bottom stations show up with good signals except obh15 and obh20 which have technical problems (see Table 6.6.5.1) and are not shown here. The seismometer channels can be easily recognized by their long-period signals, while the hydrophone channels yield only high frequent signals. Clear P and some S onsets have been observed, while obh18 shows up with a strong secondary onset one second after P, which maybe a depth phase.

JACO 2002-05-01-0518-38S.JACO_010
 Plot start time: 2002 5 1 5:21 22.992 Filt: 10.000 49.000

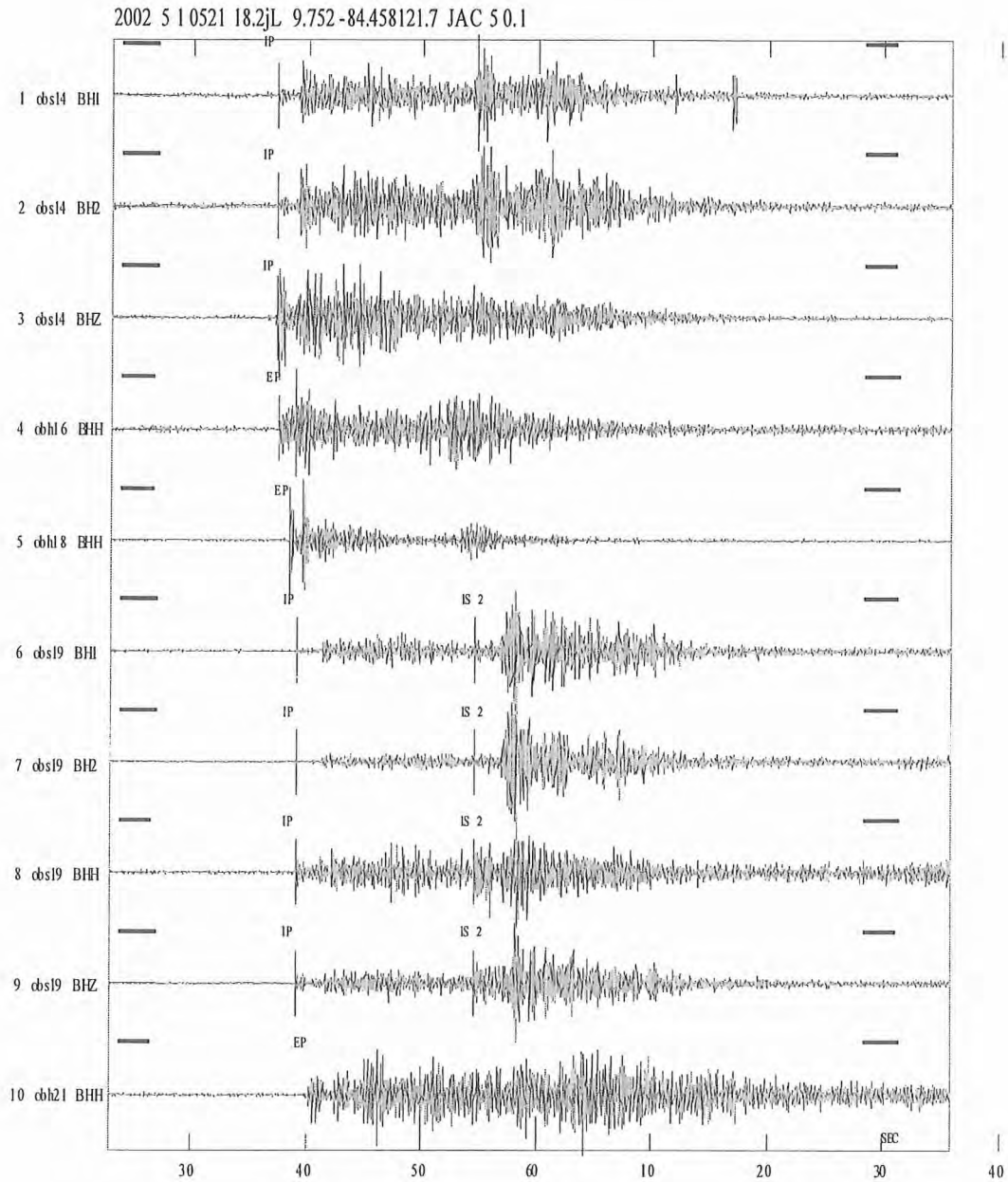


Figure 6.6.5.5: The same event from previous figure, but filtered by a 10 to 49 Hz bandpass filter. Now, all components show up with a clear first arrival and the S wave can be clearly seen on the seismometer channels, but also on some hydrophone channels.

JACO 2002-04-26-2224-57S.JACO_005

Plot start time: 2002 4 26 22:27 36.697

2002 426 2227 51.1jL 9.370 -84.857 31.2 JAC 5 0.4 1.8CJAC

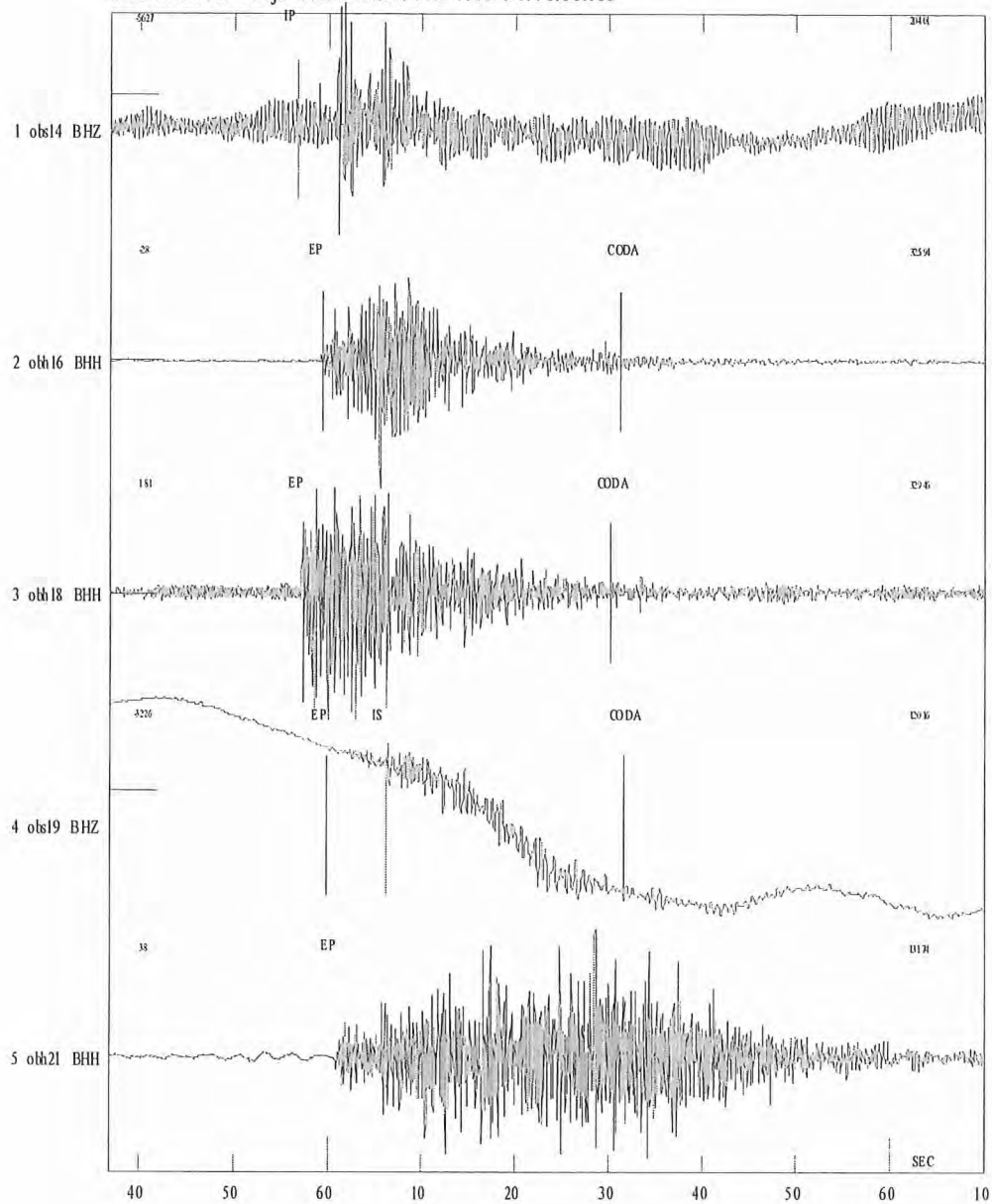


Figure 6.6.5.6: Swarm earthquake from April 26, 2002 at 22:27 UTC located where a large seamount has already entered the continental lithosphere. The focal depth of about 31 km is more than questionable, not only because only one S-pick could be made, but also because of the majority of the swarm events have shallow depths.

JACO 2002-04-26-2224-57S.JACO_005
 Plot start time: 2002 4 26 22:27 35.590 Fil: 10.000 49.000

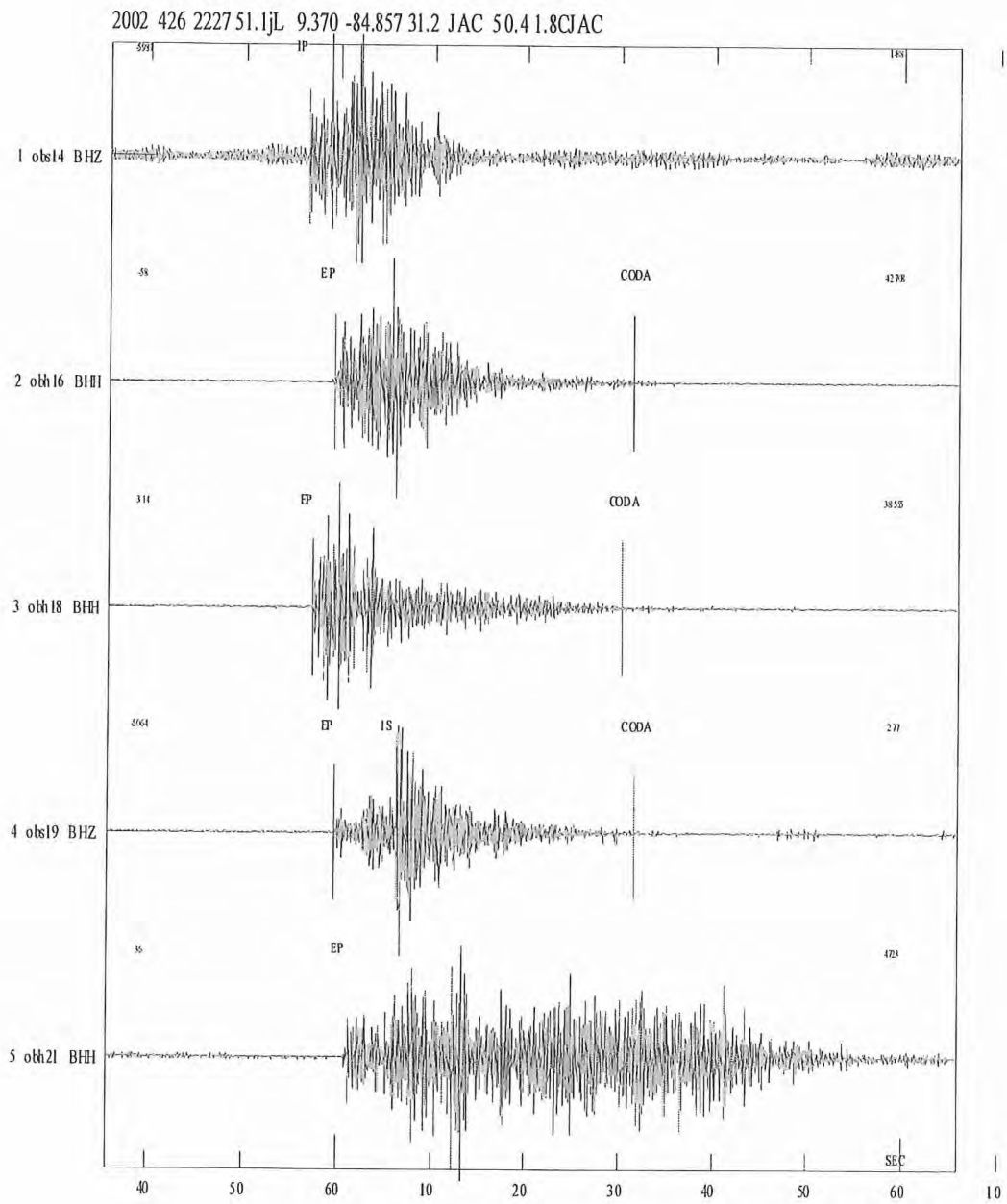


Figure 6.6.5.7: Event from previous figure, but filtered by a 10 to 49 Hz bandpass.

JACO 2002-04-26-2159-23S.JACO_005

Plot start time: 2002 4 26 22: 2 7.298

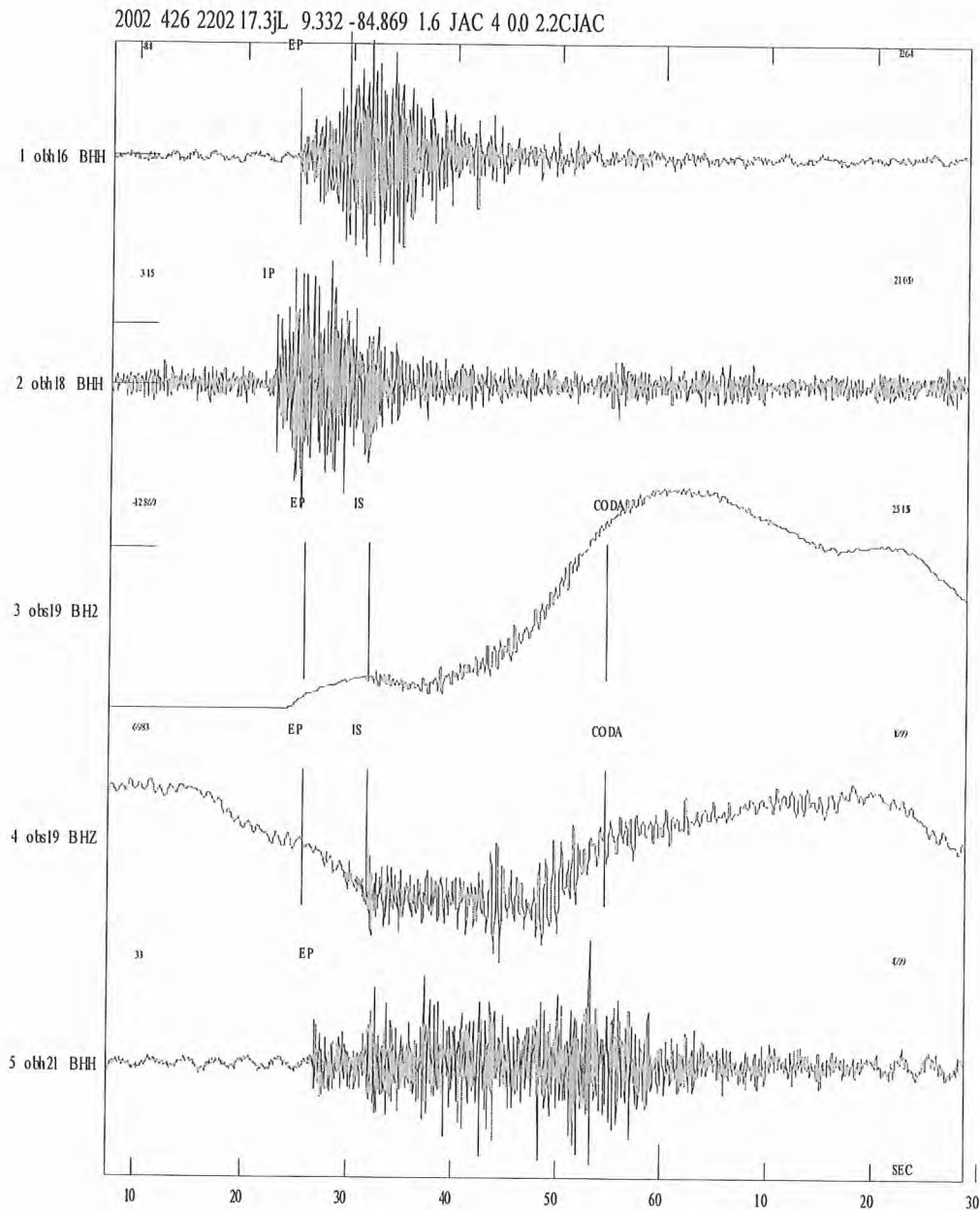


Figure 6.6.5.8: Another event from the earthquake swarm from April 26, 2002 at 22:02. Focal depth is about 2 km. Trace 3 from obs19 horizontal component demonstrate that it is sometimes possible to use corrupted seismograms, if the interesting part is not effected by distortion. Until 8 seconds before the P onset trace 3 was out of range, and recovered just in time for registering the earthquake.

SO-163-2: Profiles P38-P40 (Jaco Scar Net)

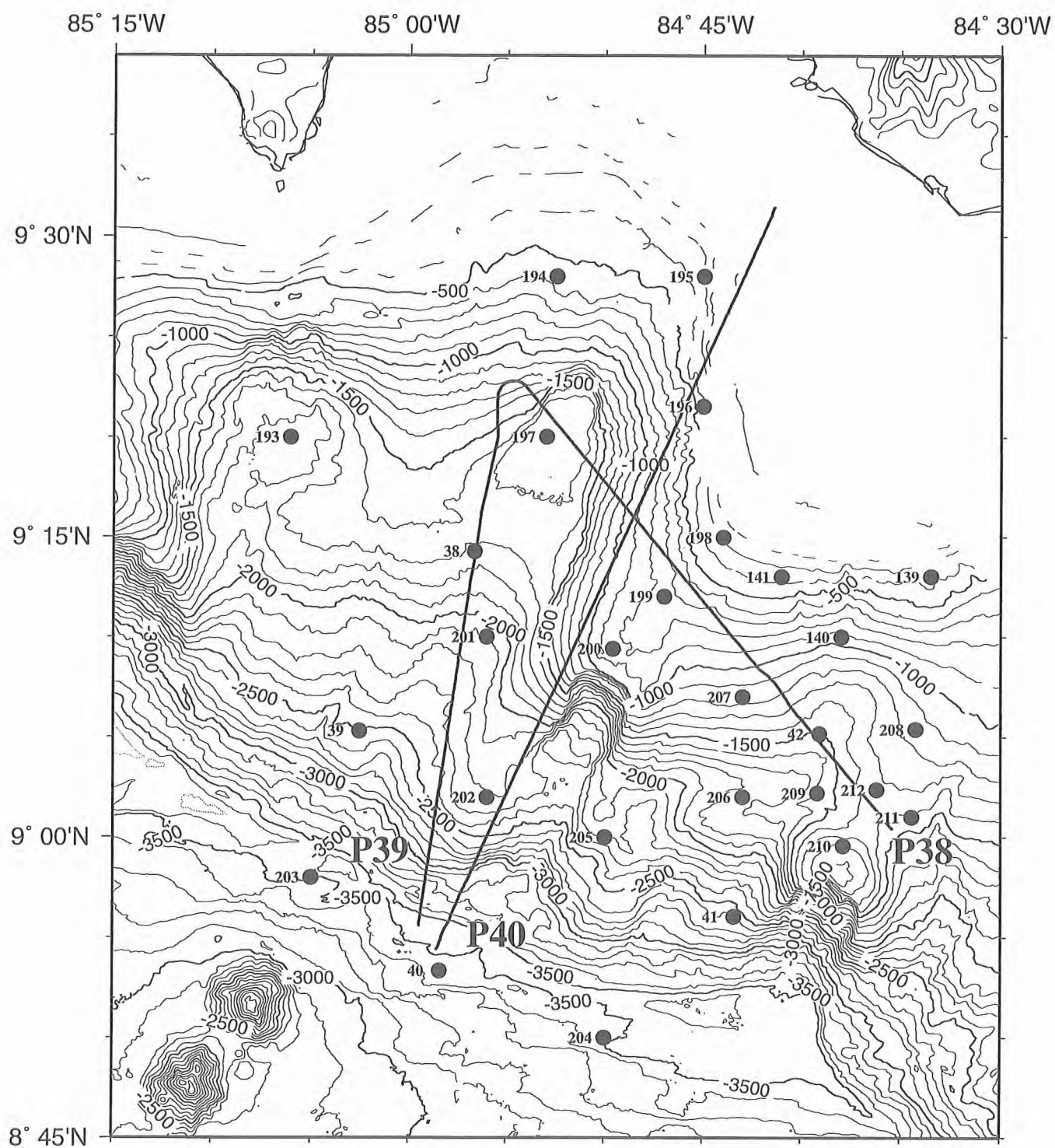


Figure 6.6.5.9: Location map of profiles P38-P40 with 100 m isolines and Jaco Scar Net stations.



Figure 6.6.5.10: Record section from str 938 (Streamer,ch2), Profile 38.

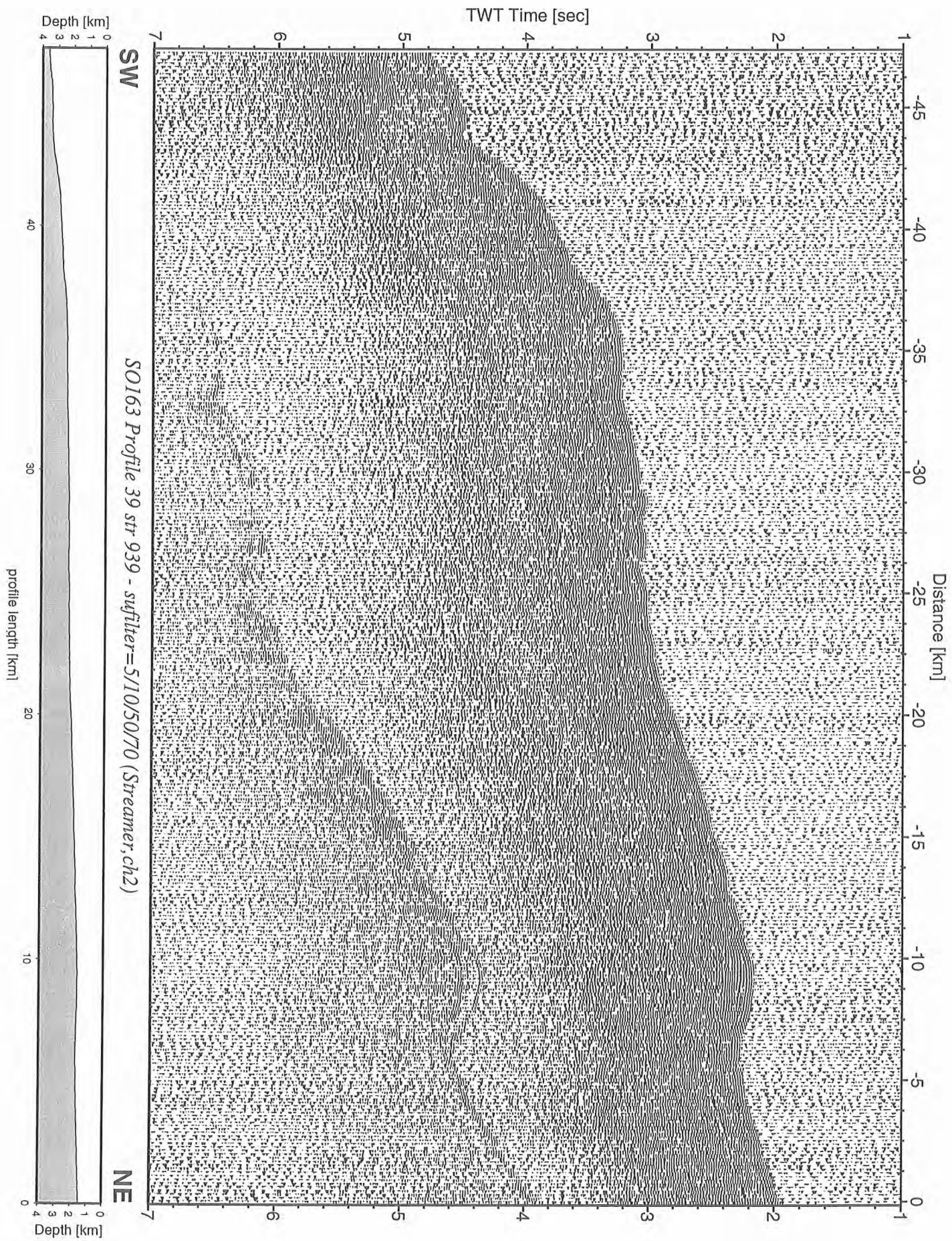


Figure 6.6.5.11: Record section from str 939 (Streamer,ch2), Profile 39.

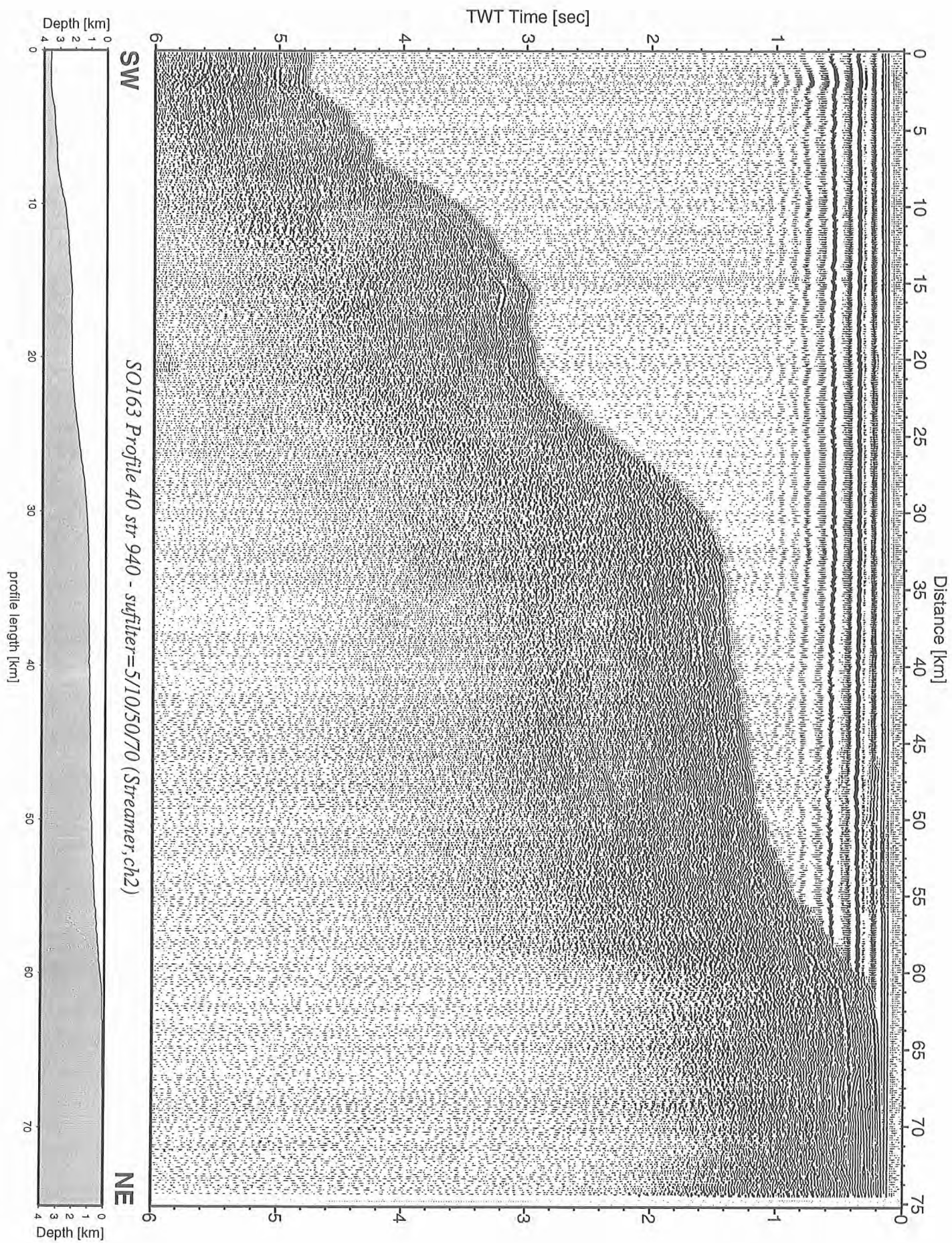


Figure 6.6.5.12: Record section from str 940 (Streamer,ch2), Profile 40.

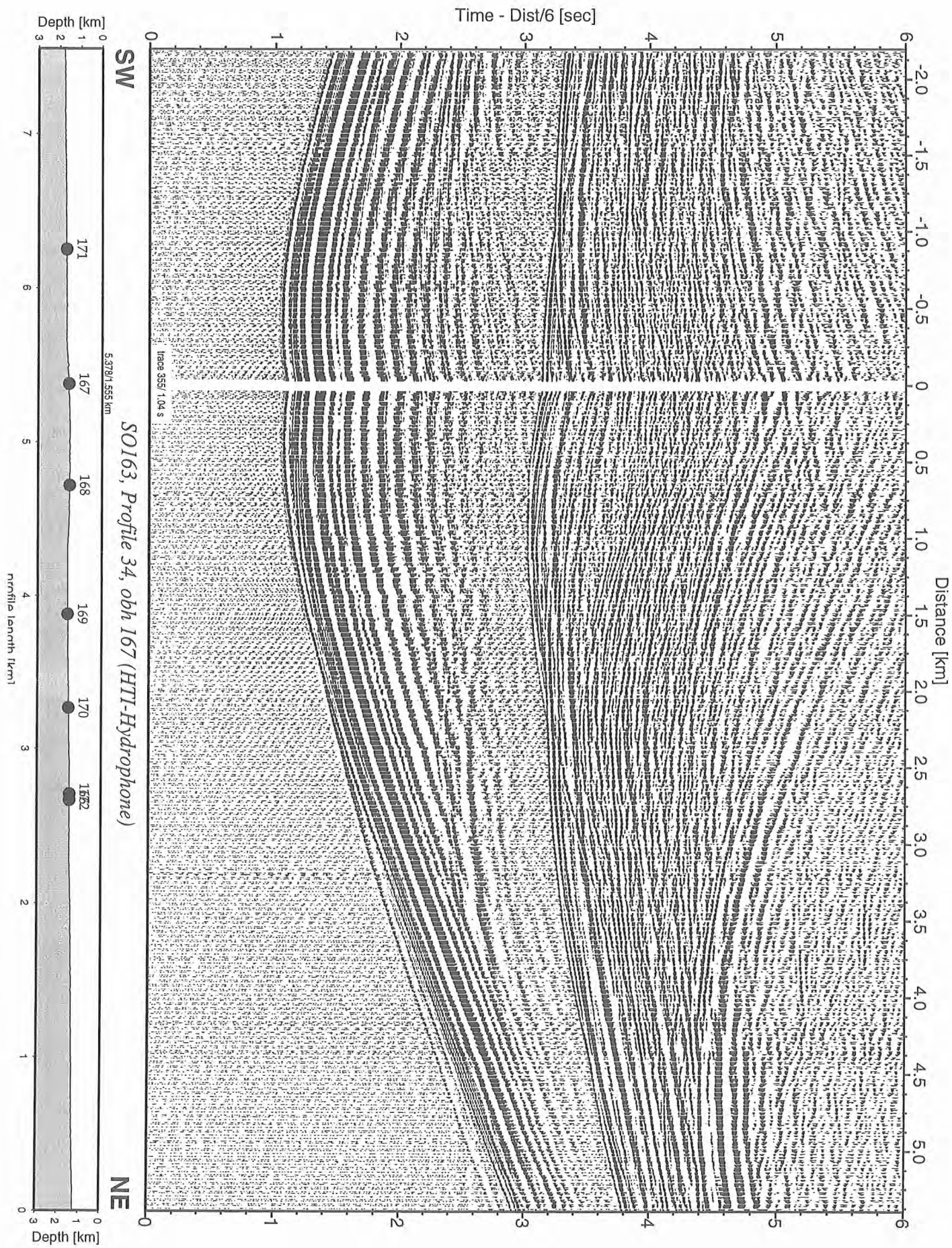


Figure 6.6.6.2: Record section from obh 167 (HTI-Hydrophone), Profile 34.

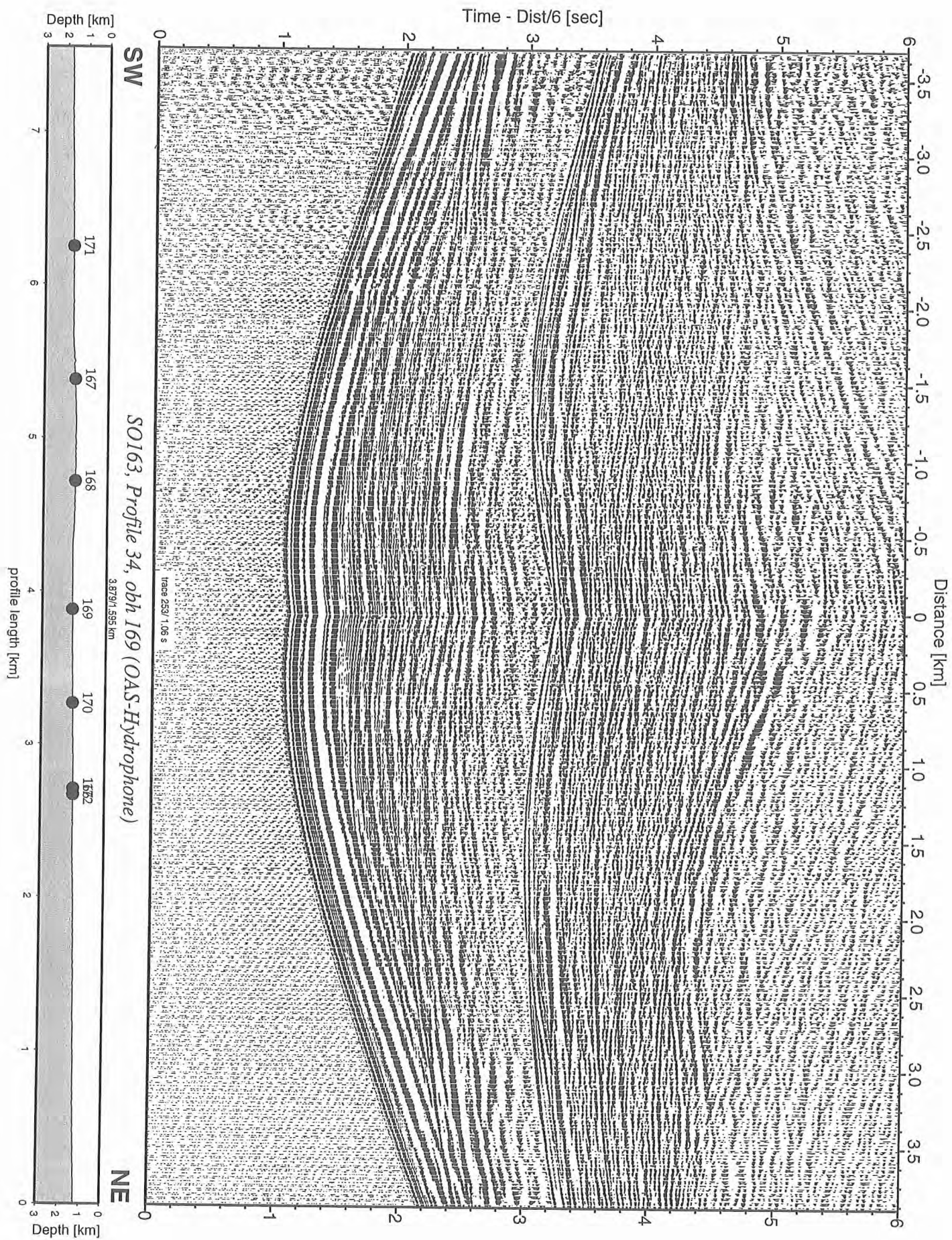


Figure 6.6.6.3: Record section from obh 169 (OAS-Hydrophone), Profile 34.

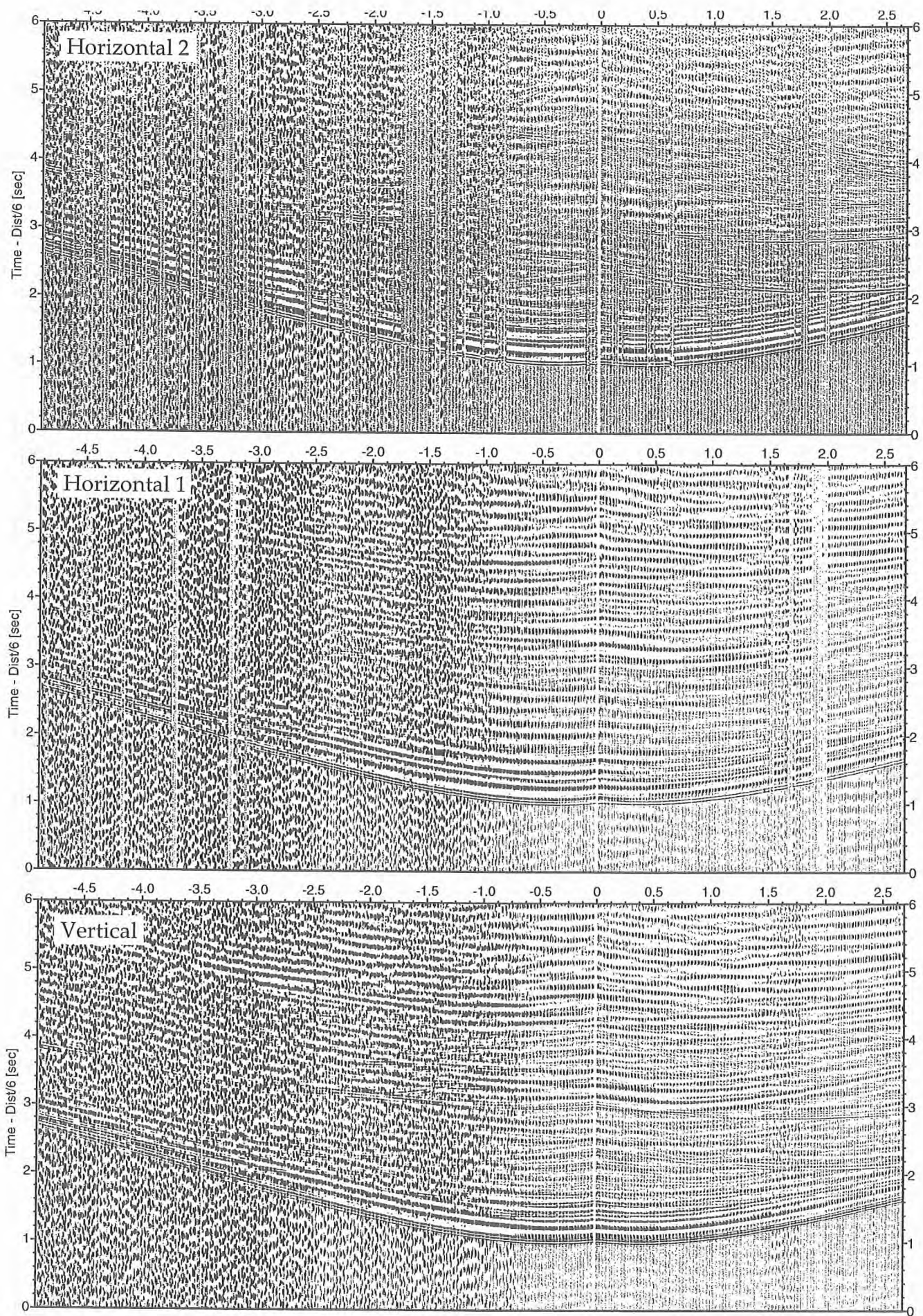


Figure 6.6.6.6: Record sections from obs 172 (PMD), SO163, Profile 34.

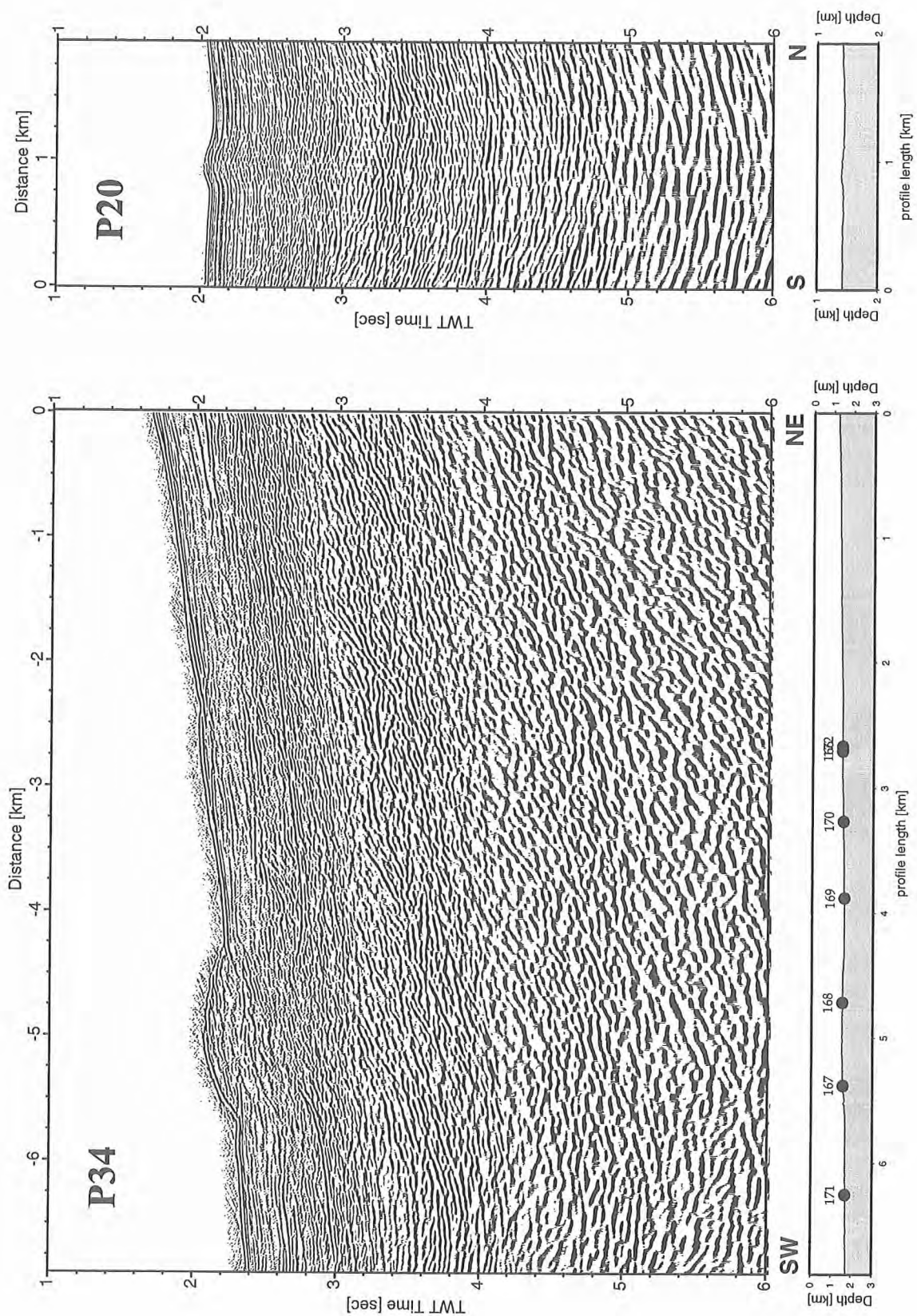


Figure 6.6.6.7: Post-stack time migration of profiles P20 and P34 across Mound Culebra and Mound Quepos.

6.7 Water column and upper sediment investigations

(G. Rehder, S. Grandel, S. Mau, and K. Stange)

6.7.1 Introduction

Water column program

An intensive sampling program was performed during SO 163/2 to investigate the methane and oxygen distribution within the water column. Methane is one of the key components cycled in the process of dewatering at active continental margins. It can be oxidized within the upper sediment, can provide the carbon and energy source for chemoautotrophic communities, or can be injected into the water column. All these processes have to be assessed to close the carbon budget at active seep sites. The methane injected into the water column will be oxidized on time scales from decades to days [Rehder et al., 1999; Valentine et al., 2001], strongly dependent on the surrounding and amount of methane available. Strong isotopic fractionation occurs during the oxidation process, which allows insight in the processes controlling the methane pathways.

Due to the relatively short lifetime and its generally low background in deep waters, the analysis of methane is ideally suited to locate currently active vent sites. Other than for the ionic constituents, methane can be transported both within the fluids in dissolved form, or by transport of free gas in case the pore waters become over-saturated. In the latter case, the transport is independent of the transport of other fluid components, and can lead to transfer to relatively shallow depth levels [Rehder et al., in press].

While today's ocean can be considered as a negligible source of methane to the atmosphere, its cycle, turnover times and transport pathways are still an important topic of marine research. Large excursions in the carbon isotopic signature of benthic and planctonic foraminifera have been suggested to be caused by gas hydrate destabilization [Dickens et al., 1995]. All models used to assess the importance of a large scale hydrate decomposition for ancient and future climate scenarios lack an understanding of the role of the ocean to mitigate the direct transfer of the gas to the atmosphere [Dickens et al., 2001]. Seepage from hydrate bearing sediments is thus the ideal setting to investigate the role of the ocean in what is now known as the "clathrate gun hypothesis".

Sediment investigations

Investigations of the geochemical composition of porewaters provide information to investigate the forces and impacts driving redox-, and remineralization processes within the upper sediment column. During the SO 163/2 cruise, pore water composition of surface sediments was investigated at about ten locations to get a first insight into sediment diagenetic processes in the subduction zone off Costa Rica. Concentration and distribution of the parameters total alkalinity, ammonia, chloride and sulphide in the porewater were determined in surface sediments. This allows to identify locations influenced by seepage and to assess the effect of methane formation and decomposition regarding sediment diagenetic processes.

Distinctive colonies of clams as well as precipitates of authigenic carbonates, resulting from biogeochemical turnover and interaction between fluids and ambient bottom water, are found at cold vent sites [e.g. Suess et al., 1985, Suess and Whiticar, 1989, Henry et al., 1992, Torres et al., 1996, Suess and Bohrmann, 1997]. Bacterial mats have also been identified and can be indicative of subsurface gas hydrate pavement [Bohrmann et al., 1998]. Preliminary OFOS-track investigations executed during SO163/1 were used to determine areas where active seepage could be expected. Sediment samples were collected at these sites, in addition to reference sites reflecting 'normal' off-shore and deep-sea situations at similar water depths. First results of the geochemical pore water studies executed during this cruise are reported.

6.7.2 Material and Methods

Water column sampling and analysis

The CTD unit as well as the 24 bottle water column carousel worked well throughout the cruise, with the exception of major problems with the performance of the oxygen sensor at the beginning of the cruise. After CTD station 10 (and using the fifth oxygen sensor), problems were resolved. At all stations, samples were taken for shipboard analysis of oxygen and methane, and 2 sets of 250ml water samples poisoned with 0.5ml of saturated HgCl were sampled into crim cap glass bottles sealed with a butyl rubber septum.

Oxygen was measured using the method developed by Winckler as described in Grasshoff et al. (1997). Usually, three Niskin bottles were sampled twice to control reproducibility of the measurements. The Seabird oxygen sensor worked well after CTD-station 10, but had been replaced several times before due to electronic noise or calibration inconsistencies. After CTD station 10, the absolute values of oxygen sensor and Winckler titration differed by less than 10 % (Fig. 6.7.2.1). However, both data sets will be further improved. The addition of the (oxygen containing) chemical reagents to the samples leads to a blind value which has not been subtracted yet. The output of the oxygen sensor is calculated from the electrode current as well as the change in current over time to adjust for the response time of the instrument. During in situ oxygen monitoring, the slope $dC_{(ox)}/dt$ can be calculated only by looking backward in the recording data. The slope can be reprocessed with the SEASAT command DERIVE, which looks at the data before and after an individual point to derive the slope. Both data sets will be reprocessed shortly after the cruise and will be incorporated into the SFB 574 data base.

For CH₄ analysis aboard, a modification of the vacuum degassing method described by *Lammers and Suess* [1994] was used [Rehder et al., 1999]. 1600 ml of water were injected into pre-evacuated 2200 ml glass bottles, which leads to almost quantitative degassing. The gas phase was subsequently recompressed to atmospheric pressure and the CH₄ mole fraction of the extracted gas was determined by gas chromatography. A Shimadzu GC14A gas chromatograph equipped with a flame ionization detector was used in connection with a Shimadzu CR6A Integrator. Nitrogen was used as carrier gas, and separation was performed using a 4m 1/8" SS column packed with Porapak Q (50/80 mesh) run isothermally at 50°C. The total gas content of the sample was calculated from the measured dissolved oxygen concentration and assuming that N₂ and argon are 100% saturated relative to their atmospheric partial pressures [Weiss, 1970]. The dissolved methane concentration was calculated as the product of the mole fraction in the extracted gas phase and the amount of total gas (STP) in the sample. For calibration, mixtures of 1.936ppm ±0.003 ppm and 9.854 ±0.006 ppm in synthetic air (Deuste Steininger, calibrated against NOAA/CMDL standards at the Institute for Environmental Physics, Heidelberg) were used.

The 2 sets of 250ml water samples will be transported to Kiel at 4 °C after SO164. One set will be used to measure methane concentration with a standard purge and trap technique for intercomparison. The other will be used for the analysis of the stable carbon isotopic ratio of methane ($\delta^{13}C$) using the new MAT 253 Mass Spectrometer which will be delivered to GEOMAR in early summer.

Sediment investigations

Sediments were sampled using a minicorer (see section 5.8). The cores, usually containing 25cm of sediment, were sliced in segments of 1cm from 0-10cm, and in segments of 3cm below that level.

Pore waters were separated from the sediment by low-pressure squeezing (1-3 atm) using argon gas, executed in the ship's cold room at 4°C. The analyses performed on the pore waters of the different

cores and the analytical techniques used are listed in Table 1. Modifications of pore water analysis are necessary for samples with high sulphide concentrations.

Table 1: Techniques used for pore water analyses.

Constituent	Method	Reference
Alkalinity	Titration	Ivanenkov and Lyakhin [1978]
Ammonium	Spectrophotometry	Grasshoff et al. [1997]
Chloride	Mohr (AgNO_3) titration	Gieskes et al. [1991]
Hydrogen sulphide	Spectrophotometry	Grasshoff et al. [1997]

Ammonia and sulphide

Concentrations of dissolved ammonia were determined using standard photometric procedures described in Grasshoff et al. [1997]. In order to remove the H_2S in samples containing high concentrations of sulphide, a certain volume (1-5ml) was acidified with 65% suprapure HCl (10 μl /ml sample), and then left open for 48 hours in the cold room.

The method for sulphide determination according to Grasshoff et al. [1997] has been adapted for pore water concentrations of S^{2-} in the millimolar range. For reliable and reproducible results, an aliquot of pore water was diluted with appropriate amounts of oxygen-free artificial seawater, the sulphide was fixed by addition of zinc acetate gelatin solution immediately after pore-water recovery. After dilution, the sulphide concentration in the sample should be less than 50 $\mu\text{mol/l}$.

Titration of Cl

High concentrations of H_2S (> 1mM) in the sample disturb the measurements. Therefore, these samples were pre-treated with a 1:1 dilution of 0.01 N suprapure HNO_3 and stored for 1-2 days, without lid, in a cool room. Subsequent titration with silver nitrate is standardized against IAPSO seawater or a 1:1 dilution of IAPSO seawater for samples containing sulphide.

Total alkalinity (TA)

Samples of the sediment pore water for total alkalinity measurements were analyzed by titration of 1 ml pore water (+ 4 ml MilliQ and 20 μl indicator solution) with 0.02 N HCl in a special titration vessel according to Ivanenkov and Lyakhin [1978]. Titration was finished until a stable pink color occurs. During titration the sample is continuously degassed with argon to remove the generated CO_2 or H_2S . The acid was standardized using an IAPSO seawater solution.

This method is well suited for samples containing H_2S and those with TA values largely exceeding 2.3 mM, because the CO_2 and H_2S is removed during the titration by the continuous stream of argon. A mixture of methylene blue and methyl red was used as indicator.

Shore based laboratory analyses

Acidified subsamples (35 μl suprapure HCl 30% + 2.5 ml sample) were prepared for ICP analyses of major cations (Na^+ , K^+ , Li^+ , Mg^{2+} , Ca^{2+} , Sr^{2+} and Mn^{2+}). Sulphate, DIC, $\delta^{18}\text{O}$ and $\delta^{13}\text{C}$ of CO_2 will be determined on selected sub-samples in the shore-based laboratories.

6.7.3 Preliminary results

6.7.3.1 Overview

Instrument deployment

A total of 24 CTD stations were deployed, of which 14 were performed with the mini corer attached underneath the CTD/rosette unit (Fig. 6.7.3.1.1 & Table 6.7.3.1.1). The investigations focussed on four key objectives. In order to understand the background distribution of methane within the water column, two hydrographic sections of about a 100 nm length were sampled in the northern and

southern part of the research area (see Fig. 6.7.3.1.1, CTD stations 1-6 and 15-20). The simultaneous deployment of the minicorer also allowed sampling of the upper sediment in areas not influenced by seepage phenomena. In addition, three areas, which appeared likely to be influenced by active seepage, were investigated. The choice of the locations was based on the results of the OFOS/CTD survey during SO144 [Bohrmann et al., 2002], and the detailed OFOS investigation during the first leg of SO 163 (see chapter by Sahling et al. in this cruise report). Five stations were dedicated to investigate the methane distribution and upper sediment pore water properties at Jaco Scarp. Four stations were used to monitor the methane distribution at Quepos Mound, a small scale mound which has been explored during SO 163/1 for the first time. Two stations were dedicated to investigate the water column methane distribution and the upper sediment layer at mound Culebra, a mound with a diameter of nearly 1nm and an elevation of more than 100m. During all of the stations at the geological structures, the ship's own SSBL transponder was attached at the cable about 40m above the CTD/rosette unit, which allowed exact positioning during the deployment of the mini corer and the water column sampling. The movement of the device at depths also gave valuable hints of the deep water current direction, which was very valuable for the choice of the following station locations.

6.7.3.2 Hydrographic Sections

There are a variety of scientific reasons that made it essential to map the methane distribution in the water column without focussing on potential seepage sites. One is that in order to quantify the excess methane in a plume caused by tectonic dewatering, one has to know the background concentration at that depth (or density) level. Very often, sharp concentration maxima are found at seepage sites, but there is also a relatively small enhancement of the CH₄ concentration in a broad layer, which might be of similar importance for the budget, but needs detailed knowledge of the background concentration to be even recognized. However, apart from defining the background for seepage investigations, the biogeochemical cycling of methane in the area is an interesting topic of marine research by its own. The geostrophic, seasonally wind enhanced upwelling west of Costa Rica and Nicaragua, known as the Costa Rica Dome [Fiedler et al., 2002], leads to enhanced biological production and a shoaling of the thermocline. Oxygen is completely depleted around the core of the oxygen minimum layer, and suboxic conditions are found nearshore at depths shallower than 100m. It has recently been shown for a section off Nicaragua that the anoxic conditions lead to methane production, generating one of the most far-reaching CH₄ plumes in the ocean [Sansone et al., in press]. The authors also showed that there is isotopic evidence for another, deeper CH₄ source on the shelf. In addition, upwelling has been shown to cause an enhanced flux of methane to the atmosphere, both due to enhanced in situ production [Bange et al., 1998] as well as due to the upwelling of water enriched in methane due to bottom sources [Rehder et al., in press].

The general distribution of methane and oxygen is illustrated in Fig. 6.7.3.2.1 by the profiles recorded at Station 18 at about 2500m water depth. A pronounced maximum in the methane concentration exists at about 400, centered around the oxygen depletion zone. A slight CH₄ oversaturation of the surface waters is driven by a narrow, sharp subsurface maximum, with concentrations up to 6x the saturation with the ambient atmosphere. Subsurface maxima are a common observation in low latitude surface waters [Karl and Tilbrook, 1994], and were shown to be generated by anoxic production within particles and fecal pellets. However, the feature is unusually pronounced in the research area. The extremely low oxygen content at shallow depth in combination with high surface production seems the most likely explanation for this finding.

Sections of the methane distribution within the upper several hundred meters along the northern and southern section are represented in Fig. 6.7.3.2.2. Methane saturation with respect to the ambient atmospheric partial pressure is plotted rather than concentrations to eliminate the effect of the temperature dependence of methane solubility [Wanninkhof, 1992]. The methane inventory seems to

diminish from the upper slope to the open ocean in the northern section (a). The southern section shows the highest subsurface concentrations towards the upper slope, while the maximum around the oxygen minimum shows no such trend. Both layers are slightly shifted to greater water depth towards the open ocean.

Sediment

Sediment cores were sampled at four stations, three of them located in the northern section (MUC-1, 2, 3) and one located in the southern Section (MUC-9) to get pore water data characterizing off-vent sites. The four stations are located in water depths similar to the sites, examined in the potential seepage areas with the aim to get information of the geochemical background situation in the area. Water depths of the investigated stations were 420m (MUC-9), 1340m (MUC-1), 2395m (MUC-3) and 3215m (MUC-2).

All cores were characterized by a dark, greenish colour. Differences were obvious in the texture. The cores of the northern section in water depths > 1300m consist mainly of silt and clay, whereas the major component of the core sampled in 400m water depth was sand, presumably due to the deposition of terrigenous material. The cores in the northern section showed many tubes from the surface down to a depth of 10cm, indicating high degree of bioturbation. MUC-9 was almost not bioturbated, the sediment contained a lot of small muscle shells (size 1-2cm) in depths > 7cm. All cores showed no visible indicators of seepage fauna.

First results of pore water analysis are shown in Fig.6.7.3.2.3. Concentrations of sulphide remains < 3 $\mu\text{mol/l}$ in all cores throughout the investigated sediment column (the data of MUC-9 is not shown; concentrations were below the detection limit). Concentrations of chloride ranged between 525-570 mmol/L for MUC 1 and 2, and between 550 – 580 mmol/L for MUC 3 and 9. The profiles of the latter scatter to a lesser extent and are characterized by slightly higher Cl^- concentrations near the surface and more or less constant values throughout the remaining sediment column. The scattering in the profiles of MUC 1 and 2 might be induced by analytical procedures. Besides these small differences and the scattering, all chloride profiles range within Cl^- concentrations of the IAPSO seawater standard (557 mmol/L), and show no anomalies, as expected for non-vent sites.

Ammonia concentrations were in general < 200 $\mu\text{mol/L}$. At the deepest site (MUC-2), NH_4^+ concentrations remain more or less constant <10 $\mu\text{mol/L}$. At the other locations ammonia concentrations raise from 5 $\mu\text{mol/L}$ in the bottom water to up to 40 $\mu\text{mol/L}$ (MUC-9) and about. 150 $\mu\text{mol/l}$ (MUC-1, MUC-3), respectively. These slight variations in the ammonia distribution might be caused by variations in bottom water oxygen concentrations and differences the content of organic matter.

Total alkalinity (TA) in the bottom water is about 2.6 mmol/L, concentrations increase up to 3 mmol/L in the cores 2 and 9, and up to 4 mmol/L at the stations located in 1400m and 2400m water depth. Lower TA values in the former stations are probably due to a lesser amount of degradable organic material at these sites, but further interpretation remain open until the analysis of organic carbon content etc. on land.

In summary all stations investigated in the northern and southern section are characterized by pore water chemistry as expected for a reference site not influenced by active seepage.

6.7.3.3. Jaco Scarp

Water column

The Jaco Scarp is created in consequence of the subduction of a seamount on the oceanic plate. This leads to a circular uplift with faulting at the upper plate (dome). Seawards, the process is associated with sedimentary landslides, leading to the formation of steep scarps. During expedition SO 144,

chemoautotrophic clam colonies had been observed at 1800 and 1900m depths, and a CTD cast revealed the highest concentrations of the entire water column program corresponding to this depths level. We reconfirmed the presence of a methane plume between 1700 and 1850m water depth with concentrations up to 270 nmol/L centered at 1780m at the station visited during SO 144 (their CTD station 2, here CTD/MUC 07 – (04). Thus, it was decided to map the CH₄ distribution along the slope of Jaco Scarp. CTD/MUC was deployed at a total of 5 stations within the Jaco Scarp area with the aim to map the extend of the plume as well as to look for indication of other vent sites within the Scarp Fig. (6.7.3.3.1).

Station CTD/MUC 08 – (05) was located at a position where the occurrence of bacterial mats and clams had been reported before [Bohrmann et al., 2002]. However, the station showed no evidence for seepage in the upper part of the slope (bottom depth 1400m). The same proved to be true at the northern sill of Jaco Scarp (900m). Sample analysis of station CTD/MUC 21 – (11) at a water depth of 2255m, however, indicate a water column enriched in methane everywhere in the depth range between 1600 and 2100m, including a narrow layer with peak concentrations near 100 nmol/L at 1930 m depth. Both plumes can still be recognized at station CTD/MUC 24 – (14) further to the south. An overview of the location of the stations along the slope and the observed methane concentration pattern is given in Fig. 6.7.3.3.2.

While the aim of the sampling was to assure that the stations with greater water depths always cover the water masses which showed an enhanced methane budget at shallower stations, it is not sure that this has been achieved at station CTD/MUC 21 – (11). Water samples were taken up to a water depths of 1600m, while the plume found at station CTD/MUC 07 – (04) had its upper boundary at 1700m. However, it appears that we did not cover the density layer of the core of the plume at station CTD/MUC 07 – (04) (Fig. 6.7.3.3.3). The reason for this phenomenon is a very different density layering in these two stations (Fig. 6.7.3.3.4), suggesting a downward sloping of the isopycnal surface $\sigma_\theta = 26.63$ of more than 200 m towards the northern wall of Jaco Scarp. This would be a very unusual finding which will need further investigation and check of data consistency. However, station CTD/MUC 24 – (14) was sampled with a detailed look at the density structure of the water column, and all density layers where plumes had been observed at the 2 shallower stations were sampled.

Sediment

Multiple cores at the subducting seamount of Jaco Scarp were sampled according to preliminary studies published by Bohrmann et al. (2002). Deployments were executed at the head of the seamount (MUC-6), the uplifted area over the scar (MUC-5) and within the slide structure (MUC-4, MUC-11), where vent fauna and carbonates had been observed during OFOS tracks. The instrument was deployed further down the slope (MUC –14) at the very end of the cruise. However, these measurements will not be finished for this report.

Figure 6.7.3.3.5 shows the results of the cores deriving from the deployments at 1400m (MUC-5) and 2200m (MUC-11). Texture and colour of all cores were similar to the ones found at the reference sites, and no vent-indicating fauna was visible. Concentration and distribution of the parameters sulphide, chloride, ammonia and total alkalinity correspond with the geochemical characteristics measured at the reference sites. The results suggest no active venting at the sampled locations. Further investigation in this area, especially by TV-guided instruments, will be necessary to get further insight into the regional distribution and scaling of the seepage activity and characteristics at Jaco Scarp, which is evident from the results of the OFOS surveys and water sampling program.

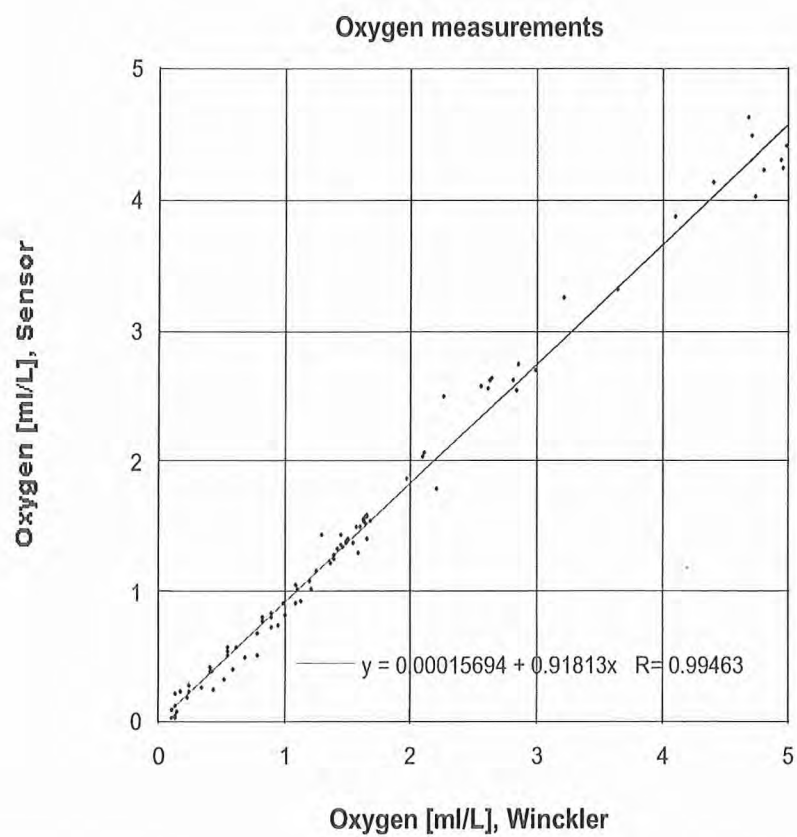


Fig 6.7.2.1 Results of the measurements of dissolved oxygen concentrations of stations 14, 17, 18, and 23 vs the output of the Beckmann oxygen sensor attached to the CTD unit.

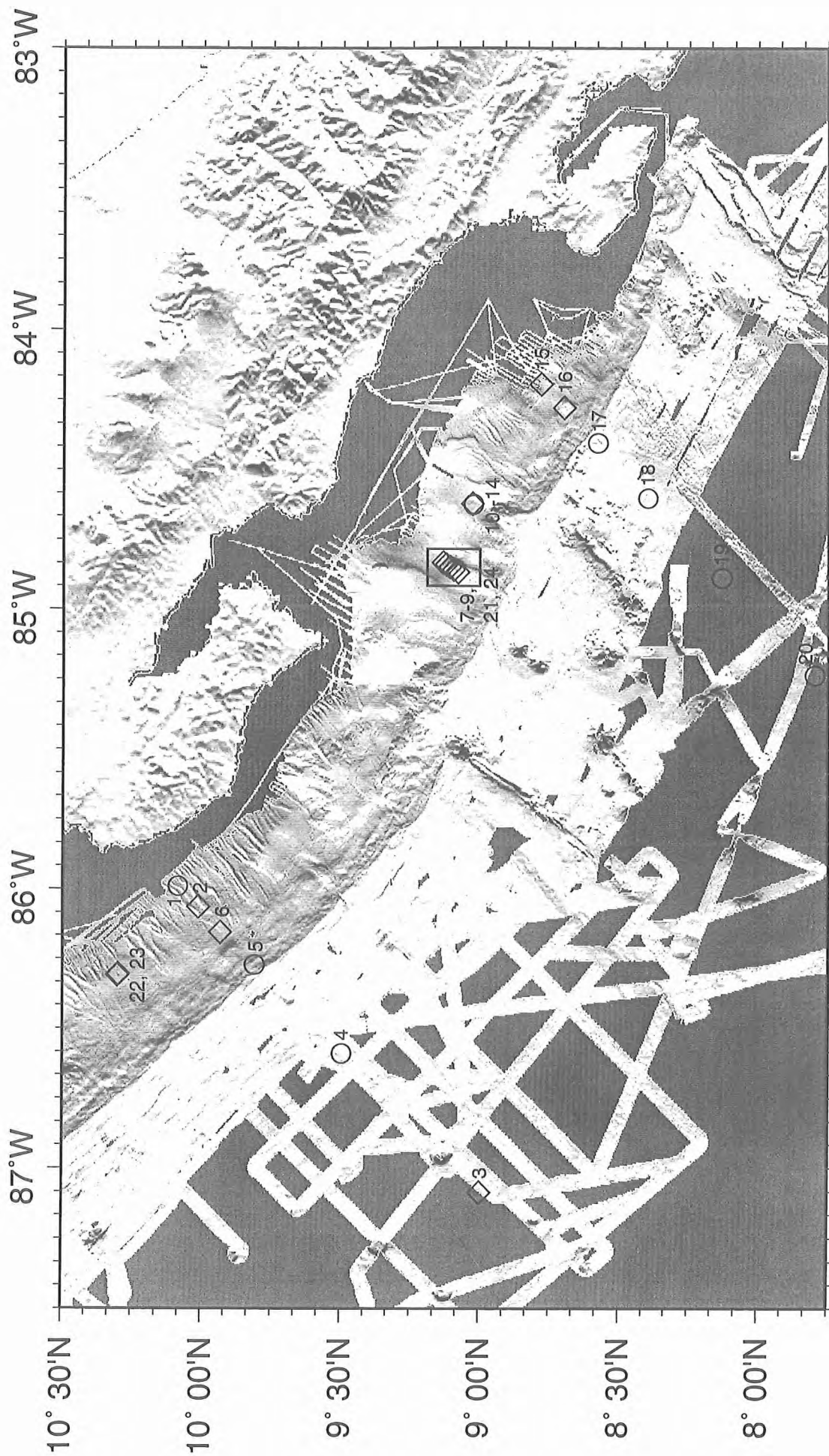


Fig. 6.7.3.1.1 Station locations of the CTD/MUC program during SO 163/2. CTD stations are indicated by circles, CTD/MUC stations by diamonds. Overlying symbols from investigations at small scale geological structures are shown in detail maps in the following chapters.

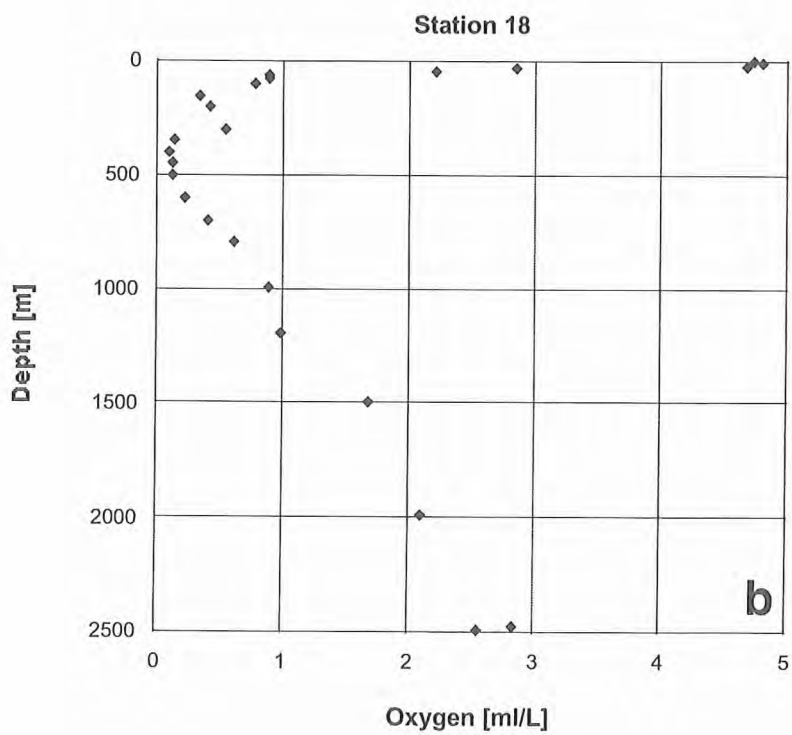
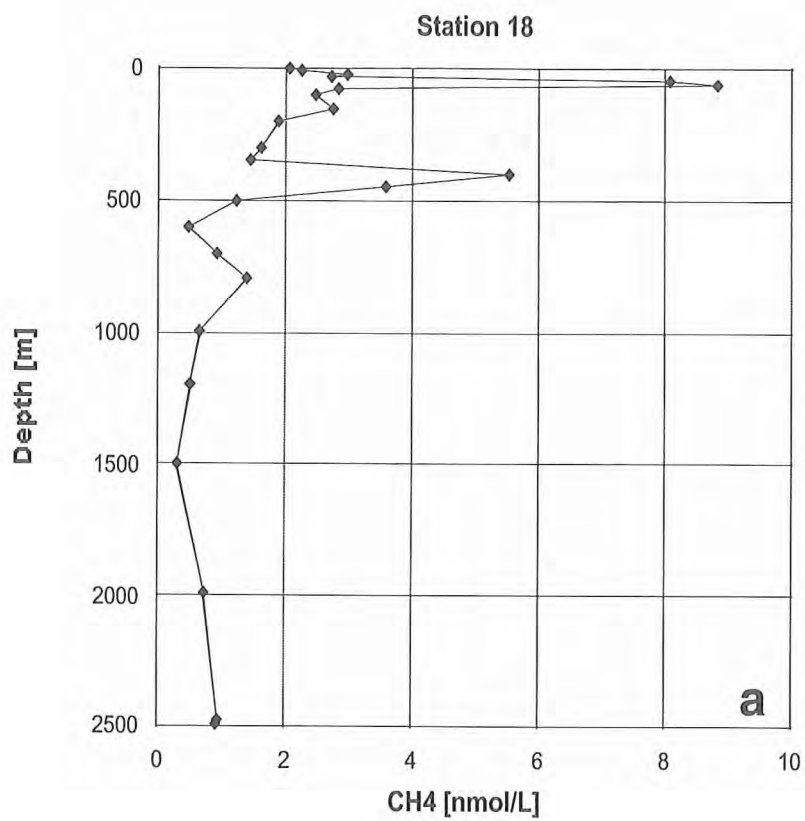


Fig. 6.7.3.2.1 Typical distribution of (a) methane and (b) oxygen within the research area (from station 18).

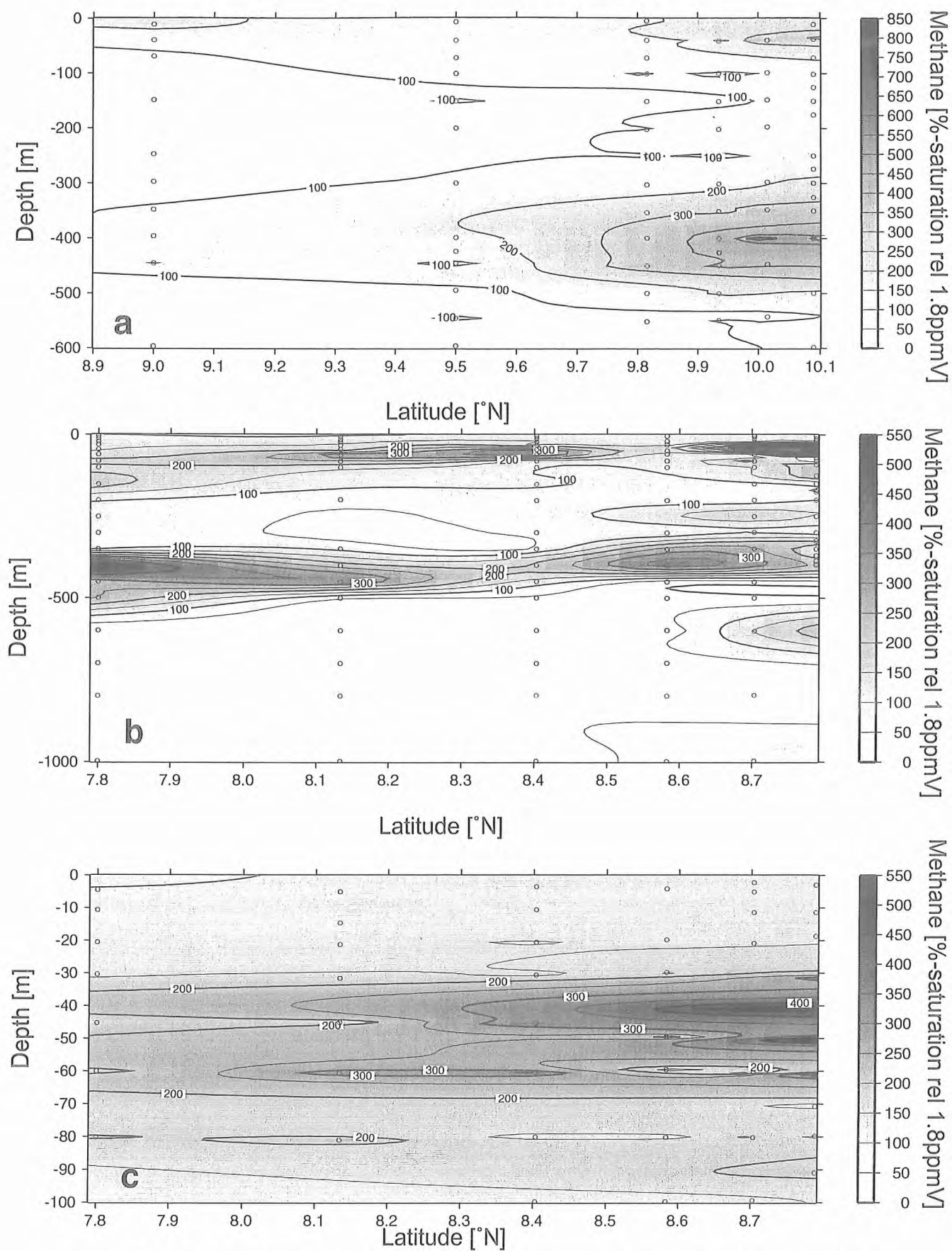


Fig. 6.7.3.2.2 Sections of methane saturation of waters along the 2 hydrographic sections: (a) Northern Transect (CTD stations 1-6); (b) Southern transect (CTD stations 15-20); (c) close-up of the upper 100m along the southern transect.

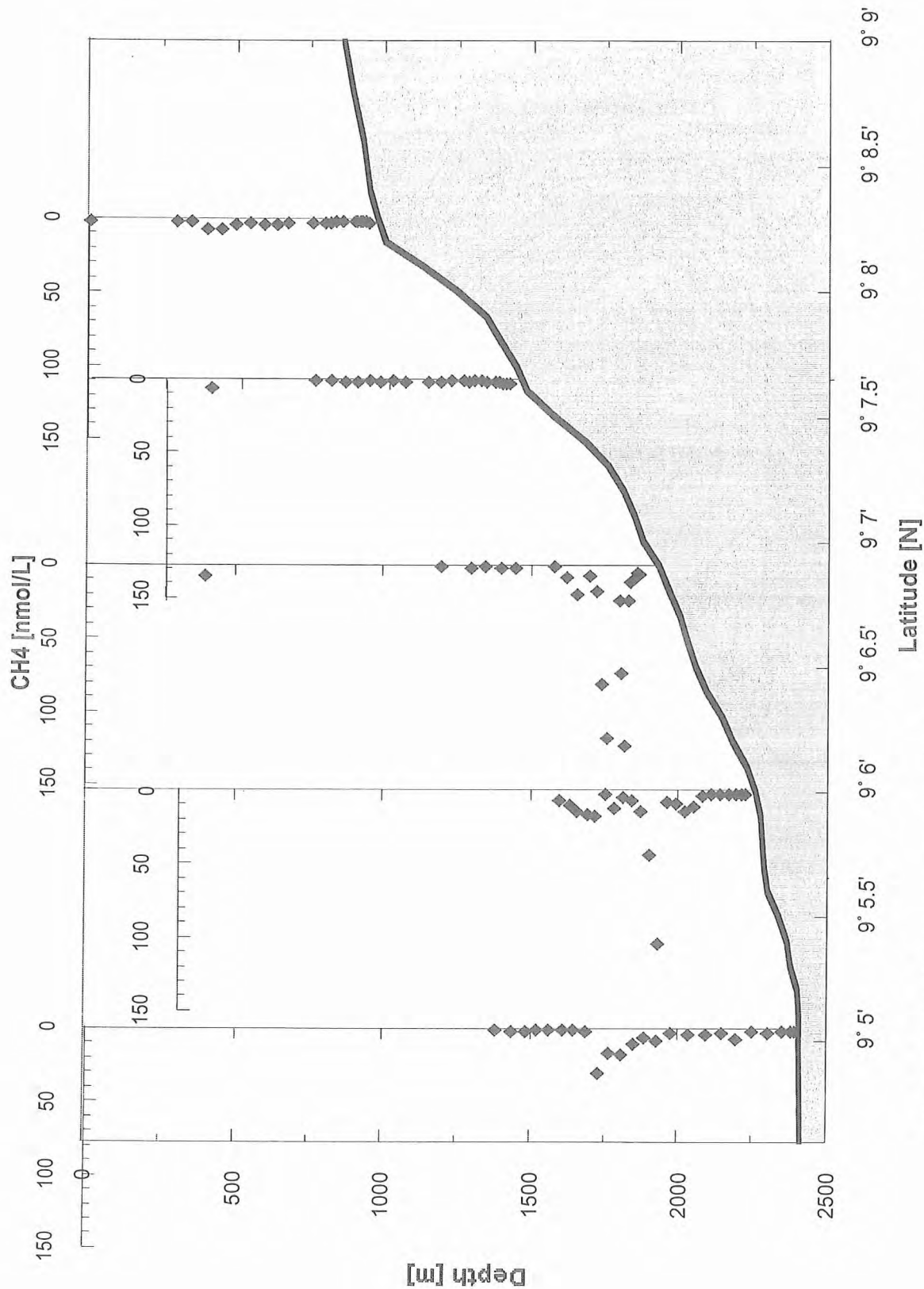


Fig. 6.7.3.3.2 Cross section of Jaco Scarp along the line defined by the CTD/MUC stations. The results of the dissolved methane analysis observed at the stations is presented.

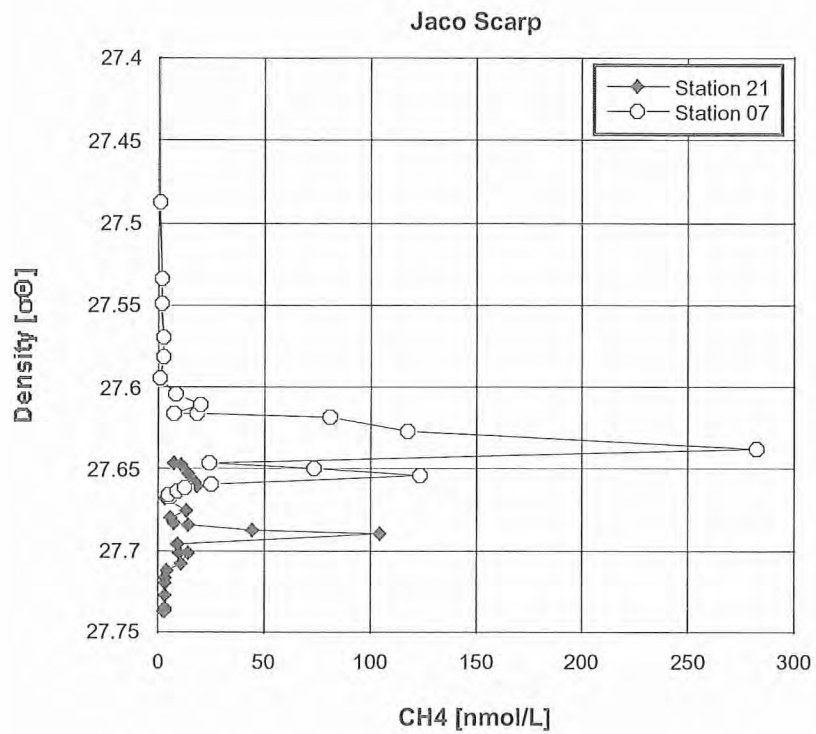


Fig. 6.7.3.3.3 Methane vs density (σ_θ) at CTD stations 7 and 21.

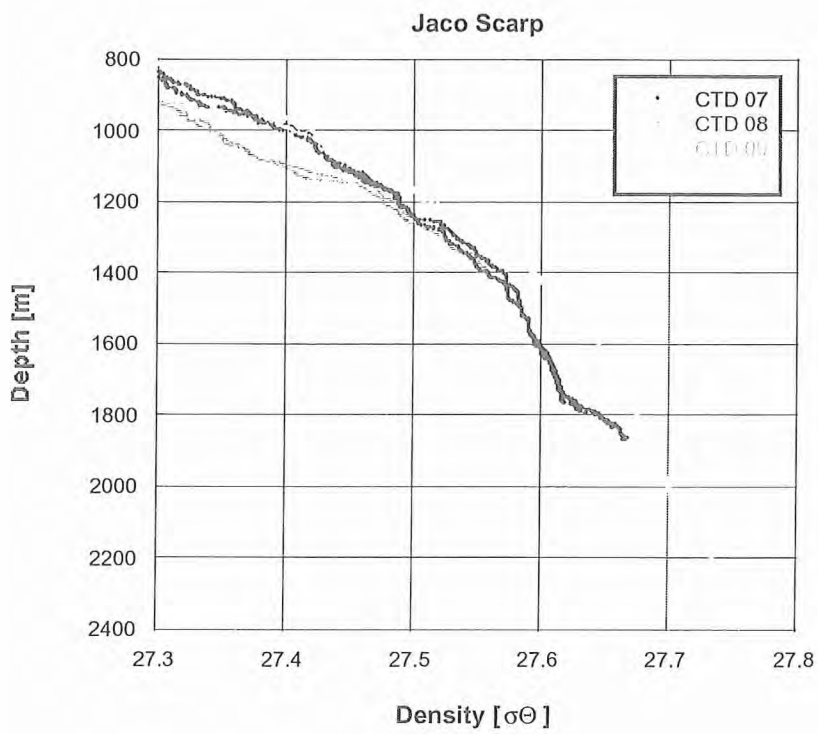


Fig. 6.7.3.3.4 Density (σ_θ) vs depth for CTD stations 7-9 and 21.

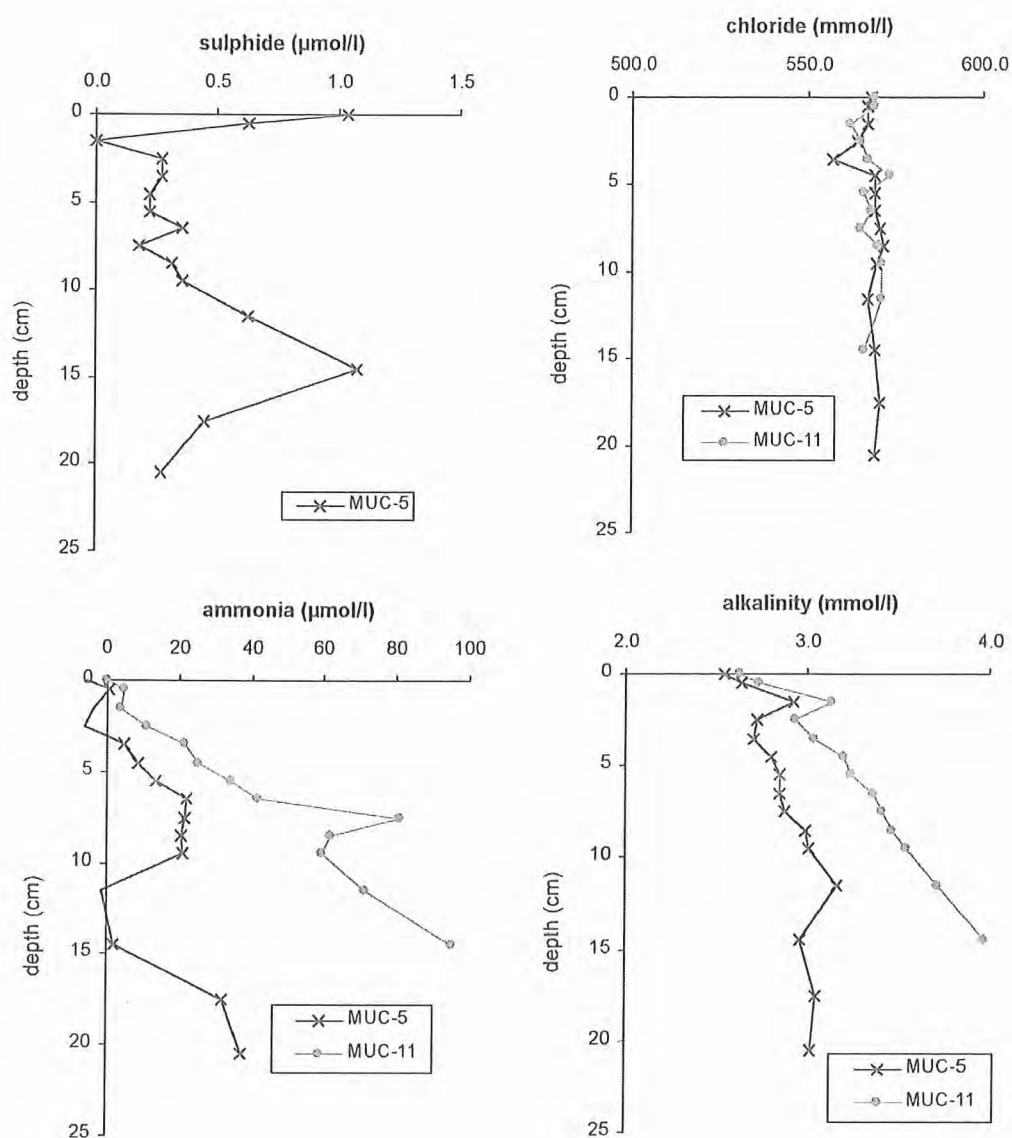


Fig. 6.7.3.3.5 Results of pore water analysis - Jaco Scarp stations.

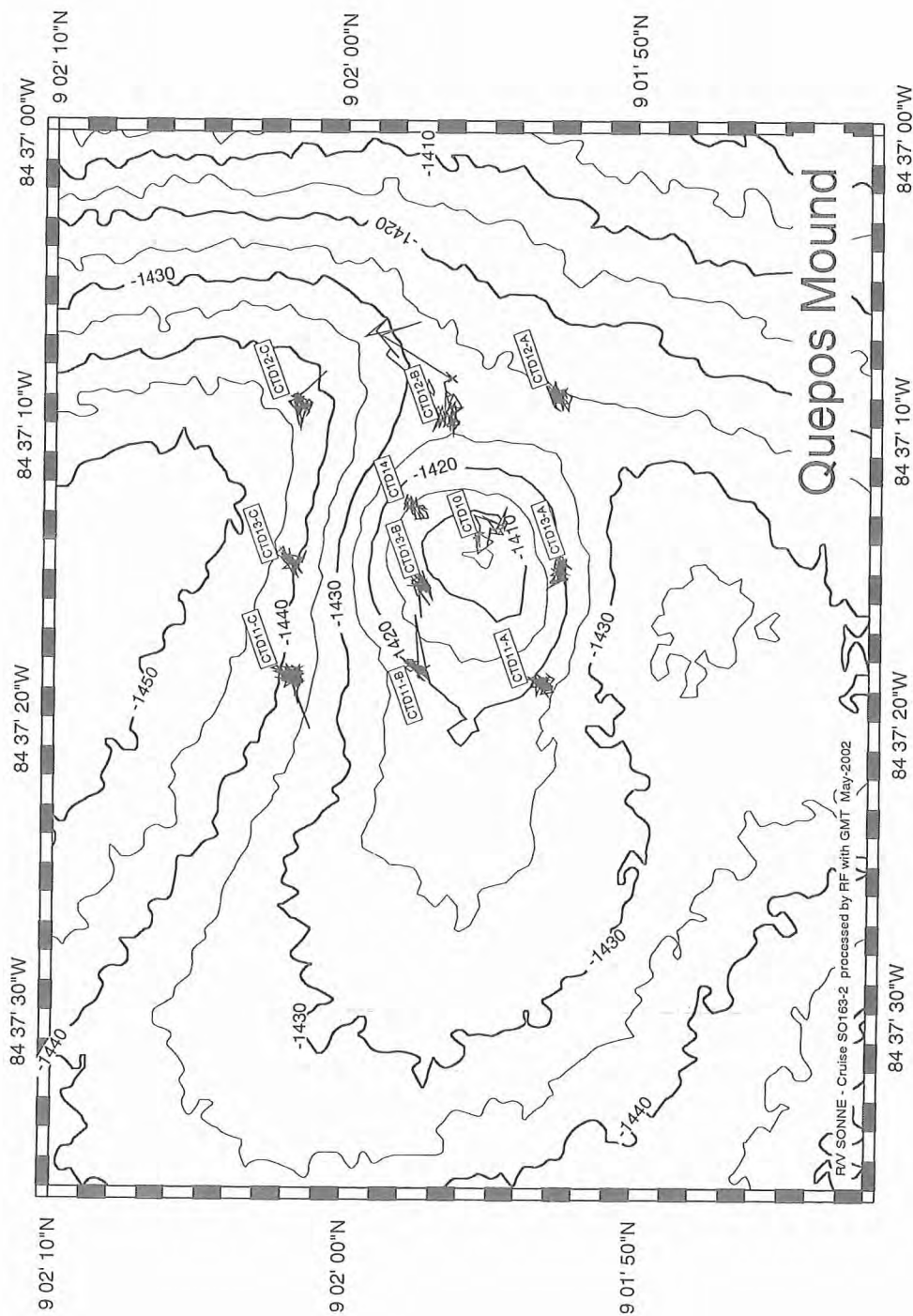


Fig 6.7.3.4.1: Bathymetric map of Quepos Mound with the positions of the SSBL (attached 40m above the CTD/rosette) during sampling.

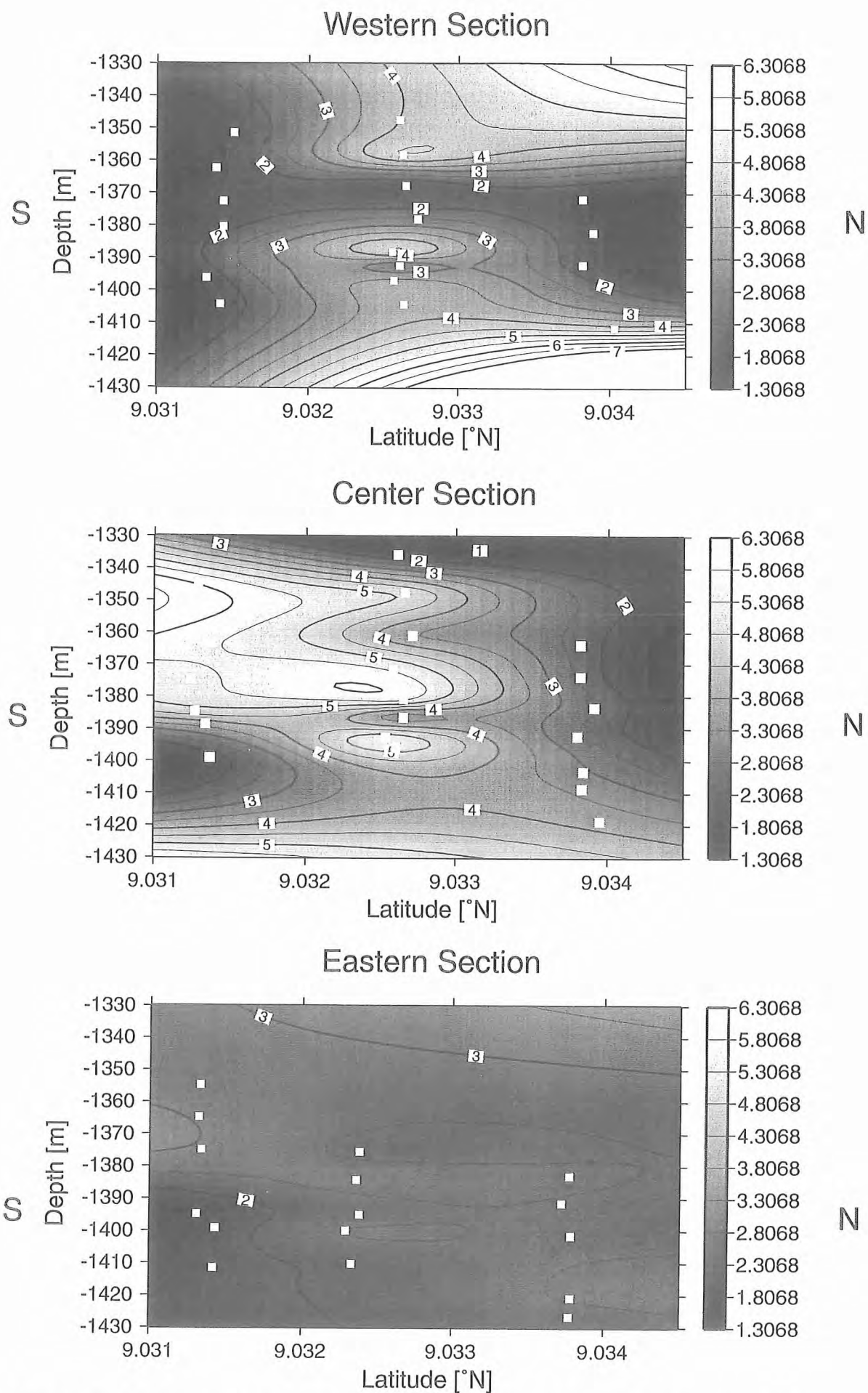


Fig 6.7.3.4.2: North-south trending cross-sections of the methane concentration in the lower water column with values given in nmol/L.

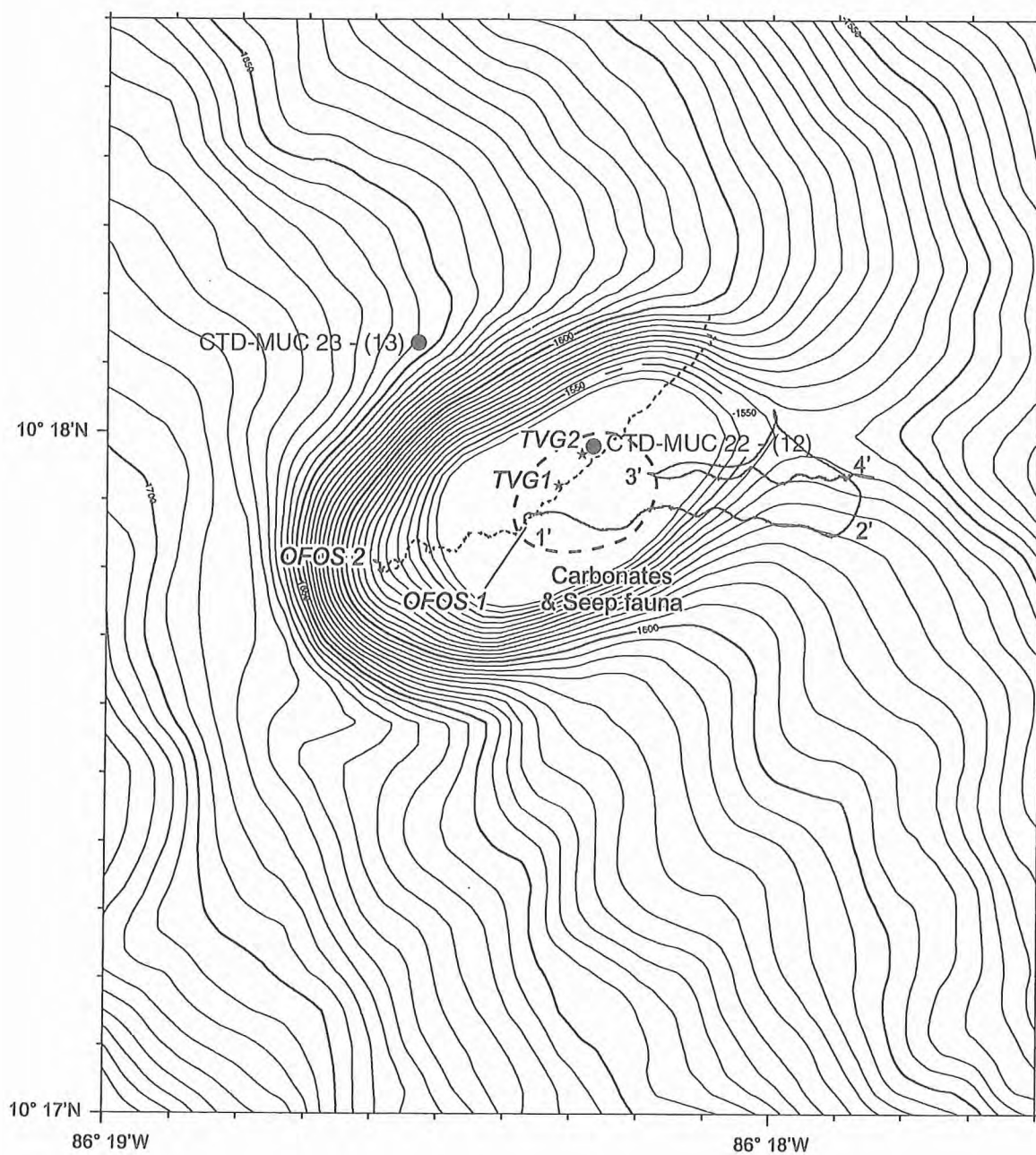


Fig 6.7.3.5.1 Bathymetric map of Mound Culebra. Indicated are the locations of the CTD/MUC deployment as well as the locations of TVG locations and OFOS surveys (SO163/1).

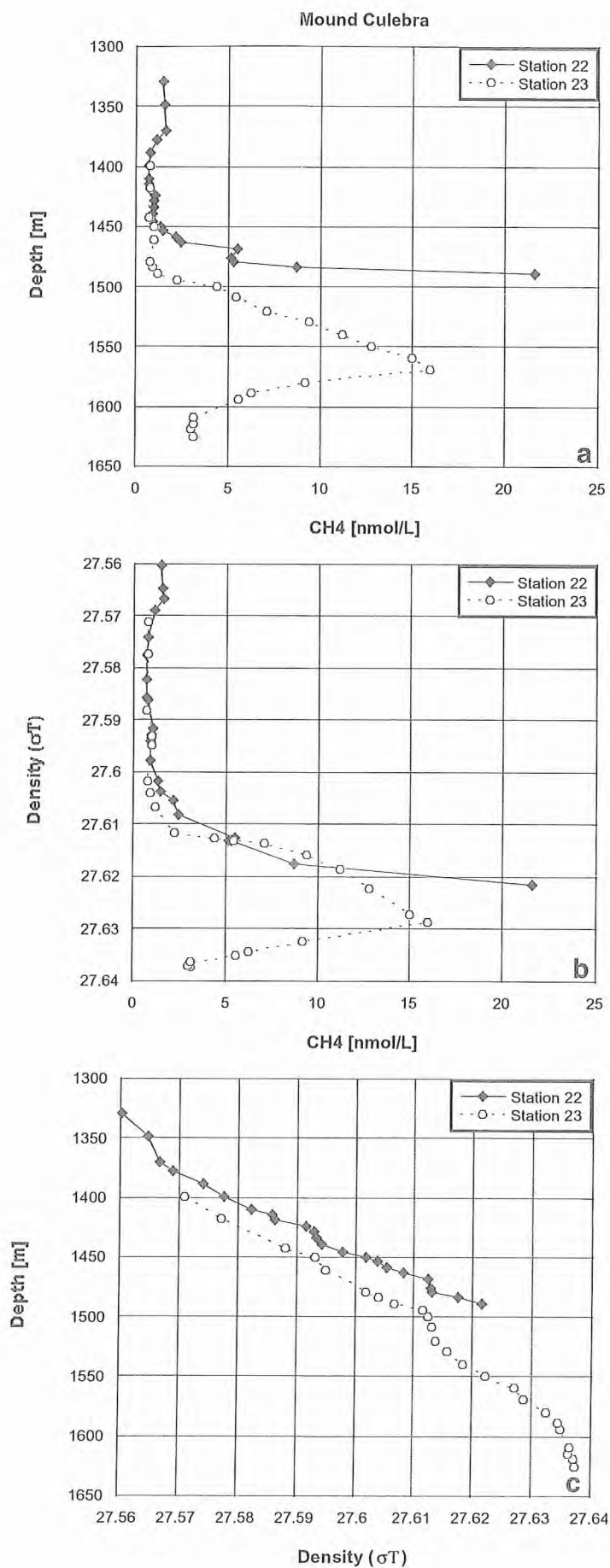


Fig 6.7.3.5.2 (a) Methane concentration vs depth, (b) methane concentration vs density, and density vs depth for stations 22 and 23 at Mound Culebra.

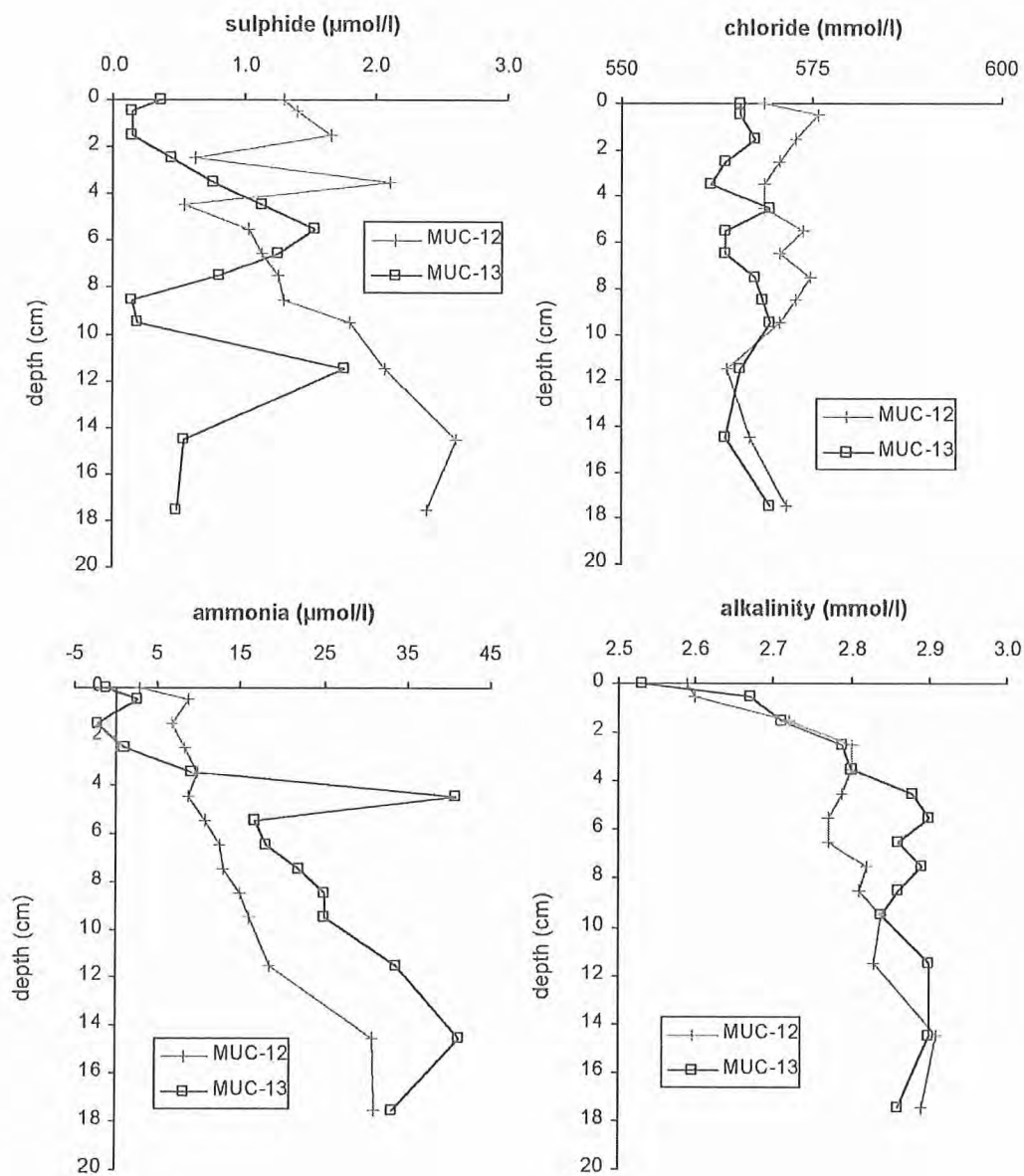


Fig 6.7.3.5.3 Results of pore water analysis from the Mound Culebra stations.

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Appendix 9.1 Mapping profiles

line number	date	time	position		course	length	remarks
EM-01	15.03.02	10:50	07°31.91 N	83°01.31 W	298		
		14:09	07°45.00 N	83°30.00 W	296	22	
		21:24	08°17.44 N	84°55.81 W	318	73	
	16.03.02	06:56	09°39.50 N	85°51.57 W		111	
EM-02	16.03.02	18:20	09°38.28 N	85°53.91 W			
		20:09	09°50.02 N	85°44.36 W	322	15	
		21:04	09°56.33 N	85°49.19 W	219	9	
		21:15	09°55.40 N	85°50.23 W	144	1	
		21:56	09°49.91 N	85°46.39 W	244	7	
		22:07	09°49.06 N	85°47.87 W	327	2	
		23:14	09°58.39 N	85°53.91 W	41	11	
		23:44	10°00.98 N	85°51.69 W	306	4	
		24:00	10°00.98 N	85°51.69 W		4	
EM-03	17.03.02	00:00	10°03.20 N	85°54.60 W			
		00:35	10°08.07 N	85°58.54 W	292	6	
		00:45	10°07.22 N	85°59.72 W	139	2	
		01:23	10°02.41 N	85°55.76 W	216	6	
		01:33	10°01.13 N	85°56.75 W	320	2	
		02:32	10°08.53 N	86°02.58 W	233	10	
		02:45	10°07.32 N	86°04.70 W	141	2	
		03:46	09°59.36 N	85°58.49 W	234	10	
		04:04	09°57.57 N	86°00.86 W	322	3	
		05:12	10°06.38 N	86°07.83 W	47	11	
		06:02	10°12.01 N	86°01.80 W	322	8	
		08:21	10°28.10 N	86°13.85 W	50	20	
		08:50	10°28.93 N	86°12.81 W	142	1	
EM-04	17.03.02	09:36	10°41.92 N	86°34.92 W	21		
		10:29	10°45.91 N	86°32.67 W	270	4	
		10:58	10°46.07 N	86°35.79 W	137	3	
		12:45	10°38.75 N	86°29.18 W		10	
EM-05	17.03.02	13:23	10°37.32 N	86°36.00 W	324		
		15:10	10°45.94 N	86°42.46 W	90	11	
		15:22	10°46.26 N	86°41.54 W	124	1	
		16:13	10°42.43 N	86°36.84 W		6	
EM-06	22.03.02	23:11	10°19.09 N	86°17.98 W	25		
	23.03.02	00:06	10°26.96 N	86°13.98 W	325	10	
		00:28	10°30.00 N	86°16.00 W	312	4	
		02:09	10°41.30 N	86°28.50 W	54	17	
		02:22	10°41.82 N	86°27.67 W	131	2	
		04:09	10°31.44 N	86°15.00 W	44	17	
		04:21	10°32.85 N	86°13.69 W	306	2	
		05:26	10°39.40 N	86°22.50 W	213	11	
		05:36	10°38.83 N	86°22.87 W	128	1	
		06:40	10°32.68 N	86°15.08 W	90	10	
		06:45	10°32.49 N	86°14.38 W	156	1	
		07:30	10°25.87 N	86°11.37 W	195	7	
		09:33	10°05.91 N	86°16.73 W		21	

line number	date	time	position		course	length	remarks
EM-07	25.03.02	12:31	09°20.70 N	86°04.00 W			
		14:21	09°32.97 N	85°50.20 W	90	18	
		15:15	09°32.00 N	85°41.22 W		9	
EM-08	29.03.02	06:44	10°00.98 N	86°11.29 W			
		11:35	09°49.13 N	86°41.88 W		32	
EM-09	29.03.02	11:35	09°49.13 N	86°41.88 W			
		12:00	09°48.01 N	86°43.26 W	134	2	
		16:58	09°30.21 N	86°24.68 W	41	26	
		17:15	09°31.18 N	86°23.65 W	314	4	
		20:57	09°44.85 N	86°37.63 W		19	
EM-10	30.03.02	09:20	09°48.22 N	86°41.97 W			
		12:34	10°20.13 N	86°41.50 W	140	27	
		17:03	09°45.00 N	86°12.00 W	131	46	
		20:27	09°21.11 N	85°45.13 W	90	36	
		21:56	09°21.01 N	85°29.86 W		15	
EM-11	02.04.02	06:47	09°09.58 N	85°41.42 W			
		08:51	09°06.10 N	85°36.00 W	56	14	
		12:07	09°19.00 N	85°36.00 W		14	
EM-12	14.04.02	14:04	09°00.00 N	84°23.00 W	35		
		14:37	09°02.32 N	84°22.05 W	129		
		15:28	08°58.40 N	84°17.20 W		11	
EM-13	14.04.02	19:51	08°46.29 N	84°14.65 W	196		
		20:45	08°51.00 N	84°22.10 W		11	
EM-14	14.04.02	22:43	08°46.29 N	84°14.65 W	304		
		00:39	08°51.00 N	84°22.10 W		11	
EM-15	18.04.02	04:00	09°06.80 N	84°51.90 W	213		
		05:01	09°02.50 N	84°54.78 W	120	5	
		05:37	09°01.01 N	84°52.20 W	34	3	
		06:47	09°05.61 N	84°48.98 W	38	7	
		07:09	09°06.92 N	84°47.97 W		2	
EM-16	18.04.02	14:11	09°01.45 N	84°38.00 W	288		
		15:22	09°04.52 N	84°46.95 W	61	10	
		15:50	09°06.23 N	84°43.96 W	108	3	
		19:00	09°00.00 N	84°25.35 W		19	
EM-17	18.04.02	19:26	09°01.56 N	84°23.00 W	288		
		21:45	09°07.55 N	84°40.81 W		19	
EM-18		22:04	09°09.61 N	84°40.62 W			
	19.04.02	00:32	09°03.52 N	84°23.00 W	63	20	
		00:40	09°03.67 N	84°21.48 W	288	1	
		03:04	09°10.37 N	84°40.00 W			

Appendix 9.2 TOBI profiles

M-O Number	File Name	Time/ Day START	Time/ Day STOP	Comments / Run #
830	TOBIB.DAT	1813/076	0523/077	Run 1 curtailed due to cable short. Day number corrected.
831	TOBI.DAT TOBIB.DAT TOBIC.DAT	2310/077 0334/078 0339/078	0329/078 0337/078 1526/078	Run 2 Day number corrected Day number correct
832	TOBI.DAT	1526/078	0733/079	
833	TOBI.DAT	0733/079	2340/079	
834	TOBI.DAT	2340/079	1548/080	
835	TOBI.DAT	1548/080	0755/081	
836	TOBI.DAT	0755/081	1419/081	End of run 2
837	TOBI.DAT	2120/082	1323/083	Start of run 3
838	TOBI.DAT	1323/083	0535/084	
839	TOBI.DAT	0535/084	1558/084	End of run 3
840	TOBI.DAT	0145/096	1752/096	Start of run 4
841	TOBI.DAT	1752/096	1000/097	
842	TOBI.DAT	1000/097	0207/098	
843	TOBI.DAT	0207/098	1815/098	
844	TOBI.DAT	1815/098	1022/099	
845	TOBI.DAT	1022/099	0229/100	
846	TOBI.DAT	0229/100	1355/100	End of run 4
847	TOBI.DAT TOBIA.DAT TOBIB.DAT	0052/102 0340/102 1016/102	0156/102 1014/102 1846/102	Run 5 Start of run 6 Reboot logging
848	TOBI.DAT	1846/102	1053/103	
849	TOBI.DAT	1053/103	2103/103	End of run 6

Appendix 9.3 DTS-1 profiles

Deployment	Average water depth	Track length	Mapped area	Time on sea-floor	Deployment	Recovering	Total
1	1300-2400	141.0 km	162.2 km ²	37 h	1 h 40	2 h	40 h 40
2	2500-3200	134.0 km	145.0 km ²	44 h	2 h 30	2 h 50	49 h 20
3	400-1900	172.4 km	175.2 km ²	44 h	35 min	2h	46 h 30
4	900-2000	120.4 km	161.0 km ²	30 h	2 h	2 h	34 h
5	1300-1500	9.0 km	1.8 km ²	2h	55 min	1h 05	4h

Appendix 9.4 OFOS und TVG deployments

Station	Area	Date	Time	at seafloor			off seafloor / grab location			Observation / recovery	
				Lat	Long	Depth	Time	Lat	Long		Depth
Mounds											
OFOS 1	Mound # 1	22.03.2002	17:41:48	10° 17.86	86° 18.37	1540	19:41:49	10° 18.01	86° 17.99	1570	Carbonates, clams and pogonophorans, No SSBL Position
OFOS 2	Mound # 1	22.03.2002	21:29:10	10° 17.84	86° 18.56	1510	23:21:10	10° 18.14	86° 18.07	1550	Carbonates, clams and pogonophorans, SSBL 50 m above OFOS
TVG 1	Mound # 1	23.03.2002	01:59:50	10° 17.88	86° 18.40	1500	02:07:00	10° 17.923	86° 18.32	1516	2 carbonate pieces on top of grab
TVG 2	Mound # 1	23.03.2002	03:26:00	10° 17.93	86° 18.33	1490	03:36:00	10° 17.987	86° 18.29	1473	Several carbonate pieces and clamshell fragments
OFOS 3	Mound # 2-6	28.03.2002	00:10:30	9° 37.12	85° 46.53	1588	08:25:26	9° 35.95	85° 50.15	1989	Crossed 4 mounts from NE to SW, some with carbonates, pogonophorans, clams; Camera B problems
OFOS 4	Mound # 7, 8	18.03.2002	11:32:36	9° 42.53	85° 56.30	2203	15:17:14	9° 41.98	85° 56.90	2240	Small aera with carbonates, pogonophorans and clams
OFOS 6	Mound # 9	29.03.2002	05:01:30	10° 02.16	86° 11.80	2240	07:24:10	10° 01.56	86° 12.23	2342	No seep indication
OFOS 7	Mound # 10	29.03.2002	09:24:10	10° 00.73	86° 11.35	2305	11:36:10	10° 00.65	86° 11.50	2341	Abundant vent indication; clams, pogonophorans, carbonate
OFOS 13	"Bright Spot W of Parita"	05.04.2002	20:19:50	9° 59.70	84° 42.50		23:35:00	8° 58.00	84° 41.65		Beside scattered clamshells no seep indications
OFOS 19	Mound # 11, 12	14.04.2002	05:44:20	8° 54.94	84° 17.91	1023	08:14:32	8° 55.95	84° 18.83	1005	Extensive carbonates, clams, mussels and pogonophorans
OFOS 21	Mound # 12	14.04.2002	15:09:30	8° 55.94	84° 18.53	998	18:40:16	8° 55.73	84° 18.95	1013	Extensive carbonates, clams, mussels and pogonophorans; Video on 16:32
TVG 4	Mound # 12	14.04.2002	22:29:35	8° 55.79	84° 18.74	1002	23:14:45	8° 55.81	84° 18.62	992	Piece of outcropping carbonate
TVG 5	Mound # 12	15.04.2002	00:16:50	8° 55.79	84° 18.75	1002	00:41:00	8° 55.80	84° 18.74	1005	Complete seep ecosystem: Acharax, Bathymodiolus, Lamellibrachia, galatheid crabs and carbonates
OFOS 23	Mound # 13	17.04.2002	05:12:40	9° 02.20	84° 37.24	1451	12:28:39	9° 01.78	84° 37.06	1403	Most extensive field of carbonates, pogonophorans, bivalves
OFOS 24	Mound # 13	17.04.2002	14:27:44	9° 01.95	84° 37.22	1406	16:41:50	9° 01.95	84° 37.13	1397	Mapping most extensive seep area
TVG 8	Mound # 13	18.04.2002	16:55:10	9° 01.96	84° 37.22	1408	17:06:40	9° 01.97	84° 37.22	1408	Carbonates
TVG 9	Mound # 13	18.04.2002	18:44:15	9° 01.96	84° 37.25	1408	18:47:35	9° 01.96	84° 37.24	1408	Vent fauna, carbonates

Station	Area	Date	Time	at seafloor		Depth	off seafloor / grab location				Observation / recovery
				Lat	Long		Lat	Long	Depth		
High heatflow area											
OFOS 8	Basement hight	30.03.2002	04:43:17	9° 40.25	86° 33.76	3275	08:20:00	9° 41.01	86° 33.17	3267	No indication for hydrothermal venting
OFOS 9	Fault	30.03.2002	10:58:04	9° 46.48	86° 40.78	3315	13:40:51	9° 46.89	86° 40.41	3346	No indication for hydrothermal venting
seamount scars											
OFOS 5	Scarp of fracture ridge trace	28.03.2002	17:42:39	9° 36.31	85° 54.28	2194	23:00:40	9° 33.94	85° 56.53	3145	very few seep indication or other geological structures
OFOS 10	Rio Bongo Scarp	02.04.2002	18:39:39	9° 18.86	85° 15.54	470	23:27:30	9° 19.16	85° 18.31	772	Carbonated top, few vent fauna
OFOS 11	Jaco Scarp	03.04.2002	04:52:39	9° 10.43	84° 47.75	731	07:16:42	9° 10.40	84° 48.66	779	Carbonates, bivalve shells
OFOS 12	Jaco Scarp	03.04.2002	09:03:40	9° 10.40	84° 50.34	928	13:15:33	9° 08.74	84° 48.69	865	Chemoherm-like carbonates, vesicomylid clamfield
TVG 3	Jaco Scarp	05.04.2002	15:44:10	9° 10.42	84° 48.25	763	15:50:57	9° 10.43	84° 48.24	763	Chemoherm-like carbonates, living solemyid bivalves, bivalve shells
OFOS 25	Jaco Scarp	18.04.2002	00:09:48	9° 11.60	84° 47.80	774	05:03:20	9° 10.78	84° 50.62	996	No vent indication
OFOS 26	Jaco Scarp	18.04.2002	06:22:55	9° 08.24	84° 49.64	953	09:08:31	9° 09.14	84° 48.04	810	No vent indication
TVG 7	Jaco Scarp	18.04.2002	14:01:32	9° 09.03	84° 49.18	834	14:32:00	9° 09.01	84° 49.13	824	Carbonates
OFOS 15	"Frontal collision"	11.04.2002	00:02:00	8° 22.20	83° 47.684	1260	02:52:24	8° 21.182	83° 48.013	1633	Steep morphology, beside clamshells no seep indication
OFOS 16	"Small seamount scarp"	11.04.2002	05:05:00	8° 26.48	83° 54.981	701	07:48:14	8° 25.893	83° 56.002	1095	No vent indication
OFOS 22	Parita	15.04.2002	09:15:58	8° 57.55	84° 36.51	1228	14:12:57	8° 55.17	84° 37.89	2338	Beside scattered clamshells no seep indications

Station	Area	Date	at seafloor			off seafloor / grab location					
			Time	Lat	Long	Depth	Time	Lat	Long	Depth	Observation / recovery
Accretionary wedge											
OFOS 14	"Crest of Cocos Ridge"	10.04.2002	19:21:24	8° 15.572	83° 34.185	1710	20:41:58	8° 15.343	83° 34.803	1902	No seep indication
OFOS 17	"Seismic Line SO 81 #5"	11.04.2002	12:34:30	8° 40.56	84° 18.35	2000	20:57:25	8° 35.93	84° 21.30	2500	No seep indication
Landslide											
OFOS 18	Quepos Landslide	13.04.2002	22:56:41	8° 51.87	84° 12.60	179	03:48:43	8° 49.50	84° 15.51	831	Larg bacterial mats at headwall, few seep indications in slide mass
OFOS 20	Quepos Landslide	14.04.2002	10:16:00	8° 50.30	84° 13.00	438	13:06:00	8° 48.61	84° 13.90	740	Beside one bacterial mat no seep indication at slide mass

SUBDUCTION

INST.	LAT (N)	LON (W)	DIST. TO NEXT (nm)	DEPTH (m)	REL. CODE	ANT. CH.	REC. NO.	SKEW (ms)	SENSOR	Other Sensors	REMARKS
	D:M	D:M							HYDRO		
OBH 01	09:39.456	85:51.494	2,45	1866	3624	C	991249	131	OAS 13		
OBH 02	09:40.605	85:53.698	2,45	1924	3669	D	991247	51	OAS 02		
OBH 03	09:38.135	85:53.489	1,20	2068	3659	C	991256	-62	OAS 30		
OBH 04	09:38.849	85:52.508	1,11	1996	3674	C	991235	63	DPG 93		
OBH 05	09:39.911	85:52.805	0,93	1940	039A + 0355	D	000713	85	DPG 92		
OBH 06	09:39.394	85:53.584	0,58	1871	6334	A	000708	-235	DPG 87		
OBH 07	09:39.930	85:53.109	0,65	1997	3609	D	991252	-60	OAS 22		
OBH 08	09:39.450	85:52.528	0,60	1964	0387 + 0355	B	991242	-43	DPG 75		
OBH 09	09:39.691	85:53.108	0,32	1951	3664	D	000707	-133	OAS 27		
OBH 10	09:39.461	85:52.905	1,03	1893	0386 + 0355	C	991237	-79	DPG 91		50 m anchor rope
OBH 22	09:38.507	85:52.366	1,95	2116	3614	C	980901	15	OAS 29		
OBH 23	09:40.222	85:53.495		1931	0386 + 0355	A	010703	-1	OAS 38		550 m anchor rope

INST.	LAT (N) D:M	LONG (W) D:M	Deployed Recovered	DEPTH (m)	REL. CODE	ANT. CH.	REC. REC.	SKEW (ms)	SENSOR HYDRO.	Other Sensors	REMARKS
OBS 14	09:19.985	84:52.996	22.04./ 09.05.02	1741	0399 + 0355	C	0706		OAS 34	PMD 540	power failure
OBS 15	09:14.661	84:36.046	22.04./ 09.05.02	114	0398 + 0355	A	0711	-18	DPG 95		
OBS 16	09:09.975	84:38.015	22.04./ 09.05.02	944	0397 + 0355	C	10405	5	HTI 302		
OBS 17	09:09.405	84:49.577	22.04./ 28.04.02	824	03BC + 0355	C	020303	-26		TM 4948	
OBS 18	09:09.397	84:49.588	23.04./ 09.05.02	827	4A44	C	010402	2976	OAS 18		
OBS 19	08:59.996	84:50.055	23.04./ 09.05.02	2206	03B9 + 355	C	010401		OAS 31	PMD 509	Rec.end: 05.05.02
OBS 20	08:50.006	84:49.996	23.04./ 09.05.02	3473	03B7 + 0355	D	991238	-19	HTI 602		
OBS 21	08:58.032	85:05.046	23.04./ 09.05.02	3602	03BB + 0355	C	991243	20	HTI 102		
OBS 38	09:14.271	84:56.637	28.04.02	1752	03BD + 0355	D	000713		DPG 93		
OBS 39	09:05.280	85:02.560	28.04.02	2496	03B6 + 0355	C	991249		DPG 75		
OBS 40	08:53.328	84:58.432	28.04.02	3420	03B5 + 0355	C	991237		DPG 91		
OBS 41	08:56.043	84:43.428	28.04.02	2615	03B3 + 0355	C	991242		DPG 87		
OBS 42	09:05.173	84:39.092	28.04.02	1481	03B8 + 0355	C	991250		DPG 92		
OBS 139	09:13.023	84:33.508	12.05.02	379	03B7 + 0355	C	991247		OAS 02		
OBS 140	09:10.017	84:37.993	12.05.02	939	03BB + 0355	A	10405		OAS 04		
OBS 141	09:13.016	84:41.022	12.05.02	351	03B9 + 0355	A	10402		OAS 29		
OBS 193	09:19.989	85:06.063	16.05.02	1880	3659	C	991238		HTI 502		
OBS 194	09:28.015	84:52.503	16.05.02	560	039A + 0355	D	707		OAS 18		
OBS 195	09:28.001	84:44.997	16.05.02	324	03B7 + 0355	B	991256		OAS 30		
OBS 196	09:21.531	84:45.034	16.05.02	413	3624	C	10401		HTI 302		
OBS 197	09:20.022	84:53.006	16.05.02	1745	03BA + 0355	C	10404		HTI 702	Owen 22 (4,5Hz)	
OBS 198	09:14.975	84:44.001	16.05.02	291	3674	C	991243		OAS 27		
OBS 199	09:12.016	84:46.990	16.05.02	734	03B1 + 0355	D	20303		HTI 802	TM 4949	
OBS 200	09:09.419	84:49.601	16.05.02	824	03B2 + 0355	C	20302		HTI 602	TM 4948	
OBS 201	09:10.013	84:56.020	16.05.02	2080	0397 + 0355	D	991252		OAS 31	LG 01 (4,5 Hz)	
OBS 202	09:01992	84:56.010	17.05.02	2237	03B6 + 0355	D	991236		HTI 102		
OBS 203	08:57.974	85:05.004	17.05.02	3561	3614	D	706		OAS 18		
OBS 204	08:50.000	84:49.995	17.05.02	3474	3629	D	991246		OAS 38		
OBS 205	08:59.985	84:49.996	17.05.02	???	3669	B	991241		HTI 902		
OBS 206	09:01.996	84:43.006	17.05.02	1726	03BC + 0355	D	10403		OAS 46	Owen 21 (4,5Hz)	
OBS 207	09:07.000	84:43.003	17.05.02	1159	3609	C	711		OAS 22		
OBS 208	09:05.376	84:34.248	17.05.02	1155	0399 + 0355	D	991248		HTI 202		
OBS 209	09:02.296	84:39.205	17.05.02	1615	3664	A	991244		OAS 75		
OBS 210	08:59.579	84:37.907	17.05.02	1202	03B4 + 0355	D	708		HTI 101		
OBS 211	09:01.004	84:34.461	17.05.02	1411	6334	C	616		OAS 09		
OBS 212	09:02.387	84:36.203	17.05.02	1326	4A44	B	971202		OAS 28		

SUBDUCTION

INST.	LAT (N) D:M	LONG (W) D:M	Dist. To Next (nm)	DEPTH (m)	REL. CODE	ANT. CH.	REC. REC.	SKEW (ms)	SENSOR	Other Sensors	REMARKS
OBS 24	09:15.938	85:37.412	0.11	2770	3614	C	98040	-6	OAS 75	LG 01 (4.5Hz)	
OBS 25	09:16.017	85:37.308	0.11	2768	03BA +0355	B	991247	2	HTI 702	PMD 539	
OBS 26	09:16.093	84:37.255	0.11	2770	03BA +0355	C	971202	-6	HTI 202	Owen 02 (4.5Hz)	
OBS 27	09:16.182	85:37.195	0.12	2738	3624	D	010703	-1	OAS 13		
OBS 28	09:16.267	85:37.079	0.16	2736	3664	C	980403	-2	OAS 27		
OBS 29	09:16.407	85:37.026	0.16	2705	03B6 +0355	D	010708	2	OAS 38		
OBS 30	09:16.519	85:36.915	0.54	2654	3669	C	980901	17	OAS 02		500m anchor rope
OBS 31	09:16.948	85:36.558	0.11	2657	03B5 +0355	A	000616	-9	OAS 46	Owen 21 (4.5Hz)	
OBS 32	09:17.043	85:36.508	4.54	2628	3659	D	991256	-2	OAS 30		
OBS 33	09:20.606	85:33.705	0.11	2260	6334	C	000707	-7	OAS 29		
OBS 34	09:20.687	85:33.625	0.11	2257	03B2 +0355	C	980402	-5	HTI 002	Owen 22 (4.5Hz)	
OBS 35	09:20.785	85:33.567	4.60	2246	03B0 +0355	B	991252	-3	HTI 802		
OBS 36	09:17.135	85:36.425	0.11	2626	03B0 +0355	B	000614	-1	HTI 302		
OBS 37	09:17.183	85:36.328		2607	3629	D	000610	-2	HTI 502		OBS-design 2002

[illegible]

SUBDUCTION

INST.	LAT (N) D:M	LONG (W) D:M	Dist. To Next (nm)	DEPTH (m)	REL. CODE	ANT. CH.	REC. REC.	SKEW (ms)	SENSOR	Other Sensors	REMARKS
OBH 76	09:03,698	84:32,899	1,53	1186	0387+0355	D	010708	0	OAS 38		
OBH 77	09:02,900	84:31,593	1,51	1238	3664	A	980901	14	OAS 27		
OBH 78	09:02,096	84:30,294	1,54	1273	3624	B	980403	-2	OAS 13		not active
OBH 79	09:01,254	84:28,970	1,47	1319	6334	D	991248	-1	OAS 29		
OBH 80	09:00,477	84:27,715	1,41	1337	3609	C	991252	-4	HTI 902		
OBH 81	08:59,741	84:26,508	1,49	1342	03BA+0355	C	000616	-7	HTI 702	Owen 14 (30 Hz)	
OBH 82	08:58,933	84:25,208	1,53	1246	03BC+0355	D	971202	-5	HTI 101	Owen 16 (30 Hz)	
OBH 83	08:58,111	84:23,913	1,51	1168	0386+0355	D	980402	-5	OAS 04	Owen 21 (4,5 Hz)	
OBH 84	08:57,313	84:22,607	1,47	1117	03B4+0355	D	980401	-5	HTI 202	Owen 02 (4,5 Hz)	
OBH 85	08:56,497	84:21,350	1,45	1092	3614	C	000611	-3	OAS 75	LG 01 (4,5 Hz)	
OBH 86	08:55,718	84:20,107	1,52	1111	3629	C	000610	-8	HTI 502		OBS-design 2002
OBH 87	08:54,903	84:18,785	1,49	1111	039A+0355	D	000707	-5	HTI 102		
OBH 88	08:53,555	84:16,626	1,07	1089	3659	D	991247	0	OAS 30		
OBH 89	08:54,104	84:17,515	0,99	1062	3674	B	000614	-6	HTI 802		
									OAS 09		
									OAS 04		
									OAS 14		
OBH 90	08:53,017	84:15,767	2,50	960	3669	B	010404	3	OAS 02		
OBH 92	09:14,922	84:15,185	1,46	150	0387+0355	B	010404	2	OAS 38		
OBH 93	09:13,72	84:16,04	1,58	186	3664	A	010708	0	OAS 27		
OBH 95	09:12,444	84:16,982	1,64	374	3629	D	000610	-2	OAS 18		
OBH 94	09:11,083	84:17,967	1,67	468	3624	D	980403	-2	OAS 13		
OBH 96	09:09,756	84:18,997	2,32	497	3614	B	000611	-4	OAS 75	LG 01 (4,5 Hz)	
OBH 97	09:07,875	84:20,329	2,11	548	03B4+0355	D	971202	-3	HTI 202	Owen 02 (4,5 Hz)	
OBH 98	09:06,187	84:21,581	2,31	649	03BA+0355	C	980401	-4	HTI 702	Owen 14 (30 Hz)	
OBH 99	09:04,285	84:22,966	1,55	840	03BC+0355	D	000616	-5	HTI 101	Owen 16 (30 Hz)	
OBH 100	09:03,013	84:23,927	1,51	1077	0386+0355	D	980402	-4	OAS 46	Owen 21 (4,5 Hz)	
OBH 101	09:01,785	84:24,812	1,79	1157	6334	D	991247	0	OAS 29		
OBH 102	09:00,351	84:25,876	1,70	1296	3609	D	980901	+9	HTI 902		
OBH 103	08:58,937	84:26,892	1,60	1429	039A+0355	B	000707	-3	OAS 22		
OBH 104	08:57,671	84:27,841	2,00	1614	3674	C	000614	-5	HTI 402		
									HTI 102		
									OAS 14		
									OAS 06		
OBH 105	08:56,043	84:29,050	2,13	1794	3659	D	991256	-1	OAS 30		
OBH 106	08:54,313	84:30,349		1960	3669	C	991248	-1	OAS 02		

INST.	LAT (N) D:M	LONG (W) D:M	DIST. TO NEXT (nm)	DEPTH (m)	REL. CODE	ANT. CH.	REC. REC.	SKEW (ms)	SENSOR	Other Sensors	REMARKS
OBH 142	10:20.969	86:07.899	1.08	207	3609	B	010404	1	HTI 902		
OBH 143	10:20.262	86:08.681	1.08	302	3669	D	010708	1	HTI 302		
OBH 144	10:19.542	86:09.433	1.08	401	0398 + 0355	C	000711	-2	DPG 95		
OBH 145	10:18.760	86:10.243	1.08	523	6334	D	000707	-3	OAS 22		
OBH 146	10:17.981	86:11.003	1.08	656	3659	D	991256		OAS 30		no data
OBH 147	10:17.209	86:11.792	1.08	804	3674	D	000614	-4	HTI 002		
									HTI 402		
									OAS 14		
									OAS 34		
OBH 148	10:16.465	86:12.560	1.08	927	0382 + 0355	C	991248	-2	HTI 602		OBS frame
OBH 149	10:15.692	86:13.383	1.08	1033	0381 + 0355	D	991252	-5	HTI 802		OBS frame
OBH 150	10:14.943	86:14.094	1.08	1153	3624	D	980901	10	OAS 13		
OBH 151	10:14.161	86:14.923	1.08	1337	4444	C	991238	-1	OAS 18		
OBS 152	10:13.402	86:15.681	1.08	1546	039A + 0355	D	010401	-5	OAS 31	PMD 509	
OBS 153	10:12.652	86:16.504	1.08	1701	3629	C	000610	-9	HTI 502	Owen 14 (30 Hz)	OBS-design 2002
OBS 155	10:11.933	86:17.267	1.08	1878	0386 + 0355	C	971202	-4	OAS 04	Owen 21 (4.5 Hz)	
OBS 157	10:11.149	86:18.016	1.08	2072	038C + 0355	C	980401	-5	HTI 202	Owen 16 (30 Hz)	
OBS 158	10:10.347	86:18.767	1.08	2224	3614	A	000611	-5	OAS 75	Owen 18 (30 Hz)	
OBS 159	10:09.599	86:19.572	1.08	2399	0399 + 0355	B	991243	0	OAS 09	PMD 540	
OBH 160	10:08.843	86:20.330	1.08	2583	0397 + 0355	D	000706	5	HTI 102		
OBH 161	10:08.120	86:21.173	1.08	2732	3664	D	980401	-2	OAS 27		1000 m anchor rope
OBH 162	10:14.241	86:19.563	1.08	1919	3659	B	010404	1	OAS 30		
OBS 154	10:13.430	86:18.820	1.08	1902	0384 + 0355	C	000101	-4	HTI 101	Owen 02 (4.5 Hz)	west of line P32
OBH 163	10:12.709	86:18.055	1.08	1886	3674	D	000614	-2	HTI 002		
									HTI 402		
									OAS 14		
									OAS 34		
OBS 155	10:11.933	86:17.267	1.08	1878	0386 + 0355	C	971202	-4	OAS 04	Owen 21 (4.5 Hz)	
OBH 164	10:11.137	86:16.530	1.08	1841	6334	D	010708	1	OAS 22		
OBS 156	10:10.361	86:15.749	1.08	1906	038A + 0355	D	000616	-9	HTI 702	Owen 22 (4.5 Hz)	east of line P32
OBH 165	10:09.576	86:14.940		1993	3669	D	000707	-3	HTI 302		

SUBDUCTION

INST.	LAT (N) D:M	LON (W) D:M	DIST. TO NEXT (nm)	DEPTH (m)	REL. CODE	ANT. CH.	REC. REC.	SKEW (ms)	SENSOR	Other Sensors	REMARKS
OBS 171	10:18.867	86:17.649	0.40	1481	0398+0355	D	991248	-1	OAS 13	Owen 02 (4.5Hz)	no data
OBH 167	10:17.691	86:18.496	0.44	1570	3669	B	991252	0	HTI 302		
OBH 168	10:18.002	86:18.265	0.40	1525	6334	D	000711	0	DPG 95		
OBH 169	10:18.348	86:18.001	0.43	1591	3664	C	991238	0	OAS 27		
OBH 170	10:18.610	86:17.799	0.34	1533	0397+0355	D	010708	0	HTI 102		
OBH 166	10:17.323	86:18.747	0.28	1686	0384+0355	A	000708	-1	HTI 101		
OBS 172	10:18.873	86:17.603		1470	039A+0355	D	991243	0	OAS 31	PMD 509	
OBH 173	09:26.944	85:39.277	0.11	2389	0398+0355	C	980901	7	OAS 01		500m anchor rope
OBH 186	09:27.019	85:39.236	0.11	2375	0397+0355	D	991256	-2	HTI 102		
OBH 187	09:27.089	85:39.136	0.11	2351	3664	A	980403	-1	OAS 27		
OBS 176	09:27.171	85:39.072	0.11	2325	3614	B	000616	-6	OAS 75	Owen 18 (30 Hz)	
OBS 177	09:27.266	85:38.998	0.11	2311	03BC+0355	D	000611	-3	HTI 202	Owen 16 (30 Hz)	
OBH 178	09:27.331	85:38.930	0.24	2323	3669	D	991238	-1	HTI 302		
OBH 179	09:27.570	85:39.018	0.11	2332	3674	C	000614	-5	OAS 14		offline
									OAS 34		
									HTI 002		
									HTI 402		
OBS 180	09:27.498	85:38.932	0.11	2335	03B4+0355	C	991252	-3	HTI 101	Owen 02 (4.5 Hz)	offline
OBH 181	09:27.420	85:38.856	0.11	2328	3659	B	991248	-1	OAS 30		
OBH 182	09:27.355	85:38.775	0.11	2323	4A44	C	010404	3	OAS 18		offline
OBS 183	09:27.268	85:38.688	0.24	2322	0386+0355	C	971202	-3	OAS 46	Owen 21 (4.5 Hz)	offline
OBS 175	09:27.490	85:38.780	0.11	2324	03BA+0355	C	980402	-4	HTI 702	Owen 22 (4.5 Hz)	
OBS 185	09:27.557	85:38.685	0.11	2318	0399+0355	D	010703	0	OAS 09	LG 01	
OBS 174	09:27.666	85:38.618	0.11	2319	039A+0355	C	991243	0	OAS 31	PMD 509	
OBS 184	09:27.740	85:38.557	0.11	2304	3629	D	000610	-7	HTI 502	Owen 14 (30 Hz)	OBS design 2002
OBH 188	09:27.813	85:38.494	0.11	2307	3624	D	000706	3	OAS 22		
OBH 189	09:27.903	85:38.403	0.11	2300	3609	D	010401	-4	HTI 902		
OBH 190	09:27.976	85:38.345	0.11	2281	0387+0355	C	980401	-4	OAS 38		
OBH 192	09:28.070	85:38.255	0.11	2283	03BA+0355	D	010708	0	HTI 802		OBS-frame

AIRGUN-SHOTS

Line	Shot Number	Date	Time UTC	Latitude N	Longitude W	Trigger interval [s]	Gun type	Remarks
P 1	1	23.04.02	14:05:38	09:40,960	85:53,952	30	32 L	at 9.4m depth
	121	23.04.02	15:05:19	09:38,342	85:52,156			
P 2	1	23.04.02	15:20:41	09:38,342	85:52,156	30	32 L	at 7.5m depth
	172	23.04.02	16:46:10	09:41,193	85:54,195			
P 3	1	23.04.02	17:03:15	09:42,097	85:54,87	30	32 L	at 5.5m depth
	196	23.04.02	18:40:52	09:38,342	85:52,156			
P 4	1	23.04.02	18:56:20	09:37,655	85:51,681	30	32 L	at 3.5m depth
	176	23.04.02	20:23:45	09:41,211	85:54,162			
P 5	1	23.04.02	20:37:45	09:41,950	85:54,682	30	32 L	at 12.5m depth
	202	23.04.02	22:17:56	09:38,223	85:52,061			
P 6	1	23.04.02	22:28:40	09:37,758	85:51,755	30	32 L	at 14.5m depth
	177	23.04.02	23:56:42	09:41,527	85:54,418			
P 9	1	26.04.02	00:10:47	09:39,630	85:18,776	60	32 L	
	1224	26.04.02	20:33:00	08:33,563	86:10,591			
P 10	1	26.04.02	21:06:00	08:33,553	86:09,592	60	G-Gun cluster	
	895	27.04.02	12:00:46	09:34,740	85:22,500			
P 11	1	27.04.02	16:51:28	09:17,160	85:36,327	20	sb+ps G-Gun	first shot both
	40	27.04.02	17:04:29					last shot both
	41	27.04.02	17:04:49			20	sb G-Gun	first shot sb
	52	27.04.02	17:08:25					last shot sb
	53	27.04.02	17:08:45			20	ps G-Gun	first shot ps
	64	27.04.02	17:12:45					last shot ps
	65	27.04.02	17:13:05			20	sb+ps G-Gun	first shot both
	76	27.04.02	17:17:05					last shot both
	77	27.04.02	17:17:25			20	sb G-Gun	first shot sb
	88	27.04.02	17:21:25	09:15,674	85:37,646			last shot sb

AIRGUN-SHOTS

Line	Shot Number	Date	Time UTC	Latitude N	Longitude W	Trigger interval [s]	Gun type	Remarks
P 12	1	27.04.02	17:41:06	09:16,393	85:35,530	20	sb+ps G-Gun	first shot both
	72	27.04.02	18:04:45			20	ps G-Gun	first shot ps
	84	27.04.02	18:08:45			20	sb G-Gun	first shot sb
	96	27.04.02	18:12:45			20	sb+ps G-Gun	first shot both
	108	27.04.02	18:16:45			20	ps G-Gun	first shot ps
	120	27.04.02	18:20:45			20	sb G-Gun	first shot sb
	132	27.04.02	18:24:48	09:15,834	85:36,540	20	sb+ps G-Gun	first shot both
P 13	1	27.04.02	18:37:25	09:15,413	85:35,875	20	sb+ps G-Gun	first shot both
	79	27.04.02	19:03:25			20	ps G-Gun	first shot ps
	91	27.04.02	19:07:25			20	sb G-Gun	first shot sb
	103	27.04.02	19:11:25			20	sb+ps G-Gun	first shot both
	115	27.04.02	19:15:25			20	ps G-Gun	first shot ps
	126	27.04.02	19:19:08	09:17,337	85:36,683			last shot ps
P 14	1	27.04.02	20:00:48	09:17,750	85:36, 044	20	ps G-Gun+32 L	first shot ps+32 L
	13	27.04.02	20:04:48			20	sb G-Gun+32 L	first shot sb+32 L
	25	27.04.02	20:08:48			20	sb+ps G-Gun+32 L	first shot both+32 L
	37	27.04.02	20:11:48			20	ps G-Gun+32 L	first shot ps+32 L
	49	27.04.02	20:16:48			20	sb G-Gun+32 L	first shot sb+32 L
	61	27.04.02	20:20:48			20	sb+ps G-Gun+32 L	first shot both+32 L
	73	27.04.02	20:24:48			20	ps G-Gun+32 L	first shot ps+32 L
	85	27.04.02	20:38:48			20	sb G-Gun+32 L	first shot sb+32 L
	97	27.04.02	20:32:48			20	sb+ps G-Gun+32 L	first shot both+32 L
	100	27.04.02	20:33:48			20	ps+sb G-Gun+32 L	last shot both+32 L
	108	27.04.02	20:36:28	09:15,831	85:37,421			last shot profile
P 15	1	29.04.02	17:46:48	08:44,004	84:34,871	30	G-Gun cluster	
	1304	30.04.02	04:44:00	09:17,417	84:10,404			
P 16	1	30.04.02	05:42:42	09:15,800	84:11,446	30	sb G-Gun+32 L	first shot sb+32 L
	147	30.04.02	06:55:31	09:11,951	84:14,242			last shot sb+32 L
	148	30.04.02	07:25:27	09:10,304	84:15,501	30	32 L	first shot 32 L
	973	30.04.02	15:31:30	08:44,20	84:37,324			last shot 32 L

AIRGUN SHOTS

Line	Shot Number	Date	Time UTC	Latitude N	Longitude W	Trigger interval [s]	Gun type	Remarks
P 17	1	02.05.02	13:36:58	08:51,360	84:08,571	30	G-Gun cluster	
	1092	02:05:02	22:42:27	09:08,358	84:35,869			
P 18	1	02.05.02	23:09:30	09:08,678	84:35,936	30	32 L	
	1048	03.05.02	07:43:00	08:52,285	84:10,007			
P 19	1	04.05.02	03:07:08	08:52,346	84:14,907	10	Prakla-array	
	2443	04.05.02	09:54:00	09:05,044	84:35,040			
P 20	1	04.05.02	09:53:15	09:02,320	84:37,254	10	Prakla-array	first shot A
	457	04.05.02	11:09:30	09:01,430	84:37,251			last shot A
	458	04.05.02	11:57:16	09:01,500	84:37,250			first shot B
	566	04.05.02	12:15:23	09:02,500	84:37,256			last shot B
P 21	1	04.05.02	13:10	09:05,100	84:35,290	30	32 L	
	876	04.05.02	20:28	08:51,547	84:13,282			
P 22	1	04.05.02	21:08:18	08:51,631	84:13,614	30	G-Gun cluster	
	604	05.05.02	04:09:38	09:05,058	84:35,075			
P 23	1	05.05.02	04:58:11	09:05,056	84:35,001	30	32 L	
	227	05.05.02	06:45:13	09:01,405	84:28,902			stop
	228	05.05.02	07:16:22	09:00,826	84:28,232	30	G-Gun cluster	start again
	816	05.05.02	12:10:08	08:51,620	84:13,524			
P 24	1	06.05.02	10:02:51	08:44,766	84:37,413	30	G-Gun cluster	
	1443	06.05.02	22:03:51	09:18,635	84:12,431			
P 25	1	06.05.02	22:36:21	09:18,475	84:12,365	30	G-Gun cluster+32 L	
	1610	07.05.02	12:00:50	08:45,471	84:36,618			
P 26	1	08.05.02	06:00:16	08:34,988	84:21,426	10	Prakla -7 guns	
	137	08.05.02	06:23:00	08:35,899	84:21,222		Prakla -6 guns	
	1465	08.05.02	10:04:16	08:42,814	84:16,922			

AIRGUN-SHOTS

Line	Shot Number	Date	Time UTC	Latitude N	Longitude W	Trigger interval [s]	Gun type	Remarks
P 27	1	08.05.02	10:46:50	08:43,207	84:16,609	10	G-Gun cluster	
	1819	08.05.02	15:48:10	08:35,762	84:21,297			
P 28	1	08.05.02	15:48:20	08:35,762	84:21,297	20	G-Gun cluster	
	686	08.05.02	19:36:12	08:42,792	84:16,929			
P 29	1	08.05.02	20:51:20	08:39,800	84:18,598	20	1 ps G-Gun	first shot 1ps
			20:53:20				2 ps G-Gun	first shot 2ps
			20:55:20				1 sb G-Gun	first shot 1sb
			20:57:00				2 sb G-Gun	first shot 2sb
			20:58:40				2 ps+2 sb G-Gun	first shot 2 ps+2 sb
			21:00:20				1 ps+1 sb G-Gun	first shot 1 ps+1 sb
			21:02:00				2 ps+1 sb G-Gun	first shot 2 ps+1 sb
			21:03:40				1 ps+2 sb G-Gun	first shot 1 ps+2 sb
P 30	1	11.05.02	07:05:04	08:33,576	84:14,000	20	G-Gun cluster	
	741	11.05.02	11:11:44	08:38,880	84:22,951			
P 31	1	11.05.02	11:31:04	08:38,171	84:22,985	20	G-Gun cluster	
	613	11.05.02	14:55:00	08:33,710	84:15,147			
P 32	1	13.05.02	00:33:17	09:41,677	86:48,186	60	G-Gun cluster+32 L	
	955	13.05.02	16:27:16	10:31,585	85:56,966			
P 33	1	13.05.02	23:12:45	10:07,772	86:13,176	10	ps G-Gun	
	320	14.05.02	01:50:15	10:16,171	86:21,501			
P 34	1	14.05.02	18:51:10	10:19,984	86:16,796	10	sb G-Gun	
	511	14.05.02	20:16:09	10:16,732	86:19,262			
P 35	1	15.05.02	18:14:35	09:31,018	85:42,630	60	G-Gun cluster	
	122	15.05.02	20:15:22	09:25,158	85:35,606			
P 36	1	15.05.02	20:16:22	09:25,158	85:35,606	60	G-Gun cluster	
	171	15.05.02	23:06:22	09:35,902	85:31,144			

AIRGUN-SHOTS

S0 163-2

Line	Shot Number	Date	Time UTC	Latitude N	Longitude W	Trigger interval [s]	Gun type	Remarks
P 37	1	15.05.02	23:06:22	09:35,902	85:31,144	60	G-Gun cluster	
	678	16.05.02	10:24:21	08:59,333	86:04,947			
P 38	1	17.05.02	14:53:24	09:00,413	84:35,406	60	G-Gun cluster	first shot G-Gun
	22	17.05.02	15:14:25	09:00,968	84:35,884		G-Gun+32 L	first shot G-Gun+32 L
	368	17.05.02	21:00:19	09:22,398	84:53,977			
P 39	1	17.05.02	21:41:19	09:21,168	84:55,528	60	G-Gun cluster+32 L	
	324	18.05.02	03:04:19	08:55,535	84:59,499			
P 40	1	18.05.02	03:30:19	08:54,384	84:58,616	60	G-Gun cluster+32 L	
	527	18.05.02	12:16:18	09:31,480	84:41,472			

Appendix 9.7.1

F.S. "S O N N E"

Reise SO 163/1

<u>Eingesetzte Geräte</u>		<u>Einsätze</u>
OBH/ OBS	Ocean Bottom Hydrophones / Seismograph	10 ausgelegt
TOBI	Towed Ocean Bottom Investigation Vehicle	
DTS	tiefgeschelepptes Sidescan Sonar	
PSD	POSIDONIO Transpondersystem zur Ortung des TOBI	
MAG	Magnetometer	
EM	SIMRAD-Fächerlot	26
OFOS	Ocean Floor Observation System	
GTV-A	Fernsehgreifer Typ A	
		09

Vermessen wurden :

Mit EM	897 sml
Mit TOBI	494 sml
Mit DTS	321 sml
Mit OFOS	052 sml
Mit MAG	612 sml

Eingesetzte Winden :

<i>Winde</i>	<i>D/M</i>	<i>Typ</i>	<i>RF-Nr</i>	<i>SO 163/1 Einsatz</i>	<i>Gesamt Einsatz</i>	<i>SO 163/1 S`länge</i>	<i>Gesamt S`länge</i>	<i>Zust.</i>
W 1	18,2	LWL	812001	0000 h	0000 h	000000 m	0000000 m	1
W 2	18,2	LWL	810001	0618	0942	093688	0199774	3
W 4	11,0	NSW	819052	0000	0000	000000	0000000	1
W 5	11,0	NSW	817164	0003	0228	002100	0194453	4
W 6	18,2	DRAKO	814150	0002	2132	001750	1856443	3

<i>Winde</i>	<i>SO 163/1 gefierte max.Länge</i>	<i>jemals gefierte max.Länge</i>
W 1	0000 m	0000 m
W 2	7202	7202
W 4	0000	0000
W 5	2100	5200
W 6	1750	7900 (SL aktuell = 6456 m)

Geräteverluste :

Keine

Abkürzungen im Stationsprotokoll:

z.W.	zu Wasser
a.D.	an Deck
Boko	Bodenkontakt
Bosi	Bodensicht
SL(max.)	(maximale)Seillänge
LT	Lottiefe nach Hydrosweep
W x	eingesetzte Winde
SM	Simrad- Multibeam-Lot
PS	Parasound
XPNDR	Transponder

Zeit : UTC – 05 Stunden**15.03.02**Profil EM-01 v = 10,0 kn

1036 – 1050 bringen Magnetometer aus

1050 Beginn Profil 07 31,91 N 83 01,31 W

Kurs = 298 Grad

1409 ä.K. auf 296 Grad 07 45,00 N 83 30,00 W 22 sml

2124 ä.K. auf 318 Grad 08 17,44 S 84 55,81 W 73 sml

2128 stellen Uhren 20 Minuten zurück

2400 08 37,84 N 84 54,78 W

16.03.02

0003 stellen Uhren 20 Minuten zurück

0430 stellen Uhren 20 Minuten zurück

0656 Ende Profil 09 39,50 N 85 51,57 W 111 sml

0703 Beginn Einholen Magnetometer

0713 Magnetometer a:D.

Bordzeit : UTC – 6 StundenReleaser-Teststation W 6

0729 Beginn Station LT = 1857 m 09 39,55 N 85 51,56 W

0731 3 Releaser z.W.

0825 S_lmax = 1750 m LT = 1860 m

0825 – 0827 Releasertest per Transducer

0827 Beginn hieven

0904 Releaser a.D./ Ende Station

Auslegung OBH/S 1-10

0920 OBH 1 z.W. LT = 1866 m 09 39,46 N 85 51,50 W

0946 OBH 2 z.W. LT = 1924 m 09 40,60 N 85 53,70 W

1017 OBH 3 z.W. LT = 2065 m 09 38,14 N 85 53,49 W

1037 OBH 4 z.W. LT = 1997 m 09 38,85 N 85 52,51 W

1056 OBH 5 z.W. LT = 1940 m 09 39,91 N 85 52,80 W

1112 OBH 6 z.W. LT = 1980 m 09 39,39 N 85 53,58 W

1132 OBH 7 z.W. LT = 1999 m 09 39,09 N 85 53,11 W

1143 OBH 8 z.W. LT = 1970 m 09 39,45 N 85 52,54 W

1204 OBH 9 z.W. LT = 1963 m 09 39,69 N 85 53,10 W

1226 OBH 10 z.W. LT = 1893 m 09 39,46 N 85 52,90 W

Station 03 CTD W 5

1308 Beginn Station LT = 2093 m 09 38,28 N 85 53,91 W

1310 CTD z.W.

1322 CTD a.D.

1339 CZD z.W. LT = 2100 m 09 38,39 N 85 53,87 W

1414 POSIDONIA-Transducer im Loschacht abgesenkt

1418 S_lmax = 1900 m LT = 2069 m 09 38,36 N 85 53,68 W

1501 CTD a.D. / Ende Station 09 38,06 N 85 53,47 W

Kalibrierung POSIDONIA

1518 OBH-K z.W. LT = 2089 m 09 38,31 85 53,89 W

1700 POSIDONIA ausgefallen

1731 OBH-K ausgelöst 1747 aufgetaucht 1757 a.D. 09 38,24 N 85 53,58 W

Profile EM-02 und 03 v= 10,0 kn

1815 Beginn Aussetzen Magnetometer

1820	Beginn EM 02	09 38,28 N 85 53,91 W	
1826	Magnetometer ausgesteckt	l = 250 m	
2009	ä.K. auf 322 Grad	09 50,02 N 85 44,36 W	15 sml
2104	ä.K. auf 219	09 56,33 N 85 49,19 W	09 sml
2115	ä.K. auf 144	09 55,40 N 85 50,23 W	01 sml
2156	ä.K. 244	09 49,91 N 85 46,39 W	07 sml
2207	ä.K. 327	09 49,06 N 85 47,87 W	02 sml
2314	ä.K. 041	09 58,39 N 85 53,91 W	11 sml
2344	ä.K. 306	10 00,98 N 85 51,69 W	04 sml
2400	Profilwechsel EM 02/03	10 03,20 N 85 54,60 W	04 sml

17.03.02

0035	ä.K. 292	10 08,07 N 85 58,54 w	06 sml
0045	ä.K. 139	10 07,22 N 85 59,72 W	02 sml
0123	ä.K. 216	10 02,41 N 85 55,76 W	06 sml
0133	ä.K. 320	10 01,13 N 85 56,75 W	02 sml
0232	ä.K. 233	10 08,53 N 86 02,58 W	10 sml
0245	ä.K. 141	10 07,32 N 86 04,70 W	02 sml
0346	ä.K. 234	09 59,36 N 85 58,49 W	10 sml
0404	ä.K. 322	09 57,57 N 86 00,86 W	03 sml
0512	ä.K. 047	10 06,38 N 86 07,83 W	11 sml
0602	ä.K. 322	10 12,01 N 86 01,80 W	08 sml
0740	v = 7,0 kn		
0807	ä.K. 050	10 28,10 N 86 13,85 W	20 sml
0821	ä.K. 142	10 28,93 N 86 12,81 W	01 sml
1050	Ende Profil EM 03		
1056 – 1106	holen Magnetometer ein		

Profile TOBI 01 – xxv = 2,5 kn W 2

1135	Beginn aussetzen TOBI		
1136	TOBI z.W.		
1211	Depressor z.W. Beginn Profil 01	10 12,15 N 86 02,54 W	
	Kurs = 318 Grad	fieren bis SL = 100 m	
1229	v = 2,5 kn		
1240 – 1248	fieren auf SL = 202 m		
1256	Magnetometer z.W.	SL = 30 m	
1559 – 1604	holen Magnetometer ein		
1800	SL = 506 m	10 22,75 N 86 13,63 W	
2400	SL = 326 m LT = 591 m	10 34,76 N 86 24,55 W	

18.03.02

0047	Abbruch Profil	10 36,50 N 86 26,12 W	35 sml
0051	hieven auf SL = 200 m		
0057	ä.K. auf 300 Grad v = 2,0 kn	warten auf Tageslicht	
0611	hieven SL = 200 m LT = 990 m	10 40,91 N 86 34,76 W	
0622	Depressor a.D.	Check der Kabelverbiindung	
0710	Beginn Einholen TOBI		
0808	TOBI a.D.		

Profil EM-04 v = 6,0 kn

0936	Beginn Profil	Kurs = 21 Grad	10 41,92 N 86 34,92 W	
1029	ä.K. 270 Grad		10 45,91 N 86 32,67 W	4 sml
1058	ä.K. 137		10 46,07 N 86 35,79 W	3 sml
1245	Ende Profil		10 38,75 N 86 29,18 W	10 sml

Profil EM-05 v = 6,0 kn

1323	Beginn Profil	Kurs = 324 Grad	10 37,32 N 86 36,00 W
1329	Magnetometer z.W.	SL = 250 m	

1510	ä.K. 090	10 45,94 N 86 42,46 W	11 sml
1522	ä.K. 124	10 46,26 N 86 41,54 W	01 sml
1612	Magnetometer a.D.		
1613	Ende Profil	10 42,43 N 86 36,84 W	06 sml

Profil TOBI 02 - xx W 2

1623	Beginn Aussetzen	10 42,48 N 86 37,01 W	
1626	TOBI z.W.		
1704	Depressor z.W. LT = 1036 m	10 42,68 N 86 36,41 W	
	Kurs 137 Grad v = 2,5 kn fieren W 2		
1800	Beginn Profil 2 SL = 755 m LT = 976 m	10 41,60 N 86 34,60 W	
2352	SL = 931 m LT = 974 m	10 32,02 N 86 24,00 W	
19.03.02			
0243	SL = 1518m LT = 1284 m		
0600	SL = 999 m LT = 986 m	10 18,52 N 86 13,51 W	
0924	Ende Profil 2 SL = 610 m LT = 914 m	10 12,31 N 86 07,92 W	39 sml
0924 – 1947 Schleife über Stb			
1047	Beginn Profil 03 SL = 794 m LT = 1179 m	10 10,68 N 86 09,98 W	
	Kurs = 318 Grad v = 2,5 kn		
1200	SL = 1631 m LT = 1129 m	10 13,02 N 86 12,09 W	
1600	SL = 2255 m LT = 1552 m	10 20,59 N 86 18,92 W	
2020	SL = 1712 m LT = 1192 m	10 28,85 N 86 26,34 W	
2400	SL = 1713 m LT = 1280 m	10 35,51 N 86 32,36 W	
20.03.02			
0234	Ende Profil 03 SL = 1713 m LT = 1265 m	10 40,20 N 86 36,60 W	39 sml
0234 – 0414 Schleife über Bb			
0414	Beginn Profil 04 SL = 1342 m LT = 1644 m	10 38,81 N 86 38,90 W	
	Kurs = 138 Grad v = 2,5 kn		
0800	SL = 2472 m LT = 1634 m	10 31,85 N 86 32,63 W	
1200	SL = 2697 m LT = 1845 m	10 24,34 N 86 25,87 W	
1600	SL = 2788 m LT = 1775 m	10 16,89 N 86 19,19 W	
2000	SL = 2178 m LT = 1626 m	10 09,34 N 86 12,39 W	
2217	Ende Profil 04 SL = 2231 m LT = 1544 m	10 05,19 N 86 08,69 W	40 sml
2246 – 2400 Schleife über Stb			
21.03.02			
0000 – 0028 Fortsetzung Schleife			
0028	Beginn Profil 05 SL = 1672 m LT = 1973 m	10 03,50 N 86 10,77 W	
	Kurs = 318 Grad v = 2,5 kn		
0348	SL = 3673 m LT = 2897 m	10 09,60 N 86 16,32 W	
0800	SL = 3223 m LT = 2139 m	10 17,15 N 86 23,21 W	
1200	SL = 3718 m LT = 2153 m	10 24,49 N 86 29,79 W	
1600	SL = 3416 m LT = 1958 m	10 32,00 N 86 36,61 W	
1950	Ende Profil 05 SL = 3004 m LT = 2166 m	10 37,29 N 86 41,47 W	45 sml
2023	SL = 3004 m LT = 2187 m	10 38,78 N 86 42,37 W	
Beginn Schleife über Bb			
2039	Ende Schleife	10 36,77 N 86 44,24 W	
2053	fieren mit 0,2 m/sec		
2107	Beginn Profil 06 SL = 3160 m LT = 2499 m	10 35,93 N 86 43,49 W	
2200	SL = 4100 m LT = 2477 m	10 34,39 N 86 42,11 W	
2400	SL = 3287 m LT = 2413 m	10 30,55 W 86 38,70 W	

22.03.02

0009	Slmax = 4352 m		
0600	Ende Profil 06 SL = 4235 m LT = 2694 m	10 19,33 N 86 26,87 W	22 sml
	beginnen hieven mit 0,5 m/sec		

0824 Depressor a.D.

0912 TOBI a.D.

OFOS-Profile 01-02 W 2

1042	stand by	LT = 1515 m	10 17,85 N 86 18,38 W
1110	OFOS z.W.		
1141	Bosi SL = 1522 m	LT = 1509 m	10 17,86 N 86 18,37 W
1234			10 17,88 N 86 17,88 W
1342	hieven ein SL = 1534 m	LT = 1553 m	10 17,94 N 86 17,83 W
1421	OFOS a.D.		
1455	OFOS z.W.	LT = 1552 m	10 17,78 N 86 18,52 W
1529	Bosi SL = 1548 m	LT = 1554 m	10 17,81 N 86 18,56 W
1721	hieven SL = 1575 m	LT = 1598 m	10 18,17 N 86 18,05 W
1810	OFOS a.D. / Ende Station		

GTV 01-02 W 2

1924	Beginn Station 01		
1926	GTV z.W.	LT = 1531 m	10 17,96 N 86 18,40 W
1959	Bosi SL = 1520 m	LT = 1529 m	10 17,93 N 86 18,35 W
2007	1. Griff SL = 1519 m	LT = 1523 m	10 17,94 N 86 18,33 W
2049	GTV a.D. / Ende Station 1		
2055	Beginn Station GTV-02		
2057	GTV z.W.		
2126	Bosi SL = 1504 m	LT = 1518 m	10 17,93 N 86 18,33 W
2134	1. Griff SL = 1515 m	LT = 1520 m	10 17,96 N 86 18,28 W
2214	GTV a.D.		
2254	Ende GTV-02		

Profil EM-06 v = 10,0 kn

2304	Magnetometer z.W.	
2311	Beginn Profil	10 19,09 N 86 17,98 W
	Kurs = 25 Grad	

23.03.02

0006	ä.K. 325 Grad	10 26,96 N 86 13,98 W	10 sml
0028	ä.K. 312	10 30,00 N 86 16,00 W	04 sml
0209	ä.K. 054	10 41,30 N 86 28,50 w	17 sml
0222	ä.K. 131	10 41,82 N 86 27,67 W	02 sml
0409	ä.K. 044	10 31,44 N 86 15,00 W	17 sml
0421	ä.K. 306	10 32,85 N 86 13,69 W	02 sml
0526	ä.K. 213	10 39,40 N 86 22,50 W	11 sml
0536	ä.K. 128	10 38,83 N 86 22,87 w	01 sml
0640	ä.K. 090	10 32,68 N 86 15,08 w	10 sml
0645	ä.K. 156	10 32,49 N 86 14,38 W	01 sml
0730	ä.K. 195	10 25,87 N 86 11,37 W	07 sml
0933	Ende Profil	10 05,91 N 86 16,73 W	21 sml

Profil TOBI 07 W 2 v = 2,5 kn

1346	Beginn Aussetzen TOBI	10 06,85 N 86 17,89 W
1349	TOBI z.W.	
1416	Depressor z.W.	10 06,69 N 86 17,56 W
1421	KL = 157 m Störung, hieven ein	
1435	Depressor a.D.	

1519	Depressor z.W.	LT = 2453 m	10 06,02 N 86 16,71 W	
	Kurs = 138 Grad	fieren mit 0,5 m/sec	erhöhen auf 2,5 kn	
1600	KL = 1205 m	LT = 2379 m	10 04,80 N 86 15,61 W	
1955	ä.K. 147 SL = 3847 m	LT = 2319 m	09 57,41 N 86 09,01 W	11 sml
2400	KL = 4245 m	LT = 2456 m	09 48,81 N 86 03,44 W	
24.03.02				
0230	ä.K. 123 SL = 4128 m	LT = 2469 m	09 43,51 N 86 00,00 W	
0938	Ende Profil SL = 2474 m	LT = 2109 m	09 33,57 N 85 44,90 W	18 sml

Profil TOBI 08-09 v = 2,5 kn W 2

0938 – 1056 Schleife über Bb

1130	Beginn Profil	SL = 2338 m	LT = 1590 m	09 36,57 N 85 44,69 W	
1200	SL = 2338 m	LT = 1464 m	09 37,24 N 85 45,67 W		
1600	SL = 2884 m	LT = 1918 m	09 43,00 N 85 54,08 W		
1700	Ende Profil 08	SL = 3152 m	LT = 2081 m	09 44,39 N 85 56,11 W	14 sml
1700	Beginn Schleife über Stb				
2052	Beginn Profil 09	SL = 2087 m	LT = 1564 m	09 44,93 N 85 50,53 W	
	Kurs = 206 Grad				
2400	SL = 2953 m	LT = 2181 m	09 37,99 N 85 53,95 W		

25.03.02

0116	ä.K. 227 Grad	SL = 3606 m	LT = 2742 m	09 35,00 N 85 55,50 W	11 sml
0727	Ende Profil 09	SL = 5964 m	LT = 3703 m	09 25,58 N 86 05,70 W	14 sml
	Beginn hieven				
1005	Depressor a.D.	1040 Tobi a.D.			

Profil EM-07 v = 10,0 kn

1150	Magnetometer z.W.	SL = 250 m		
1231	Beginn Profil		09 20,70 N 86 04,00 W	
1421	ä.K. 90		09 32,97 N 85 50,20 W	18 sml
1515	Ende Profil EM-07		09 32,00 N 85 41,22 W	09 sml
1529	Magnetometer ein			

Kalibrierung POSIDONIA W 5

1550	OBH-P z.W.	LT = 2052 m	09 33,01 N 85 41,21 W
1659 - 1728	Transponder z.W.	W 5	
1748	OBH ausgelöst		
1816	OBH a.D.		09 33,00 N 85 41,34 W
1827	OBH-P z.W.		09 33,01 N 85 41,20 W
1952 – 2127	Achterfahrt um den OBH		
2146	OBH ausgelöst		
2215	OBH a.D.		09 33,06 N 85 41,24 W

Profile DTS-01 bis xx v = 3,5 kn W 2

2226	Beginn Aussetzen Sidescan	09 33,26 N 85 41,48 W
2230	Depressor z.W. , fieren W 2 mit 0,3 m/sec	

26.03.02

0015	Beginn Profil 01	SL = 3228 m	LT = 1726 m	09 36,13 N 85 46,80 W	
	Kurs = 303 Grad				
0408	Ende Profil 01	SL = 4983 m	LT = 2196 m	09 42,66 N 85 58,20 W	13 sml
0538	Beginn Profil 02	SL = 3842 m	LT = 2383 m	09 40,97 N 85 57,53 W	
0932	Ende Profil 02	SL = 4325 m	LT = 2158 m	09 33,40 N 85 46,04 W	13 sml
0953 – 1045	Schleife über Bb				
1120	Beginn Profil 03	SL = 3730 m	LT = 1802 m	09 35,56 N 85 47,07 W	
1523	Ende Profil 03	SL = 5133 m	LT = 2286 m	09 43,45 N 85 59,00 W	13 sml
1558 – 1645	Schleife				
1800	Beginn Profil 04	SL = 5562 m	LT = 2524 m	09 40,43 N 85 57,83 W	

2114 Ende Profil 04 SL = 4543 m LT = 2007 m 09 34,09 N 85 48,2 W 11 sml
 2212 – 2245 Schleife
 2342 Beginn Profil 05 SL = 4073 m LT = 1874 m 09 35,11 N 85 47,52 W
27.03.02
 0340 Ende Profil 05 SL = 5232 m LT = 2406 m 09 42,75 N 85 59,25 W 13 sml
 0429 – 0512 Schleife
 0639 Beginn Profil 06 SL = 4216 m LT = 2170 m 09 42,94 N 85 56,01 W
 0853 Ende Profil 06 SL = 3801 m LT = 1781 m 09 38,80 N 85 49,60 W 08 sml
 0853 – 1053 Schleife
 1117 Beginn Profil 07 SL = 2504 m LT = 1715 m 09 36,44 N 85 40,95 W
 1310 Ende Profil 07 SL = 4022 m LT = 2217 m 09 36,07 N 85 52,56 W 07 sml
 hieven ein
 1454 Depressor a.D.
 1503 Sidescan a.D.

OFOS 03 W 2

1630 Beginn Station LT = 1591 m 09 37,37 N 85 46,58 W
 1635 OFOS z.W.
 1642 SL = 210 m Ausfall Datenstrecke, hieven ein
 1652 OFOS a.D.
 1721 OFOS z.W.
 1810 Bosi SL = 1595 m LT = 1603 m 09 37,08 N 85 46,53 W
 Kurs = 235 v = 0,8 kn
 1930 ä.K. 251 SL = 1762 m LT = 1760 m 09 36,46 N 85 47,37 W 1,0 sml
 1958 ä.K. 301 SL = 1794 m LT = 1742 m 09 36,33 N 85 47,81 W 0,5 sml
 2133 ä.K. 236 SL = 1830 m LT = 1832 m 09 36,98 N 85 48,51 W 0,9 sml
 2339 ä.K. 270 SL = 1874 m LT = 1905 m 09 36,23 N 85 49,35 W 1,4 sml

28.03.02

0118 ä.K. 206 SL = 1983 m LT = 1973 m 09 36,30 N 85 49,89 W 1,4 sml
 0154 ä.K. 225 SL = 1905 m LT = 1904 m 09 36,10 N 85 50,00 W 0,5 sml
 0225 hieven SL = 1991 m LT = 2002 m 09 35,92 N 85 50,19 W 1,5 sml
 0317 OFOS a.D.
 0320 Ende Station

OFOS 04 W 2

0420 Beginn Station LT = 2222 m 09 42,71 N 85 56,40 W
 0444 OFOS z.W.
 0532 Bosi SL = 2208 m LT = 2214 m 09 42,50 N 85 56,29 W
 Kurs = 231 Grad v = 0,8 kn
 0705 ä.K. 257 SL = 2178 m LT = 2238 m 09 41,85 N 85 56,98 W 1,0 sml
 0755 ä.K. 090 SL = 2233 m LT = 2256 m 09 41,82 N 85 57,38 W 0,5 sml
 0902 ä.K. 045 SL = 2235 m LT = 2258 m 09 41,95 N 85 56,89 W 0,5 sml
 0918 hieven SL = 2303 m LT = 2184 m 09 42,17 N 85 56,72 W 0,3 sml
 1002 OFOS a.D.
 1012 Ende Station

OFOS-05 W 2

1102 OFOS z.W. LT = 2208 m 09 36,31 N 85 54,29 W
 1143 Bosi SL = 2200 m LT = 2208 m 09 36,29 N 85 54,29 W
 Kurs = 237 v = 0,8 kn
 1316 ä.K. 210 SL = 2433 m LT = 2433 m 09 35,80 N 85 55,10 W 1,5 sml
 1608 ä.K. 233 SL = 3080 m LT = 3067 m 09 33,97 N 85 56,14 W 2,0 sml
 1700 hieven ein SL = 3250 m LT = 3232 m 09 33,53 N 85 56,74 W 0,6 sml
 1810 OFOS a.D. Ende Station

OFOS-06 W 2

2043 OFOS z.W. LT = 2248 m 10 02,25 N 86 11,77 W
 2123 – 2249 Ausfall Spulgeschirr
 2302 Bosi SL = 2243 m LT = 2254 m 10 02,18 N 86 11,77 W
 Kurs = 213 Grad v = 0,8 kn

29.03.02

0124 hieven ein SL = 2354 m LT = 2370 m 10 01,43 N 86 12,25 W 1,0 sml
 0223 OFOS a.D.

OFOS-07 W 2

0240 OFOS z.W. LT = 2326 m 10 00,86 N 86 11,41 W
 0324 Bosi SL = 2312 m LT = 2330 m 10 00,71 N 86 11,35 W
 Kurs = 191 Grad v = 0,8 kn
 0432 ä.K. 011 SL = 2319 m LT = 2370 m 10 00,22 N 86 11,45 W 0,5 sml
 0536 hieven SL = 2350 m LT = 2350 m 10 00,66 N 86 11,49 W 0,5 sml
 0634 OFOS a.D.
 0644 Ende Station

Profil EM-08 v = 8,0 kn

0644 Beginn Profil 10 00,98 N 86 11,29 W
 1135 Ende Profil 09 09 49,13 N 86 41,88 W 32 sml

Profil EM-09 v = 5,0 kn

1135 Beginn Profil 09 49,13 N 86 41,88 W
 1200 ä.K. 134 09 48,01 N 86 43,26 w 02 sml
 1658 ä.K. 041 09 30,21 N 86 24,68 W 26 sml
 1715 ä.K. 314 09 31,18 N 86 23,65 W 04 sml
 2057 Ende Profil 09 44,85 N 86 37,63 W 19 sml

OFOS 08 und 09 W 2

2144 Beginn Station LT = 3304 m 09 40,28 N 86 33,81 W
 2146 OFOS z.W.
 2243 Bosi SL = 3284 m LT = 3302 m 09 40,26 N 86 33,78 W
 Kurs = 312 Grad v = 1,0 kn
 2400 SL = 3280 m LT = 3311 m 09 40,78 N 86 34,25 W
30.03.02
 0014 ä.K. 90 SL = 3292 m LT = 3299 m 09 40,74 N 86 34,16 W 1,0 sml
 0119 ä.K. 45 SL = 3272 m LT = 3286 m 09 40,75 N 86 33,52 W 1,0 sml
 0220 hieven SL = 3359 m LT = 3311 m 09 41,39 N 86 32,89 W 1,0 sml
 0342 OFOS a.D.
 0344 Ende Station
 0455 OFOS z.W. LT = 3342 m 09 40,26 N 86 40,84 W
 0558 Bosi SL = 3328 m LT = 3348 m 09 46,49 N 86 40,80 W
 Kurs = 45 Grad v = 1,0 kn
 0641 hieven SL = 3380 m LT = 3382 m 09 47,01 N 86 40,26 w 1,0 sml
 0850 OFOS a.D. Ende Station

Profil EM-10 v = 10,0 kn

0858- 0905 bringen Magnetometer aus
 0920 Beginn Profil LT = 3354 m 09 48,22 N 86 41,97 W
 Kurs = 360
 1234 ä.K. 140 LT = 43100 m 10 20,13 N 86 41,50 W 27 sml
 1703 ä.K. 131 LT = 3373 m 09 45,00 N 86 12,00 W 46 sml
 2027 ä.K. 090 09 21,11 N 85 45,13 W 36 sml
 2156 Ende Profil LT = 1650 m 09 21,01 N 85 29,86 W 15 sml
 2212 Magnetometer a.D.

Kalibrierung POSIDONIA

2217 Transducer im Lotschacht z.W.
 2238 OBH-T z.W. LT = 1656 m 09 21,02 N 85 30,03 W
 2258 Beginn Schleife
31.03.02
 0030 Ende Kalibrierung
 0041 OBH ausgelöst 0106 aufgetaucht 0117 a.D. 09 21,10 N 85 30,14 W

Profile DTS-08 bis 13 W 2 v = 3,5 kn

0224 SideScanSonar z.W.
 0232 Depressor z.W.
 Kurs = 294 Grad v = 2,5 kn fieren W 2 mit 0,5 m/sec
 0504 Beginn Profil 08 09 18,60 N 85 33,95 W
 Kurs = 290 Grad v = 3,5 kn SL = 5117 m LT = 2668 m
 0916 Ende Profil 08 09 23,79 N 85 47,73 W 17 sml
 1200 Beginn Profil 09 09 21,20 N 83 45,80 W
 Kurs = 110 Grad v = 3,0 kn SL = 6144 m LT = 3142 m
 1740 Ende Profil 09 09 15,33 N 85 30,10 W 17 sml
 SL = 3944 m LT = 2718 m
 1935 Beginn Profil 10 09 18,04 N 85 34,12 W
 Kurs = 290 v = 3,0 kn SL = 4551 m LT = 2969 m
01.04.03
 0018 Ende Profil 10 09 23,50 N 85 48,73 W 15 sml
 SL = 6321 m LT = 3272 m
 0434 Beginn Profil 11 09 20,70 N 85 46,05 W
 Kurs = 110 v = 3,0 kn SL = 6809 m LT = 3136 m
 0930 Ende Profil 11 09 15,60 N 85 32,54 W 15 sml
 SL = 5840 m LT = 2714 m
 1147 Beginn Profil 12 09 17,48 N 85 34,33 W
 Kurs = 290 v = 3,2 kn SL = 4937 m LT = 2676 m
 1629 Ende Profil 12 09 22,87 N 85 48,50 W 18 sml
 SL = 6543 m LT = 3294 m
 2011 Beginn Profil 13 09 20,26 N 85 46,45 W
 Kurs = 111 Grad v = 3,2 kn SL = 6872 m LT = 3199 m
02.04.02
 0058 Ende Profil 13 09 14,84 N 85 32,26 W 15 sml
 SL = 6323 m LT = 2743 m
 0059 hieven mit 0,7 m/sec
 0340 Depressor a.D.
 0348 Sidescan a.D.

Profil EM 13 v = 7,5 kn

0647 Beginn Profil 09 09,58 N 85 41,42 W
 0851 ä.K. 56 09 06,10 N 85 36,00 W 14 sml
 1207 Ende Profil 09 19,00 N 85 36,00 W 14 sml

OFOS-10 W 2

1215 Beginn Station LT = 504 m 09 18,82 N 85 15,39 W
 1221 OFOS z.W.
 1239 Bosi SL = 472 m LT = 483 m 09 18,36 N 85 15,50 W
 Kurs = 270 Grad v = 0,8 kn
 1500 hieven ein wegen Angelleinen
 1515 OFOS Wasseroberfläche, ok
 1530 Bosi SL = 544 m LT = 549 m 09 18,91 N 85 17,03 W
 Kurs = 280 Grad v = 0,7 kn
 1727 hieven SL = 780 m LT = 790 m 09 19,18 N 85 18,33 W 1 sml

1752 OFOS a.D.

OFOS-11 W 2

2234 Beginn Station LT = 744 m09 10,42 N 84 50,89 W

2235 OFOS z.W.

2255 Bosi SL = 737 m LT = 743 m09 10,42 N 84 44,71 W

Kurs = 270 v = 0,5 kn

03.04.02.

0117 hieven SL = 773 m LT = 783 m09 10,39 N 84 48,70 W 1 sml

0133 OFOS a.D.

OFOS-12 W 2

0237 Beginn Station LT = 940 m09 10,47 N 84 50,35 W

0239 OFOS z.W.

0303 Bosi SL = 948 m LT = 919 m09 10,36 N 84 50,25 W

Kurs = 132 v = 0,6 kn

0715 hieven SL = 865 m LT = 871 m09 08,70 N 84 48,61 W 3 sml

0750 OFOS a.D.

DTS-Profile 14 bis 22

0908 LT = 1089 m 09 08,00 N 84 41,77 W

0911 Sidescan z.W.

0915 Depressor z.W. fieren mit 0,3 m/s

0941 SL = 582 m fieren mit 0,5 m/s

1012 Beginn DTS-14 SL = 1390 m LT = 1055 m 09 08,01 N 84 45,02 W

Kurs = 270 Grad v = 3,5 kn

1336 Ende DTS-14 SL = 3426 m LT = 2200 m 09 08,00 N 84 56,70 W 12 sm

1522 Beginn DTS-15 SL = 4132 m LT = 1842 m 09 10,40 N 84 53,01 W

Kurs = 90 Grad v = 3,5 kn

1847 Ende DTS-15 SL = 1365 m LT = 760 m09 10,39 N 84 43,33 W 12 sm

2006 Beginn DTS-16 SL = 1325 m LT = 989 m09 08,61 N 84 45,02 W

Kurs = 270 v = 3,5 kn

2333 Ende DTS-16 SL = 3472 m LT = 2217 m 09 08,56 N 84 57,17 W 11 sm

04.04.03

0112 Beginn DTS-17 SL = 3863 m LT = 1921 m 09 11,00 N 84 55,00 W

Kurs = 90 v = 3,5 kn

0422 Ende DTS-17 SL = 1247 m LT = 684 m09 11,00 N 84 43,59 W 12 sm

0552 Beginn DTS-18 SL = 1130 m LT = 923 m09 09,20 N 84 45,00 w

Kurs = 270 v = 3,5 kn

0913 Ende DTS-18 SL = 3565 m LT = 2172 m 09 09,19 N 84 56,69 W 12 sm

1047 Beginn DTS-19 SL = 3471 m LT = 1880 m 09 11,59 N 84 55,06 W

1436 Ende DTS-19 SL = 682 m LT = 538 m09 11,60 N 84 41,81 W 13 sm

1600 Beginn DTS-20 SL = 1044 m LT = 861 m09 09,79 N 84 44,94 W

1923 Ende DTS-20 SL = 3569 m LT = 2144 m 09 09,79 N 84 56,74 W 12 sm

2053 Beginn DTS-21 SL = 3462 m LT = 1818 m 09 12,20 N 84 55,02 W

05.04.02

0037 Ende DTS-21 SL = 565 m LT = 480 m09 12,20 N 84 42,08 W 12 sm

0253 Beginn DTS-22 SL = 1094 m LT = 710 m09 12,40 N 84 46,70 W

Kurs = 206 Grad

0411 ä.K. 215 SL = 935 m LT = 835 m09 09,08 N 84 48,40 W

0613 Ende DTS-22 SL = 3457 m LT = 2376 m 09 04,38 N 84 51,77 W 10 sm

0746 Depressor a.D.

0754 Sidescan a.D.

GTV-03 W 2

0924 Beginn Station LT = 754 m09 10,28 N 84 48,27 W

0925	GTV z.W.				
0944	Bosi	SL = 754 m	LT = 763 m	09 10,42 N	84 48,25 W
0951	Griff	SL = 777 m	LT = 762 m	09 10,44 N	84 48,25 W
0955	hieven				
1013	GTV a.D.				
1044	Ende Station				

OFOS 13 W 2

1240	Beginn Station		LT = 1839 m	08 59,87 N	84 42,22 W
1244	OFOS z.W.				
1252	SL = 291 m	Datenausfall hieven			
1337	OFOS z.W.		LT = 1868 m	08 59,73 N	84 42,57 W
1420	Bosi	SL = 1862 m	LT = 1876 m	08 59,69 N	84 42,51 W
	Kurs = 156	v = 0,9 kn			
1700	ä.K. 90	SL = 2109 m	LT = 2099 m	08 58,62 N	84 42,03 W
1735	hieven	SL = 2033 m	LT = 1999 m	08 58,63 N	84 41,61 W 2 sm
1827	OFOS a.D.				

Profile TOBI 10 bis xx W 2

1909	TOBI z.W.			09 09,95 N	84 41,32 W
1939	Depressor z.W.			09 00,59 N	84 41,15 W
	Kurs = 158	v = 1,5 kn			
2049	v = 2,5 kn				
2400		SL = 5651 m	LT = 3064 m	08 51,16 N	84 37,44 W
06.04.02					
0110	v = 2,2 kn				
0251	ä.K. 121	SL = 5612 m	LT = 3391 m	08 45,02 N	84 34,92 W 18 sm
0400		SL = 5721 m	LT = 3248 m	08 43,65 N	84 32,78 W
0800				08 36,68 N	84 21,34 W
1200		SL = 4280 m	LT = 2585 m	08 34,38 N	84 17,63 W
1535	v = 2,5 kn				
1600		SL = 3752 m	LT = 2391 m	08 29,72 N	84 09,90 W
2000	ä.K. 107	SL = 3855 m	LT = 2216 m	08 24,45 N	84 01,29 W 39 sm
2400		SL = 2891 m	LT = 83 51,51 W	08 21,26 N	83 51,51 W
07.04.02					
0400		SL = 0770 m	LT = 1695 m	08 18,13 N	83 41,81 W
0800		SL = 2790 m	LT = 1657 m	08 14,91 N	83 31,82 W
1205	Ende TOBI-10	SL = 2614 m	LT = 1882 m	08 11,85 N	85 22,39 W 39 sm
1410	Beginn TOBI-11	SL = 2366 m	LT = 1846 m	08 14,80 N	83 22,50 W
	Kurs = 287	v = 2,5 kn			
1800		SL = 1772 m	LT = 1464 m	08 17,71 N	83 31,74 W
2400		SL = 1793 m	LT = 1180 m	08 22,38 N	83 46,49 W

08.04.02

0522	ä.K. 301	SL = 2009 m	LT = 1496 m	08 26,50 N	83 59,40 W 38 sml
1200		SL = 3195 m	LT = 2145 m	08 35,32 N	84 13,74 W
1800		SL = 3627 m	LT = 2316 m	08 43,14 N	84 26,53 W
2109	Ende Tobi-11	SL = 3959 m	LT = 2574 m	08 47,35 N	84 33,34 W 39.5 sm
2324	Beginn TOBI-12	SL = 3780 m	LT = 2378 m	08 48,54 N	84 30,35 W

09.04.02

0600		SL = 2693 m	LT = 1725 m	08 39,88 N	84 16,16 W
1200		SL = 1006 m	LT = 1048 m	08 31,81 N	84 02,94 W
1800		SL = 150 m	LT = 316 m	08 26,08 N	83 48,03 W
2400		SL = 1228 m	LT = 1017 m	08 21,41 N	83 33,98 W

10.04.02

0700 Ende TOBI-12 SL = 1300 m LT = 1008 m 08 16,25 N 83 17,25 W 43 sm
 hieven ein
 0800 Depressor a.D.
 0823 Tobi a.D.

OFOS-14 W 2

1243 OFOS z.W. LT = 1734 m 08 15,56 N 83 34,11 W
 1321 Bosi SL = 1722 m LT = 1738 m 08 15,60 N 83 34,22 W
 Kurs = 248 Grad v = 0,5 kn
 1442 hieven SL = 1915 m LT = 1922 m 08 15,35 N 83 34,80 W 1 sm
 1521 OFOS a.D. / Ende Station

OFOS-15 W 2

1725 Beginn Station LT = 1528 m 08 22,21 N 83 47,54 W
 1727 OFOS z.W.
 1800 Bosi SL = 1267 m LT = 1288 m 08 22,20 N 83 47,70 W
 Kurs = 180 v = 1,0 kn
 1840 ä.K. 143 SL = 1435 m LT = 1480 m 08 21,66 N 83 47,69 W
 1930 stoppen auf SL = 1505 m LT = 1545 m 08 21,19 N 83 47,40 W
 1940 Kurs = 270 v = 0,5 kn
 2052 hieven ein SL = 1624 m LT = 1648 m 08 21,18 N 83 48,10 W 2 sm
 2127 OFOS a.D. Ende Station

OFOS-16 W 2

2248 Beginn Station LT = 709 m 08 26,48 N 83 55,00 W
 2249 OFOS z.W.
 2305 Bosi SL = 699 m LT = 712 m 08 26,49 N 83 54,99 W
 Kurs = 258 v = 0,7 kn

11.04.02

0057 ä.K. 180 SL = 860 m LT = 875 m 08 26,31 N 83 56,00 W
 0148 hieven ein SL = 1113 m LT = 1150 m 08 25,83 N 83 56,01 W 2 sm
 0214 OFOS a.D. Ende Station

OFOS-17 W 2

0545 Beginn Station LT = 1782 m 08 40,63 N 84 18,20 W
 0553 OFOS z.W.
 0634 Bosi SL = 1797 m LT = 1817 m 08 40,60 N 84 18,32 W
 Kurs = 212 v = 0,5 – 1,0 kn
 1000 SL = 2345 m LT = 2357 m 08 38,51 N 84 19,63 W
 1200 SL = 2578 m LT = 2582 m 08 37,28 N 84 20,39 W
 1400 SL = 2806 m LT = 2795 m 08 36,23 N 84 21,06 W
 1525 Ende Bosi SL = 2922 m LT = 2863 m 08 35,46 N 84 21,55 W
 1527 OFOS a.D. Ende Station

TOBI-13 W 2

1820 LT = 1895 m 08 48,50 N 84 24,98 W
 1823 TOBI z.W.
 1851 Depressor z.W. 08 48,15 N 84 24,42 W
 1930 SL = 952 m Datenausfall ; hieven ein
 2002 Depressor a.D.
 2032 TOBI a.D. 08 46,82 N 84 22,71 W
 2116 TOBI z.W.
 2138 Depressor z.W. LT = 1890 m 08 48,40 N 84 24,85 W
 Kurs = 122 v = 1,5 kn
 2200 SL = 624 m LT = 1889 m 08 48,07 N 84 24,35 W
 2241 v = 2,5 kn

2400 SL = 2670 m LT = 1602 m 08 45,65 N 84 20,55 W

12.04.02

0400 SL = 1399 m LT = 1062 m 08 40,26 N 84 12,08 W

0800 SL = 199 m LT = 533 m 08 34,79 N 84 03,55 W

0833 Ende TOBI-13 SL = 151 m LT = 465 m 08 33,97 N 84 02,26 W 27 sm

0833 – 0937 Schleife über Bb

TOBI-14 W 2

0937 Beginn Profil SL = 151 m LT = 366 m 08 35,63 N 84 01,20 W

Kurs = 304 v = 2,5 kn

1200 SL = 200 m LT = 492 m 08 39,22 N 84 06,54 W

1600 SL = 894 m LT = 991 m 08 44,67 N 84 14,67 W

1957 Ende Profil SL = 2085 m LT = 1613 m 08 50,25 N 84 22,96 W 25 sm

TOBI-15 W 2

2210 Beginn Profil SL = 1796 m LT = 1330 m 08 51,86 N 84 20,26 W

Kurs = 127 v = 2,5 kn

0221 Ende Profil SL = 325 m LT = 629 m 08 45,41 N 84 11,88 W 11 sm

TOBI-16 W 2

0333 Beginn Profil SL = 151 m LT = 335 m 08 47,30 N 84 10,70 W

Kurs = 311 v = 311

0943 Ende Profil SL = 1172 m LT = 1126 m 08 57,68 N 84 22,34 W 16 sm

TOBI-17 W 2

1108 Beginn Profil SL = 956 m LT = 863 m 08 59,27 N 84 20,69 W

Kurs = 133 v = 2,5 kn

1447 Ende Profil SL = 442 m LT = 625 m 08 52,97 N 84 13,97 W 09 sm

1510 Depressor a.D.

1537 TOBI a.D.

OFOS-18 W 2

1646 Beginn Station LT = 184 m 08 51,94 N 84 12,60 W

1647 OFOS z.W.

1653 Bosi SL = 178 m LT = 185 m 08 51,88 N 84 12,60 W

Kurs = 216 v = 0,8 kn

1814 ä.K. 227 SL = 397 m LT = 403 m 08 51,00 N 84 13,23 W 1,2

2048 ä.K. 292 SL = 702 m LT = 708 m 08 49,39 N 84 14,92 W 1,8

2134 ä.K. 180 SL = 828 m LT = 848 m 08 49,62 N 84 15,47 W 0,6

2155 hieven ein SL = 815 m LT = 808 m 08 49,39 N 84 15,50 W 0,4

2215 OFOS a.D.

2220 Ende OFOS 18

OFOS-19 W 2

2322 Beginn Station LT = 1036 m 08 54,98 N 84 17,82 W

2323 OFOS z.W.

2344 Bosi SL = 1024 m LT = 1037 m 08 54,94 N 84 17,93 W

Kurs = 318 v = 0,5 kn

14.04.02

0217 hieven ein SL = 1005 m LT = 1010 m 08 56,00 N 84 18,86 W 2

0240 OFOS a.D.

0242 Ende Station

OFOS-20 W 2

0352 Beginn Station LT = 429 m 08 50,40 N 84 12,98 W

0403 OFOS z.W.
 0417 Bosi SL = 436 m LT = 444 m 08 50,30 N 84 12,99 W
 Kurs = 223 v = 0,7 kn
 0600 ä.K. 180 SL = 641 m LT = 651 m 08 49,32 N 84 13,92 W 1
 0707 hieven ein SL = 721 m LT = 774 m 08 48,58 N 84 13,90 W 1
 0730 OFOS a.D. / Ende Station

OFOS-21 W 2

0849 Beginn Station LT = 1004 m 08 55,96 N 84 18,57 W
 0851 OFOS z.W.
 0910 Bosi SL = 1000 m LT = 1005 m 08 55,94 N 84 18,53 W
 Kurs = 219 v = 0,6 kn
 0938 ä.K. 042 SL = 992 m LT = 1013 m 08 55,71 N 84 18,71 W 0,4
 1031 div. Kurse 08 55,90 N 84 18,55 W 1,0
 1249 hieven ein SL = 1015 m LT = 1022 m 08 55,69 N 84 19,00 W 0,6
 1313 OFOS a.D.
 1214 Ende Station

Profil EM-12

1404 Beginn Profil 08 23,64 N 84 23,95 W
 Kurs = 35
 ä.K. 20 09 00,00 N 84 23,00 W
 1437 ä.K. 129 09 02,32 N 84 22,05 W
 1528 Ende Profil 08 58,40 N 84 17,20 W 11 sm

GTV-04 und 05

1600 Beginn Station LT = 1002 m 08 55,81 N 84 18,73 W
 1602 GTV z.W.
 1630 Bosi SL = 996 m LT = 1002 m 08 55,80 N 84 18,74 W
 1705 1. Griff SL = 986 m LT = 980 m 08 55,81 N 84 18,62 W
 1715 2. Griff SL = 990 m LT = 990 m 08 55,81 N 84 18,62 W
 1721 hieven ein
 1746 GTV a.D.
 1748 Ende Station
 1751 GTV z.W.
 1816 Bosi SL = 993 m LT = 1002 m 08 55,80 N 84 18,75 W
 1841 Griff SL = 1001 m LT = 999 m 08 55,80 N 84 18,73 W
 1844 hieven ein
 1922 GTV a.D.
 1930 Ende Station

Profil EM-13

1951 Beginn Profil 08 53,58 N 84 19,33 W
 Kurs = 132 v = 8 kn
 2045 Ende Profil 08 49,02 N 84 16,41 W 7

Station GTV-06 W 2

2112 Beginn Station LT = 399 m 08 51,23 N 84 13,11 W
 2114 GTV z.W.
 2124 Bosi und Griff SL = 391 m LT = 399 m 08 51,22 N 84 13,10 W
 2126 hieven ein
 2150 GTV a.D.
 2213 Ende Station

Profil EM-14

2243 Beginn Profil 08 46,29 N 84 14,65 W
 Kurs = 196 v = 8 kn

2330 ä.K. 304
0039 Ende Profil

08 46,29 N 84 14,65 W
08 51,00 N 84 22,10 W 11

OFOS-22 W 2

0244	Beginn Station	LT = 1242 m	08 57,63 N 84 36,51 W	
0248	OFOS z.W.			
0316	Bosi	SL = 1238 m LT = 1240 m	08 57,51 N 84 36,50 W	
	Kurs = 180	v = 0,8 kn		
0540	ä.K. 235	SL = 1610 m LT = 1636 m	08 55,85 N 84 36,49 W	1,5
0755	ä.K. 180	SL = 2283 m LT = 2339 m	08 55,21 N 84 37,91 W	1,4
0813	hieven ein	SL = 2364 m LT = 2373 m	08 55,00 N 84 37,91 W	0,2
0858	OFOS a.D. / Ende Station			

Profile DTS 23 bis xx W 2

1104		LT = 2234 m	09 01,94 N 84 51,07 W	
1108	Sidescan z.W.			
1113	Depressor z.W.			
	Kurs = 108	v = 3,0 kn		
1200		SL = 1579 m LT = 2187 m	09 01,27 N 84 48,83 W	
1303	Beginn DTS-23	SL = 4000 m LT = 2029 m	09 00,16 N 84 45,42 W	
	Kurs = 108	v = 3,5 kn		
1708	Ende DTS-23	SL = 4063 m LT = 2047 m	08 55,55 N 84 31,64 W	14
1921	Beginn DTS-24	SL = 3344 m LT = 1907 m	08 57,72 N 84 32,40 W	
	Kurs = 288	v = 3,5 kn		
2332	Ende DTS-24	SL = 3851 m LT = 2040 m	09 02,37 n 84 46,43 W	15
16.04.02				
0110	Beginn DTS-25	SL = 3568 m LT = 1833 m	09 00,70 N 84 45,23 W	
	Kurs = 108	v = 3,5 kn		
0542	Ende DTS-25	SL = 4028 m LT = 1852 m	08 55,94 N 84 30,91 W	16
0735	Beginn DTS-26	SL = 3218 m LT = 1848 m	08 58,27 N 84 32,31 W	
	Kurs = 288	v = 3,5 kn		
1155	Ende DTS-26	SL = 3515 m LT = 1977 m	09 02,98 N 84 46,36 W	15
1412	Beginn DTS-27	SL = 3858 m LT = 1797 m	09 01,28 N 84 45,04 W	
	Kurs = 108	v = 3,5 kn		
1842	Ende DTS-27	SL = 4168 m LT = 1793 m	08 56,46 N 84 30,55 W	15
	hieven ein			
2026	Depressor a.D.			
2045	Sidescan a.D.			

OFOS-23

2242	Beginn Station	LT = 1452 m	09 02,19 N 84 37,22 W	
2244	OFOS z.W.			
2312	Bosi	SL = 1436 m LT = 1449 m	09 02,20 N 84 37,24 W	
	Kurs = 180	v = 0,5 kn		
2400		SL = 1423 m LT = 1432 m	09 02,02 N 84 37,20 W	
17.04.02				
0200		SL = 1412 m LT = 1424 m	09 01,97 N 84 37,32 W	
0400		SL = 1397 m LT = 1415 m	09 01,91 N 84 37,21 W	
0628	hieven ein	SL = 1400 m LT = 1403 m	09 01,77 N 84 37,03 W	4
0703	OFOS a.D. / Ende Station			

OFOS-24

0758	Beginn Station	LT = 1416 m	09 01,93 N 84 37,22 W	
0800	OFOS z.W.			
0828	Bosi	SL = 1396 m LT = 1407 m	09 01,95 N 84 37,23 W	
	diverse Kurse ; v = 0,3 kn			

1042 hieven SL = 1400 m LT = 1410 m 09 01,82 N 84 37,06 W 1
 1113 OFOS a.D. / Ende Station

Profil DTS-28 W 2

1208 Sidescan z.W. LT = 1538 m 08 59,78 N 84 31,87 W
 1212 Depressor z.W.
 Kurs = 292 v = 2,5 kn
 1300 SL = 1854 m LT = 1480 m 09 00,53 N 84 33,80 W
 1400 SL = 2191 m LT = 1363 m 09 01,40 N 84 36,19 W
 1506 hieven ein SL = 2209 m LT = 1539 m 09 02,48 N 84 38,72 W 6
 1604 Depressor a.D.
 1620 Sidescan a.D.
 1625 POSIDONIA a.D.

OFOS-25 W 2

1744 OFOS z.W. LT = 774 m 09 11,62 N 84 47,79 W
 1810 Bosi SL = 769 m LT = 774 m 09 11,59 N 84 47,79 W
 Kurs = 264 Grad v = 0,7 kn
 2140 ä.K. 225 09 11,36 N 84 50,05 W
 2302 hieven ein SL = 1018 m LT = 999 m 09 10,73 N 84 50,67 W 4
 2330 OFOS a.D. / Ende Station

OFOS-26 W 2

2355 Beginn Station LT = 953 m 09 08,25 N 84 49,65 W
 2400 OFOS z.W.
18.04.02
 0023 Bosi SL = 946 m LT = 954 m 09 08,22 N 84 49,63 W
 Kurs = 90 v = 0,6 kn
 0114 ä.K. 50 SL = 946 m LT = 931 m 09 08,25 N 84 49,10 W
 0309 hieven ein SL = 812 m LT = 810 m 09 09,17 N 84 48,02 W 2
 0329 OFOS a.D.
 0330 Ende Station

Profil EM-15 v = 5,0 kn

0400 Beginn Profil Kurs = 213 09 06,80 N 84 51,90 W
 0501 ä.K. 120 09 02,50 N 84 54,78 W 5
 0537 ä.K. 034 09 01,01 N 84 52,20 W 3
 0647 ä.K. 038 09 05,61 N 84 48,98 W 7
 0709 Ende Profil 09 06,92 N 84 47,97 W 2

Station GTV-07 W 2

0736 Beginn Station LT = 851 m 09 09,04 N 84 49,23 W
 0739 GTV z.W.
 0800 Bosi SL = 836 m LT = 830 m 09 09,04 N 84 49,18 W
 0907 Griff SL = 828 m LT = 825 m 09 09,02 N 84 49,13 W
 0907 GTV a.D.
 0915 Ende Station

Station GTV-08 W 2

1025 Beginn Station LT = 1410 m 09 01,97 N 84 37,21 W
 1027 GZV z.W.
 1055 Bosi SL = 1399 m LT = 1404 m 09 01,94 N 84 37,23 W
 1106 Griff SL = 1405 m LT = 1408 m 09 01,94 N 84 37,23 W
 1107 hieven
 1143 GTV a.D.

Station GTV-09 W 2

1218	GTV z.W.		LT = 1413 m	09 01,97 N 84 37,24 W
1244	Bosi	SL = 1400 m	LT = 1409 m	09 01,95 N 84 37,24 W
1248	Griff	SL = 1408 m	LT = 1409 m	09 01,95 N 84 37,23 W
1339	GTV a.D.			
1400	Ende Station			

Profil EM-16 v = 8,0 kn

1411	Beginn Profil	Kurs = 288	09 01,45 N 84 38,00 W
1522	ä.K. 061		09 04,52 N 84 46,95 W 10
1550	ä.K. 108		09 06,23 N 84 43,96 W 03
1900	Ende Profil		09 00,00 N 84 25,35 W 19

Profil EM-17 v = 8,0 kn

1926	Beginn Profil EM-17	Kurs = 288	09 01,56 N 84 23,00 W
2145	Ende Profil EM-17		09 07,55 N 84 40,81 W 19
2204	Beginn Profil EM-18		09 09,61 N 84 40,62 W

19.04.02

0032	ä.K. 063		09 03,52 N 84 23,00 W 20
0040	ä.K. 288		09 03,97 N 84 21,48 W 1
0304	Ende Profil EM-18		09 10,37 N 84 40,00 W
	Ende der Auszeichnung		

F.S. "S O N N E"

Reise SO 163/2

Eingesetzte Geräte		Einsätze
OBH/ OBS	Ocean Bottom Hydrophones / Seismograph (Ausgesetzt)	: 196
OBH/OBS	dto (Aufgenommen)	: 177
MAG	Magnetometer	: 2
CTD	CTD Rosette	: 8
CTD/MUC	CTD Rosette mit Multicorer	: 16
32 Ltr Airgun, 250 m Streamer, aktiv, 4 Kanal		
G-Gun		
Prakla Airgun-Array		
DR	Kettensackdredge	: 4

Vermessen wurden :

Mit Magnetometer	223 sm
Mit Seizmik	826 sm

Eingesetzte Winden :

<i>Winde</i>	<i>D/M</i>	<i>Typ</i>	<i>RF-Nr</i>	<i>SO 163/2 Einsatz</i>	<i>Gesamt Einsatz</i>	<i>SO 163/2 S'länge</i>	<i>Gesamt S'länge</i>	<i>Zust.</i>
W 1	18,2	LWL	812001	0000 h	0000 h	000000 m	0000000 m	1
W 2	18,2	LWL	810001	0000	0942	093688	0199774	3
W 4	11,0	NSW	819052	0000	0000	000000	0000000	1
W 5	11,0	NSW	817164	0041	0269	045177	0239630	4
W 6	18,2	DRAKO	814150	0016	2148	013641	1870084	6

(Die W6 wird mit einem neuen Draht in Balboa am 20.05.02 bestückt)

<i>Winde</i>	<i>SO 163/2 gefierte max.Länge</i>	<i>jemals gefierte max.Länge</i>
W 1	0000 m	0000 m
W 2	0000	7202
W 4	0000	0000
W 5	4197	5200
W 6	3000	7900

Geräteverluste :

1 Kettensackdredge (Bundeseigentum)

Abkürzungen im Stationsprotokoll:

z.W.	zu Wasser
a.D.	an Deck
Boko	Bodenkontakt
Bosi	Bodensicht
SL(max.)	(maximale)Seillänge
LT	Lottiefe nach Hydrosweep
W x	eingesetzte Winde
SM	Simrad- Multibeam-Lot
PS	Parasound
XPNDR	Transponder

Zeit : UTC – 06 Stunden**22.04.02**

1310	Beginn Aufzeichnungen mit SM		09-28.85N 84-51.00W
1400	OBS 14 z/W	LT = 1748 m	09-19.99N 84-53.00W
1535	OBH 15 z/W	LT = 115 m	09-14.66N 84-36.05W
1608	OBH 16 z/W	LT = 945 m	09-09.98N 84-38.01W

Teststation 01 Releaser W6

1755	Arbeitskorb mit Releasern z/W	LT = 3317 m	08-53.10N 84-47.93W
1855	Slmax 3000m	LT = 3346 m	08-53.04N 84-48.00W
1941	Hieven		
2045	Arbeitskorb a/D		
2234	OBH 19 z/W	LT = 2205 m	09-00.00N 84-50.05W
2338	OBH 18 z/W	LT = 825 m	09-09.39N 84-49.60W
2340	OBH 17 z/W ^	LT = 825 m	09-09.39N 84-49.60W

23.04.02

0121	OBH 21 z/W	LT = 3603 m	08-58.02N 85-05.02W
0558	OBH 22 z/W	LT = 2019 m	09-38.58N 85-52.35W
0659	OBH 23 z/W	LT = 1932 m	09-40.20N 85-53.45W

Seizmik Profil 01

0729	Airgun z/W		
0742	Streamer ausgesteckt, L 300m		
0804	Beginn Profil; V 3.5Kn; Kurs 145°		09-40.98N 85-53.97W
0858	Ende Profil, 3.2sm		09-38.34N 85-52.15W
0909	Stb Airgun vorgehievt		
0917	Airgun z/W		
0920	Drehen über Stb		

Seizmik Profil 02

0953	Beginn Profil, V 4 Kn; Kurs 325°		09-38.34N 85-52.15W
1043	Ende Profil; 3.2 sm		09-40.99N 85-55.98W
1047	Stb Airgun vorgehievt		
1103	Airgun z/W		
1104	Drehen über Stb		

Seizmik Profil 03

1142	Beginn Profil; V 4Kn; Kurs 145°		09-40.99N 85-55.98W
1238	Ende Profil; 3.2sm		09-38.34N 85-52.15W
1248	Stb Airgun vorgehievt		
1253	Airgun z/W		
1300	Drehen über Stb		

Seizmik Profil 04

1335	Beginn Profil; V 4 Kn; Kurs 325°		09-38.34N 85-52.15W
1420	Ende Profil, 3.2sm		09-40.99N 85-53.98W
1428	Stb Airgun vorgehievt		
1435	Airgun z/W		
1438	Drehen über Stb		

Seizmik Profil 05

1519	Beginn Profil; V 4Kn; Kurs 145°	09-40.99N 85-53.98W
1615	Ende Profil; 3.2sm	09-38.34N 85-52.15W
1625	Stb Airgun vorgehievt	
1630	Airgun z/W	
1632	Drehen über Stb	

Seizmik Profil 06

1705	Beginn Profil; V 4Kn; Kurs 325°	09-38.34N 85-52.15W
1748	Ende Profil; 3.2sm	09-40.99N 85-53.98W
1803	Airgun a/D	
1820	Streamer a/D	

OBH 06

1827	Release Command	
1851	OBH gesichtet; Release Command OBH 03	
1904	OBH a/D LT = 1980 m	09-39.37N 85-53.76W

OBH 03

1917	Weitere Release Commands	
1941	OBH gesichtet	
1950	OBH a/D LT = 2078 m	09-38.13N 85-53.66W

Station CTD 01 W5

2213	Beginn Station LT = 757 m	10-05.45N 85-59.32W
2215	CTD z/W	
2220	CTD a/D (Probleme Bodenmelder)	
2228	CTD z/W	
2248	Boko, Slmax 731m LT = 739 m	10-05.40N 85-59.28W
2330	CTD a/D	
2332	Ende Station	

24.04.02Station CTD/MUC 02 W5

0030	Beginn Station LT = 1357 m	10-00.84N 86-03.84W
0121	MUC z/W	
0135	CTD z/W	
0210	Boko Slmax 1334m LT = 1345 m	10-00.84N 86-03.83W
0318	CTD a/D	
0326	MIC a/D	
0330	Ende Station	

Profil MAG 01

0357	MAG z/W	10-00.57N 86-06.68W
0403	MAG ausgesteckt; L 250m	
0420	Beginn Profil; V 11Kn; Kurs 225°	10-00.00N 86-10.00W
0956-1136	Fahren div. Ausweichkurse (Longline Fischerei)	
1318	Ende Profil; 91 smEinhieven MAG	09-02.56N 87-05.36W
1330	MAG a/D	

Station CTD/MUC 03 W5

1348	Beginn Station LT = 3191 m	09-00.02N 87-03.15W
1350	MUC z/W	
1357	CTD z/W	

FS SONNE

Stationsprotokoll SO 163-2

1525	Boko, SImax 3181m	LT = 3195 m	09-00.05N 87-05.04W
1655	CTD a/D		
1705	MIC a/D		
1707	Ende Station		

Test G-Guns

1717	GI Guns z/W; V 2Kn	09-00.49N 87-04.86W
1737	GI Guns a/D	

Station CTD 04 W5

2110	Beginn Station	LT = 3135 m	09-30.00N 86-35.29W
2112	CTD z/W		
2120	CTD a/D (Techn. Probleme)		
2149	CTD z/W		
2236	SImax 2080m	LT = 3131 m	09-30.00N 86-35.29W
2338	CTD a/D		
2346	Ende Station		

Seizmik Profil 07

0330	G-Gun z/W	09-42.52N 85-55.09W
0407	Beginn Profil; V 4Kn; Kurs 145°	
0455	Ende Profil; 3 sm	09-38.34N 85-52.15W
0505	G-Gun a/D	

Seizmik Profil 08

0510	G-Gun z/W	
0515	Beginn Profil; V 4Kn; Kurs 325°	09-38.34N 85-52.15W
0600	Ende Profil; 3 sm	09-40.98N 85-53.98W
0611	G-Gun a/D	

OBH 02

0615	Release Command	
0636	OBH aufgetaucht	
0644	OBH a/D	09-40.69N 85-53.85W

OBH 23

0627	Release Command	
0645	OBH aufgetaucht	
0654	OBH a/D	09-40.39N 85-53.90W

OBH 09

0647	Release Command	
0710	OBH aufgetaucht	
0715	OBH a/D	09-39.65N 85-53.25W

OBH 05

0702	Release Command	
0730	OBH aufgetaucht	
0733	OBH a/D	09-39.90N 85-52.87W

OBH 07

0715	Release Command	
0742	OBH aufgetaucht	
0747	OBH a/D	09-39.04N 85-53.25W

OBH 10

0725 Release Command
 0755 OBH aufgetaucht
 0802 OBH a/D

09-39.46N 85-53.20W

0728 Schlauchboot z/W (Aufnahme der OBH's)

OBH 08

0733 Release Command
 0810 OBH aufgetaucht
 0815 OBH a/D

09-39.46N 85-52.83W

OBH 04

0747 Release Command
 0817 OBH aufgetaucht
 0831 OBH a/D

09-38.92N 85-53.03W

OBH 22

0800 Release Command
 0828 OBH aufgetaucht
 0841 OBH a/D

09-38.85N 85-53.09W

0845 Schlauchboot z/W

OBH 01

0840 Release Command
 0903 OBH aufgetaucht
 0914 OBH a/D

09-39.48N 85-51.74W

0918 Schlauchboot a/D

1202	OBS 24 z/W	LT = 2771 m	09-15.92N 85-37.40W
1225	OBS 25 z/W	LT = 2769 m	09-16.00N 85-37.31W
1248	OBH 26 z/W	LT = 2756 m	09-16.08N 85-37.25W
1303	OBH 27 z/W	LT = 2756 m	09-16.18N 85-37.17W
1315	OBH 28 z/W	LT = 2739 m	09-16.27N 85-37.08W
1327	OBH 29 z/W	LT = 2703 m	09-16.40N 85-37.00W
1351	OBH 30 z/W	LT = 2665 m	09-16.52N 85-36.91W
1417	OBH 31 z/W	LT = 2658 m	09-16.94N 85-36.56W
1431	OBH 32 z/W	LT = 2630 m	09-17.03N 85-36.50W
1505	OBH 33 z/W	LT = 2263 m	09-20.60N 85-33.68W
1516	OBH 34 z/W	LT = 2249 m	09-20.69N 85-33.61W
1529	OBH 35 z/W	LT = 2247 m	09-20.78N 85-33.55W

Seizmik/Mag Profil 09

1801 Airgun z/W
 1825 Streamer ausgesteckt L 350m
 Beginn Profil; V 4Kn; Kurs 218°

09-38.50N 85-16.30W

26.04.02

1240 MAG ausgesteckt; L 250m
 1430 Ende Profil 84sm

08-33.63N 86-10.44W

1442 Airgun a/D

Seizmik/MAG Profil 10

1456 G-Gun z/W; Stb Seite

1504 G-Gun z/W; Bb Seite

1506 Beginn Profil; V 4 Kn; Kurs 38°

08-33.53N 86-09.60W

27.04.02

0040 MAG a/D

0600 Ende Profil 76 sm

09-34.66N 85-22.64W

0606 Stb G-Gun a/D

0610 Bb G-Gun a/D

0622 Streamer a/D

OBH 35

0748 Release Command

0811 Div. Release Commands

0835 OBH geortet

0847 OBH a/D

09-21.22N 85-33.87W

OBH 34

0749 Release Command

0815 OBH geortet

0832 OBH a/D

09-21.24N 85-34.20W

OBH 33

0749 Release Command

0815 OBH geortet

0827 OBH a/D

09-21.08N 85-34.13W

0822 Schlauchboot z/W

0850 Schlauchboot a/D

0935 OBH 36 z/W

LT = 2613 m

09-17.13N 85-36.42W

Schwimmtest OBH (Neu)

0954 OBH z/W

09-17.09N 85-36.41W

0956 Ende Schwimmtest

1026 OBH z/W

LT = 2612 m

09-17.18N 85-36.33W

Seizmik Profil 11

1048 G-Gun , Bb/Stb-Seeite z/W

Beginn Profil; V 4Kn; Kurs 217°

09-17.30N 85-36.29W

1100 Streamer z/W, L350m

1117 Ende Profil 2 sm

09-15.85N 85-37.40W

Seizmik Profil 12

1131 Stb Airgun vorgehievt

1137 Stb Airgun z/W

1203 Beginn Profil; V 4Kn, Kurs 129°

09-16.77N 85-37.65W

1223 Ende Profil 2 sm

09-15.86N 85-36.54W

1229 Bb Airgun a/D

Seizmik Profil 13

1235	Bb Airgun z/W	
1305	Beginn Profil; V 4Kn; Kurs 309°	09-16.77N 85-35.97W
1322	Ende Profil; 2 sm	09-17.47N 85-36.84W

Seizmik Profil 14

1337	32 Ltr Airgun z/W	
1405	Beginn Profil; V 4Kn; Kurs 218°	09-17.40N 85-36.19W
1435	Ende Profil ; 2 sm	09-15.85N 85-37.40W
1447	Airgun a/D	
1451	G-Gun Stb Seite a/D	
1454	G-Gun Bb Seite a/D	
1504	Streamer a/D	

OBS 24

1510	Release Command	
1548	OBS gesichtet	
1606	OBS a/D	09-16.55N 85-37.48W

OBS 25

1515	Release Command	
1550	OBS gesichtet	
1610	OBS a/D	09-16.55N 85-37.48W

OBS 26

1517	Release Command	
1555	OBS gsichtet	
1618	OBS a/D	09-16.86N 85-37.46W

OBH 27

1518	Release Command	
1556	OBH gesichtet	
1627	OBH a/D	09-16.65N 85-37.70W

OBH 28

1519	Release Command	
1610	OBH gesichtet	
1637	OBH a/D	09-16.77N 85-37.78W

OBH 29

1519	Release Command	
1620	OBH gesichtet	
1700	OBH a/D	09-17,77N 85-37.74W

1600 Schlauchboot z/W

OBH 30

1610	Release Command	
1640	OBH gesichtet	
1722	OBH a/D	09-17.57N 85-37.47W

OBS 31

1624	Release Command		
1650	OBS gesichtet		
1731	OBS a/D		09-18.07N 85-37.09W

OBH 32

1640	Release Command		
1730	OBH gesichtet		
1801	OBH a/D		09-17.86N 85-37.00W

OBH 36

1646	Release Command		
1730	OBH gesichtet		
1801	OBH a/D		09-17.86N 85-37.00W

1700	Schlauchboot a/D		
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1735	Schlauchboot z/W		
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1822	Schlauchboot a/D		
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OBH 37

1750	Release Command		
1844	OBH gesichtet		
1853	OBH a/D		09-17.79N 85-36.65W

Station CTD 05 W5

2328	Beginn Station	LT = 4227 m	09-48.88N 86-16.23W
2329	CTD z/W		

28.04.02

0054	Ausfall Winde		
0115	Winde ok (Sicherung)		
0118	Slmax 4197m	LT = 4228 m	09-48.90N 86-16.23W
0322	CTD a/D		
0325	Ende Station		

Station CTD/MUC 06 W5

0435	Beginn Station	LT = 2403 m	09-56.10N 86-09.00W
0447	MUC z/W		
0500	CTD z/W		
0551	Boko Slmax 2406m	LT = 2408 m	09-56.10N 86-09.00W
0704	CTD a/D		
0715	MUC a/D		
0716	Ende Statiopn		

OBH 17

1515	Release Command		
1523	OBH gesichtet		
1535	OBH a/D		09-09.55N 84-49.51W

1627	OBH 38 z/W	LT = 1745 m	09-14.27N 84-56.65W
1730	OBH 39 z/W	LT = 2495 m	09-05.26N 85-02.56W
1848	OBH 40 z/w	LT = 3420 m	08-53.33N 84-58.43W

2001	OBH 41 z/W	LT = 2614 m	08-56.04N 84-43.43W
2058	OBH 42 z/W	LT = 1485 m	09-05.17N 84-39.09W

Station CTD/MUC 07 W5

2216	Beginn Station	LT = 1888 m	09-06.91N 84-50.87W
2218	MUC z/W		
2225	CTD z/W		
2315	Boko; SImax 1871m	LT = 1906 m	09-06.91N 84-50.86W

29.04.02

0033	CTD a/D
0042	MUC a/D
0043	Ende Station

0428	OBH 43 z/W	LT = 93 m	09-17.09N 84-10.52W
0443	OBH 44 z/W	LT = 105 m	09-16.12N 84-11.22W
0500	OBH 45 z/W	LT = 120 m	09-14.92N 84-12.10W
0517	OBH 46 z/W	LT = 148 m	09-13.52N 84-13.14W
0538	OBH 47 z/W	LT = 205 m	09-11.96N 84-14.26W
0600	OBS 48 z/W	LT = 360 m	09-10.59N 84-15.29W
0618	OBH 49 z/W	LT = 459 m	09-08.99N 84-16.48W
0639	OBS 50 z/W	LT = 480 m	09-08.04N 84-17.15W
0700	OBH 51 z/W	LT = 502 m	09-06.50N 84-18.30W
0720	OBS 52 z/W	LT = 567 m	09-05.05N 84-19.37W
0737	OBH 53 z/W	LT = 666 m	09-03.54N 84-20.50W
0755	OBS 54 z/W	LT = 763 m	09-02.02N 84-21.61W
0824	OBH 55 z/W	LT = 966 m	09-00.04N 84-23.03W
0847	OBH 56 z/W	LT = 1192 m	08-58.14N 84-24.46W
0916	OBH 57 z/W	LT = 1578 m	08-56.01N 84-26.05W
0938	OBH 58 z/W	LT = 1819 m	08-54.52N 84-27.14W
1000	OBH 59 z/W	LT = 1975 m	08-53.04N 84-28.30W

Seizmik Profil 15

1127	G-Gun Bb Seite z/W	
1134	G-Gun Stb Seite z/W	
1134	Streamer z/W L 350m	
1156	Beginn Profil; V 4Kn; Kurs 036°	08-44.56N 84-34.48W
2240	Ende Profil 43sm	09-17.09N 84-10.56W

Seizmik Profil 16

2250	G-Gun Bb Seite vorgehievt	
2255	G-Gun Bb Seite z/W	
2301	G-Gun Stb Seite vorgehievt	
2309	G-Gun Stb Seite z/W	
2323	Beginn Profil; V 4Kn; Kurs 216°	09-16.81N 84-10.74W
2329	32 Ltr. Airgun z/W	
2345	G-Gun Bb Seite eingeholt	

30.04.02

0113	G-Gun Stb Seite a/D	
0928	Ende Profil 40 sm	08-44.26N 84-34.69W
0914	Sreamer a/D	
0938	Airgun a/D	

OBH 59

1027	Release Command	
1052	OBH gesichtet	
1102	OBH a/D	08-52.99N 84-28.17W

OBH 58

1050	Release Command	
1113	OBH gesichtet	
1126	OBH 58 a/D	08-54.59N 84-26.99W

OBS 57

1115	Release Command	
1135	OBH gesichtet	
1155	OBS a/D	08-55.98N 84-25.89W

OBS 56

1145	Release Command	
1210	OBS aufgetaucht	
1224	OBH a/D	08-58.13N 84-24.53W

OBH 55

1224	Release Command	
1242	OBH gesichtet	
1253	OBH a/D	09-00.04N 84-23.08W

OBS 54

1254	Release Command	
1308	OBH gesichtet	
1320	OBS a/D	09-02.05N 84-21.66W

OBH 53

1322	Release Command	
1330	OBH gesichtet	
1344	OBH a/D	09-03.59N 84-20.55W

OBS 52

1350	Release Command	
1357	OBS gesichtet	
1413	OBS 52 a/D	09-05.14N 84-19.39W

OBH 51

1415	Release Command	
1425	OBH gesichtet	
1436	OBH a/D	09-06.60N 84-18.30W

OBS 50

1441	Release Command	
1448	OBS gesichtet	
1500	OBS a/D	09-08.11N 84-17.17W

OBH 49

1502	Release Command	
1510	OBH gesichtet	
1524	OBH a/D	09-08.97N 84-16.48W

Test G-Gun

1530	G-Gun z/W	09-08.97N 84-16.48W
1535	G-Gun a/D	

OBS 48

1542	Release Command	
1550	OBS gesichtet	
1604	OBS a/D	09-10.61N 84-15.29W

OBS 47

1605	Release Command; keine Anzeige, daß steigt; div. Ortungsversuche
1704	Div. Simrad Profile über Pos. OBS

OBH 46

1142	Release Command	
1650	OBH gesichtet	
1704	OBH a/D	09-13.46N 84-13.22W

1. Bergungsversuch OBS 47 W6

1755	Beginn Versuch; DR z/W	LT = 203 m	09-12.03N 84-14.36W
1756-1809	Setzen SL 80m alle 30m ein Eisenbahnstück bis SL 200		
1811	Boko; SL 235m; Beginn Vollkreisfahren um Pos. OBS ; Abstand 0.15sm V 0.8Kn; Fieren 0.5msec		
1910	Erreichen Aussetzpunkt; Vollkreis gefahren; SL 2178m Fahren in NE#L Richtung		
1930	Stoppen Winde SImax 2637 m		09-12.25N 84-14.27W
1931	Beginn Hieven, 0.8m/sec		
1945	Draht aus Seilrolle Schiebebalken gesprungen: SL 2038m		
2024	Draht klariert		
2134	DR a/D		

2. Bergungsversuch OBS 47 W6

2139	Beginn Versuch; DR z/W	LT = 203 m	09-12.02N 84-14.36W
2143-2154	Setzen SL 80m alle 30m ein Eisenbahnstück		
2155	Boko; SL 235m; Beginn Vollkreisfahren um Pos. OBS ; Abstand 0.15sm; V 0.8Kn; Fieren 0.8m/sec		
2217	Verringern Vollkreis um OBS auf 0.05sm		09-11.90N 84-14.21W
2214	Fahren div. Ausweichkurse wegen Fischerei		
2230	Erreichen Aussetzpos.; drehen über Stb auf 135°		
2249	Seillänge 1800m; Stoppen Winde; V 0.8Kn		
2256	Stoppen Schiff; Hieven mit 0.6m/sec		
2354	DR a/D		

01.05.02OBH 45

0017	Release Command	
0020	OBH gesichtet	
0034	OBH a/D	09-14.98N 84-12.16W

OBH 44

0046	Release Command	
0048	OBH gesichtet	
0100	OBH a/D	09-16.13N 84-11.27W

OBH 43

0112	Release Command		
0114	OBH gesichtet		
0125	OBH a/D		09-17.08N 84-10.61W

Station CTD/MUC 08 W5

0450	Beginn Station	LT = 1417 m	09-07.68N 84-50.35W
0523	MUC z/W		
0530	CTD z/W		
0535	XPNDR z/W		
0613	Boko; SImax 1407m	LT = 1439 m	09-07.68N 84-50.33W
0717	XPNDR a/D		
0721	CTD a/D		
0730	MUC a/D		
0731	Ende Station		

Station CTD/MUC 09 W5

0919	Beginn Station	LT = 967 m	09-08.30N 84-49.93W
0922	MUC z/W		
0929	CTD z/W		
0936	XPNDR z/W		
0954	Boko; SImax 948m	LT = 967 m	09-08.30N 84-49.93W
1037	XPNDR a/D		
1048	CTD a/D		
1055	MUC a/D		
1056	Ende Station		

3. Bergungsversuch OBS 47 W6

1405	Beginn Versuch; DR z/W	LT = 203 m	09-11.86N 84-14.22W
1407-1414	Setzen alle SL 20m von DR z/W je ein Eisenbahnstück		
1420	Boko; SL 239m; fahren in W'l Richtung S'l an Pos. OBS vorbei		
1430	Beginn Vollkreis zu fahren V 0.8Kn; Fieren 0.4m/sec; Abstand Pos. OBS 0.11s		09-12.01N 84-14.38W
1453	Verringern Vollkreis auf 0.05sm		
1500	Div. Ausweichen wegen Fischerei		
1525	Vollkreis abgefahren SL 1530m; fahren in Richtung NE		
1535	Seillänge 1807m; Stoppen Fieren; V 0.8KN		
1615	Stoppen Schiff; Beginn Hieven 0.6m/sec;		
1625	OBS geortet; Tension auf 5.5 to; dann Entlastun		
1630	OBS gesichtet		
1651	Drahtende a/D; verloren : 1 Kettensackdredge; 811m Draht; 5 Eisenbahnstücke		

OBS 47

1630	OBS gesichtet		
1710	OBS a/D		09-12.24N 84-13.77W
2048	OBT 60 z/W	LT = 1839 m	08-36.01N 84-35.77W
2056	OBT 61 z/W	LT = 1840 m	08-36.00N 84-35.49W

Station CTD/MUC 10 W5

2306	Beginn Station	LT = 1411 m	09-01.96N 84-57.22W
2351	MUC z/W		
2400	CTD z/W		

02.05.02

0006	XPNDR z/W		
0035	Boko, SImax 1390m	LT = 1406 m	09-01.93N 84-37.24W
0135	XPNDR a/D		
0140	CTD a/D		
0150	MUC a/D		
0226	OBH 62 z/W	LT = 1035 m	09-06.99N 84-33.56W
0244	OBH 63 z/W	LT = 1009 m	09-05.98N 84-31.98W
0301	OBH 64 z/W	LT = 1031 m	09-04.97N 84-30.40W
0318	OBS 65 z/W	LT = 1128 m	09-03.98N 84-28.79W
0335	OBH 66 z/W	LT = 1173 m	09-03.01N 84-27.23W
0353	OBH 67 z/W	LT = 1191 m	09-02.01N 84-25.59W
0410	OBS 68 z/W	LT = 1157 m	09-01.01N 84-24.03W
0427	OBH 69 z/W	LT = 929 m	09-59.95N 84-22.35W
0512	OBS 70 z/W	LT = 980 m	08-58.91N 84-20.67W
0530	OBH 71 z/W	LT = 904 m	08-57.88N 84-19.05W
0547	OBS 72 z/W	LT = 930 m	08-56.85N 84-17.40W
0604	OBH 73 z/W	LT = 944 m	08-55.79N 84-15.65W
0632	OBS 74 z/W	LT = 792 m	08-54.70N 84-13.98W
0648	OBH 75 z/W	LT = 274 m	08-53.64N 84-12.20W

Seizmik Profil 17

0725	G-Gun Bb Seite z/W		
0729	G-Gun Stb Seite z/W		
0738	Streamer z/W, L 350m		
0751	Beginn Profil, V 3.5Kn; Kurs 302°		08-51.83N 84-07.33W
1640	Ende Profil 30sm,		09-08.33N 84-35.72W

Seizmik Profil 18

1648	G-Gun Stb Seite a/D		
1652	G-Gun Bb Seite a/D		
1706	32 Ltr Airgun z/W		
1710	Beginn Profil; V 3.5Kn; Kurs 122°		09-08.65N 84-35.91W

03.05.02

0142	Ende Profil 30sm		
0151	Airgun a/D		
0200	Streamer a/D		

OBH 75

0223	Release Command		
0226	OBH gesichtet		
0239	OBH a/D		08-53.58N 84-12.24W

OBS 74

0248	Release Command		
0304	OBS gesichtet		
0312	OBS a/D		08-54.60N 84-13.95W

OBH 73

0315	Release Command		
0328	OBH gesichtet		

0337	OBH a/D	08-55.77N 84-15.63W
<u>OBS 72</u>		
0339	Release Command	
0351	OBS gesichtet	
0400	OBS a/D	08-56.87N 84-17.29W
<u>OBH 71</u>		
0402	Release Command	
0414	OBH gesichtet	
0426	OBH a/D	08-57.95N 84-18.99W
<u>OBS 70</u>		
0427	Release Command	
0440	OBS gesichtet	
0452	OBS a/D	08-58.95N 84-20.56W
<u>OBH 69</u>		
0452	Release Command	
0503	OBH gesichtet	
0517	OBH a/D	08-59.97N 84-22.20W
<u>OBS 68</u>		
0518	Release Command	
0535	OBH gesichtet	
0544	OBS a/D	09-01.01N 84-23.99W
<u>OBH 67</u>		
0545	Release Command	
0558	OBH gesichtet	
0606	OBH a/D	09-02.09N 84-25.61W
<u>OBH 66</u>		
0606	Release Command	
0621	OBH gesichtet	
0627	OBH a/D	09-03.10N 84-27.23W
<u>OBS 65</u>		
0627	Release Command	
0638	OBH gesichtet	
0647	OBH a/D	09-04.02N 84-28.77W
<u>OBH 64</u>		
0650	Release Command	
0700	OBH gesichtet	
0706	OBH a/D	09-04.99N 84-30.42W
<u>OBH 63</u>		
0708	Release Command	
0719	OBH gesichtet	
0725	OBH a/D	09-06.02N 84-31.96W
<u>OBH 62</u>		
0725	Release Command	
0737	OBH gesichtet	

0744 OBH a/D

09-07.01N 84-33.59W

Station CTD 11 W5

0930	Beginn Station	LT = 1428 m	09-01.87N 84-37.32W
0936	CTD z/W		
0943	SL 40 XPNDR z/W		
1021	Slmax 1408m	LT = 1429 m	09-01.87N 84-37.32W
1033	SL 1354 m, verholen 0.08sm N'l		
1059	Slmax 1409m	LT = 1419 m	09-01.95N 84-37.31W
1111	SL 1350m verholen 0.07sm N'l		
1132	Slmax 1421m	LT = 1433 m	09-02.02n (\$-\$/:"w
1220	XPNDR a/D		
1230	CTD a/D		
1232	Ende Station		

Station CTD 12 W5

1230	Beginn Station	LT = 1428 m	09-02.02N 84-37.32W
1333	CTD z/W		
1339	XPNDR z/W SL 40m		
1415	Slmax 1415m	LT = 1426 m	09-02.02N 84-37.32W
1425	SL 1357m verholen 0.08sm N'l		
1450	Slmax 1410m	LT = 1417 m	09-01.95N 84-37.17W
1510	SL 1301m verholen 0.07sm N'l		
1535	Slmax 1435m	LT = 1441 m	09-02.02N 84-37.17W
1620	XPNDR a/D		
1625	CTD a/D		
1625	Ende Station		

1658	OBH 76 z/W	LT = 1186 m	09-03.69N 84-32.88W
1713	OBH 77 z/W	LT = 1244 m	09-02.89N 84-31.58W
1727	OBH 78 z/W	LT = 1273 m	09-02.08N 84-30.28W
1742	OBH 79 z/W	LT = 1315 m	09-01.25N 84-28.96W
1754	OBH 80 z/W	LT = 1340 m	09-00.46N 84-27.70W
1807	OBS 81 z/W	LT = 1340 m	08-59.74N 84-26.51W
1822	OBH 82 z/W	LT = 1242 m	08-58.93N 84-25.19W
1833	OBS 83 z/W	LT = 1172 m	08-58.11N 84-23.92W
1851	OBS 84 z/W	LT = 1117 m	08-57.31N 84-22.60W
1905	OBS 85 z/W	LT = 1099 m	08-56.50N 84-21.36W
1918	OBS 86 z/W	LT = 1118 m	08-55.72N 84-20.12W
1933	OBH 87 z/W	LT = 1117 m	08-54.91N 84-18.81W
1947	OBH 88 z/W	LT = 1077 m	08-54.11N 84-17.51W
1959	OBH 89 z/W	LT = 1089 m	08-53.56N 84-16.62W
2011	OBH 90 z/W	LT = 958 m	08-53.01N 84-15.77W

Seizmik Profil 19

2026	Streamer z/W L 350m	
2106	Prakla Airgun z/W Bb Seite	
2113	Beginn Profil, V 3.5Kn, Kurs 302°	08-52.58N 84-15.14W

04.05.02

0354 Ende Profil 23 sm

Seizmik Profil 20

0453	Beginn Profil, V 3.5<Kn, Kurs 180°	09-02.50N 84-37.25W
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0509 Ende Profil 1 sm

Seizmik Profil 20B

0600 Beginn Profil, V 3.5Kn, Kurs 360° 09-01.50N 84-37.25W
 0616 Ende Profil 1sm ; Prakla Airgun Array a/D 09-02.57N 84-37.25W

Seizmik Profil 21

0703 32 Ltr Airgun Stb Seite z/W
 0724 Beginn Profil, V 3.5Kn, Kurs 122° 09-04.80N 84-34.62W
 1426 Ende Profil 34sm
 1436 Airgun a/D, drehen über Stb
 1450 G-Gun Bb Seite z/W
 1455 G-Gun Stb Seite z/W

Seizmik Profil 22

1502 Beginn Profil, V 3.5Kn, Kurs 302° 08-51.68N 84-13.63W
 2209 Ende Profil 25sm 09-05.03N 84-35.00W
 2215 G-Gun Bb Seite a/D
 2218 G-Gun Stb Seite a/D, drehen über Stb
 2254 32 Ltr Airgun z/W

Seizmik Profil 23

2255 Beginn Profil, V 3.5Kn; Kurs 122° 09-05.10N 85-35.09W

05.05.02

0100 G-Gun Bb Seite z/W
 0103 G-Gun Stb Seite z/W
 0108 Airgun a/D
 0607 Ende Profil 25sm 08-51.70N 84-13.66W
 0612 G-Gun Bb Seite a/D
 0620 G-Gun Stb Seite a/D
 0629 Streamer a/D

OBH 90

0656 Release Command
 0707 OBH gesichtet
 0712 OBH a/D 08-53.04N 84-15.78W

OBH 88

0707 Release Command
 0750 OBH gesichtet
 0759 OBH a/D 08-54.33N 84-17.37W

OBH 89

0727 Release Command
 0744 OBH gesichtet
 0747 OBH a/D 08-53.61N 84-16.62W

OBH 87

0800 Release Command
 0816 OBH gesichtet
 0827 OBH a/D 08-54.96N 84-18.72W

OBH 86

0801 Release Command

FS SONNE

Stationsprotokoll SO 163-2

0834	OBH gesichtet	
0848	OBH a/D	08-55.86N 84-19.99W
<u>OBH 85</u>		
0845	Release Command	
0901	OBH gesichtet	
0910	OBH a/D	08-56.57N 84-21.35W
<u>OBS 84</u>		
0905	Release Command	
0919	OBH gesichtet	
0937	OBH a/D	08-57.40N 84-22.64W
<u>OBH 83</u>		
0920	Release Command	
0945	OBH gesichtet	
0956	OBS a/D	08-58.21N 84-23.92W
<u>OBS 82</u>		
0948	Release Command	
1007	OBS gesichtet	
1016	OBS a/D	08-58.99N 84-25.20W
<u>OBS 81</u>		
1016	Release Command	
1032	OBS gesichtet	
1037	OBS a/D	08-59.78N 84-26.52W
<u>OBH 80</u>		
1032	Release Command	
1055	OBH gesichtet	
1057	OBH a/D	09-00.51N 84-27.69W
<u>OBH 79</u>		
1055	Release Command	
1106	OBH gesichtet	
1115	OBH a/D	09-01.32N 84-28.97W
<u>OBH 78</u>		
1109	Release Command	
1125	OBH gesichtet	
1131	OBH a/D	09-02.08N 84-30.30W
<u>OBH 77</u>		
1125	Release Command	
1145	OBH geischte	
1151	OBH a/D	09-02.90N 84-31.64W
<u>OBH 76</u>		
1145	Release Command	
1200	OBH gesichtet	
1228	OBH a/D	09-03.66N 84-32.96W

Station CTD 13 W5

1308	Beginn Station	LT = 1419 m	09-01.87N 84-37.25W
1310	CTD z/W		
1318	XPND z/W SL 40m		
1347	Slmax 1403m	LT = 1418 m	09-01.87N 84-37.25W
1400	SL 1328m, verholen 0.08sm N'l		
1423	Slmax 1401m	LT = 1412 m	09-01.95N 84-37.26W
1433	SL 1339m, verholen 0.07sm N'l		
1455	Slmax 1423m	LT = 1431 m	09-02.02N 84-37.25W
1535	XPNDR a/D		
1540	CTD a/D		
1541	Ende Station		

Station CTD/MUC 14 W5

1645	Beginn Station	LT = 1402 m	09-01.95N 84-37.22W
1650	MUC z/W		
1700	CTD z/W; SL 10m über MUC		
1708	XPNDR z/W		
1733	Boko, Slmax 1397m	LT = 1403 m	09-01.95N 84-37.22W
1817	XPNDR a/D		
1829	CTD a/D		
1830	Ende Station		

Test OBS 2002

2117	Beginn Test	LT = 62 m	09-22.11N 84-13.50W
2121	OBS z/W (Fangleine und Transducer)		
2123	Release Command OBS		
2138	OBS a/D		
2149	Ende Test		
2242	OBH 92 z/W	LT = 151 m	09-14.92N 84-15.18W
2257	OBH 93 z/W	LT = 186 m	09-13.72N 84-16.04W
2312	OBH 94 z/W	LT = 374 m	09-12.44N 84-16.99W
2325	OBH 95 z/W	LT = 467 m	09-11.08N 84-17.96W
2341	OBS 96 z/W	LT = 497 m	09-09.74N 84-19.03W
2400	OBS 97 z/W	LT = 554 m	09-07.85N 84-20.34W

06.05.02

0019	OBS 98 z/W	LT = 651 m	09-06.15N 84-21.61W
0038	OBS 99 z/W	LT = 847 m	09-04.27N 84-22.97W
0054	OBS 100 z/W	LT = 1082 m	09-03.01N 84-23.93W
0108	OBH 101 z/W	LT = 1160 m	09-01.77N 84-24.83W
0124	OBH 102 z/W	LT = 1297 m	09-00.34N 84-25.88W
0138	OBH 103 z/W	LT = 1423 m	08-58.95N 84-26.88W
0152	OBH 104 z/W	LT = 1614 m	08-57.66N 84-27.85W
0210	OBH 105 z/W	LT = 1789 m	08-56.04N 84-29.05W
0227	OBH 106 z/W	LT = 1961 m	08-54.31N 84-30.33W

Seizmik Profil 24

0345	Streamer z/W; L 350m		
0355	G-Gun Bb Seite z/W		
0359	G-Gun Stb Seite z/W		
0420	Beginn Profil, V 3.5Kn, Kurs 36°		08-45.56N 84-36.77W
1600	Ende Profil, 41sm; drehen über Stb		

Seizmik Profil 25

1630	32 Ltr Airgun z/W	
1635	Beginn Profil, V 3.5Kn; Kurs 216°	09-18.44N 84-12.45W

07.05.02

0600	Ende Profil 41sm	08-45.56N 84-36.77W
0610	Airgun a/D	
0615	G-Gun Stb Seite a/D	
0620	G-Gun Bb Seite a/D	
0626	Streamer a/D	

OBH 106

0716	Release Command	
0736	OBH gesichtet	
0745	OBH a/D	08-54.28N 84-30.28W

OBH 105

0736	Release Command	
0801	OBH gesichtet	
0812	OBH a7d	08-56.02N 84-28.99E

OBH 104

0801	Release Command	
0827	OBH gesichtet	
0836	OBH a/D	08-57.70N 84-27.85W

OBH 103

0828	Release Command	
0850	OBH gesichtet	
0858	OBH a/D	08-58.94N 84-26.91W

OBH 102

0850	Release Command	
0911	OBH gesichtet	
0919	OBH a/D	09-00.39N 84-25.92W

OBH 101

0910	Release Command	
0930	OBH gesichtet	
0942	OBH a/D	09-01.84N 84-24.84W

OBS 100

0935	Release Command	
0951	OBS gesichtet	
1003	OBS a/D	09-03.11N 84-23.90W

OBS 99

0955	Release Command	
1014	OBS gesichtet	
1023	OBS a/D	09-04.42N 84-22.97W

OBS 98

1030	Release Command	
1041	OBS gesichtet	

1048 OBS a/D 09-06.17N 84-21.55W

OBS 97

1055 Release Command

1106 OBS gesichtet

1113 OBS a/D 09-07.95N 84-20.29W

OBS 96

1122 Release Command

1130 OBS gesichtet

1136 OBS a/D 09-07.75N 84-18.98W

OBH 94

1143 Release Command

1151 OBH aufgetaucht

1159 OBH a/D 09-11.12N 84-17.99W

OBH 95

1206 Release Command

1213 OBH gesichtet

1222 OBH a/D 09-12.42N 84-17.01W

OBH 93

1233 Release Command

1237 OBH gesichtet

1247 OBH a/D 09-13.77N 84-16.07W

OBH 92

1255 Release Command

1258 OBH gesichtet

1309 OBH a/D 09-14.96N 84-15.19W

Station CTD/MUC 15 W5

1550 Beginn Station LT = 420 m 08-47.24N 84-11.60W

1555 MUC z/W

1603 CTD z/W SL 20m

1623 Boko, Slmax 401m LT = 421 m 08-47.23N 84-11.62W

1700 CTD a/D

1707 MUC a/D

1710 Ende Station

Station CTD/MUC 16 W5

1800 Beginn Station LT = 1474 m 08-42.20N 84-16.70W

1835 Muc z/W

1841 CTD z/W SL 20m

1913 Boko, Slmax 1449m LT = 1466 m 08-42.20N 84-16.71W

2008 CTD a/D

2016 MUC a/D

2025 Ende Station

2107 OBH 107 z/W (1000m Seil) LT = 1839 m 08-40.51N 84-18.32W

2114 OBH 108 z/W LT = 1865 m 08-40.31N 84-18.48W

2122 OBH 109 z/W LT = 1011 m 08-40.08N 84-18.61W

2127 OBH 110 z/W LT = 1963 m 08-39.93N 84-18.71W

2133	OBS 111 z/W	LT = 2042 m	08-39.76N 84-18.81W
2141	OBS 112 z/W	LT = 2098 m	08-39.61N 84-18.91W
2146	OBS 113 z/W	LT = 2127 m	08-39.45N 84-19.02W
2152	OBS 114 z/W	LT = 2168 m	08-39.30N 84-19.11W
2205	OBS 115 z/W	LT = 2198 m	08-39.14N 84-19.19W
2211	OBH 116 z/W	LT = 2228 m	08-38.99N 84-19.29W
2217	OBH 117 z/W	LT = 2272 m	08-38.82N 84-19.41W
2222	OBH 118 z/W	LT = 2304 m	08-38.64N 84-19.51W
2226	OBH 119 z/W	LT = 2353 m	08-38.50N 84-19.60W
2232	OBH 120 z/W	LT = 2385 m	08-38.35N 84-19.71W
2301	OBH 121 z/W (1000m Seil)	LT = 2425 m	08-38.21N 84-19.80W

Seizmik Profil 26

2337	Streamer z/W, L 350m		
2400	Beginn Profil, V 2.0Kn, Kurs 32°; Pr. Airgun Array z/W	08-35.27N	84-21.68W

08.05.02

0403	Ende Profil, 9sm	08-42.80N	84-16.92W
0420	Airgun Array a/D		
0440	G-Gun Bb Seite z/W		
0442	G-Gun Stb Seite z/W		

Seizmik Profil 27

0445	Beginn Profil, V 2Kn, Kurs 212°	08-43.27N	84-16.55W
0911	Ende Profil 9 sm	08-35.89N	84-21.24W

Seizmik Profil 28

0950	Beginn Profil, V 2.0Kn, Kurs 32°	08-35.89N	84-21.24W
1335	Ende Profil 9sm,	08-42.80N	84-16.92W

Kalibrierungsprofil 29

1455	Beginn Profil, V 3.0Kn; div. Kurse	08-39.17N	84-18.81W
1505	Ende Profil 5sm	08-39.55N	84-19.24W
1515	G-Gun Bb Seite a/D		
1516	G-Gun Stb Seite a/D		
1525	Streamer a/D		

OBH 107

1650	Release Command, Schlauchboot z/W (Aufnahme OBH/S)		
1700	OBH gesichtet		
1706	OBH a/D	08-40.59N	84-18.19W

OBH 108

1652	Release Command		
1717	OBH gesichtet		
1723	OBH a/D	08-40.35N	84-18.34W

OBH 109

1700	Release Command		
1717	OBH gesichtet		
1730	OBH a/D	08-40.35N	84-18.34W

OBH 110

1706	Release Command		
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1732	OBH gesichtet	
1741	OBH a/D	08-39.98N 84-18.49W
<u>OBH 111</u>		
1718	Release Command	
1759	OBH gesichtet	
1800	OBH a/D	08-39.85N 84-18.62W
<u>OBH 112</u>		
1733	Release Command	
1800	OBH gesichtet	
1818	OBH a/D; Schlauchboot a/D (Riß in Stb Kammer)	08-39.75N 84-18.57W
<u>OBS 113</u>		
1743	Release Command	
1819	OBS gesichtet	
1834	OBS a/D	08-39.51N 84-18.86W
<u>OBS 114</u>		
1750	Release Command	
1823	OBS gesichtet	
1844	OBSa/D	08-39.52N 84-18.71W
<u>OBS 115</u>		
1805	Release Command	
1830	OBS gesichtet	
1857	OBS a/D	08-39.32N 84-18.79W
<u>OBH 116</u>		
1819	Release Command	
1846	OBH gesichtet	
1903	OBH a/D	08-39.14N 84-18.95W
<u>OBH 117</u>		
1855	Release Command	
1924	OBH gesichtet	
1932	OBH a/D	08-38.83N 84-19.27W
<u>OBH 118</u>		
1906	Release Command	
1940	OBH gesichtet	
1944	OBH a/D	08-38.63N 84-19.41W
<u>OBH 119</u>		
1919	Release Command	
1947	OBH gesichtet	
1952	OBH a/D	08-38.48N 84-19.51W
<u>OBH 120</u>		
1933	Release Command	
2007	OBH gesichtet	
2011	OBH a/D	08-38.34N 84-19.58W

OBH 121

1952 Release Command
 2011 OBH aufgetaucht
 2020 OBH a/D

08-38.21N 84-19.65W

Station CTD 17 W5

2057 Beginn Station
 2109 CTD z/W
 2211 Slmax 2810m

LT = 2833 m

08-34.99N 84-24.18W

LT = 2835 m

08-35.00N 84-24.20W

09.05.02

0040 CTD a/D
 0041 Ende Station

Station CTD 18 W5

0210 Beginn Station
 0213 CTD z/W
 0314 Slmax 2497m
 0435 CTD a/D
 0436 Ende Station

LT = 2515 m

08-24.20N 84-35.95W

LT = 2515 m

08-24.20N 84-36.00W

OBT 60

0535 Release Command
 0552 OBT gesichtet
 0602 OBT a/D

08-36.03N 84-35.37W

OBT 61

0545 Release Command
 0602 OBT gesichtet
 0613 OBT a/D

08-35.98N 84-35.41W

OBH 20

0749 Release Command
 0818 OBH gesichtet
 0828 OBH a/D

08-49.92N 84-49.79W

OBH 21

0946 Release Command
 1023 OBH gesichtet
 1035 OBH a/D

08-58.09N 85-05.02W

OBS 19

1223 Release Command
 1308 OBS gesichtet
 1316 OBS a/D

09-19.99N 84-53.12W

OBH 18

1410 Release Command
 1421 OBH gesichtet
 1431 OBH a/D

09-09.40N 84-49.60W

OBH 15

1545 Release Command
 1547 OBH gesichtet
 1555 OBH a/D

09-14.64N 84-36.11W

OBH 16

1620	Release Command		
1630	OBH gesichtet		
1639	OBH a/D		09-10.01N 84-38.10W

OBS 19

1740	Release Command		
1826	OBS gesichtet		
1845	OBS a/D		09-00.06N 84-50.19W

Profilfahrt mit Simrad

1845	Beginn Profil, V 11Kn, Kurs div		09-00.06N 84-50.19W
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10.05.02

0718	Ende Profil 135sm		08-07.87N 84-52.98W
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Station CTD 19 W5

0738	Beginn Station	LT = 2573 m	08-07.98N 84-53.20W
0755	CTD z/W		
0851	Slmax 2564m	LT = 2579 m	08-07.99N 84-53.24W
1010	CTD a/D		
1011	Ende Station		

Station CTD 20 W5

1235	Beginn Station	LT = 2467 m	07-47.96N 85-13.97W
1236	CTD z/W		
1328	Slmax 2470m	LT = 2481 m	07-47.99N 85-13.99W
1450	CTD a/D		
1452	Ende Station		

2109	OBH 122 z/W	LT = 2581 m	08-34.80N 84-17.37W
2131	OBS 123 z/W	LT = 2624 m	08-35.25N 84-18.05W
2146	OBH 124 z/W	LT = 2630 m	08-35.68N 84-18.77W
2156	OBH 125 z/W	LT = 2618 m	08-36.00N 84-19.31W
2202	OBS 126 z/W	LT = 2647 m	08-36.18N 84-19.64W
2212	OBH 127 z/W	LT = 2689 m	08-36.45N 84-20.15W
2221	OBS 128 z/W	LT = 2686 m	08-36.83N 84-20.66W
2232	OBH 129 z/W	LT = 2718 m	08-37.21N 84-21.28W
2242	OBH 130 z/W	LT = 2688 m	08-37.53N 84-21.89W
2253	OBH 131 z/W	LT = 2571 m	08-38.07N 84-21.48W
2304	OBS 132 z/W	LT = 2589 m	08-37.69N 84-20.83W
2312	OBH 133 z/W	LT = 2620 m	08-37.37N 84-20.31W
2325	OBH 134 z/W	LT = 2603 m	08-36.90N 84-19.52W
2335	OBH 135 z/W	LT = 2526 m	08-36.46N 84-18.78W
2350	OBS 136 z/W	LT = 2446 m	08-35.96N 84-17.99W

11.05.02

0003	OBH 137 z/W	LT = 2501 m	08-35.53N 84-17.26W
0014	OBH 138 z/W	LT = 2485 m	08-35.17N 84-16.64W

Seizmik Profil 30

0048	Streamer z/W L350m
0101	G-Gun Bb Seite z/W

0104	G-Gun Stb Seite z/W	
0134	Beginn Profil, V 2.5Kn, Kurs 301°	08-34.26N 84-15.12W
0508	Ende Profil 9sm	

Seizmik Profil 31

0525	Beginn Profil, V 2.5Kn; Kurs 121°	08-38.38N 84-33.22W
0853	Ende Profil 9sm	08-33.70N 84-15.47W
0902	G-Gun Bb Seite a/D	
0907	G-Gun Stb Seite a/D	
0915	Streamer a/D	

OBH 138

0924	Release Command	
0950	OBH gesichtet	
1000	OBH a/D	08-35.18N 84-16.55W

OBH 137

0924	Release Command	
1008	OBH gesichtet	
1017	OBH a/D	08-35.58N 84-17.19W

OBS 136

0932	Release Command	
1024	OBS gesichtet	
1033	OBS a/D	08-36.00N 84-17.71W

OBH 135

1023	Release Command	
1053	OBH gesichtet	
1057	OBH a/D	08-36.42N 84-18.66W

OBH 134

1024	Release Command	
1103	OBH gesichtet	
1111	OBH a/D	08-36.92N 84-19.35W

OBH 133

1036	Release Command	
1116	OBH gesichtet	
1124	OBH a/D	08-37.35N 84-20.02W

OBS 132

1102	Release Command	
1134	OBS gesichtet	
1139	OBS a/D	08-37.67N 84-20.69W

OBH 131

1141	Release Command	
1210	OBH gesichtet	
1225	OBH a/D	08-38.04N 84-21.33W

OBH 130

1145	Release Command	
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1200	OBH gesichtet		
1239	OBH a/D		08-37.53N 84-21.65W
<u>OBH 129</u>			
1145	Release Command		
1246	OBH gesichtet		
1302	OBH a/D		08-37.32N 84-20.87W
<u>OBS 128</u>			
1238	Release Command		
1315	OBS gesichtet		
1324	OBS a/D		08-36.79N 84-20.46W
<u>OBH 127</u>			
1249	Release Command		
1330	OBH gesichtet		
1343	OBH a/D		08-36.50N 84-20.00W
<u>OBS 126</u>			
1315	Release Command		
1358	OBS gesichtet		
1405	OBS a/D		08-36.17N 84-19.52W
<u>OBH 125</u>			
1334	Release Command		
1410	OBH gesichtet		
1424	OBH a/D		08-36.16N 84-19.20W
<u>OBH 124</u>			
1353	Release Command		
1432	OBH gesichtet		
1443	OBH a/D		08-35.82N 84-18.65W
<u>OBS 123</u>			
1430	Release Command		
1520	OBS gesichtet		
1536	OBS a/D		08-35.58N 84-18.12W
<u>OBH 122</u>			
1435	Release Command		
1505	OBH gesichtet		
1515	OBH a/D		08-34.91N 84-17.31W
1848	OBH 139 z/W	LT = 380 m	09-13.02N 84-33.50W
<u>Test OBS 2002</u>			
1918	OBS z7W	LT = 96 m	09-15.72N 84-32.78W
1923	Hydrophon z/W		
1929	Release Command		
1931	OBS aufgetaucht		
1936	Ende Station, OBS a/D		
2026	OBH 140 z/W	LT = 934 m	09-10.00N 84-38.00W
2057	OBH 141 z/W	LT = 352 m	09-13.01N 84-41.02W

Station CTD/MUC 21 W5

2205	Beginn Station	LT = 2252 m	09-06.03N 84-51.40W
2210	MUC z/W		
2217	CTD z/W		
2220	XPNDR z/W SL 50m		
2312	Boko, SImax 2240m	LT = 2256 m	09-06.03N 84-51.40W

12.05.02

0033	XPNDR a/D
0040	CTD a/D
0045	MUC a/D
0046	Ende Station

0843	OBH 142 z/W	LT = 207 m	10-20.98N 86-07.91W
0858	OBH 143 z/W	LT = 302 m	10-20.25N 86-08.68W
0917	OBH 144 z/W	LT = 401 m	10-19.55N 86-09.44W
0933	OBH 145 z/W	LT = 525 m	10-18.76N 86-10.25W
0951	OBH 146 z/W	LT = 657 m	10-17.99N 86-11.01W
1009	OBH 147 z/W	LT = 795 m	10-17.23N 86-11.78W
1022	OBH 148 z/W	LT = 932 m	10-16.46N 86-12.57W
1037	OBH 149 z/W	LT = 1031 m	10-15.69N 86-13.38W
1050	OBH 150 z/W	LT = 1157 m	10-14.92N 86-14.11W
1102	OBH 151 z/W	LT = 1330 m	10-14.17N 86-14.91W
1129	OBS 152 z/W	LT = 1541 m	10-13.42N 86-15.68W
1142	OBS 153 z/W	LT = 1697 m	10-12.66N 86-16.50W
1201	OBS 154 z/W	LT = 1902 m	10-13.43N 86-18.82W
1223	OBS 155 z/W	LT = 1874 m	10-11.89N 86-17.23W
1250	OBS 156 z/W	LT = 1897 m	10-10.35N 86-15.74W
1312	OBS 157 z/W	LT = 2074 m	10-11.13N 86-18.01W
1326	OBS 158 z/W	LT = 2226 m	10-10.34N 86-18.78W
1350	OBS 159 z/W	LT = 2393 m	10-09.59N 86-19.58W
1406	OBH 160 z/W	LT = 2583 m	10-08.84N 86-20.34W
1449	OBH 161 z/W (Ankers. 1000m)	LT = 2731 m	10-08.10N 86-21.14W

Seizmik Profil 32

1800	Streamer z/W L 350m	
1815	G-Gun Bb Seite z/W	
1821	G-Gun Stb Seite z/W	
1832	32 Ltr Airgun Stb Seite z/W	
1835	Beginn Profil, V 4.5Kn, Kurs 045°	09-41.74N 86-48.12W

13.05.02

1015	Ende Profil 70sm	10-31.30N 85-57.25W
1030	Streamer a/D	
1035	G-Gun Bb Seite a/D	
1049	Airgun a/D	
1045	G-Gun Stb Seite a/D	

OBH 142

1200	Release Command	
1215	OBH gesichtet	
1223	OBH a/D	10-21.17N 86-08.06W

OBH 143

1230	Release Command		
1235	OBH gesichtet		
1242	OBH a/D		10-20.39N 86-08.83W

OBH 144

1250	Release Command		
1257	OBH gesichtet		
1304	OBH a/D		10-19.74N 86-09.63W

OBH 145

1310	Release Command		
1316	OBH gesichtet		
1324	OBH a/D		10-18.99N 86-10.55W

OBH 146

1330	Release Command		
1338	OBH gesichtet		
1347	OBH a/D		10-18.23N 86-11.33W

OBH 147

1354	Release Command		
1404	OBH gesichtet		
1415	OBH a/D		10-17.54N 86-12.21W

1505	OBH 162 z/W	LT = 1919 m	10-14.25N 86-19.59W
1530	OBH 163 z/W	LT = 1892 m	10-12.70N 86-18.07W
1555	OBH 164 z/W	LT = 1845 m	10-11.14N 86-16.53W
1616	OBH 165 z/W	LT = 1992 m	10-09.57N 86-14.95W

Seismik Profil 33

1653	Streamer z/W L350m		
1704	G-Gun Bb Seite z/W		
1708	G-Gun Stb Seite		
1721	Beginn Profil, V 4.5Kn, Kurs 315°		10-08.26N 86-13.68W
1941	Ende Profil 10sm		10-15.74N 86-21.10W
1949	Streamer a/D		
1959	G-Gun Bb Seite a/D		
2000	G-Gun Stb Seite a/D		

Station CTD/MUC 22 W5

2043	Beginn Station	LT = 1521 m	10-17.99N 85-18.29W
2047	MUC z/W		
2055	CTD z/W		
2200	Boko, SImax 1497m	LT = 1519 m	10-17.99N 86-18.28W
2246	CTD a/D		
2251	MUC a/D		
2255	Ende Station		

OBH 148

2326	Release Command		
2332	OBH gesichtet		
2346	OBH a/D		10-16.84N 86-12.98W

OBH 149

2346 Release Command
2350 OBH gesichtet

14.05.02

0015 OBH a/D 10-16.08N 86-14.18W

OBH 150

0022 Release Command
0037 OBH gesichtet
0054 OBH a/D

10-15.25N 86-14.61W

OBH 151

0040 Release Command
0102 OBH gesichtet
0117 OBH a/D

10-14.59N 86-15.58W

OBS 152

0103 Release Command
0135 OBS gesichtet
0150 OBS a/D

10-13.75N 86-16.40W

OBS 153

0136 Release Command
0218 OBS gesichtet
0235 OBS a/D

10-13.12N 86-17.40W

OBS 154

0218 Release Command
0245 OBS gesichtet
0330 OBS a/D

10-14.23N 86-20.22W

OBH 162

0300 Release Command
0330 OBH gesichtet
0345 OBH a/D

10-14.69N 86-20.46W

OBH 163

0333 Release Command
0400 OBH gesichtet
0425 OBH a/D

10-13.31N 86-19.13W

OBS 155

0400 Release Command
0425 OBS gesichtet
0450 OBS a/D

10-12.72N 86-18.44W

OBS 157

0435 Release Command
0502 OBS gesichtet
0516 OBS a/D

10-11.61N 86-18.70W

OBS 158

0502 Release Command
0550 OBS gesichtet

0612 OBS a/D 10-11.28N 86-19.82W

OBS 159

0530 Release Command
 0629 OBS gesichtet
 0648 OBS a/D 10-10.63N 86-20.94W

OBH 160

0612 Release Command
 0650 OBH gesichtet
 0713 OBH a/D 10-09.70N 86-21.64W

OBH 161

0653 Release Command
 0723 OBH gesichtet
 0743 OBH a/D 10-08.72N 86-22.33W

OBH 164

0819 Release Command
 0848 OBH gesichtet
 0855 OBH a/D 10-11.54N 86-17.34W

OBS 156

0850 Release Command
 0930 OBS gesichtet
 0945 OBS a/D 10-11.54N 86-17.34W

OBH 165

0933 Release Command
 0959 OBH gesichtet
 1015 OBH a/D 10-11.54N 86-17.34W

1106	OBS 166 z/W	LT = 1689 m	10-17.32N 86-18.76W
1116	OBH 167 z/W	LT = 1571 m	10-17.69N 86-18.50W
1125	OBH 168 z/W	LT = 1517 m	10-17.99N 86-18.27W
1135	OBH 169 z/W	LT = 1605 m	10-18.33N 86-18.00W
1145	OBH 170 z/W	LT = 1538 m	10-18.61N 86-17.80W
1157	OBH 171 z/W	LT = 1481 m	10-18.87N 86-17.65W
1202	OBS 172 z/W	LT = 1480 m	10-18.86N 86-17.60W

Seizmik Profil 34

1230 Streamer z/W
 1239 G-Gun Bb Seite z/W
 1250 Beginn Profil, V 3.0Kn, Kurs 216° 10-20.00N 86-16.79W
 1405 Ende Profil 4sm 10-17.03N 86-19.00W
 1415 Streamer a/D
 1420 G-Gun Bb Seite a/D

OBH 166

1429 Release Command
 1449 OBH gesichtet
 1456 OBH a/D 10-17.55N 86-19.14W

OBH 167

1450	Release Command		
1510	OBH gesichtet		
1518	OBH a/D		10-17.85N 86-18.84W

OBH 168

1511	Release Command		
1532	OBH gesichtet		
1540	OBH a/D		10-18.18N 86-18.64W

OBH 169

1526	Release Command		
1550	OBH gesichtet		
1600	OBH a/D		10-18.53N 86-18.40W

OBH 170

1545	Release Command		
1605	OBH gesichtet		
1614	OBH a/D		10-18.85N 86-18.30W

OBH 171

1600	Release Command		
1618	OBH gesichtet		
1632	OBH a/D		10-19.17N 86-18.30W

OBS 172

1608	Release Command		
1637	OBS gesichtet		
1648	OBS a/D		10-19.19N 86-18.10W

Station DR 01 W6

1720	Beginn Station	LT = 1583 m	10-17.71N 86-18.53W
1725	DR z/W		
1831	Boko; SL 2194m	LT = 1503 m	10-17.87N 86-18.32W
1831	Slmax 2194 m	LT = 1504 m	10-17.94N 86-18.27W
1939	DR a/D		
1940	Ende Station		

Station DR 02

1942	Beginn Station	LT = 1588 m	10-18.00N 86-18.34W
1943	DR z/W		
2005	Boko, SL 1602m	LT = 1593 m	10-17.99N 86-18.53W
2033	Slmax 2203m	LT = 1522 m	10-17.83N 86-18.23W
2136	DR a/D		
2137	Ende Station		

Station CTD/MUC 23 W5

2150	Beginn Station	LT = 1514 m	10-18.02N 86-18.30W
2200	MUC z/W		
2206	CTD z/W		
2242	Boko, Slmax 1633m	LT = 1657 m	10-18.14N 86-18.54W
2335	CTD a/D		
2342	MUC a/D		
2345	Ende Station		

15.05.02

Profil 205 Hydroakustik

0315	Beginn Profil, V 5.0Kn, Kurs 222°		09-41.02N 86-09.64W
0418	Ende Profil 5sm		09-36.96N 86-13.27W
0754	OBH 173 z/W (Ankers.500m)	LT = 2384 m	09-26.95N 85-39.30W
0800	OBH 174 z/W	LT = 2375 m	09-27.01N 85-39.23W
0809	OBH 175 z/W	LT = 2348 m	09-27.08N 85-39.13W
0819	OBH 176 z/W	LT = 2326 m	09-27.17N 85-39.07W
0827	OBH 177 z/W^	LT = 2324 m	09-27.27N 85-39.00W
0836	OBH 178 z/W	LT = 2323 m	09-27.30N 85-38.94W
0848	OBH 179 z/W	LT = 2332 m	09-27.33N 85-39.01W
0854	OBH 180 z/W	LT = 2326 m	09-27.50N 85-38.94W
0902	OBH 181 z/W	LT = 2330 m	09-27.42N 85-38.86W
0908	OBH 182 z/W	LT = 2323 m	09-27.35N 85-38.77W
0915	OBH 183 z/W	LT = 2327 m	09-27.26N 85-38.68W
0926	OBH 184 z/W	LT = 2325 m	09-27.49N 85-38.77W
0950	OBS 185 z/W	LT = 2317 m	09-27.56N 85-38.68W
1011	OBS 186 z/W	LT = 2304 m	09-27.66N 85-38.61W
1042	OBS 187 z/W	LT = 2304 m	09-27.74N 85-38.56W
1048	OBH 188 z/W	LT = 2301 m	09-27.81N 85-38.49W
1055	OBH 189 z/W	LT = 2307 m	09-27.90N 85-38.40W
1100	OBH 190 z/W	LT = 2285 m	09-27.98N 85-38.35W
1107	OBS 191 z/W	LT = 2283 m	09-28.07N 85-38.25W
1113	OBH 192 z/W	LT = 2281 m	09-28.12N 85-38.18W

Seismik Profil 35

1155	Streamer z/W		
1205	G-Gun Bb Seite z/W		
1208	G-Gun Stb Seite z/W		
1240	Beginn Profil, V 4.5Kn, Kurs 132°		09-29.64N 85-41.24W
1408	Ende Profil 7sm		09-25.21N 85-36.45W

Seismik Profil 36

1408	Beginn Profil, V 4.5Kn, Kurs 026°		09-25.21N 85-36.45W
1642	Ende Profil 10sm		09-34.58N 85-30.64W

Seismik Profil 37

1718	Beginn Profil, 4.5Kn, Kurs 223°		09-35.31N 85-31.57W
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16.05.02

0415	Ende Profil 48sm		08-59.72N 86-04.75W
0417	Streamer a/D		
0430	G-Gun Bb Seite a/D		
0438	G-Gun Stb Seite a/D		

OBH 173

0717	Release Command		
0740	OBH gesichtet		
0754	OBH a/D		09-27.09N 85-39.58W

OBH 186

0725	Release Command		
0758	OBH gesichtet		
0804	OBH a/D		09-27.17N 85-39.56W

OBH 187

0735	Release Command	
0813	OBH gesichtet	
0822	OBH a/D	09-27.28N 85-39.55W

OBH 176

0810	Release Command	
0820	OBH gesichtet	
0830	OBH a/D	09-27.40N 85-39.53W

OBS 177

0736	Release Command	
0833	OBS gesichtet	
0841	OBS a/D	09-27.55N 85-39.56W

OBH 178

0824	Release Command	
0903	OBH gesichtet	
0909	OBH a/D	09-27.57N 85-39.31W

OBS 180

0852	Release Command	
0920	OBS gesichtet	
026	OBS a/D	09-27.69N 85-39.16W

OBH 181

0907	Release Command	
0935	OBH gesichtet	
0941	OBH a/D	09-27.54N 85-39.08W

OBH 182

0924	Release Command	
0950	OBH gesichtet	
0958	OBH a/D	09-27.53N 85-39.04W

OBS 183

0928	Release Command	
1001	OBS gesichtet	
1011	OBS a/D	09-27.53N 85-39.08W

OBH 179

0955	Release Command	
1023	OBH gesichtet	
1031	OBH a/D	09-27.80N 85-39.27W

OBS 175

1011	Release Command	
1048	OBS gesichtet	
1103	OBS a/D	09-27.80N 85-39.17W

OBS 185

1034	Release Command	
1111	OBS gesichtet	
1119	OBS a/D	09-27.80N 85-38.83W

OBS 174

1034	Release Command	
1119	OBS gesichtet	
1127	OBS a/D	09-27.89N 85-38.79W

OBS 184

1051	Release Command	
1145	OBS gesichtet	
1200	OBS a/D	09-28.13N 85-39.60W

OBH 188

1119	Release Command	
1142	OBH gesichtet	
1218	OBH a/D	09-28.46N 85-39.33W

OBH 189

1127	Release Command	
1200	OBH gesichtet	
1238	OBH a/D	09-28.71N 85-39.34W

OBH 190

1141	Release Command	
1200	OBH gesichtet	
1255	OBH a/D	09-28.94N 85-39.49W

OBH 191

1245	Release Command	
1325	OBH gesichtet	
1350	OBH a/D	09-28.78N 85-38.66W

OBH 192

1245	Release Command	
1320	OBH gesichtet	
1335	OBH a/D	09-28.26N 85-38.66W

1655	OBH 193 z/W	LT = 1879 m	09-19.99N 85-06.02W
1825	OBH 194 z/W	LT = 567 m	09-28.02N 84-52.40W
1911	OBH 195 z/W	LT = 323 m	09-28.00N 84-45.00W
2002	OBH 196 z/W	LT = 414 m	09-21.53N 84-45.03W
2049	OBS 197 z/W	LT = 1739 m	09-20.02N 84-53.00W
2153	OBH 198 z/W	LT = 289 m	09-14.99N 84-44.01W
2223	OBT 199 z/W	LT = 733 m	09-12.01N 84-46.99W
2251	OBT 200 z/W	LT 823 m	09-09.42N 84-49.60W
2333	OBS 201 z/W	LT = 2079 m	09-10.01N 84-56.01W

17.05.02

0022	OBH 202 z/W	LT = 2230 m	09-02.00N 84-56.01W
0116	OBH 203 z/W	LT = 3535 m	08-57.99N 85-04.99W
0246	OBH 204 z/W	LT = 3472 m	08-50.00N 84-50.00W
0343	OBH 205 z/W	LT = 2207 m	09-00.00N 84-50.00W
0426	OBS 206 z/W	LT = 1727 m	09-02.00N 84-43.00W
0503	OBH 207 z/W	LT = 1168 m	09-07.00N 84-43.00W
0600	OBS 208 z/W	LT = 1155 m	09-05.37N 84-34.24W

0639	OBH 209 z/W	LT = 1613 m	09-02.29N 84-39.18W
0703	OBH 210 z/W	LT = 1207 m	08-59.58N 84-37.91W
0730	OBH 211 z/W	LT = 1412 m	09-01.00N 84-34.46W
0754	OBH 212 z/W	LT = 1327 m	09-02.38N 84-36.20W

Seizmik Profil 38

0838	Streamer z/W, L350m		
0840	G-Gun Bb Seite z/W		
0845	G-Gun Stb Seite z/W		
	Beginn Profil, V 4.5Kn, Kurs 320°		09-00.31N 84-35.34W
0906	32 Ltr Airgun z/W		
1500	Ende Profil 29sm		09-22.49N 84-54.05W

Seizmik Profil 39

1528	Beginn Profil, V 4.5Kn; Kurs 320°		09-21.94N 84-55.51W
2115	Ende Profil 29sm		

Seizmik Profil 40

2138	Beginn Profil V 4.5Kn, Kurs 025°		08-55.02N 84-58.33W
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18.05.02

0600	Ende Profil 40sm		09-31.10N 84-41.55W
0619	Streamer a/D		
0630	Airgun a/D		
0634	G-Gun Bb Seite a/D		
0638	G-Gun Stb Seite a/D		

Station CTD/MUC 24

0925	Beginn Station	LT = 2413 m	09-05.02N 84-51.96W
0930	MUC z/W		
0935	CTD z/W		
0939	XPNDR z/W		
1024	Boko, SImax 2395m	LT = 2412 m	09-05.02N 84-51.94W
1150	XPNDR a/D		
1154	CTD a/D		
1200	MUC a/D; Ende Station		

Profil MAG 04

1530	MAG z/W		
1606	Beginn Profil, V 11.0Kn, Kurs 120°		08-20.00N 84-55.00W

19.05.02

0416	Ende Profil, 132sm		07-12.50N 83-00.00W
0430	MAG a/D		

Appendix 9.8

Press Clippings

Barco científico realiza estudios geológicos y sismológicos en aguas ticas

Valioso equipo previene amenaza sísmica

Estudios ayudan a instituciones como el ICE para diseñar obras de ingeniería

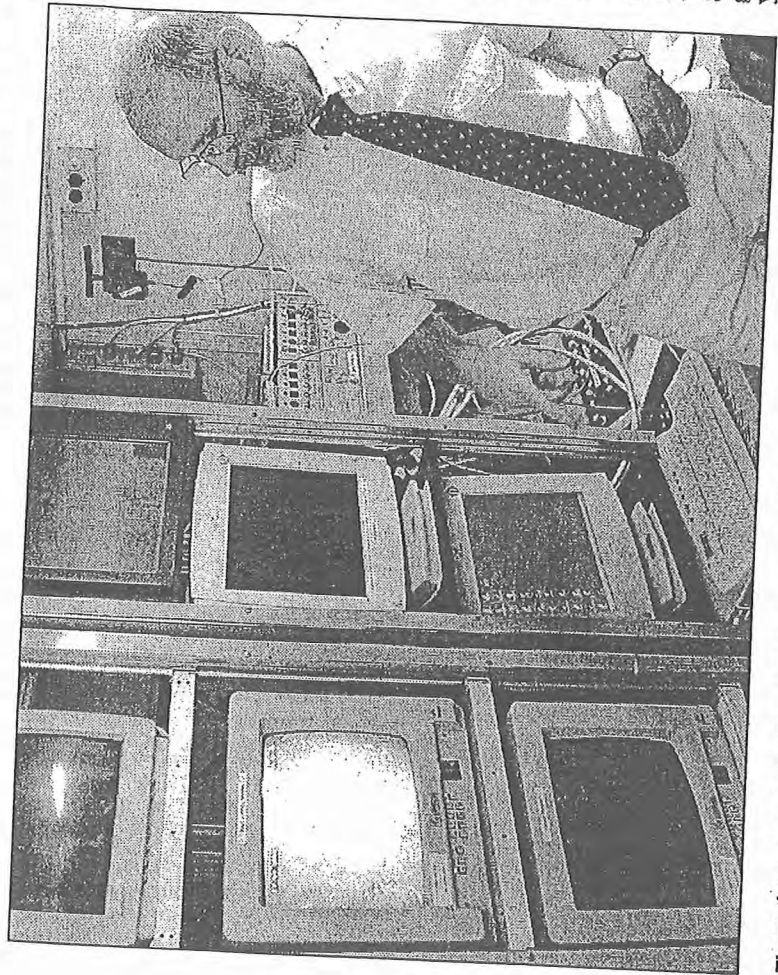
RODRIGO DIAZ LORIA
La República

Un grupo de expertos alemanes y un millonario equipo científico previenen al país de eventuales desastres naturales, al hacer diversos estudios geológicos, sismológicos, vulcanológicos y biológicos en aguas territoriales.

En este momento se encuentra en el país el RV Sonne, una de las dos embarcaciones alemanas que desarrollan una serie de investigaciones en la zona desde 1991. Desde entonces se han efectuado estudios de vulcanología submarina y terrestre, sobre sismología y tectónica, así como peligros por maremotos y biología marina, además se inició la confección de un mapa de relieve combinado.

Para estas investigaciones, el barco laboratorio cuenta con varias grúas, sismómetros o estaciones sismológicas (con un costo de \$700 mil cada una), hidrófonos, una cámara especial para hacer el barrido fotográfico de las profundidades (cuyo costo supera los \$500 mil), un sofisticado equipo de cómputo y demás equipo científico. El barco cuenta con el sistema GPS, de posicionamiento global y un sonar que le permite hasta detectar otras embarcaciones cercanas.

Su tripulación es de 55 personas, entre personal técnico, náutico y científico.



El geofísico alemán Willhem Weinrebe explicó las técnicas que utilizan para estudiar el fondo marino.

co. "Se han reconocido 89 volcanes, hay más de 1.000 toneladas de roca que son estudiadas en Alemania, y hay estudios muy importantes sobre la sismicidad terrestre", precisó el ingeniero Guillermo Alvarado, del Instituto Costarricense de Electricidad (ICE).

En un recorrido por la embarcación, el geofísico alemán Willhem Wein-

rebe, mostró los principales avances que se han hecho y explicó las relaciones que pueden existir en el cambio climático y su relación con una serie de fumarolas que se forman entre las placas. También se estudian la fauna y flora marina que rodean dichos focos volcánicos.

Este año desarrolla un proyecto de investigación de la zona de subducción

o de contacto de las placas Cocos y Caribe, frente a las costas del Pacífico tarricense, y la relación que esto puede tener con los cambios climáticos y desastres naturales. Además se continuará con las labores de mapeo de la periferia bajo las aguas.

El presupuesto de esta inversión es de \$5 millones, financiados por el gobierno de Alemania, con el apoyo de la Universidad Christian Brecht en Kiel y su Instituto de Investigaciones Geomar.

En nuestro país, ICE brinda su apoyo en investigaciones en tierra que hacen los especialistas. Precisamente, la institución es una de las más beneficiadas con el proyecto, debido a la ejecución de sus investigaciones en proyectos geotérmicos e hidroeléctricos que construye.

"Esta información ha sido fundamental para el diseño y todo lo que tiene que ver con seguridad operativa varias obras importantes como Boruca, Savegre, Brujos y Pirris", expresó el ingeniero Salvador López, director de proyectos del ICE.

López reconoció que si no fuera por el apoyo de los alemanes, el país no tendría posibilidad de contar con una información precisa sobre las fuentes sísmicas más importantes.

Para estas investigaciones, el Instituto Geomar también cuenta con otro barco, que hace estudios similares en otras partes del mundo.

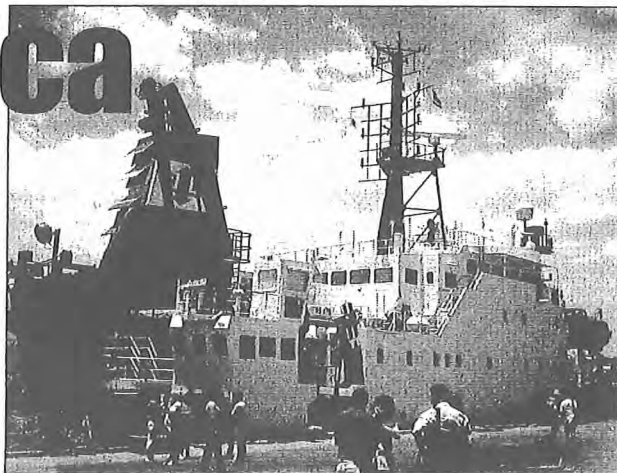
El Sonne permanecerá en nuestro país hasta finales de agosto y posteriormente continuará estudios en Panamá y Suramérica.

Científicos alemanes

Estudian amenaza sísmica

A través de un proyecto conjunto con la Universidad de Kiel, la UCR y el ICE, el barco alemán Sonne estudia causas de los desastres naturales, como los sismos que afectan a toda Centroamérica.

FABIOLA POMAREDA G.
redactora
fpomareda@hotmail.com



El barco Sonne se construyó en 1969 y ha sido reformado paulatinamente a un barco científico.

EN EL "SONNE" no falta el sol, porque el nombre con que bautizaron a este barco significa eso mismo en alemán. Los científicos que viajan en la embarcación, que proviene de la Universidad de Kiel, en Alemania, se encuentran frente a Puerto Caldera, en Puntarenas, realizando investigaciones geológicas, sísmológicas y vulcanológicas.

No es la primera vez que visita el país, pero este año los investigadores iniciaron el proyecto "Volátiles y fluidos en zonas de subducción: retroalimentación del clima y mecanismo disparador de desastres naturales", que desarrolla en conjunto con el Instituto Costarricense de Electricidad (ICE) y la Universidad de Costa Rica (UCR).

El objetivo del proyecto es estudiar las zonas donde se unen las placas Cocos y Caribe, frente a las costas del Pacífico de Costa Rica y Nicaragua, donde circulan fluidos y materias volátiles. Estos procesos están relacionados con el clima en la Tierra, la evolución geoquímica de la hidrosfera y la atmósfera y con las causas de los desastres naturales.

Para emprender esta investigación cuentan con un presupuesto de alrededor de \$5 millones, financiados por el Gobierno de Alemania y el Instituto de Investigaciones Geomar, de la Universidad de Kiel. El acuerdo duraría cinco años y eventualmente puede ser renovado.

TEMA DE PRIORIDAD

En una visita realizada el pasado 20 de abril al barco, los encargados explicaron los aportes del estudio. Participaron Andrés Papenhagen, capitán del Sonne; Guillermo Ruiz, presidente

del Instituto Costarricense de Puertos del Pacífico y el embajador de Alemania, Friedrich Gröning.

También los acompañaron Pablo Cob, presidente del ICE y Guillermo Alvarado, coordinador del Área de Amenazas y Auscultación Sísmovolcánica del ICE y profesor de geología en la UCR.

Como recordó el capitán Papenhagen, desde 1991 el ICE y la UCR participan en estudios similares, que han generado datos relevantes para elaborar estudios geológicos y de amenaza sísmica, para los proyectos de generación eléctrica y la infraestructura del ICE.

El embajador alemán planteó preguntas críticas con respecto a los fines de las investigaciones científicas que se realizan en el ámbito global, por ejemplo: ¿se fijan en la investigación científica las prioridades correctas o nos dejamos llevar por determinadas modas?

"Ciertamente estudiar las amenazas que causan problemas climatológicos o desastres naturales, como en este proyecto, debe ser una prioridad. No es un lujo, son temas que nos conciernen a todos y que son vitales para la supervivencia de muchos seres humanos."

ESCUCHAR A LA TIERRA

En esta etapa de la investigación se han generado datos muy detallados y también se destacó el aporte de la Escuela Centroamericana de Geología, de la UCR.

Guillermo Alvarado dijo que, por ejemplo, hay rocas que se son estudiadas en Alemania para determinar cómo se formaron y cuál es la relación que tienen en la incidencia de los sismos. Agregó que la UCR ha aportado con la

participación de geólogos en los estudios.

El ICE, por su parte, está involucrado por los proyectos hidroeléctricos y geotérmicos que se ubican cerca de volcanes activos. Consiste en instalar 15 estaciones sísmológicas en la tierra, y 25 en el mar, para entender por qué se generan sismos y cómo se pueden prevenir.

"Se trata de determinar cuándo y en qué zona esperar un fuerte movimiento sísmico,".

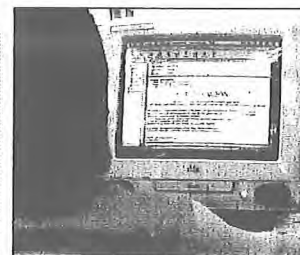
En el Sonne, construido en 1969, viajan 25 científicos y 30 personas que forman la tripulación. Uno de los investigadores, el biólogo Hei-bo Sahling, junto con Germán Leandro, que colabora con el ICE, guiaron el recorrido a través del barco.

Los sismómetros del fondo oceánico -en forma de grandes boyas- son los instrumentos que conforman las estaciones por instalar y son capaces de detectar las fricciones de la tierra. El movimiento del fondo marino se registra en un sensor y luego se transmiten las señales acústicas a computadores en el centro de información.

Además, se cuenta con un sistema "multi-beam", que consiste en cien sonares ubicados en el casco del barco, que barren el fondo oceánico y envían ondas que son interpretadas con un software especial. Este sirve para mostrar la estructura del suelo del mar.

Los investigadores pueden usar una cámara filmadora gigante, que pesa aproximadamente 1 tonelada métrica y costó unos \$200.000, según Sahling. Esta rueda por el fondo oceánico, se maneja desde el barco y permite ubicar organismos vivos, como bivalvos y gusanos que se acumulan entre las placas tectónicas.

Meteor es el otro barco que llegará al país en setiembre, según dijo su capitán, Michael Berkenheger. □



Un pirata informático ("hacker") habría provocado daños en los programas de correo electrónico.

Cariari sufre averías

EDUARDO RAMÍREZ FLORES
redactor
erflores@cariari.ucr.ac.cr

EL SERVIDOR DE CORREO electrónico "Cariari" de la Universidad de Costa Rica (UCR), ha sufrido desde el 18 de abril averías que lo han mantenido prácticamente fuera de servicio y con ello se han visto afectados sus casi 6.000 usuarios.

Problemas originados aparentemente en los programas ("software") que se utilizan para prestar el servicio de correo, impiden desde la fecha mencionada enviar mensajes; no obstante los usuarios sí los reciben.

Según el encargado en el Centro de Informática de administrar el servidor -Bernal Calderón-, al pasado 26 de abril no sabían con precisión la causa del problema, aunque sospechan que podrían ser culpa de un pirata informático ("hacker") que alteró los programas.

Esta es la primera vez -comentó- que un "hacker" habría entrado a la plataforma mediante la cual se opera el correo electrónico de la UCR.

A los obstáculos citados -agregó Calderón-, se suma el saturamiento en el uso de ese servidor, al tener en algunos momentos de 200 a 300 usuarios simultáneos, lo cual bloquea la salida de mensajes hacia servidores externos.

Para corregir los daños, el Centro de Informática contrató a un asesor especializado en el programa Linux y a partir del 29 de abril comenzarían a instalar un paquete con nuevos "software", incluido un cambio de la versión del sistema operativo, lo cual le dará mayor velocidad al servicio.

Calderón aseguró que hacer estos cambios es más práctico y rápido que tratar de detectar lo que provoca la caída del servidor.

Dejó abierta la posibilidad de que puedan restablecer el correo electrónico en los primeros días de la semana; en caso contrario, a más tardar el 2 de mayo lo tendrían listo.

De última hora (martes 28), mediante un correo electrónico a los usuarios, el director del Centro de Informática -Guillermo Lora- avisó que los problemas los resolvieron el viernes 26.

Asimismo, informó que entre el sábado 4 y domingo 5 de mayo interrumpirán periódicamente el servicio, con el fin de trasladar los archivos de cada usuario a un nuevo servidor. □

Pescando Terremotos

GUILLERMO CUBILLOS
EN NUESTRO AGRO

El RV Sonne ancló en Puerto Caldera. Su tripulación, compuesta por científicos alemanes y algunos costarricenses, visita por segunda vez costas centroamericanas, desde Nicaragua hasta Panamá.

Pero no son los atunes, los pargos o los caracoles el motivo de su interés. Son los terremotos a los que pretenden "pescar" antes de que ocurran.

El Gobierno alemán donó \$5 millones para desarrollar esta investigación, en la que participan la Universidad de Costa Rica y el Instituto Costarricense de Electricidad (ICE).

El presidente ejecutivo del ICE, Ing. Pablo Cob, indicó que para esta institución es importante este tipo de investigaciones, pues es un factor de consideración a la hora de desarrollar proyectos como plantas hidroeléctricas y geotermicas.

En general, Centroamérica es propensa a sufrir con regularidad este tipo de fenómenos naturales. La gran actividad volcánica y el constante choque de las placas Cocos y Caribe mantienen vibrando estas tierras a un ritmo más fuerte que la salsa, el merengue o la cumbia.

"Lo cierto es que existe una cordillera volcánica en las profundidades marinas —aseguró el ingeniero geólogo Guillermo Alvarado, de la Universidad de Costa Rica—. Se han registrado más de 100 bajo las aguas del Pacífico costarricense, y de ellos más de 80 se encuentran activos".

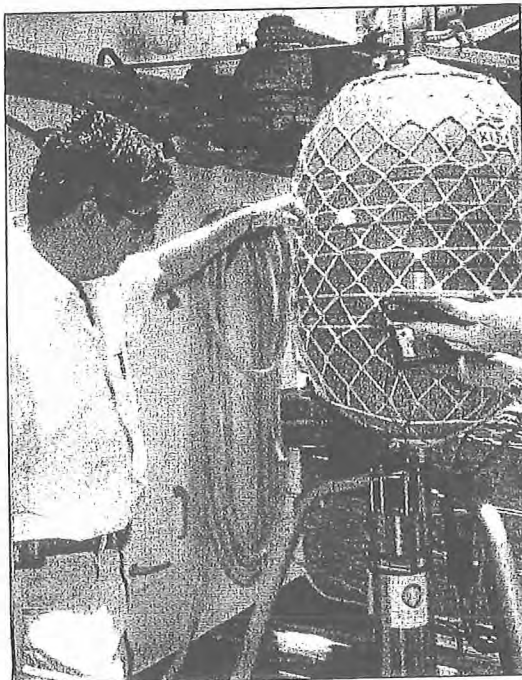
Subducción

La palabra puede parecer familiar, pero no tiene nada que ver con la seducción. Se refiere más bien al movimiento por el que una placa se desliza por debajo de otra, avanzando constantemente.

En nuestro caso, la placa Caribe es subducida por la Cocos. Este movimiento constante y en la mayoría de los casos imperceptible arrastra tras sí todo cuanto en ella se encuentra.



El presidente ejecutivo del INCOPE, don Guillermo Ruiz, formó parte de la comitiva invitada a conocer el Sonne, en el puerto de Caldera. (Levi Vega M.-Nuestro Agro)



El experto Guillermo Alvarado, de la Universidad de Costa Rica, muestra una de las boyas sísmicas que se han distribuido en diversos sectores del Golfo de Nicoya, en el marco de las investigaciones que se realizan en cooperación con el gobierno alemán. (Levi Vega M.-Nuestro Agro)

Pero, ¿qué pasa si también es subducido un volcán? Esta es precisamente una de las interrogantes que los científicos alemanes y costarricenses buscan evacuar.

Y aún más allá. La investigación trata de determinar si estas zonas de subducción pueden ser la causa de los cambios climáticos, la evolución geoquímica de la hidrosfera y la atmósfera y, por supuesto, los desastres naturales.

Laboran dentro de las investigaciones científicas encabezados por los geólogos alemanes, los Drs. Hans-Ulrich Schmincke, Kaj Hoernle, Folmar Hauff, Reinhard Werner, y por el geólogo costarricense, Dr. Guillermo Alvarado de la UCR.

La investigación

El RV Sonne está equipado con el instrumental tecnológico más moderno. Mediante él se extraen muestras del fondo del mar, para lo cual se sumergen a profundidades que superan los 2 mil metros. Toneladas de rocas son extraídas y estudiadas para determinar su edad.

Además, una cámara especialmente diseñada contra la presión hace tomas de suelo marino, registrando así no solo información geológica, sino también biológica.

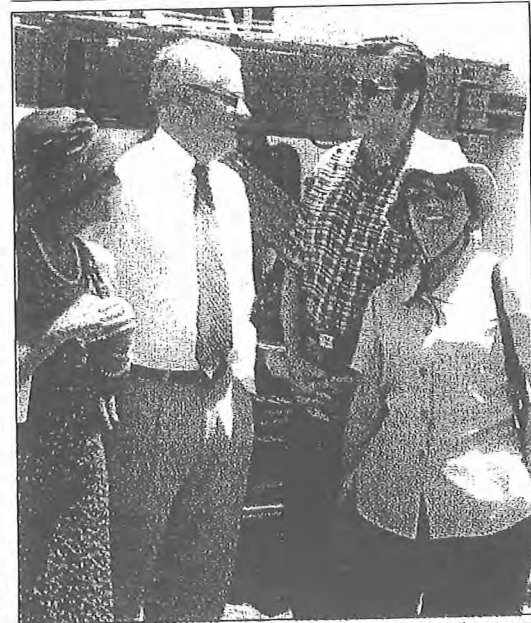
De paso se elaboran planos que permiten un mejor reconocimiento y caracterización de los lechos marinos.

En esta oportunidad la meta es colocar al menos unas 50 estaciones sismológicas, la mitad en tierra y el resto en el mar, que permitan el monitoreo total de cualquier movimiento telúrico.

Consultado el embajador Alemán Friedrich Gröneng sobre el interés de su gobierno en realizar esa investigación en nuestras aguas, indicó que una de las principales razones son las características geológicas de la región, y en segundo lugar, la estabilidad político-social de Costa Rica.



El capitán del Sonne dio a conocer algunas propiedades de la nave y otros detalles sobre su operación y versatilidad. A su lado, la periodista de la embajada alemana, Belina Susin, traductora que ya ha hecho anteriormente la travesía de la investigación. (Levi Vega M.-Nuestro Agro)



El embajador Friedrich Gröneng y el Ing. Pablo Cob Saborío, con sus respectivas esposas, Sra. de Gröneng y Roxana de Cob, recorrieron el barco para conocer las maravillas que en él mismo hay. (Levi Vega M.-Nuestro Agro)

NUESTRO AGRO

Editor:

Levi Vega M.
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Colaborador:

Guillermo Cubillos

Diseño:

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Escaneo:

Fabián Araya y
Wálter Ballester



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